

1 Towards a data-effective calibration of a fully distributed catchment water quality model

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13 **Abstract**

14 Distributed hydrological water quality models are increasingly being used to manage natural
15 resources at the catchment scale but there are no calibration guidelines for selecting the most useful
16 gauging stations. In this study, we investigated the influence of calibration schemes on the
17 spatiotemporal performance of a fully distributed process-based hydrological water quality model
18 (mHM-Nitrate) for discharge and nitrate simulations at Bode catchment in central Germany. We used
19 a single- and two multi-site calibration schemes where the two multi-site schemes varied in number
20 of gauging stations but each subcatchment represented different dominant land uses of the
21 catchment. To extract a set of behavioral parameters for each calibration scheme, we chose a
22 sequential multi-criteria method with 300.000 iterations.

23 For discharge (Q), model performance was similar among the three schemes (NSE varied from 0.88 to
24 0.92). However, for nitrate concentration, the multi-site schemes performed better than the single
25 site scheme. This improvement may be attributed to that multi-site schemes incorporated a broader
26 range of data, including low Q and NO₃⁻ values, thus provided a better representation of within-
27 catchment diversity. Conversely, adding more gauging stations in the multi-site approaches did not
28 lead to further improvements in catchment representation but showed wider 95% uncertainty
29 boundaries. Thus, adding observations that contained similar information on catchment
30 characteristics did not seem to improve model performance and increased uncertainty. These results
31 highlight the importance of strategically selecting gauging stations that reflect the full range of
32 catchment heterogeneity rather than seeking to maximize station number, to optimize parameter
33 calibration.

34 **Keywords**

35 Multi-Multi-site calibration, Spatiotemporal validation, Hydrological water quality model, Uncertainty,
36 Parameter transferability

37 **Highlights:**

- 38 • Single- and multi-site calibration approaches generally led to similar model performance for
39 discharge (Q) at the catchment outlet.
- 40 • Influence of calibration stations on the spatiotemporal performance of a fully distributed
41 process-based hydrological water quality model.
- 42 • The quality of the nitrate model simulation depends less on the number of calibration
43 stations than on their representativeness of the catchment characteristics.

44

45 1. Introduction

46 Distributed hydrological water quality models provide crucial support for water management
47 decisions. The models include many parameters that represent spatial variability in hydrological and
48 biogeochemical processes at the catchment scale that cannot be measured directly in the field (Li et
49 al., 2010). Thus, parameters must be calibrated to optimize model performance (Engel et al., 2007;
50 Moriasi et al., 2012; Saraswat et al., 2015).

51 Most commonly, hydrological water quality models are calibrated using measurements made at the
52 catchment outlet and may thus poorly simulate dynamics at sites within catchments, given spatial
53 variability in conditions (Cao et al., 2006; Refsgaard et al., 2016, Refsgaard et al. 2022). As spatially
54 structured discharge and water quality data become increasingly available, researchers are calling for
55 multi-objective calibration strategies that allow for the inclusion of multiple sites, variables, and
56 criteria (Daggupati et al., 2015; Efstratiadis and Koutsoyiannis, 2010; Khu et al., 2008).

57 However, to date, findings are mixed regarding the performance of single- versus multi-site
58 calibration techniques. Many studies have found that, for catchment outlets, multi-site calibration
59 yields more accurate results than does single-site calibration (e.g., Ghaffar et al., 2021; Her and
60 Chaubey, 2015; Jiang et al., 2015; Zhang et al., 2008). For example, Shrestha et al. (2016) found such
61 to be the case for a SWAT model (Arnold et al., 2012; Arnold et al., 1998) simulating total nitrogen
62 (TN) and total phosphorus (TP) loads. Ghaffar et al. (2021) reported the same for a HYPE model
63 (Lindström et al., 2010) seeking to replicate nitrate (NO_3^-) and TP concentrations across a suite of
64 monitoring stations in central Germany's Selke catchment.

65 In contrast, several other studies have found that performance was largely equivalent for multi-site
66 and single-site calibration techniques (e.g., Franco et al., 2020; Lerat et al., 2012; Wu et al., 2022a).
67 They explained the unimproved model performance with high degree of similarity between flow data
68 used to evaluate the model performance (Lerat et al., 2012), errors in boundary conditions as well as
69 in representations of spatially structured hydrogeological properties (Wang et al., 2012) and
70 hydrological processes (Wu et al., 2022a). However, it is important to note that previous studies have

71 largely utilized semi-distributed hydrological and water quality models (e.g., SWAT: (Leta et al., 2017;
72 Zhang et al., 2008) and HYPE: (Ghaffar et al., 2021; Jiang et al., 2015) and that station choice has
73 frequently been driven by availability. Guidance is lacking when it comes to selecting the most useful
74 gauging stations when calibrating fully distributed hydrological water quality models.

75 Compared to their lumped and semi-distributed counterparts, fully distributed hydrological water
76 quality models incorporate detailed spatial information for sites within catchments while also
77 including a broader range of parameters (Khu et al., 2008; Refsgaard, 1997). The applicability of
78 parameters across spatial and temporal scales (i.e., parameter transferability) presents a major
79 challenge for the construction of distributed hydrological water quality models (Beven, 2001;
80 Samaniego et al., 2010). Parameters defined using information from calibration locations can be
81 applied to other locations using a process called regionalization, as per Blöschl and Sivapalan (1995).
82 Regionalization can be based on spatial proximity (Oudin et al., 2008a; Parajka et al., 2005), similarity
83 in climatic and catchment characteristics (Beck et al., 2016; Merz and Blöschl, 2004; Oudin et al.,
84 2008b; Parajka et al., 2005), and non-linear transfer functions that relate the parameters to
85 catchment characteristics (e.g., land use, soil type, and geological type) (Hundecha and Bárdossy,
86 2004; Pokhrel et al., 2008; Wagener and Wheeler, 2006). Samaniego et al. (2010) specifically
87 developed a multi-scale parameter regionalization (MPR) method, whose appeal stems from the fact
88 that only the coefficients in the transfer functions (i.e., the global parameters) need calibration, and
89 not the parameters for each grid, substantially reducing the dimensionality of the calibrated
90 parameters (Parajka et al., 2013; Singh et al., 2014). When model parameters are tied to catchment
91 characteristics, calibration data drawn from diverse gauging stations are assumed to better represent
92 within-catchment heterogeneity and to enhance model performance at spatial scales. However, little
93 is known about the impact of different calibration schemes on the spatial and temporal performance
94 of fully distributed hydrological water quality models.

95 Hydrological water quality models are typically developed using current knowledge about the physical
96 and chemical processes taking place in the focal catchment, an endeavor that inherently involves

97 simplifications and assumptions (Beven, 2007; Gupta et al., 2005). Uncertainty in model simulations is
98 rooted in uncertainty from the measurement data, used as input and for calibration, as well as from
99 model structure and parameterization (Vrugt et al., 2005; Wagener and Gupta, 2005). Such is
100 especially true for spatially distributed hydrological water quality models, which contain more
101 parameters than those of a lumped or semi-distributed model. While the hydrological modelling
102 community has spent considerable time and effort designing uncertainty analysis techniques, the
103 latter are rarely applied to distributed process-based hydrological water quality models, perhaps due
104 to model complexity (Wellen et al., 2015). In addition, contrasting estimates of model simulation
105 uncertainty have been obtained with single- versus multi-site calibration techniques. Jiang et al.
106 (2015) found that, compared to single-site calibration, multi-site calibration reduced the uncertainty
107 around estimates of Q and NO_3^- concentrations in the HYPE model. In contrast, Her and Chaubey
108 (2015) found the opposite effect for Q estimates from a SWAT model: better performance was
109 obtained using single-site than multi-site calibration. Finally, Shrestha et al. (2016) reported mixed
110 results: for a SWAT model, single-site calibration resulted in less uncertainty for simulated Q values,
111 while multi-site calibration accomplished the same for simulated TN and TP loading values. Thus,
112 there is a pressing need to explore the impact of multi-calibration techniques on the uncertainty
113 associated with fully distributed models.

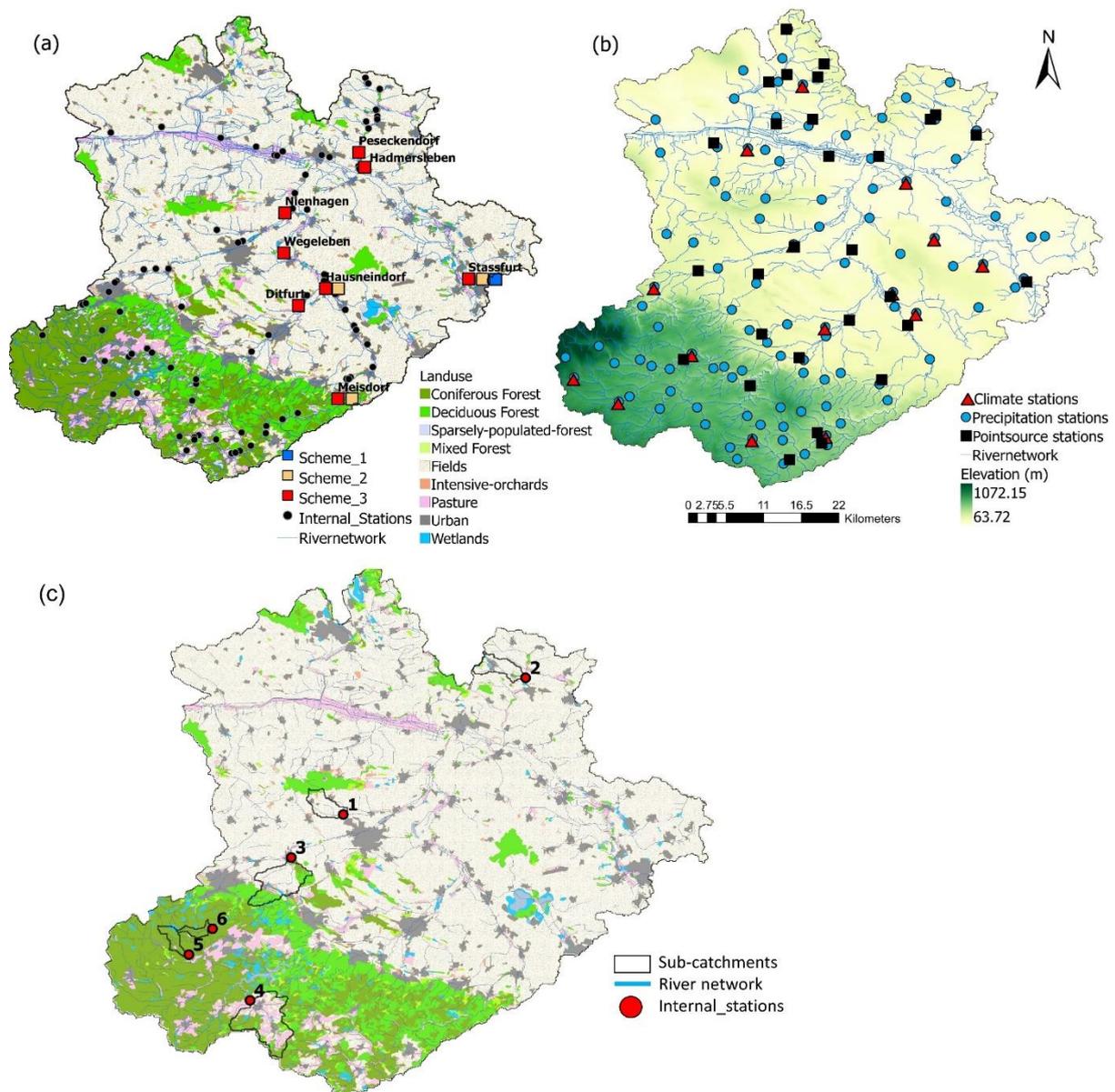
114 Recently, Yang et al. (2018) developed a fully distributed hydrological water quality model (mHM-
115 Nitrate) that is based on both the mesoscale hydrological model (mHM) (Samaniego et al., 2010) and
116 the HYPE model (Lindström et al., 2010). The mHM-Nitrate model appears to successfully handle
117 different catchment characteristics (Wu et al., 2022b; Yang et al., 2019a), but it is unknown how well
118 it deals with parameter transferability across space. Our study's overarching aim was to evaluate the
119 effects of different calibration schemes on the spatiotemporal performance of the mHM-Nitrate
120 model. The specific objectives were as follows: (i) to evaluate and compare three calibration schemes
121 that differed in gauging station number and representation of within-catchment diversity (e.g., land
122 use and stream order); (ii) to assess parameter transferability across space under the three calibration

123 schemes using NO_3^- data from a large number of sampling locations; and (iii) to examine the effects of
124 the three calibration schemes on the degree of uncertainty associated with simulated NO_3^-
125 concentrations. Ideally, the study's results should help guide the choice of effective calibration
126 schemes, depending on the availability of Q and water quality data.

127 2. Study area and methods

128 2.1 Study area

129 The Bode catchment has a area of 3,200 km² and is located in central Germany (Figure 1). It is part of
130 the Harz/Central German Lowland Observatory, within the broader TERENO Earth observation
131 network focused on integrated, multi-scale monitoring and intensive research (Wollschläger et al.,
132 2016). There is dramatic spatial heterogeneity across the catchment, which extends from the Harz
133 Mountains in the southwest to the lowlands of central Germany in the northeast. There is also a
134 marked elevational gradient, ranging from 1,142 m above sea level (a.s.l.) at Brocken, the highest
135 peak in the Harz Mountains, to 70 m a.s.l. in the central lowlands. These extremes are reflected in
136 dramatic differences in mean annual precipitation at these two locations, equal to 1,500 mm and 500
137 mm, respectively (climatic data: 1990–2019). In the mountains, mean monthly temperature ranges
138 from -0.4°C in January to 16.6°C; for the lowlands, these figures are 1.3°C and 18.9°C, respectively. In
139 the mountains, land surfaces are dominated by forests, with some pastures (10%), agricultural fields
140 (8%), and urban areas and lakes (7%). In the lowlands, land surfaces are largely dedicated to
141 cultivating crops (81%), primarily winter wheat, winter barley, rapeseed, and sugar beet. There is
142 much less representation of other land use categories: forests (7%), pastures (3%), and urban areas
143 and small lakes (9%) (Figure 1a). The predominant soil types in the mountains and lowlands are
144 cambisols and chernozems, respectively.



145
146

147 **Figure 1.** Maps of the Bode catchment showing (a) land use, the gauging stations, and the spatially
148 distributed sampling locations as well as (b) elevation and the meteorological stations and (c) location
149 of 6 internal stations presented in section 3.2.

150 We gathered observations of daily precipitation, daily temperature (maximum, mean, and minimum),
151 and potential evapotranspiration to use as model input. These measurements spanned the period
152 between 1993–2019 and were provided by the German Weather Service (DWD); they came from 78
153 rain gauges and 13 climate stations within the study area. To create the meteorological forcing
154 dataset for the model, the daily precipitation and temperature data were spatially interpolated to 1
155 km × 1 km grid data using the External Drift Kriging method. This interpolation approach uses

156 elevation, an external variable, to predict orographic effects on precipitation and temperature
157 (Hundecca and Bárdossy, 2004). The daily potential evapotranspiration values were calculated using
158 the Hargreaves and Samani (1985) method and interpolated at the same scale of spatial resolution.

159 To set up the mHM-Nitrate model, several sources of geographical data were used. Elevation
160 measurements (spatial resolution: 90 m × 90 m) were obtained from the Shuttle Radar Topography
161 Mission (SRTM) (Jarvis, 2008). The digitized geological map and the soil map (scale: 1:1,000,000) were
162 provided by the German Federal Institute for Geosciences and Natural Resources (BGR)
163 (<https://produktcenter.bgr.de>; last accessed 1 June 2020). The land cover data came from CORINE
164 Land Cover 2012, which contains information on land cover/land use in the year 2012
165 (<https://gdz.bkg.bund.de/index.php/default/open-data.html>; last accessed 1 June 2020). These
166 datasets were resampled to generate model input (spatial resolution: 100 m × 100 m).

167 For model calibration and validation, we used measurements of Q and NO_3^- concentrations from
168 eight gauging stations. Daily measurements of Q at these stations were provided by the State Agency
169 for Flood Protection and Water Management of Saxony-Anhalt (LHW) ([http://gldweb.dhi-](http://gldweb.dhi-wasy.com/gld-portal/)
170 [wasy.com/gld-portal/](http://gldweb.dhi-wasy.com/gld-portal/); last accessed 10 April 2020). High-frequency (15 minutes) NO_3^- concentrations
171 for four stations (Meisdorf, Hausneindorf, Hadmersleben, and Stassfurt) between 2010 and 2019
172 were obtained from the Helmholtz Center for Environmental Research—UFZ; we aggregated these
173 high-frequency measurements to daily values. For the other four stations (Ditfurt, Wegeleben,
174 Nienhagen, and Peseckendorf), the NO_3^- data were low-frequency measurements collected every two
175 weeks to every two months from 1994 to 2019 by LHW (<http://gldweb.dhi-wasy.com/gld-portal/>; last
176 accessed 10 April 2020). Finally, we also gathered low-frequency NO_3^- measurements from 94
177 sampling locations to spatially validate the mHM-Nitrate model. The catchment characteristics at
178 these sites are described in the Supplementary Materials (Table S1).

179 **2.2 mHM-Nitrate model**

180 The mHM-Nitrate model takes a grid-based approach and seeks to reliably represent complex
181 processes (Yang et al., 2018). It includes the following hydrological processes: canopy interception,
182 snow accumulation and melt, evapotranspiration, infiltration, soil moisture dynamics, runoff
183 generation, percolation, and flood routing along the river network. The model incorporates nitrate
184 processes described in the HYPE model (Lindström et al., 2010) as well as others: NO_3^- retention in
185 deep groundwater, NO_3^- dynamics associated with spatially distributed crop rotations, and temporally
186 variable point-source inputs of NO_3^- . These processes are fully integrated into hydrological cycling.
187 Major N inputs include wet atmospheric deposition via precipitation, fertilizer and manure
188 application, and plant/crop residues. For each soil layer, four N pools are defined—active solid organic
189 N, inactive solid organic N, dissolved organic N, and dissolved inorganic N, along with soil N processes,
190 namely denitrification, plant/crop uptake, and transformations among the four N pools. In-stream N
191 transformations include denitrification, primary production, and mineralization. A more detailed
192 description of the mHM-Nitrate model can be found in Yang et al. (2018), and the source code can be
193 found in Yang and Rode (2020).

194 **2.3 Model set-up**

195 The mHM-Nitrate model was set up using available hydrometeorological and geographical data for
196 1993–2019 and was run at a daily time step (Table 1). To exclude the effects of a reservoir in the Harz
197 Mountains, we used daily Q and NO_3^- concentrations measured at a downstream gauging station
198 (Thale) as input.

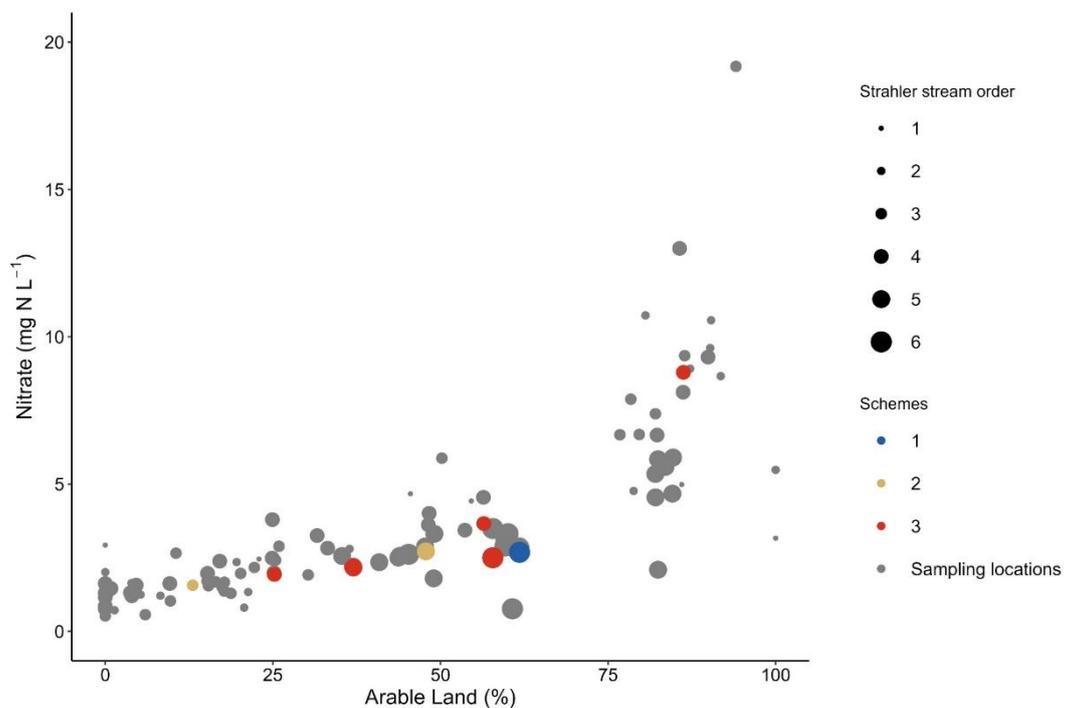
199 **Table 1.** Description of the spatiotemporal data from the Bode catchment used as input for mHM-
200 Nitrate model set-up.

General data type	Specific data type	Resolution	Source
Geographical	Digital elevation model	100 m × 100 m	SRTM
	Land use		CORINE Land Cover 2012
	Geological history		BGR

	Soil type		
Meteorological	Daily precipitation and mean air temperature	1km×1 km	DWD
Agricultural practices	Manure and inorganic fertiliser application, timing and amount of fertilization, sowing and harvesting	Land-use dependent	Field survey and scientific literature
Soil nitrogen content	Initial N storage		Scientific literature
Sewage treatment plants	N load	Daily time step	Operating reports from sewage treatment plants

201 2.4 Calibration schemes

202 The parameters of the mHM-nitrate model were related to catchment characteristics. Based on
 203 catchment characteristics, land use, mean NO_3^- concentration, and stream order, three calibration
 204 schemes were designed. In scheme. Scheme 1 used only data from the catchment outlet station
 205 (Stassfurt). Scheme 2 used data from Stassfurt and two gauging stations upstream (Meisdorf and
 206 Hausneindorf) (Table 1 and Figure 2). Scheme 3 used data from Stassfurt and seven gauging stations
 207 upstream (Figures 1a and 2).



208

209 **Figure 2.** Relationship between nitrate concentration and share of arable land use with information on
 210 stream order of the sub-catchments represented by the eight gauging stations and the 94 spatially
 211 distributed sampling locations. Station inclusion within the calibration schemes is indicated (with
 212 higher-level schemes including the stations found in lower-level schemes).

213 The eight gauging stations used in scheme 3 reflect different combinations of land use and
 214 meteorological conditions found in the Bode catchment (Table 2). Compared to scheme 2, scheme 3
 215 includes data from five additional gauging stations that are associated with larger streams (stream
 216 order: 4–6) (Krabbenhof et al., 2022). There are four main gauging stations along the Bode River:
 217 Ditfurt (upstream), Wegeleben (intermediate stream), Hadmersleben (downstream), and Stassfurt
 218 (catchment outlet). Ditfurt and Wegeleben are in a forest-dominated subcatchment, while
 219 Hadmersleben and Stassfurt locate in an area dominated by farmlands. The headwaters of the Selke
 220 and Holtemme Rivers are located in the mountains, a region with extensive forests (71.9%) and low
 221 NO_3^- concentrations. In contrast, the lowlands are covered by agricultural fields, and NO_3^-
 222 concentrations are high. The Meisdorf station is located in the mountainous Upper Selke, while the
 223 Hausneindorf station is the Selke’s outlet, an area with a mixture of forests and farms. The Nienhagen
 224 station is the Holtemme outlet, whose upstream and downstream areas are dominated by forest and
 225 agricultural surfaces (Ehrhardt et al., 2019), respectively. At Nienhagen, Q values are heavily affected
 226 by the presence of weirs (Kunz et al., 2017). The Peseckendorf station is the outlet of the Geesgraben
 227 stream, which merges into the Bode after Hadmersleben; the surrounding area is predominantly
 228 covered by crops (88.8%).

229 **Table 2.** Subcatchment characteristics for the eight gauging stations. Abbreviations: Subcatch =
 230 subcatchment; Precip = precipitation; Q = discharge; and NO_3^- = nitrate concentration range (mean).

Station	Subcatch	Area (km^2)	Elevation (m)	Precip (mm y^{-1})	% Forest	% Farm land	Stream order	Q (mm y^{-1})	NO_3^- (mg N L^{-1})
Meisdorf	Selke	180	199–597	690	73.1	12.8	3	186	0.01–5.14 (1.57)
Hausneind.	Selke	458	106–597	590	37.8	48.5	5	99	0.44–8.55

									(2.73)
Ditfurt	Bode	714	107–1072	783	56.4	25.3	4	211	1.30–2.90 (1.93)
Wegeleben	Bode	1230	94–1072	698	46.9	36.9	5	166	1.10–4.75 (2.24)
Nienhagen	Holtemme	260	94–931	678	31.6	54.2	4	162	1.22–10.4 (4.59)
Peseckend.	Geesgraben	137	76–200	546	3.0	88.8	4	58	0.77–17.0 (8.80)
Hadmersl.	Bode	2620	76–1072	639	29.2	56.6	6	132	0.47–11.0 (2.51)
Stassfurt	Bode	3179	66–1072	617	24.7	61.6	6	114	0.46–8.10 (2.68)

231 2.5 Model calibration and validation

232 Parameter sensitivity analysis was performed using the Morris method (Morris, 1991). We calculated
 233 the elementary effect (EE) of each parameter using the Sensitivity Analysis For Everybody toolbox
 234 (SAFE; (Pianosi et al., 2015)). We identified the eight most sensitive hydrological parameters and the
 235 six most sensitive water quality parameters (Table S1) based on the ranked values of the sensitivity
 236 indices (absolute mean and standard deviation of EE). This suite of parameters was then used in
 237 mHM-Nitrate model calibration. A more detailed description of the parameter sensitivity analysis is
 238 available in Zhou et al. (2022).

239 Instead of using an optimization algorithm, like a dynamically dimensioned search (DDS) (Tolson and
 240 Shoemaker, 2007), we opted for a sequential multi-criteria method (Wu et al., 2021) to filter out sets
 241 of behavioral parameters for each calibration scheme. This process involved two steps. During the
 242 first step, 300,000 parameter sets were created for the eight sensitive hydrological parameters. Next,
 243 the best 100 parameter sets were selected for each calibration scheme, a decision guided by the
 244 ranks of both the Nash-Sutcliffe coefficient (NSE) and percent bias (PBIAS) values for Q at the relevant
 245 gauging stations. During the second step, 300,000 parameter sets were generated for the six sensitive
 246 water quality parameters, which were combined with the 100 best Q parameter sets. For each
 247 calibration scheme, we selected the best 100 parameter sets from this second step based on the

248 ranks of the NSE and PBIAS values for Q and NO_3^- concentrations for the relevant gauging stations.
 249 The preliminary calibration results revealed that 300,000 iterations allowed the objective function
 250 values to converge upon minimum values. This procedure made it possible to compare the three
 251 calibration schemes, as this allows each calibration scheme to achieve its own best performance from
 252 the same parameter space.

253 Following the split-sample test, this calibration procedure was applied to the mHM-Nitrate model
 254 incorporating Q and NO_3^- concentrations from 2011 to 2014. Each calibration scheme was validated
 255 (time period: 2015–2019) at all eight gauging stations for both Q and NO_3^- concentrations (Table 3).
 256 NSE and PBIAS were used as performance evaluation criteria. However, it is difficult to draw
 257 conclusions about the relative performance of calibration schemes when sample size is small.
 258 Therefore, we carried out spatiotemporal validation of the model using NO_3^- data from the 94
 259 spatially distributed sampling locations (i.e., low-frequency measurements for 1994–2019). In this
 260 case, only PBIAS was used to evaluate model performance, which is satisfactory when values are less
 261 than 35%, according to Moriasi et al. (2015).

262 **Table 3.** Discharge (Q) and nitrate (NO_3^-) concentration data used in model calibration and validation
 263 for the three calibration schemes.

Scheme	Calibration		Validation	
	2011–2014		Q and NO_3^- (2015–2019)	NO_3^- (1994–2019)
1	Q and NO_3^- at Stassfurt			
2	Q and NO_3^- at Stassfurt, Meisdorf, Hausneindorf		Q and NO_3^- at Stassfurt, Meisdorf, Hausneindorf, Nienhagen, Peseckendorf, Ditfurt, Wegeleben, Hadmersleben	NO_3^- at 94 sampling locations
3	Q and NO_3^- at Stassfurt, Meisdorf, Hausneindorf, Nienhagen, Peseckendorf, Ditfurt, Wegeleben, Hadmersleben			

264 2.6 The value of added calibration stations on parameter distributions and model performance

265 To assess the value of additional calibration stations on the identification of the model, the
 266 cumulative parameter distributions were computed for all calibration schemes utilizing the top 100

267 model runs from the second calibration phase of calibration schemes. To the extent that additional
268 calibration stations change the cumulative distribution function of the individual model parameters
269 defined due to model calibration. Significant differences in these cumulative distribution functions
270 can be tested statistically and should allow an assessment of the added value of a modified data set
271 for model identification. In this study, we determined the statistical significance of the differences in
272 these cumulative distribution functions between calibration schemes using the two-sample
273 Kolmogorov-Smirnov (Conover, 1999) test (D):

$$274 \quad D = \max |F(\theta_i) - G(\theta_i)| \quad (1)$$

275 where $F(\theta_i)$ and $G(\theta_i)$ are the empirical cumulative distribution functions of the parameter θ_i for
276 calibration scheme 1(2) and 2(3). The null hypothesis is that the two samples are from the same
277 continuous distribution. If D is closer to zero, it indicates that the probability of the two samples being
278 drawn from the same population is higher. Moreover, the two-sample Kolmogorov-Smirnov test
279 generates a p-value that corresponds to the calculated D statistic. A higher p-value (> 0.05) provides
280 stronger support for the null hypothesis. The relative occurrences of certain, significant, KS statistics
281 can be inspected by means of cumulative frequency plots. As different calibration stations result in
282 varying levels of model parameters, distinct cumulative frequency curves of model performance will
283 be observed.

284 **2.7 Uncertainty analysis**

285 To compare model uncertainty among the three calibration schemes, 95% uncertainty boundaries
286 were calculated based on the 2.5th and 97.5th percentiles of the cumulative distributions for the best
287 100 model runs from the second calibration step. The R-factor quantifies differences between
288 observed and simulated data and is calculated by dividing the average distance between the upper
289 and lower 95% uncertainty boundaries by the standard deviation of the observed data (Abbaspour et
290 al., 2007). The R-factor expresses the width of the 95% uncertainty and a value less than 1 is being
291 desirable. The uncertainty analysis was performed for both Q and NO_3^- concentrations at all the

292 gauging stations included in schemes 2 and 3. We compared model uncertainty for schemes 2 and 3
293 by comparing results for the stations shared by the schemes (Stassfurt, Hausneindorf, and Meisdorf).

294 3. Results

295 The mHM-Nitrate model was calibrated using the three schemes, resulting in different patterns of
296 performance (parameter description: Table S1).

297 3.1 Model performance at gauging stations

298 The model performance of discharge (Q) for at the catchment outlet (Stassfurt station) was similar
299 across the three calibration schemes (NSE—scheme 1: 0.82, scheme 2: 0.87, and scheme 3: 0.88;
300 PBIAS—scheme 1: 0.30%, scheme 2: 0.0%, and scheme 3: -8.60%; Table 4). During the calibration
301 period, at the Meisdorf and Hausneindorf stations, performance was lower for scheme 3 than for
302 scheme 2 (NSE—scheme 2: 0.58 to 0.69 vs. scheme 3: 0.53 to 0.66; PBIAS—scheme 2: -7.80% to -
303 23.5% vs. scheme 3: -20.2% to -32.0%). During the validation period, water balance was well captured
304 across all the calibration schemes and gauging stations, with the exception of Nienhagen (PBIAS—
305 scheme 1: -3.7% to 7.1%, scheme 2: -7.7% to 2.6%, and scheme 3: -12.7% to 1.4%). Performance was
306 lowest at the Peseckendorf and Nienhagen stations across the three schemes, albeit lower for
307 scheme 1 than for schemes 2 and 3 (NSE—scheme 1: -0.34 to 0.13 vs. scheme 2: 0.17 to 0.29 and
308 scheme 3: 0.36 to 0.45; Table 4). It was also better at the Stassfurt, Meisdorf, and Hausneindorf
309 stations during the validation period than during the calibration period across all calibration schemes
310 (NSE—lower ranges: 0.53–0.88 and upper ranges: 0.71–0.92). The mean absolute PBIAS values for Q
311 at all validation stations were 8.4%, 7.5%, and 9.2% for scheme1, scheme2, and scheme3,
312 respectively.

313 Model performance of NO_3^- concentration at the catchment outlet Stassfurt station decreased from
314 Scheme 1 to 2 and 3 during the calibration period (NSE—scheme 1: 0.67, scheme 2: 0.64, and
315 scheme 3: 0.62; PBIAS—scheme 1: 0.40%, scheme 2: -6.90%, and scheme 3: 7.10%). Also during the
316 calibration period, model performance at the Meisdorf station was better at scheme 2 (PBIAS: -

317 2.60%) than scheme 3 (PBIAS: -23.2%). At Hausneindorf, scheme 3 yielded better performance than
318 did scheme 2 (PBIAS: -7.90% vs. 1.20%, respectively). During the validation period, performance was
319 better at scheme 2 than at scheme 1 for all the gauging stations except for Nienhagen station, with
320 PBIAS values in ranges —scheme 2: 1.8–33.9% and scheme 1: -10.1–23.3%, respectively. While NO_3^-
321 concentration model performance decreased from Scheme 2 to 3 at all gauging stations except
322 Nienhagen station, with larger absolute PBAIS values in Scheme 3 than Scheme 2. The mean absolute
323 PBIAS values for NO_3^- were 15.7%, 9.5%, and 13.8% for scheme1, scheme2, and scheme3,
324 respectively. These findings provide evidence that scheme 2 is the most promising option.
325 Additionally, the results indicate that the model performance is categorized as good for Q and very
326 good for NO_3^- .

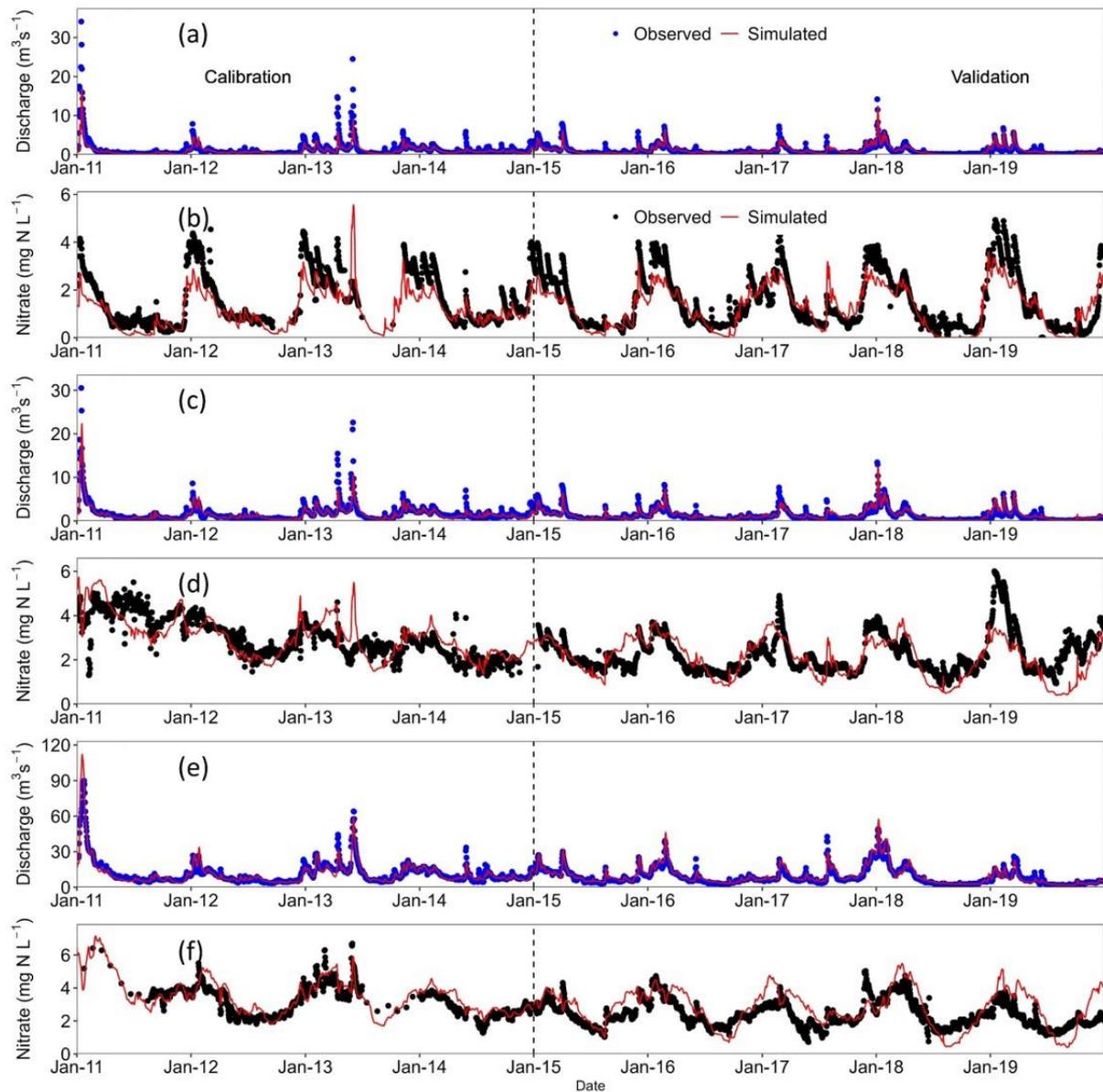
327 **Table 4.** Model performance for discharge (Q) and nitrate (NO_3^-) concentrations during the calibration
328 and validation periods across the three calibration schemes and their associated gauging stations.

Schemes	Stations	Q				NO_3^-			
		Calibration		Validation		Calibration		Validation	
		NSE	PBIAS (%)	NSE	PBIAS (%)	NSE	PBIAS (%)	NSE	PBIAS (%)
1	Stassfurt	0.82	0.30	0.92	4.20	0.67	0.40	0.33	12.5
	Meisdorf	-	-	0.71	7.10	-	-	0.32	33.9
	Hausneindorf	-	-	0.77	0.70	-	-	-0.08	8.10
	Wegeleben	-	-	0.92	-3.70	-	-	-1.19	16.4
	Hadmersleben	-	-	0.93	2.90	-	-	0.01	20.9
	Peseckendorf	-	-	-0.34	-3.50	-	-	-3.84	23.0
	Ditfurt	-	-	0.97	-0.20	-	-	-4.36	8.80
	Nienhagen	-	-	0.13	44.7	-	-	-0.66	1.80
2	Stassfurt	0.87	0.00	0.88	0.60	0.64	-6.90	0.23	9.30
	Meisdorf	0.58	-23.5	0.72	-1.70	0.66	-2.60	0.67	-10.1
	Hausneindorf	0.69	-7.80	0.76	-3.00	0.27	-7.90	0.31	-4.00
	Wegeleben	-	-	0.92	-3.10	-	-	-0.14	4.00
	Hadmersleben	-	-	0.92	2.60	-	-	0.26	14.3
	Peseckendorf	-	-	0.17	-7.70	-	-	-2.72	23.3
	Ditfurt	-	-	0.96	2.20	-	-	-2.18	1.60
	Nienhagen	-	-	0.29	38.8	-	-	-0.17	-9.10
3	Stassfurt	0.88	-8.60	0.90	1.40	0.62	7.10	-0.33	16.8
	Meisdorf	0.53	-32.0	0.71	-12.0	0.53	-23.2	0.71	-14.0
	Hausneindorf	0.66	-20.2	0.73	-12.7	0.31	1.20	0.20	-7.80

Wegeleben	0.87	-12.6	0.92	-5.60	0.37	-9.50	-1.39	11.0
Hadmersleben	0.87	-9.10	0.92	-0.90	0.21	14.0	-0.49	23.8
Peseckendorf	0.56	-21.6	0.45	-9.80	-0.44	-15.6	-1.70	24.7
Ditfurt	0.94	-3.40	0.96	1.00	0.35	-9.80	-3.56	9.20
Nienhagen	0.68	6.00	0.36	29.9	0.59	-14.2	0.39	-3.20

329

330 The seasonal dynamics of Q were captured by scheme 2 at its three gauging stations during both the
331 calibration and validation periods as well as during low- and high-flow conditions (Figures 3a, 3c, and
332 3e). The same was true for the seasonal dynamics of NO_3^- concentrations (i.e., high values during
333 high-flow periods and low values during low-flow periods; Figures 3b, 3d, and 3f). In addition, over the
334 period from 2011 to 2019, NO_3^- concentrations followed a constant seasonal pattern at the Meisdorf
335 station (Figure 3b) but tended to decline at the Hausneindorf and Stassfurt stations (Figures 3d and
336 3f), which were well captured by the model. Model performance for NO_3^- concentrations was
337 greatest at the Meisdorf station (NSE—calibration: 0.66 and validation: 0.67; Table 4). It was lowest at
338 the Hausneindorf station (NSE—calibration: 0.27 and validation: 0.31; Table 4). At Stassfurt, Meisdorf,
339 and Hausneindorf, model performance for NO_3^- concentrations were satisfactory (PBIAS ranged
340 between -7.9% and 9.3% during calibration and validation).



341

342 **Figure 3.** Observed and simulated Q and NO_3^- concentration at (a-b) Meisdorf, (c-d) Hausneindorf and
 343 (e-f) Stassfurt stations for calibration Scheme 2.

344

345 3.2 Model performance at spatially distributed sampling locations

346 We further tested how the calibration schemes affected model performance using NO_3^- data from the

347 94 spatially distributed sampling locations. Performance was generally better for scheme 2 than for

348 scheme 1 (PBIAS $\leq 15.0\%$: 34 vs. 9 sampling stations, respectively, and PBIAS $> 45\%$: 12 vs. 65

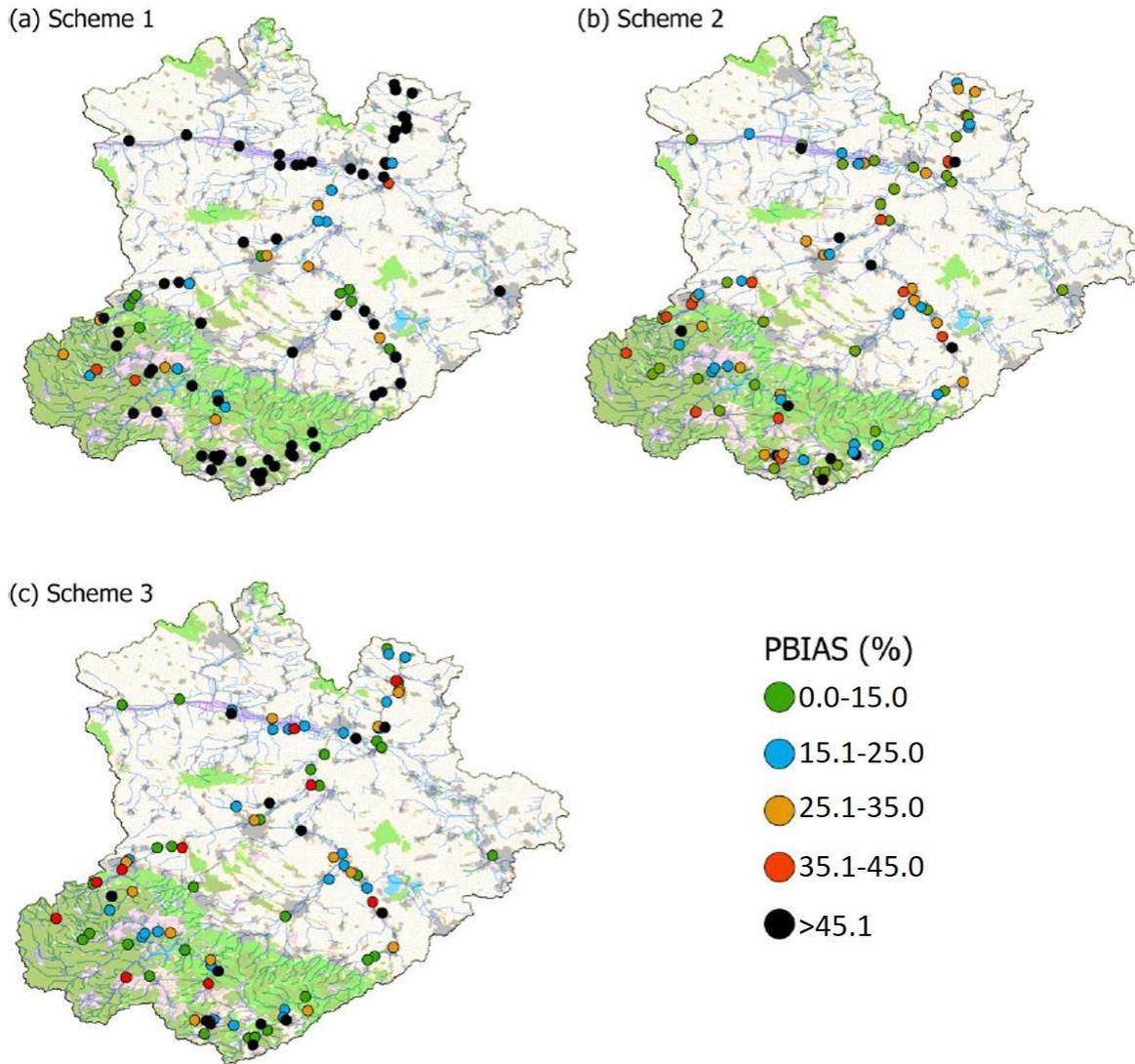
349 sampling stations, respectively) (Table 5). Performance was similar for schemes 2 and 3 (PBIAS \leq

350 15.0%: 34 vs. 35 sampling locations, respectively).

351 **Table 5.** Frequency of sampling locations associated with different PBIAS ranges across the three
 352 calibration schemes.

PBIAS (%)	Scheme 1	Scheme 2	Scheme 3
0.00–15.0	9	34	35
15.1–25.0	9	19	16
25.1–35.0	8	17	20
35.1–45.0	3	12	10
> 45.1	65	12	13

353 We also examined how catchment characteristics might influence model performance by looking at
 354 the spatial distributions of the PBIAS values for all 94 sampling locations across the three calibration
 355 schemes (Figure 4). The model performance for NO_3^- concentration at each stream order and land
 356 use (farmland vs. forest) are shown in Figure S1. Overall, more locations showed a good level of
 357 performance ($\text{PBIAS} \leq 15.0\%$) at scheme 2 versus scheme 1; no such difference was seen between
 358 schemes 2 and 3. For example, in forested areas, scheme 2 demonstrated considerable improvement
 359 compared to scheme 1. By visually inspecting, there was no noticeable distinction between scheme 2
 360 and 3 (Figure 4). More specifically, performance was better at scheme 2 than scheme 1 in areas
 361 dominated by farmlands for all stream orders (Figure S1). Additionally, performance was better for
 362 scheme 2 than scheme 3 except in the case of stream orders 2 and 4 in agricultural areas and stream
 363 order 5 in forested areas (Figure S1).



364

365 **Figure 4.** Performance of the mHM-Nitrate model for NO_3^- concentrations at the 94 spatially
 366 distributed sampling locations across the three calibration schemes.

367 We used the optimized parameter sets for scheme 2 to explore model performance in greater detail

368 at six spatially distributed sampling locations that displayed distinct characteristics (map: Figure 1c;

369 observed and simulated NO_3^- concentrations: Figure 6; PBIAS: Table 6). There was variation in the

370 duration and frequency of the validation data for the six sampling locations. Seasonal patterns of NO_3^-

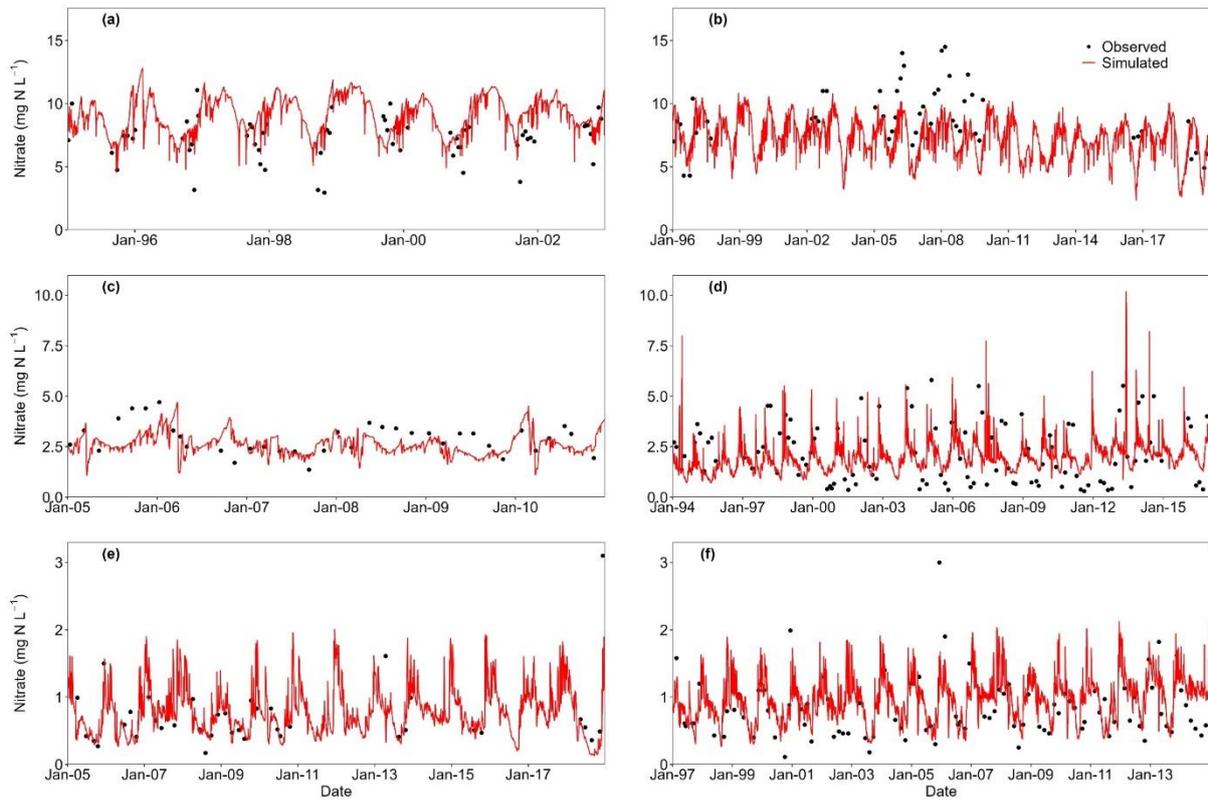
371 concentrations were well captured by the model over different levels of NO_3^- (Figure 5), with PBIAS

372 values ranging from -17.1% to 14.5% (Table 6). This result indicates that the mHM-Nitrate model was

373 capable of representing NO_3^- dynamics within different subcatchments when scheme 2 was applied.

374 The largest difference between mean observed and simulated NO_3^- concentrations occurred at NO_3^-

375 sampling location 2 (Figure 5c) with PBIAS value of -17.1%, which represents an arable dominated
 376 sub-catchment. The best fit between mean observed and simulated NO_3^- concentration was found at
 377 NO_3^- sampling location 4 (Figure 5d; PBIAS = -9.3%), which is found in a mountainous sub-catchment
 378 that contains a mixture of farmland and pasture (Figure 1c).



379

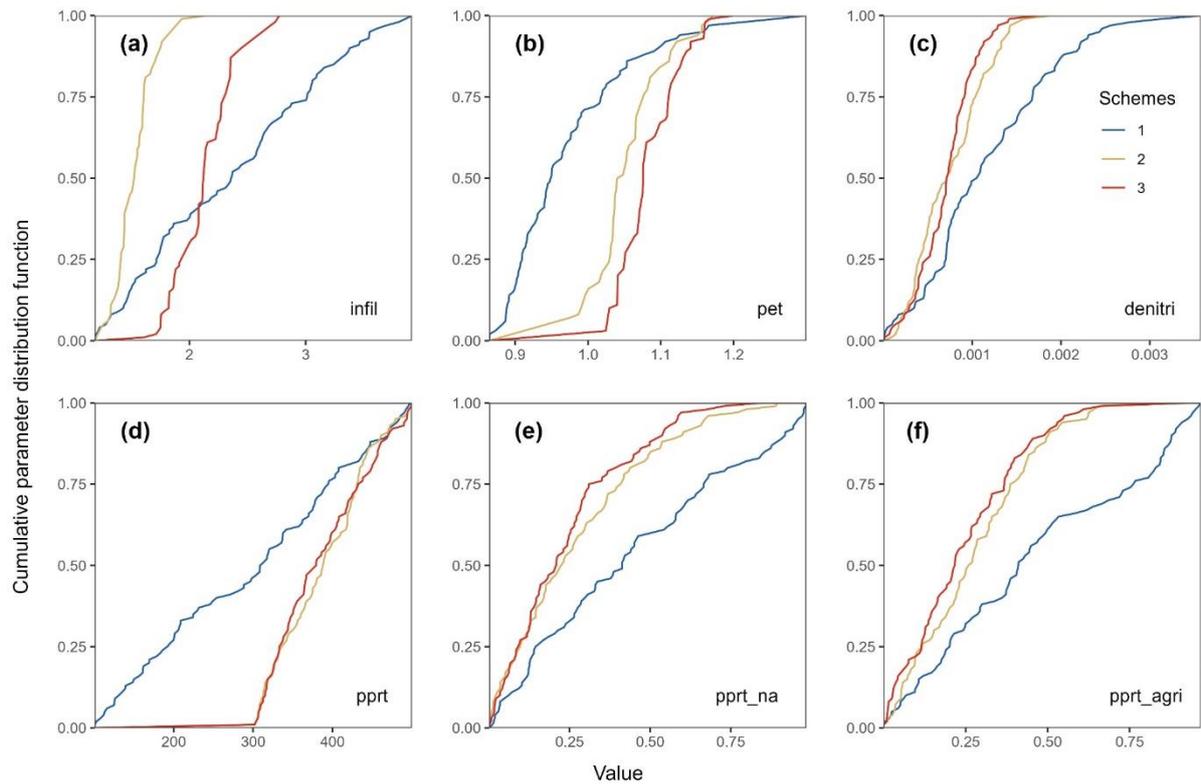
380 **Figure 5.** Observed and simulated nitrate (NO_3^-) concentrations (calibration scheme 2) for the six
 381 sampling locations displaying distinct characteristics.

382 **Table 6.** Summary of catchment characteristics represented by the six sampling locations, model
 383 performance for nitrate (NO_3^-) concentrations (PBIAS values), minimum and maximum values of
 384 simulated and observed NO_3^- concentrations at the sampling locations, and range (mean) of NO_3^-
 385 concentrations.

Sampling location	Sub-catchment area (km ²)	Dominant land use	PBIAS (%)	Simulated NO_3^- concentration (mg N L ⁻¹)	Observed NO_3^- concentration (mg N L ⁻¹)
1=a	11.8	Arable (87.2%)	12.8	4.6-12.9 (9.1)	2.9-11.1 (7.4)
2=b	12.6	Arable (78.3%)	-17.1	2.4-10.8 (7.4)	4.3-14.5 (8.9)
3=c	26.4	Arable (53.6%) Forest (40.1%)	-12.6	1.1-4.7 (2.6)	1.4-4.7 (2.9)
4=d	37.1	Arable (22.2%) Pasture (29.0%)	-9.3	0.8-10.2 (2.0)	0.3-5.8 (2.2)
5=e	6.1	Forest (96.0%)	-11.7	0.1-2.0 (0.7)	0.2-3.1 (0.7)
6=f	3.9	Forest (100%)	14.5	0.3-2.1 (1.0)	0.1-3.0 (0.8)

386 3.3 Model parameter distributions

387 For the three calibration schemes, we constructed cumulative distribution functions for the most
388 sensitive hydrological and water quality parameters using the best 100 model runs (Figure 6). From
389 the results, it is clear that the hydrological parameters—infiltration shape factor (infil) and potential
390 evapotranspiration (pet) differ significantly between schemes 1 and 2 as well as between schemes 2
391 and 3 ($p < 0.01$) (Figure 6 and Table 7). In the mHM-Nitrate model, soil infiltration is parameterized
392 using the power function of soil saturation, whose exponent is determined by the infiltration shape
393 factor (infil). Cuntz et al. (2015) reported that, as a parameter, infil is highly related to soil saturation,
394 where higher infiltration occurs in mountain soils than in lowland soils. Because the Meisdorf station
395 was included in scheme 2, a greater range of soil types were represented, allowing infil to be better
396 defined. In contrast, scheme 1 averaged all the soil types present in the catchment, as reflected by
397 the narrower ranges of infil for scheme 2 versus 1 (Figure 6). The cumulative distributions of four
398 water quality parameters, namely in-stream denitrification rate (denitri), primary production rate
399 (pprt), primary production coefficient in non-agriculture stream (pprt_na), and primary production
400 coefficient in agriculture stream (pprt_agri), showed dissimilarities between scheme 1 and schemes 2
401 and 3. However, there were no differences in the cumulative parameter distributions between
402 scheme 2 and scheme 3 ($p > 0.05$) (Figure 6 and Table 7). The four water quality parameters were
403 better constrained for scheme 2 than scheme 1, as reflected by their narrower ranges in the former
404 versus the latter (Figure 6). Yang et al. (2019b) found the control factors for denitri and pprt varied
405 between the Meisdorf and Hausneindorf stations. At Meisdorf, both parameters have a strong
406 correlation with stream discharge and benthic area, while at Hausneindorf they are highly correlated
407 with terrestrial flows and fluxes. In summary, parameter distributions were dramatically affected by
408 the increase in station number between scheme 1 and scheme 2. In contrast, the additional stations
409 added in scheme 3 had little to no effect.



410

411 **Figure 6.** Cumulative distributions for the hydrological parameters infil (a) and pet (b) and the
 412 water quality parameters, in-stream denitrification rate (denitri) (c), primary production rate (pprt)
 413 (d), primary production coefficient in non-agriculture stream (pprt_na) (e), and primary production
 414 coefficient in agriculture stream (pprt_agri) (f) across the three calibration schemes.

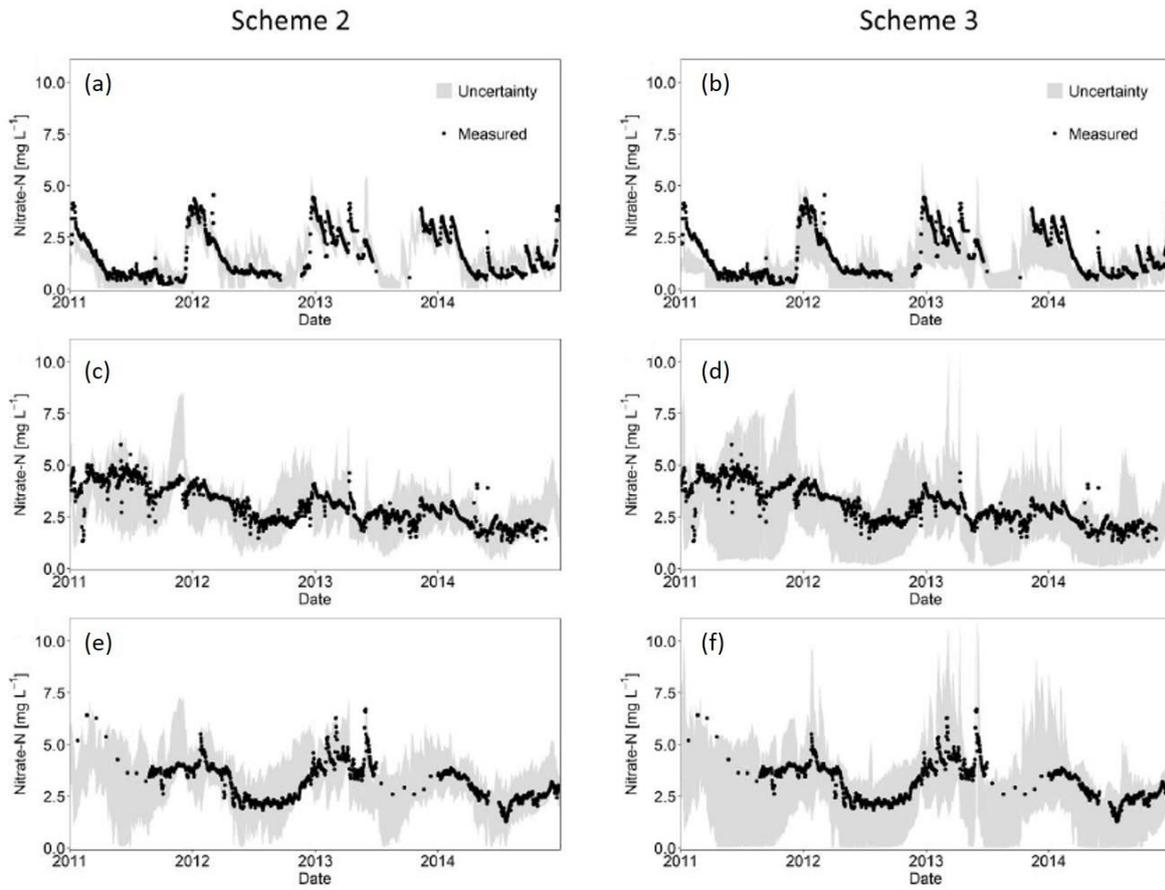
415 Table 7. Kolmogorov-Smirnov (KS) statistics and significance estimates for cumulative parameter
 416 distributions between calibration schemes.

Parameters	KS statistic and p-value		
	Scheme 1 and scheme 2	Scheme 1 and scheme 3	Scheme 2 and scheme 3
infil	0.64 (p<0.01)	0.38 (p<0.01)	0.88 (p<0.01)
pet	0.61 (p<0.01)	0.78 (p<0.01)	0.39 (p<0.01)
denitri	0.32 (p<0.01)	0.38 (p<0.01)	0.16 (p>0.05)
pprt	0.46 (p<0.01)	0.46 (p<0.01)	0.09 (p>0.05)
pprt_na	0.28 (p<0.01)	0.34 (p<0.01)	0.12 (p>0.05)
pprt_agri	0.31 (p<0.01)	0.37 (p<0.01)	0.14 (p>0.05)

417

418 3.4 Uncertainty analysis—nitrate concentrations

419 We calculated the 95% uncertainty boundaries for simulated daily NO_3^- concentrations at the
420 Meisdorf, Hausneindorf, and Stassfurt stations for schemes 2 and 3 (Figure 7). The associated R-
421 factors are given in Table 8. The 95% uncertainty boundaries for simulated daily Q associated with
422 schemes 2 and 3 are available in the Supplementary Materials (Figure S3). Whether under low- or
423 high-flow conditions, 95% uncertainty boundaries for daily NO_3^- concentrations were narrower for
424 scheme 2 than for scheme 3 (Figure 7). For instance, they were nearly twice as wide for scheme 3
425 than scheme 2 at Hausneindorf (R-factor = 4.13 vs. 2.18, respectively) and Stassfurt (R-factor = 4.52
426 vs. 2.79, respectively) (Table 8). Furthermore, over 60% of the observed NO_3^- concentrations lay
427 within the 95% uncertainty boundaries for scheme 2. When scheme 2 was used, the Meisdorf station,
428 located in a forested subcatchment, displayed lower levels of uncertainty than did the Hausneindorf
429 and Stassfurt stations, which are found in a subcatchment dominated by farmland. The same was also
430 true for scheme 3. This finding was reflected in the narrower 95% uncertainty boundaries for
431 Meisdorf versus Hausneindorf and Stassfurt (Figures 7a-b vs. 7c-f), as well as in the lower R-factor
432 values for Meisdorf (scheme 2 = 0.92; scheme 3 = 1.08; Table 8).



433

434 **Figure 7.** Comparison of 95% uncertainty boundaries for the simulated nitrate (NO_3^-) concentrations
 435 obtained with schemes 2 and 3 for three gauging stations: Meisdorf (a-b), Hausneindorf (c-d), and
 436 Stassfurt (e-f).

437 **Table 8.** R-factor values for nitrate (NO_3^-) concentrations at three gauging stations for schemes 2 and
 438 3.

Stations	Scheme 2	Scheme 3
Meisdorf	0.92	1.08
Hausneindorf	2.18	4.13
Stassfurt	2.79	4.52

439 4. Discussion

440 4.1 Evaluation of model performance for different calibration schemes

441 We evaluated the ability of the mHM-Nitrate model to simulate discharge and nitrate concentrations
442 at eight gauging stations. We specifically examined the transferability of hydrological and water
443 quality parameters at spatial scales.

444 4.1.1 Model performance for discharge under three calibration schemes

445 During model validation, simulated discharge at the catchment outlet was similar whether the
446 calibration data came from a single site (scheme 1: catchment outlet station) or multiple sites
447 (scheme 2: 3 stations and scheme 3: 8 stations) (Table 4). This result suggests that, for discharge, the
448 number of stations used during calibration did not affect model performance at the catchment outlet.
449 Our finding is consistent with those of Chiang et al. (2014); Wang et al. (2012); Wu et al. (2022a).

450 That said, performance was better with scheme 2 than scheme 1 when discharge was simulated for
451 all eight gauging stations, except in the case of Hausneindorf (Table 4). This result could have arisen
452 because multi-site calibration better constrains model parameters by including information on
453 catchment characteristics (e.g., land use and soil types) at upstream stations (here, Meisdorf and
454 Hausneindorf); these characteristics are frequently heterogeneous in space and shape hydrological
455 parameters (e.g., infil and pet, Figures 6a and 6b). Jiang et al. (2015) reported that, compared to
456 single-site calibration, multi-site calibration may better capture dynamics in large, diverse catchments
457 because it accounts for the effects of different hydrological processes (e.g., slow groundwater
458 dynamics and quick interflows). For example, in the Bode catchment, interflow is the primary form of
459 runoff in mountainous areas (Jiang et al., 2014), while the share of groundwater increases from the
460 mountains to the lowlands (Zhou et al., 2022).

461 In contrast, model performance was similar for schemes 2 and 3 (NSE values for the eight gauging
462 stations; Table 4), which suggests that adding more sites does not always improve simulations for
463 upstream stations. This finding is consistent with those of previous studies (Her and Chaubey (2015);

464 Wang et al. (2012); Xie et al. (2021)) and could potentially be explained by station choice and the
465 failure of scheme 3 to introduce any new catchment characteristics. As a result, schemes 2 and 3
466 displayed similar cumulative distributions for their hydrological parameters (Figures 6a-b). Therefore,
467 during calibration, it may be challenging to optimize model parameters by relying on station number
468 only.

469 4.1.2 Model performance for NO_3^- concentration under three calibration schemes

470 Simulated nitrate concentrations were significantly better for all gauging stations (with the exception
471 of Nienhagen) when scheme 2 versus scheme 1 was used (Table 4). This could be due to the fact that
472 the inclusion of Meisdorf in scheme 2 results in additional parameter constraint. The station is found
473 in a forested subcatchment, which likely led to changes in the values of land-use-dependent
474 parameters (e.g., pppt_na, pppt_agri; Figures 6e and 6f). These parameters were optimized in scheme
475 2 and, additionally, improved model performance at non-calibrated stations, such as Wegeleben and
476 Ditfurt. Both stations are located in subcatchments with intermediate levels of forest cover (> 30%
477 and > 56.4%, respectively). Similarly, the inclusion of Hausneindorf in scheme 2 improved model
478 performance at Hadmersleben, which had not been part of the calibration process, because the two
479 stations occur in regions with similar levels of farmlands (Table 2). This finding indicates that utilizing
480 multi-site calibration schemes that capture diverse catchment characteristics can improve simulated
481 nitrate concentrations even at locations that were not included in the calibration process. This result
482 concurs with those of previous studies (Chiang et al., 2014; Jiang et al., 2015; Shrestha et al., 2016),
483 which found that such improvements result from the fact that multi-site calibration schemes can
484 account for dramatic variability in observed nitrate concentrations and hydrological regimes across
485 catchments. These schemes can thus better constrain parameters associated with nitrate transport
486 and transformation.

487 In contrast, model performance was slightly lower at all stations (except Nienhagen) for scheme 3
488 than scheme 2 (PBIAS values; Table 4), which suggests that adding more gauging stations to the
489 calibration process cannot, by itself, result in further improvements to simulations of nitrate

490 concentrations. This finding may have two explanations. First, the five additional gauging stations
491 (Wegeleben, Hadmersleben, Peseckendorf, Ditfurt, and Nienhagen) included in scheme 3 did not
492 introduce additional diversity in catchment characteristics, which was the case when Meisdorf and
493 Hausneindorf were included in scheme 2 (Figure 2). For instance, except for Peseckendorf, four of the
494 five additional stations have farmland surface areas and mean nitrate concentrations that are similar
495 to that of Hausneindorf, which led to similar model parameter distributions for schemes 2 and 3
496 (Figures 6c-f). Second, three of the five additional stations have low-frequency measurements of
497 nitrate concentrations (i.e., once or twice per month). Jiang et al. (2019) found that, when the HYPE
498 model was applied to the Selke catchment, performance was better when calibration used nitrate
499 concentrations that were collected daily versus every two weeks. The slight decline in performance
500 from scheme 2 to scheme 3 could be affected by the model's attempt to satisfyingly balance the large
501 number of additional observations resulting from site addition (Jiang et al., 2015). In other words,
502 multi-site calibration approaches try to identify the parameter set that represents the best
503 compromise given the presence of multiple subcatchments, which is a more intensive task than
504 simply focusing on a single catchment outlet.

505 4.1.3 Comparison of hydrological and water quality model performance

506 In brief, the model's accuracy for predicting both discharge and NO_3^- concentration improved when
507 using Scheme 2 compared to Scheme 1. However, while the model's accuracy for discharge remained
508 consistent between Scheme 2 and 3, its accuracy for nitrate decreased in Scheme 3. On one hand,
509 hydrology is a physical process that is well understood and can be easily quantified through
510 measurements and modeling. On the other hand, nitrate dynamics are much more complex and can
511 be influenced by a variety of specific factors that are unique to a particular location, such as the
512 amount of fertilizer applied and the level of moisture in the soil. Nitrogen fertilizer application rates
513 are often uncertain and can vary depending on crop type and management practices. Nitrate uptake
514 by plants is also difficult to predict, as it is influenced by a range of factors such as soil moisture,

515 temperature, and nutrient availability. Overall, nitrate simulations are likely to be more accurate in
516 mountainous regions where quick flowing systems lead to less storage and transformation of nitrate
517 (Table 4). In lowland agricultural systems, nitrate can persist in soils for several years and in
518 groundwater for even longer time scales, leading to legacy effects that can complicate stream nitrate
519 dynamics (Wriedt and Rode 2006, Ehrhardt et al., 2019; Hrachowitz et al., 2015).

520 **4.2 Simulating nitrate concentrations across space**

521 Scheme 1, which solely utilized data from the catchment outlet, was unable to accurately simulate
522 nitrate dynamics at upstream sites within the large, heterogeneous Bode catchment. Indeed, PBIAS
523 values were high (> 45%) for many of the 94 spatially distributed sampling locations when scheme 1
524 was used (Figure 4a and Table 5). The model performed much better when scheme 2 was employed.
525 Its addition of two gauging stations to the calibration process thus appeared to greatly influence
526 model performance at the catchment scale.

527 However, little to no further improvement was seen with scheme 3 and its five additional gauging
528 stations. This assertion has two sources of support: schemes 2 and 3 had similar numbers of sampling
529 locations within the different PBIAS ranges (Table 5) and displayed similar cumulative distributions for
530 their parameters (Figure 6). Comparing cumulative parameter distributions can help identify
531 informative calibration stations. It can also determine whether adding or removing calibration
532 stations would improve.

533 Further results of the model performance of NO_3^- concentration at Scheme 2 shows varying
534 performances among NO_3^- sampling locations that represent different catchment characteristics (e.g.,
535 precipitation, land use, and fertilizer inputs) (Figure 5). At sampling location 4, NO_3^- concentration
536 was overestimated in summer, but the PBIAS value of the whole period was negative, it means that
537 the model underestimated NO_3^- concentrations during other times of the year. This could be due to
538 errors in the representation of hydrological processes, such as groundwater recharge, which can
539 affect NO_3^- transport and concentration in the groundwater. This suggests that spatial representation

540 of groundwater processes (such as groundwater NO_3^- concentration) are needed to be refined to
541 obtain better model performance for small sub-catchments. Faramarzi et al. (2015) and Gao et al.
542 (2016) concluded that the hydrological and water quality models that only rely on calibration without
543 refining internal process representation (e.g., groundwater NO_3^- concentration) will often not result in
544 further improvement. Nevertheless, the above analysis indicates that Scheme 2 is sufficient to ensure
545 the satisfactory model performance at NO_3^- sampling locations, since 75% of the NO_3^- sampling
546 locations showed absolute PBIAS $\leq 35\%$ (Figure 4b and Table 5) and the mHM-Nitrate model was
547 capable to present different magnitudes of NO_3^- levels for different sub-catchments which differ in
548 their catchment characteristics (Figure 5). These findings are in line with Ghaffar et al. (2021), where
549 they found that considering archetypal gauging stations in the calibration process leads better spatial
550 validation of the model at internal locations that were not originally considered in calibration. These
551 stations represent the maximum catchment characteristics in heterogeneous catchments in terms of
552 dominant land-use and meteorological features. This highlights the need for multiple internal
553 stations/locations to validate the model's capacity to accurately capture the complexity of natural
554 processes and identify which process needs to be improved (Beven, 2001; Daggupati et al., 2015).

555 **4.3 Impact of calibration approaches on model uncertainty**

556 For the three gauging stations, there was more uncertainty around simulated nitrate concentrations
557 for scheme 3 than for scheme 2 (Figure 7), likely because scheme 3 included stations with low-
558 frequency measurements. This result highlights the effect of measurement frequency on simulation
559 uncertainty. Indeed, low-frequency measurements may not capture the full range of variability in
560 NO_3^- dynamics. Furthermore, multi-site calibration approaches that rely on low-frequency data may
561 give rise to spatial representation issues, given that water quality can vary widely across
562 heterogeneous catchments and be influenced by local factors, such as land use and soil type. This
563 finding is in line with those of previous studies (Jiang et al., 2019; Khorashadi Zadeh et al., 2019;
564 Ullrich and Volk, 2010). For example, Jiang et al. (2019) found that, for the HYPE model, uncertainty

565 was reduced when the calibration process used NO_3^- concentrations that had been collected daily
566 versus every two weeks.

567 Once the catchment function is well captured by representative key stations (Scheme 2), additional
568 measurements may not be cost-effective and could increase model simulation uncertainty (Scheme
569 3). Therefore, it is essential to consider the specific requirements of the study and the desired level of
570 accuracy in model simulation. Depending on the goals and context, it may be necessary to find a
571 balance between cost-effectiveness and model performance by considering the spatial distribution of
572 measurements.

573 **4.4 Implication of spatial evaluation of distributed hydrological water quality model**

574 Improving the performance of hydrological water quality models has become a critical concern as
575 these models grow more complex (Beven, 2001; Refsgaard et al., 2016; Refsgaard et al., 2022).
576 Calibration using multiple sites is a crucial step in this process as it enables a better representation of
577 the spatial variability of hydrological and water processes. It is also equally essential to extend the
578 evaluation beyond calibration in order to gain insights into the spatial variability of hydrological and
579 water processes and understand the underlying processes that govern the behavior of the system
580 (Efstratiadis and Koutsoyiannis, 2010; Koch et al., 2015).

581 It is possible to use remote sensing data such as soil moisture (Mei et al., 2023; Rajib et al., 2016) and
582 evapotranspiration (Rajib et al., 2018; Zhang et al., 2021) to evaluate spatial performance of
583 distributed models for water quantity, but this approach cannot be applied to spatially evaluate
584 models for NO_3^- and other chemicals. To effectively evaluate the spatial performance of distributed
585 models for water quality, water quality monitoring or sampling is always necessary. This study
586 highlights the significance of long-term and spatially distributed monitoring water quality data, which
587 is readily accessible from authorities.

588 5. Conclusion

589 Using three different approaches, we calibrated a fully distributed process-based mHM-Nitrate model
590 that was then validated spatially and temporally at 8 gauging stations (discharge and NO_3^- -
591 concentrations) and 94 spatially distributed sampling locations (NO_3^- - concentrations) within the
592 heterogeneous Bode catchment in central Germany. Scheme 1 used only data from the catchment
593 outlet; scheme 2 used data from the catchment outlet and two upstream stations; and scheme 3 used
594 data from the catchment outlet and seven additional upstream stations. Our study found that, for
595 simulated discharge, model performance was similar at the catchment outlet for the three calibration
596 schemes. Furthermore, model performance did not improve consistently across the upstream
597 gauging stations.

598 In contrast, for NO_3^- - concentrations, scheme 2 was better than scheme 1 when it came to
599 simulating dynamics at sampling locations that had not been part of the calibration process. That said,
600 model performance across the sampling locations was similar for schemes 2 and 3. Our results
601 indicate that increasing the number of stations used in calibration does not necessarily improve
602 simulations of NO_3^- - concentrations. Additionally, we found that the use of low-frequency
603 calibration data may increase the degree of model uncertainty.

604 In conclusion, this research provides guidance on selecting gauging stations for the purposes of model
605 calibration: differences in cumulative parameter distributions should signal which stations can add
606 helpful additional representation. Furthermore, our work highlights that this selection process must
607 account for diversity in catchment characteristics, such as land use, meteorological patterns, and
608 elevation. In this way, the calibration data will better represent spatial patterns, and the model will
609 yield more accurate predictions. Overall, this study provides valuable insights into calibration-related
610 decision-making when carrying out fully distributed hydrological water quality models to simulate
611 dynamics within spatially heterogeneous catchments. This study also highlights the value of using

612 readily available water authorities monitoring data with high spatial resolution but low temporal
613 resolution for validating fully distributed models, even in the absence of discharge measurements.

614 Data Availability Statement

- 615 • The high-frequency monitoring data used are available at Zhang et al. (2022)
616 via <https://doi.org/10.48758/ufz.12911>
- 617 • The discharge and low frequency monitoring data are available at the data portal
618 (Datenportal) of the State Agency for Flood Protection and Water Management of Saxony
619 Anhalt, Germany (LHW, 2022) <https://gld.lhw-sachsen-anhalt.de/>
- 620 • The high-frequency monitoring data is available at TERENO (TERrestrial ENVironmental
621 Observatories) Data Discovery Portal <https://ddp.tereno.net/ddp/> (TERENO, 2020).

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