

1           **Enhancing quantitative precipitation estimation of  
2           NWP model with fundamental meteorological variables  
3           and Transformer based deep learning model**

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10          **Key Points:**

- 11          • We employ a transformer based deep learning model to improve the accuracy of  
12           precipitation estimation in NWP models.
- 13          • Various training strategies were implemented to manage the highly skewed pre-  
14           cipitation data leading to improvements in heavy rainfall events.
- 15          • Evaluation conducted with multiple metrics including skill score, quantile and spa-  
16           tial distribution as well as two case studies.

17 **Abstract**

18 Quantitative precipitation forecasting in numerical weather prediction (NWP) models  
 19 rely on physical parameterization schemes. However, these schemes involve considerable  
 20 uncertainties due to limited knowledge of the mechanisms involved in the precipitating  
 21 process, ultimately leading to degraded precipitation forecasting skills. To address this  
 22 issue, our study proposes using a Swin-Transformer based deep learning (DL) model to  
 23 quantitatively map fundamental variables solved by NWP models to precipitation maps.  
 24 Our results show that the DL model effectively extracts features over meteorological vari-  
 25 ables, leading to improved precipitation skill scores of 21.7%, 60.5%, and 45.5% for light  
 26 rain, moderate rain, and heavy rain, respectively, on an hourly basis. We also evaluate  
 27 two case studies under different driven synoptic conditions and show promising results  
 28 in estimating heavy precipitation during strong convective precipitation events. Over-  
 29 all, the proposed DL model can provide a vital reference for capturing precipitation-triggering  
 30 mechanisms and enhancing precipitation forecasting skills. Additionally, we discuss the  
 31 sensitivities of the fundamental meteorological variables used in this study, training strate-  
 32 gies, and performance limitations.

33 **Plain Language Summary**

34 Numerical weather prediction (NWP) models depend on certain empirical formu-  
 35 lations known as parameterizations to estimate precipitation. However, these methods  
 36 often fall short due to the intricate dynamics of rainfall, which involves numerous small-  
 37 scale interactions that these models are unable to fully capture. To counteract these lim-  
 38 itations, our study deploys a form of machine learning known as deep learning (DL) to  
 39 predict precipitation. This DL model utilizes fundamental weather variables derived from  
 40 NWP models to make its estimations, serving as a remedy for the inherent weaknesses  
 41 of traditional models caused by the uncertainties in their parameterization schemes. The  
 42 implementation of our deep learning model resulted in a significant enhancement in rain-  
 43 fall prediction accuracy, particularly in the case of extreme precipitation events. This  
 44 suggests that the application of machine learning strategies could be a promising approach  
 45 to improve the reliability of rainfall forecasts, a crucial element for effective weather pre-  
 46 diction and water resource management.

47 **1 Introduction**

48 Accurate quantitatively forecasting precipitation is essential for future planning and  
 49 very helpful for minimizing human lives and property damage beforehand of extreme events,  
 50 especially under the current rapidly changing climate. NWP models have been playing  
 51 an increasingly important role in all operational centres and academia for understand-  
 52 ing our earth system. It relies on discretizing a full set of governing equations includ-  
 53 ing the Navier-Stokes equations, ideal gas law and thermodynamics and solving them  
 54 numerically (Kalnay, 2003). With computing these prognostic equation sets over phys-  
 55 ical grids across different scales, the spatiotemporal evolving of the meteorological vari-  
 56 ables such as the temperature, wind speed and direction, air pressure and density are  
 57 represented under the rotating earth coordinates. Building upon the enrichment of sci-  
 58 entific knowledge in fundamental physics, and accelerated with the advances in technol-  
 59 ogy such as computational power and numerous sources of observational data, NWP is  
 60 showing quite a revolution over the past decades (Bauer et al., 2015).

61 As a result, the forecasting of fundamental meteorological variables under a so-called  
 62 resolved scale of motion is readily available and more reliable in terms of its accuracy.  
 63 However, many processes under the unresolved scales of motion also enter the equations,  
 64 such as the moist processes involving condensation and evaporation, turbulence, convec-  
 65 tive activities and cloud microphysics, which are tightly related to precipitation forma-  
 66 tion and need to be parameterized to describe their relations with the states in resolved

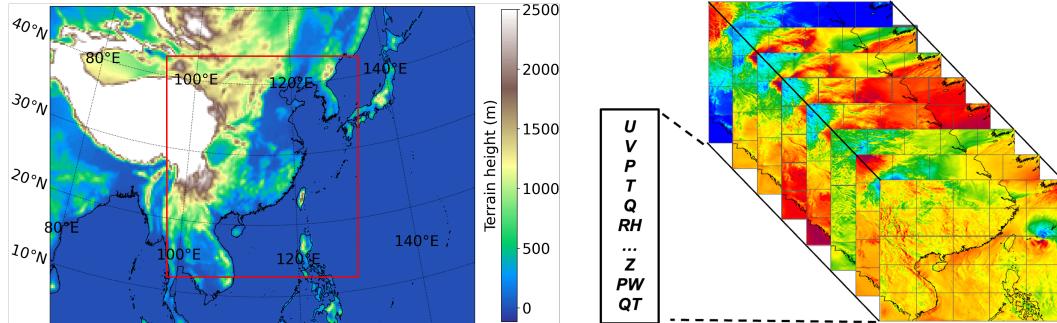
scales (Bauer et al., 2015). These parameterization schemes are generally based on the simplification and approximation of the physic laws to facilitate the numerical solutions, hence carefully chosen and sensitivity tests for the parameterization schemes will considerably affect the precipitation forecasting skills. Moreover, with an insufficient understanding of the underlying physics and some inherent uncertainties, using parameterization schemes will intrinsically bottleneck the further improvements of the performance for quantitatively estimating the precipitation (Zhou et al., 2022).

With the blossoming of artificial intelligence, many researchers have demonstrated the great ability of deep learning models in handling geoscience and remote sensing tasks, including precipitation estimation and forecasting. Shi et al. (2015) proposed and evaluated a series of spatial-temporal models dealing with precipitation nowcasting problems by extrapolating radar echoes and achieved better performance compared to traditional optical flow method (Shi et al., 2015, 2017). Ravuri et al. (2021) proposed to use a generative model with the stochastic method to extend nowcasting leading time without resorting to blurring. Sønderby et al. (2020) constructed a deep learning predictive model that uses satellite, radar and precipitation data and achieved a forecast leading time of 8 hours with a high spatiotemporal resolution, and outperforms the High-Resolution Rapid Refresh (HRRR) in terms of its accuracy. Other than precipitation nowcasting tasks, machine learning tools are also commonly applied to satellite images for precipitation estimation. Tao et al. (2017) proposed a deep learning model to extract features from bispectral satellite infrared (IR) and water vapour (WV) channels for detecting rain areas. Chen et al. (2019) proposed a two-stage hybrid neural network to estimate precipitation using ground-based radar and satellite observations. Wang et al. (2021) proposed a transfer learning based method, which uses data-riched Continental US (CONUS) IR dataset from the Geostationary Operational Environment Satellite (GOES) for pre-training of the model, and then transferred to China through re-training with multi-band IR signals from Chinese Fengyun (FY) satellite. Gao et al. (2022) used a U-Net model combined with the attention mechanism to directly retrieve precipitation maps using multi-band FY satellite images at a near real-time scale.

Many recent studies attempt to use data-driven models directly perform the NWP tasks in favour of their computational efficiencies compared to state-of-the-art NWP models. These data-driven models are generally trained on climate model outputs, general circulation models (GCM)(Scher & Messori, 2019; Chattopadhyay et al., 2020), or trained on reanalysis products such as ECMWF Reanalysis v5 (ERA5) dataset (Rasp et al., 2020; Rasp & Thuerey, 2021). Dueben and Bauer (2018) presented a “toy model” to identify challenges and fundamental design choices for deep learning based forecasting systems. Arcomano et al. (2020) designed a deep learning model and performed a 20-day global forecast. Evaluation results indicate that the DL model outperforms the NWP models for those state variables most affected by parameterization processes. Other than convolutional based deep learning models, Pathak et al. (2022) built a Fourier operator based transformer network to perform weather forecasting at globally  $0.25^\circ$  resolution and achieved matched accuracy with the state-of-the-art NWP model system the ECMWF Integrated Forecasting System (IFS).

Another prominent application of machine learning and deep learning techniques for NWP tasks is post-processing and bias correction. Grönquist et al. (2021) applied a convolutional neural network for bias correction of ensemble NWP predicted temperature field at various pressure levels, and achieved 14% improvement of ensemble forecast skill (CRPS) with a considerable reduction of computational cost owing to reduce the usage of trajectories. Taillardat and Mestre (2020); Li et al. (2022); Hess and Boers (2022) used machine learning and deep learning frameworks for post-processing quantitative precipitation forecasting results on the ensemble NWP models and achieved promising results on estimating heavy rainfall events located at long tails of the distribution curve.

With many machine learning methods have achieved remarkable results for nowcasting tasks or forecasting basic meteorological variables mentioned above. Directly map-



**Figure 1.** Left: Study domain and terrain height for input WRF meteorological data.

Right: Input meteorological data simulated by WRF model.

ping basic meteorological variables to precipitation amounts using a machine learning model instead of parameterization schemes has been rarely explored. Therefore, in this study, we aim to develop a deep learning method for extracting rainfall features, from basic meteorological variables including temperature, water vapour, and atmospheric movements simulated by Weather Research and Forecasting Model (WRF) model at 27-km resolution. The basic variables were fed to an attention mechanism based Shift Window Vision Transformer (Swin Transformer) neural network (Dosovitskiy et al., 2021; Liu et al., 2021), and targeted to reproduce the high-resolution satellite rainfall product, the Climate Prediction Center morphing method (CMORPH) data (Xie, Pingping et al., 2019). This deep learning method will circumvent uncertainties of the physical parameterization scheme owing to incompletely understood physical processes and capture the non-linearity relationship between the predictors and labels.

## 2 Methodology

We consider the task of quantitative precipitation estimation using a deep learning model as an optimization problem, which can be formulated as follows:

$$\boldsymbol{\Theta} = \underset{\boldsymbol{\theta}}{\operatorname{argmin}} \int \mathcal{L}(\boldsymbol{\Psi}(\mathbf{X}; \boldsymbol{\theta}), \mathbf{y}) dy \quad (1)$$

In this formulation, the model takes pairs of input consisting of basic atmospheric variables  $\mathbf{X}$  and precipitation observational data  $\mathbf{y}$ . The mapping function  $\boldsymbol{\Psi}$  is used to relate these inputs, and it has trainable parameters  $\boldsymbol{\theta}$ . The goal is to find the optimal parameters  $\boldsymbol{\Theta}$  by minimizing a set of loss functions  $\mathcal{L}$  using optimization algorithms.

### 2.1 Dataset preparing

#### 2.1.1 Training predictors:

To generate a dataset for high-resolution precipitation maps, we conducted a long-term dynamical simulation of 5 years (2017 – 2021) over the wettest season of south-east China, as shown in the left panel in Fig.1. The simulation was performed using the Weather Research and Forecasting (WRF) model with driven data from ERA5. The raw resolution for ERA5 data is  $0.25^\circ \times 0.25^\circ$ , and we used the meteorological fields from the WRF domain 1 simulation with 27-km resolution as predictors.

For the numerical simulation, we used a vertical grid with 38 levels to accurately represent the atmospheric system. We selected four model layers of three-dimensional

(3D) basic variables, including wind velocity ( $U$ ,  $V$ ,  $W$ ), pressure ( $P$ ), temperature ( $T$ ), geopotential height ( $z$ ), and humidity ( $Q$ ), and stacked them with the corresponding variable at the surface level. We also included a two-dimensional (2D) diagnostic variable, total precipitable water (TPW), in our analysis. The vertical wind speed ( $W$ ) at a height of 10 meters was not available in our model output, so the number of layers for this variable is limited to four. These variables were combined to form a 34-layer feature map, which we will refer to as 34 channels, as shown in the right panel Fig.1.

In atmospheric modelling, it is common practice to use pressure levels due to the decrease in pressure with height in the atmosphere. However, interpolating values from the model layers to the pressure levels can sometimes result in missing values of the basic variables due to varying terrain heights. To avoid this issue and minimize the memory requirements for training a deep learning model, we chose to directly use the values from the model layers rather than interpolating to pressure levels. This allowed us to accurately represent the atmospheric system while minimizing the computational resources needed for the simulation.

### **2.1.2 Observational precipitation data:**

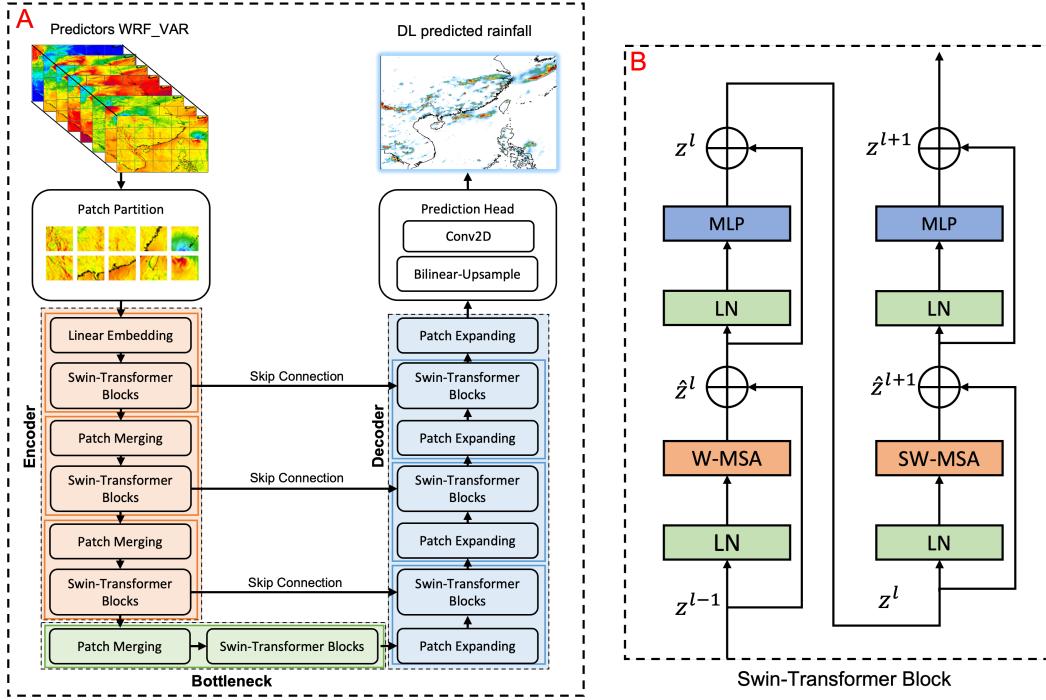
To obtain observational precipitation data as the reference for ground truth, we used CMORPH (NOAA CPC Morphing Technique), a high-resolution global satellite precipitation product. The data is created by combining passive microwave and infrared wave radiance measurements from multiple satellite instruments and adjusted using daily rain gauge analysis. The full-resolution CMORPH data used in this study has a high spatial resolution of 8 km and a frequency of 30 minutes. To align it with the meteorological data from the model used in this study, the CMORPH data was resampled to an hourly frequency, and its pixel values were matched to the corresponding grid points in the model.

## **2.2 Network Architecture**

The architecture of the proposed model shown in Fig.2 A is based on the classical encoder-decoder framework, which has been successfully applied to many semantic segmentation tasks in the computer vision field. The model is inspired by the original UNet model (Ronneberger et al., 2015). The gridded meteorological data generated by the WRF model are first divided into  $4 \times 4$  non-overlapping patches by using a 2D convolutional layer with a stride and kernel size equal to the patch size. The patches are then transformed into sequence embeddings and fed into the encoder.

In the encoder, we replace the CNN backbone network in the original UNet model with the Swin-transformer (Liu et al., 2021). Each encoder block consists of a patch merging layer and four Swin-transformer blocks. The patch merging layer performs downsampling, similar to the pooling operation in CNN-based models, while the Swin-transformer block extracts features, similar to a convolutional operation. The input data passed through the patch partition layer has a size of  $\frac{H}{4} \times \frac{W}{4}$  with  $C$  embedding channels, where  $H$  and  $W$  are the height and width of the input data. With each pass through an encoder block, the height and width are halved and the number of channels doubles.

In the decoder block, the patch-expanding layer performs upsampling using bilinear interpolation to restore the feature map to its original resolution, doubling the size of the feature map and reducing the feature dimension to half of its original dimension. To maintain the information lost during downsampling in the encoder, the expanded feature maps are fused with the downsampled features through a skip connection structure. This allows the Swin-transformer blocks in the decoder to receive inputs with the same size as the corresponding level of the encoder, but with features crossing multiple dimensions.



**Figure 2.** Architecture of the Swin-Transformer-Unet model (A), and the basic computational principle for a Swin Transformer block (B).

The bottom levels of the encoder and decoder are connected by a bottleneck, which has the same structure as the encoder but only with two Swin-transformer blocks. This hierarchical architecture design is rooted in the principle of enhancing the network's ability to learn features at multiple scales, which is particularly important for meteorological data due to its multi-scale nature.

The computation route of the Swin-transformer block, as depicted in Fig.2 B, is designed to reduce the computational cost compared to traditional multi-head self-attention (MSA) modules through the implementation of a sliding window mechanism. The Swin-transformer structure (Liu et al., 2021), accomplishes this reduction by composing a Swin-transformer block of two consecutive attention modules. Each attention module is composed of two LayerNorm (LN) layers, a multi-head self-attention module, a residual connection shortcut, and a 2-layer multilayer perceptron (MLP) with Gaussian Error Linear Units (GELU) nonlinearity.

Before the MSA module and the MLP module, the LayerNorm (LN) is applied, and the two attention modules differ in the type of multi-head self-attention employed. The first MSA layer uses a regular Window-Based MSA (W-MSA), while the second MSA layer adopts a Shifted Window-Based MSA (SW-MSA) module. The calculation of the Swin-transformer block is described by the following equations:

$$\hat{z}^l = W\text{-MSA}(\text{LN}(z^{l-1})) + z^{l-1} \quad (2)$$

$$z^l = \text{MLP}(\text{LN}(\hat{z}^l)) + \hat{z}^l \quad (3)$$

$$\hat{z}^{l+1} = \text{SW-MSA}(\text{LN}(z^l)) + z^l \quad (4)$$

$$z^{l+1} = W\text{-MSA}(\text{LN}(\hat{z}^{l+1})) + \hat{z}^{l+1} \quad (5)$$

218 where  $\hat{z}^l$  is the output features of the (S)W-MSA module,  $z^l$  is the output features of  
 219 the MLP module, and  $l$  represents the number of blocks.

## 220 2.3 Experiment details

### 221 2.3.1 Loss functions:

In meteorology, most variables, such as temperature and wind speed, are typically represented as continuous values over the model grids. These variables are usually smooth and evenly distributed unless there are significant changes in terrain or extreme weather conditions. Precipitation data, however, is often underestimated during extreme precipitation events due to its imbalanced distribution, making it a challenging problem to accurately predict from an engineering perspective. To address this issue, we propose the use of the Tversky (Salehi et al., 2017) loss function in our model. The Tversky loss function is defined as:

$$\mathcal{L}_{Tversky} = 1 - \frac{\sum_i^N p_i g_i}{\sum_i^N p_i g_i + \alpha \sum_i^N p_i (1 - g_i) + \beta \sum_i^N (1 - p_i) g_i} \quad (6)$$

where  $p_i$  and  $g_i$  are the predicted and ground truth values at pixel  $i$ , respectively, and  $\alpha$  and  $\beta$  are the weighting factors for the false positive and false negative terms. It can be viewed as a generalization version of the Dice similarity coefficient, which is widely used in image segmentation tasks. The advantage of using the Tversky loss function is that it provides better control over the trade-off between precision and recall by allowing for the adjustment of the false positive and false negative weighting factors. This is particularly useful in our situations where the precipitation data is extremely imbalanced, with the background pixels represented by 0 indicating no rainfall. To encourage the model to better predict extreme precipitation, the Tversky loss function has been adjusted to assign higher penalties to false negative predictions. This is achieved by setting  $\beta$  to a higher value than  $\alpha$  in our experiments, making the model more inclined to accurately predict extreme precipitation values. In addition to the Tversky loss function, we also use the mean squared error (MSE) loss function, which is a commonly used loss function for regression problems, as our precipitation prediction task can also be viewed as a pixel-wise regression problem. The combined loss function can be expressed as:

$$\mathcal{L}(y, \hat{y}) = 0.5\mathcal{L}_{Tversky}(y, \hat{y}) + 0.5\mathcal{L}_{MSE}(y, \hat{y}) \quad (7)$$

### 222 2.3.2 Data transform:

To enhance the network's convergence, we introduce data transformations to the original input data. Initially, a center-cropping operation is executed on the meteorological data to mitigate boundary errors. This step is crucial as the WRF model, being a regional weather forecasting model, necessitates input from the global background field from ERA5. This requirement can potentially introduce errors at the lateral boundaries of the study domain. Next, we apply mean-std normalization to the input data to scale each predictor. This helps to bring the data into a similar range and prevent one predictor from dominating the other. Finally, we take a log transformation of the observational label data, the precipitation map, to reduce the skewness of its distribution. With the log transformation, the distribution of data is better represented and improves the convergence in the training as well as the accuracy while inferencing.

### 234 2.3.3 Evaluation metrics:

To evaluate the performance of the model, various metrics are calculated including the probability of detection (POD), threat score (TS), equitable threat score (ETS), false alarm ratio (FAR), and BIAS. These metrics are defined as follows:

$$POD = \frac{h}{h + miss} \quad (8)$$

$$TS = \frac{h}{h + miss + f} \quad (9)$$

$$ETS = \frac{h - h_{random}}{h + miss + f - h_{random}} \quad (10)$$

$$BIAS = \frac{h + f}{h + miss} \quad (11)$$

$$FAR = \frac{f}{h + f} \quad (12)$$

The POD primarily focuses on the number of hits, while the TS, FAR, and BIAS evaluate the combined impact of hits and false alarms. The ETS takes into account the possibility of a hit by chance by calculating  $h_{random} = \frac{h+f}{h+m+f+cn}$ , where  $cn$  is the number of correct negatives. A higher POD, TS, and ETS, and a lower FAR or a BIAS closer to 1 are considered an indicator of a more accurate prediction.

### 240 3 Analysis of experiment results

#### 241 3.1 Overall performance

In our study, we utilized a neural network that was trained using 6 years of WRF simulation data, and the WRF simulation was driven by the ERA5 reanalysis data, which is the foundation of our deep-learning model. Therefore, the performance of our deep-learning model is ultimately dependent on the original ERA5 background and our WRF simulation. Despite the advancements in weather forecasting, it remains challenging to reproduce the heavy precipitation events that are generated by intense convective systems. By utilizing the Swin-transformer Unet to process fundamental meteorological variables, we observed significant improvements in the prediction of heavy precipitation events on the tail, both in terms of accuracy in rainfall quantities and the location of the rainfall. The overall evaluation scores calculated from July to December 2021, with all leading times from 0 to 72 hours, are listed in Table 1.

(1) Our deep learning model for the prediction of drizzles with a rainfall intensity greater than  $1\text{mmh}^{-1}$  has been shown to enhance POD, TS, and ETS, while simultaneously reducing the FAR. Additionally, the FIAS is closer to 1 when compared with the results obtained from a pure WRF simulation. The improvement in the TS score is particularly significant, reaching as high as 21.7%.

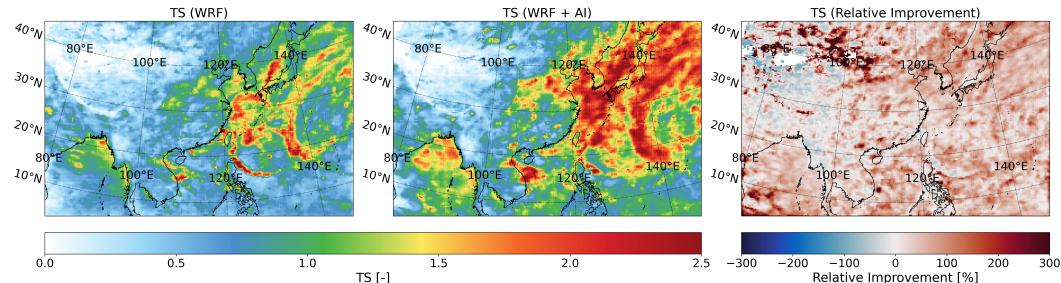
(2) In the prediction of moderate rainfall with an hourly intensity of  $3\text{mmh}^{-1}$  and  $5\text{mmh}^{-1}$ , the POD has increased from 0.145 to 0.218 and from 0.088 to 0.161 respectively, while the FAR has slightly decreased. As a result, the TS and ETS scores have also seen considerable increases, with the TS score increasing from 0.117 to 0.164 for the  $3\text{mmh}^{-1}$  threshold and from 0.076 to 0.122 for the  $5\text{mmh}^{-1}$  threshold. The ETS score has increased from 0.106 to 0.151 for the  $3\text{mmh}^{-1}$  threshold and from 0.071 to 0.114 for the  $5\text{mmh}^{-1}$  threshold. The relative improvement ratio for both the TS and ETS scores is as high as 60.5% for the  $5\text{mmh}^{-1}$  threshold.

(3) For the heavy rainfall events with an hourly precipitation intensity exceeding 10mm, our model has demonstrated the ability to significantly enhance the POD, TS, and ETS scores. However, in detecting heavy rainfall, there is a trade-off of introducing higher FAR. This indicates that the meteorological background data may not fully match the observational precipitation data, causing some intensive weather systems simulated by the WRF model with possibilities of heavy precipitation to be mistakenly placed. The improvement in POD for heavy rainfall, with the addition of both 10mm and 20mm intensities, is around 50%, increasing from 0.06 to 0.12. Similarly, the improvement in the TS and ETS scores ranges from 0.055 to 0.08, with an improvement rate of 45.5%. The pure WRF simulation has a BIAS of around 0.1 for heavy rainfall, which suggests that the detected rainfall area is sub-optimally small. Our deep-learning model improved

**Table 1.** The evaluation results for the WRF simulation and WRF + Swin-transformer Unet for the hourly rainfall intensity with forecasting leading time of 0 – 72h, the evaluation period is July to December 2021. Evaluation metrics including POD (probability of detection), FAR (false alarm ratio), TS (threat score), ETS (equitable threat score), and BIAS ratio are listed.

	Precipitation ( $mmh^{-1}$ )	POD	FAR	TS	ETS	FBIAS
<b>WRF</b>	0.1	0.508	0.568	0.295	0.213	1.263
	1.0	0.326	0.611	0.209	0.172	0.902
	3.0	0.145	0.610	0.117	0.106	0.378
	5.0	0.088	0.605	0.076	0.071	0.223
	10.0	0.045	0.563	0.040	0.039	0.123
	20.0	0.016	0.337	0.014	0.014	0.073
<b>WRF + AI</b>	0.1	0.525	0.475	0.359	0.289	1.013
	1.0	0.346	0.523	0.254	0.222	0.726
	3.0	0.218	0.611	0.164	0.151	0.551
	5.0	0.161	0.667	0.122	0.114	0.462
	10.0	0.086	0.755	0.063	0.061	0.391
	20.0	0.030	0.695	0.021	0.020	0.356

this score to nearly 0.4, demonstrating a better ability to detect heavy rainfall. Improving the accuracy of heavy rainfall detection and reducing the FAR is an important area for further research.

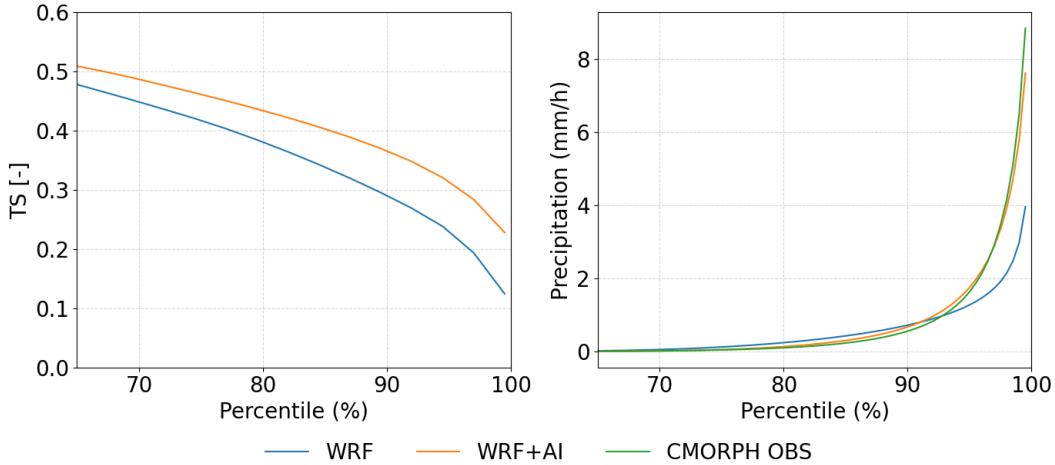


**Figure 3.** Spatial distribution of Threat Score (TS) for the baseline WRF, WRF + AI framework, and its relative improvements over the evaluation dataset.

(4) In terms of spatial distribution, our deep learning framework exhibits a substantial enhancement in forecast skill, as quantified by the cumulative TS score across all precipitation intensity thresholds. Notably, this improvement is observed in areas prone to heavy precipitation events, such as potential tropical cyclone pathways and associated rainbands, in addition to Mei-Yu frontal systems during monsoon seasons. The relative advancements in these regions amount to several-fold increases, while minor degradation is detected in an insignificant fraction of pixels, ensuring overall model performance remains robust mesoscale application.

### 3.2 Quantile distribution

Quantile distribution stands as a pivotal measure for evaluating the predictive proficiency of heavy rainfall events. Fig. 4 demonstrates a monotonic decline in model per-



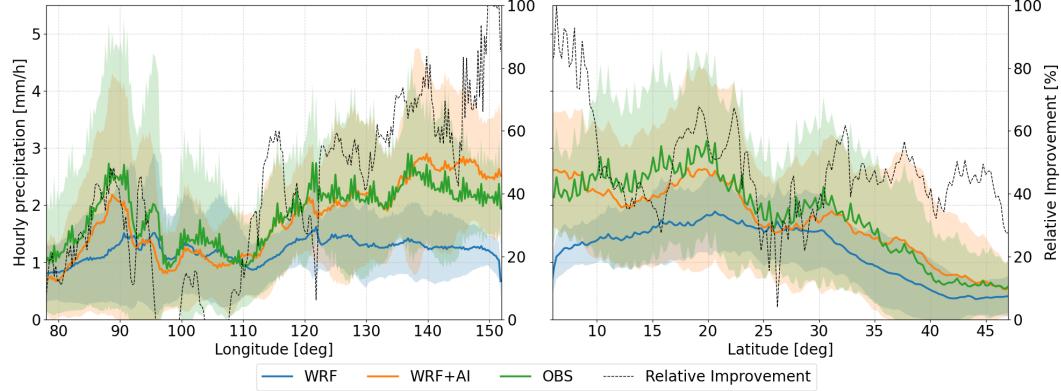
**Figure 4.** Threat Score (TS) for precipitation events above the percentile thresholds for WRF and WRF + AI framework (left), and hourly precipitation intensity for WRF, WRF + AI, CMORPH observation at each corresponding percentiles.

291 performance across varying quantile thresholds, signifying the escalating challenge of accu-  
292 rately predicting more extreme events.

293 When examining lower percentiles, our baseline WRF simulation, which employs  
294 the Kain-Fritsch (KF) cumulus parameterization scheme (Kain & Fritsch, 1993; Kain,  
295 2004), displays performance characteristics akin to our deep learning framework in terms  
296 of its TS, albeit with a slight overestimation of light rainfall between the 70th and 90th  
297 percentile with an hourly rainfall intensity less than 1mm. This is primarily due to the  
298 fact that the KF scheme is well-suited to accommodate light precipitation events, which  
299 are less intricate to predict. This finding is consistent with several dynamic downscal-  
300 ing studies (Ma & Tan, 2009; Otieno et al., 2020) that have indicated a tendency of the  
301 KF scheme to produce a wet bias during light rainfall while demonstrating limited pre-  
302 dictability for heavier rainfall events.

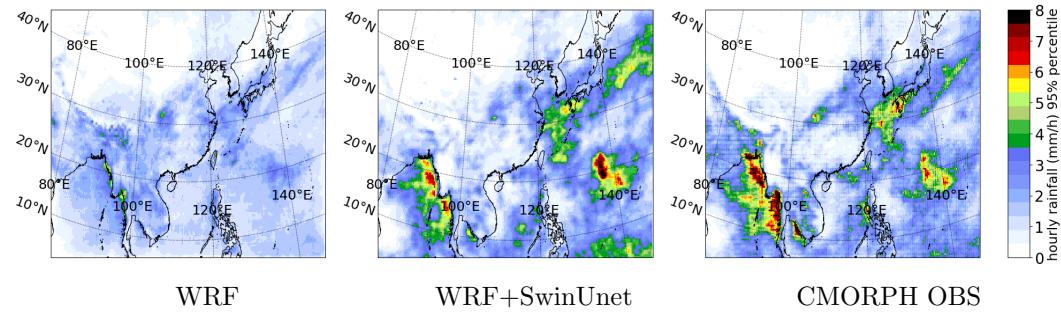
303 When we examine more intense rainfall events that surpass the 95th percentile thresh-  
304 old, our proposed deep learning framework begins to exhibit superiority in its capacity  
305 to replicate these extreme events and align the distribution more closely with observed  
306 ground truth. Within this percentile range, the TS score for the deep learning frame-  
307 work is approximately double that of the baseline WRF model, and the estimated rain-  
308 fall intensity is closely tracks the CMORPH observation until the 99.5th extreme per-  
309 centile is reached. At this point, the largest bias is about 20%, whereas the baseline WRF  
310 model can only capture less than 50% of the intensity. This result underscores the en-  
311 hanced performance of this deep learning framework in the accurate prediction and rep-  
312 resentation of intense rainfall events.

313 In addition to the above, we present the meridional and zonal averages of hourly  
314 precipitation intensity at the 95th percentile in Fig. 5, as well as its spatial distribution  
315 in Fig. 6. The application of our deep learning framework enables us to refine the es-  
316 timated rainfall intensity at the 95th percentile, resulting in a significantly improved align-  
317 ment with the CMORPH observation, both spatially and in terms of its zonal and merid-  
318 ional mean. The spread property, represented by the shadowed area, also aligns more  
319 closely with the observations.



**Figure 5.** Meridionally (left) and Zonally (right) averaged precipitation intensity at 95th percentile.

The highest relative improvements are evident in the eastern part and southwest quadrant of our experimental domain, with the maximum relative improvement exceeding 100%. Some minor fluctuations and degradation are noticeable around 100°–110° Longitude and approximately 25° Latitude, the mountainous southwest part of China. This suggests that our deep learning framework may demonstrate lower confidence when dealing with orographic precipitation. Similar observations can be made from the spa-

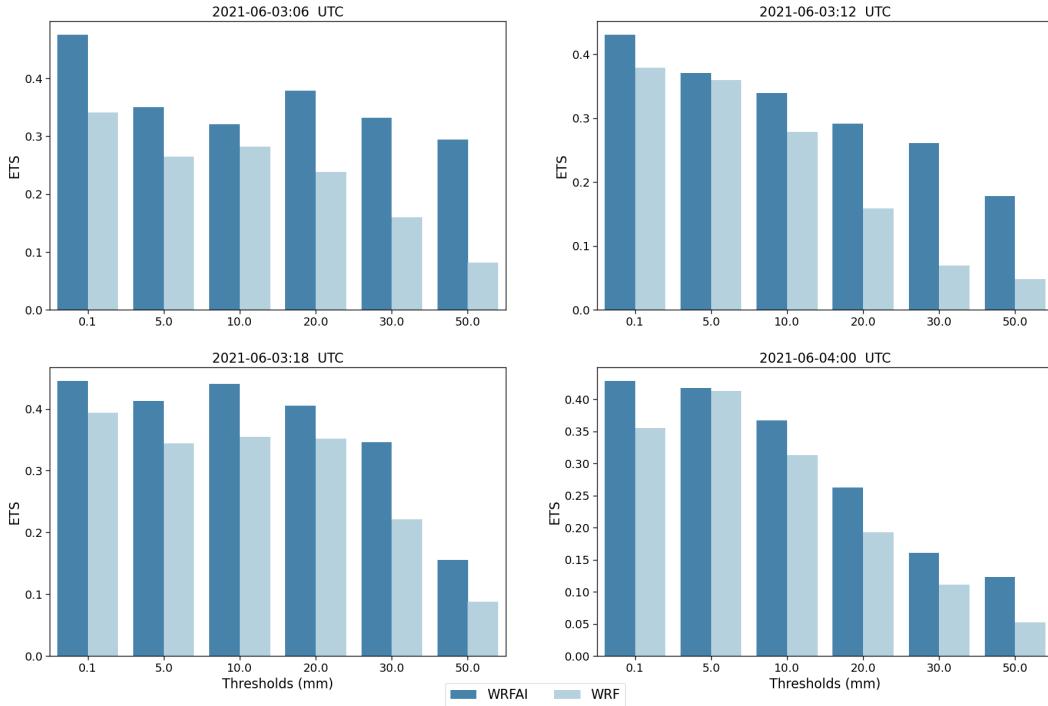


**Figure 6.** Spatial distribution of 95th hourly precipitation intensity for the baseline WRF (left), WRF + AI (mid) and CMORPH observation (right).

tial distribution plot in Fig. 6. The largest improvements are seen over potential tropical cyclone initiation areas in the western Pacific Ocean, the eastern China region influenced by the Mei-Yu frontal rain belts around the middle and lower reaches of the Yangtze River, the southern part of Japan and the Korean Peninsula, as well as the Southeast Asia region around the Bay of Bengal. Other regions also show varying degrees of improvement in terms of spatial distribution, indicating that our deep learning framework more accurately represents extreme precipitation patterns compared to the baseline WRF simulation, which exhibits limited predictability over these heavily precipitating areas during the wet season.

### 3.3 Case study

In this section, we present two case studies of heavy precipitation recorded in the study domain, as shown in Fig. 8 and Fig. 10. The precipitation maps depict the ac-



**Figure 7.** Equivalent Threat Score (ETS) of 6-hour accumulated precipitation for the case study on June 3rd, 2021 at 06:00 UTC, 12:00 UTC, 18:00 UTC, and 24:00 UTC

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cumulated intervals of 6 hours and cover a period of one day for each case. The first case study occurred on June 3rd, 2021, while the second case study occurred on August 20th, 2021.

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### 3.3.1 2021-06-03:

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On June 3rd, 2021, a frontal rain band was observed to be moving southeastward. In our default WRF simulation, results are generally indicating weak signals for several precipitation hot spots over both the continent and the ocean side.

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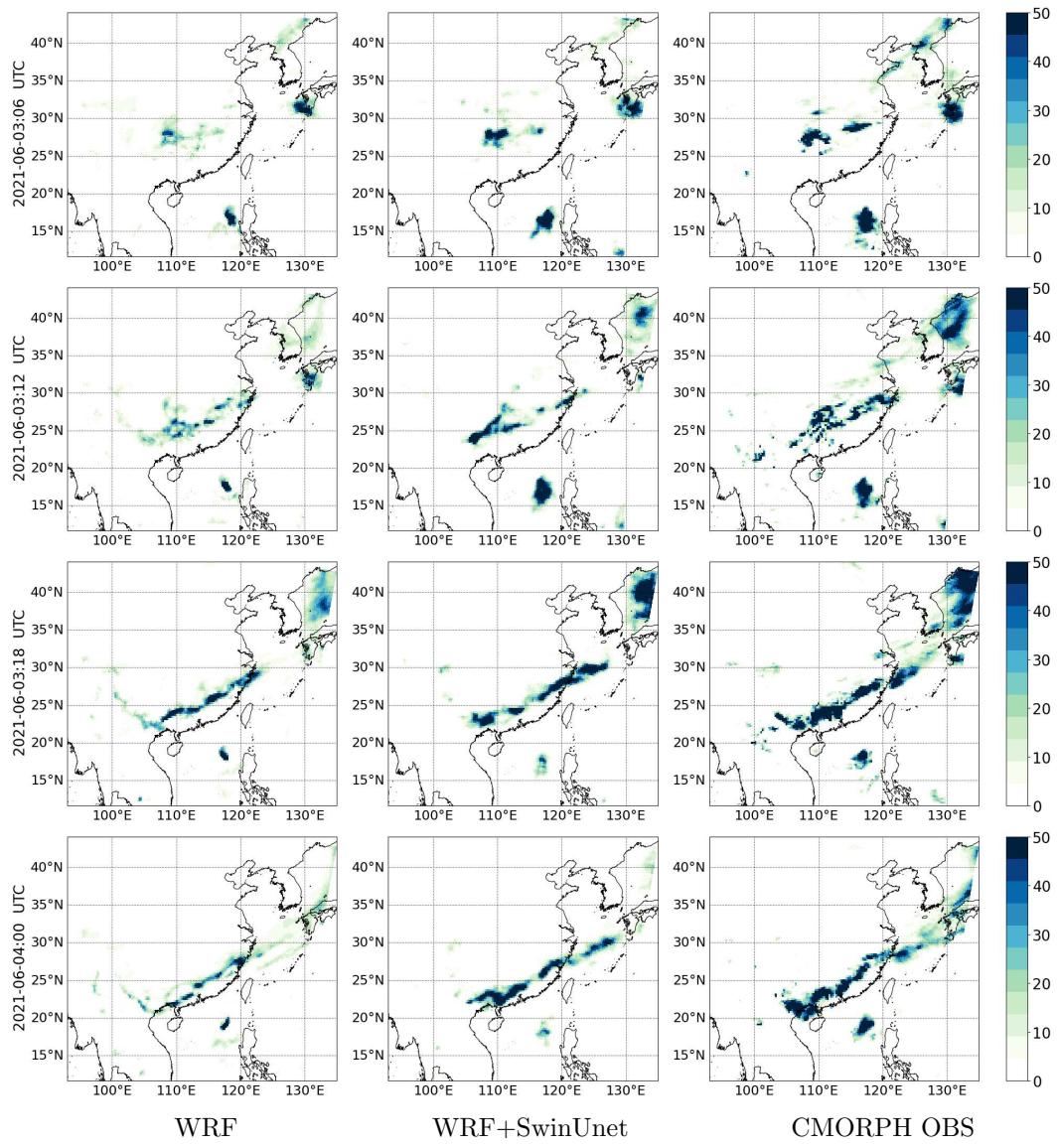
To improve the accuracy of our predictions, we utilized the Swin-Transformer-Unet model in conjunction with the basic meteorological variables predicted by WRF. As a result, the overall precipitation patterns observed are exhibiting more consistency with the CMORPH observational dataset for all four intervals during this case.

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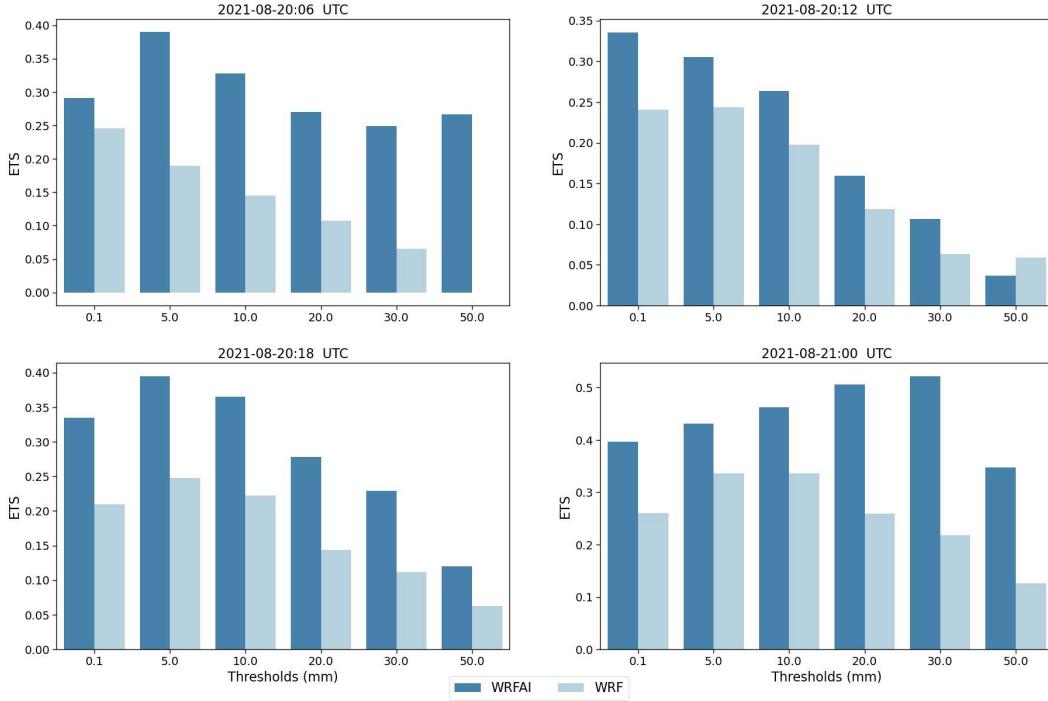
The frontal rain band were initialized around the centre of the study domain (Latitude 27°N Longitude 110°E) before 06 UTC (Fig.8) and started moving southeastward driven by the low-pressure centre located on the northeast corner of the study area, the movement and the structure of the rain band is well preserved in the prediction results at 12 UTC by our deep learning model. In the following sequences of time, this frontal rain band further extended in length and almost covered the whole south and southeast part of China while approaching the coast.

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Based on the quantitative evaluation shown in Fig.7, the performance of the default WRF model and the deep learning model were compared in terms of predicting drizzle and light rainfall with thresholds less than 10mm and heavier rainfall areas with thresholds exceeding 20mm or even 50mm.



**Figure 8.** 6-h accumulated precipitation predicted by WRF, WRF + Swin-Transformer Unet, and observational data from CMORPH datasets on 2021-06-03.



**Figure 9.** Equivalent Threat Score (ETS) of 6-hour accumulated precipitation for the case study on Aug 20th, 2021 at 06:00 UTC, 12:00 UTC, 18:00 UTC, and 24:00 UTC

For predicting drizzle and light rainfall with thresholds less than 10mm within a 6 hours interval, the ETS scores of both models were relatively close, except for the first time interval 06 UTC at 0.1mm thresholds, where the ETS score increased by nearly 40% from 0.34 to 0.5 in the deep learning model.

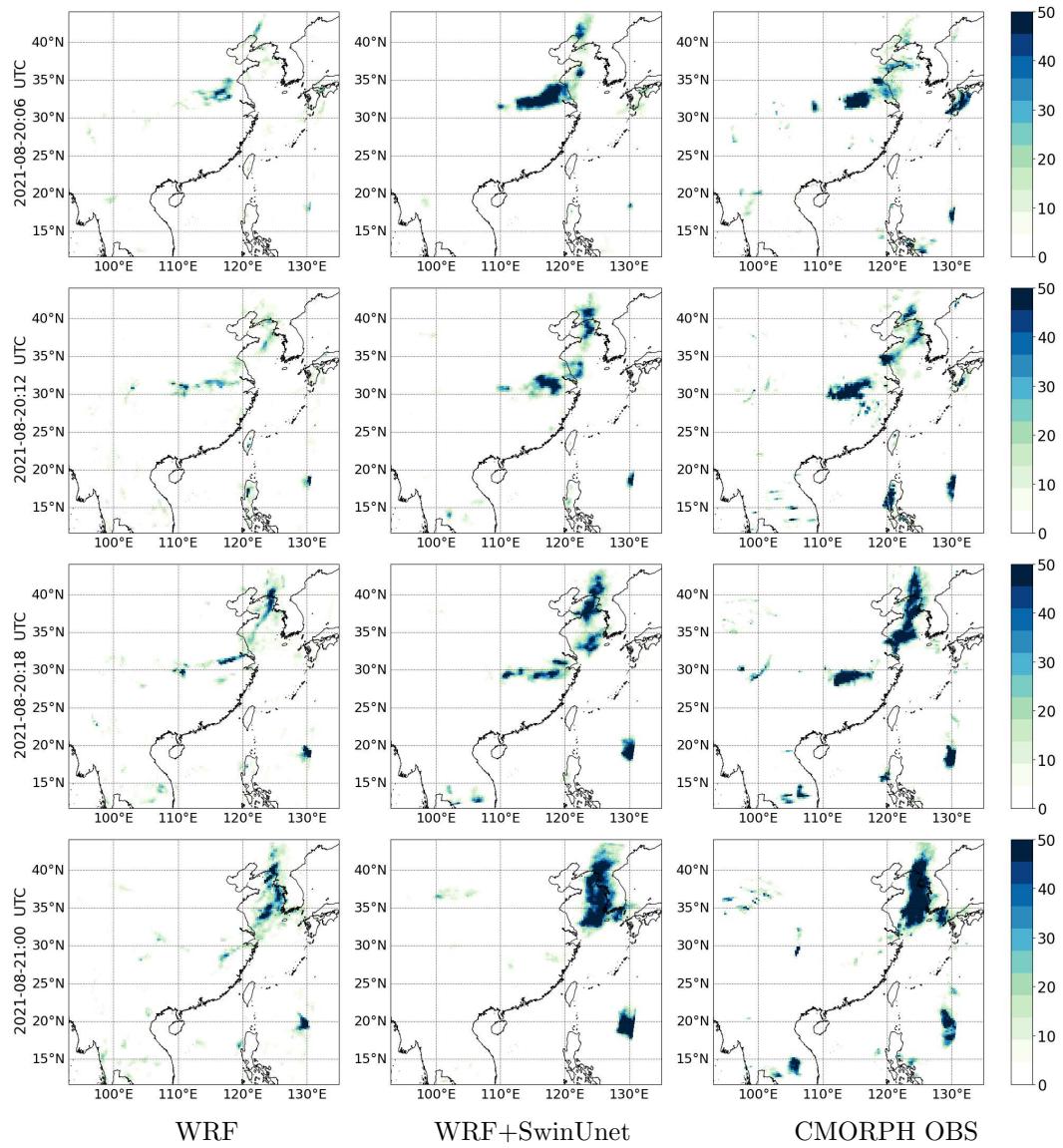
On the other hand, for heavier rainfall areas with thresholds exceeding 20mm or even 50mm, the deep learning model outperformed the default WRF model by doubling or even trebling the ETS score, as indicated by the results at 06 UTC and 12 UTC. Moreover, the decrease of ETS for the deep learning enhanced prediction was less steep, which indicates a more stable performance in estimating precipitation for all ranges compared to the default WRF model.

It is worth noting that NWP models like the WRF model generally suffer from sharp degradation in performance when moving from the synoptic scale to the convective scale. Therefore, the superior performance of the deep learning model in estimating precipitation for heavier rainfall areas suggests its potential for improving the prediction accuracy of convective precipitation in NWP models.

### 3.3.2 2021-08-20:

On August 20th, 2021, Central and Northeast China experienced extreme precipitation, accompanied by thunderstorms and strong convective weather. Heavy rainfall of over 100mm was initially observed before 12:00 UTC (Fig.10) in the Central China region (located at Latitude 33°N and Longitude 115°E). Subsequently, later in the day, extreme was observed in the Yellow Sea, Northeast China, and the Korean Peninsula.

As shown in the precipitation map in Fig.10, our baseline WRF simulation captures only a limited signal of the strong convective rainfall due to the coarse domain grid



**Figure 10.** 6-h accumulated precipitation predicted by WRF, WRF + Swin-Transformer Unet, and observational data from CMORPH datasets on 2021-08-20.

size and the limitations of the parameterization scheme for cumulus and cloud micro-physics. However, by enhancing the estimation of precipitation with our deep learning model, we were able to bridge the gap between the WRF simulation and the observations, caused by intrinsic limitations rooted in the parameterization schemes. This considerably reduces the negative bias and facilitates the estimation of extreme precipitation, both in its maximum precipitation amount and in reducing errors in its spatial distribution.

Fig.9 presents statistical evidence demonstrating the effectiveness of our deep learning framework. The results indicate that the most significant improvement over the baseline WRF simulation occurs at 20th 06 UTC and 21st 00 UTC, where the ETS score increases by an average of 30% for events with light to moderate rainfall (precipitation amount less than 10mm). Additionally, the baseline WRF simulation's performance degrades rapidly as the precipitation threshold increases, failing to detect precipitation exceeding 50mm in a 6-hour interval. In contrast, our deep learning model enhances the estimation results, maintaining a relatively good performance with an ETS score above 0.25 for all thresholds at 06 UTC and 00 UTC, while the baseline WRF model's performance drops to less than 0.1. These results indicate that our deep learning framework provides a significant improvement over the baseline WRF model, even though the grid size and parameterization schemes are not ideally suited for capturing strong convective precipitation during the monsoon season.

## 4 Conclusion and discussion

Regional numerical weather prediction (NWP) models like the WRF model are known to be sensitive to domain grid size (Jee & Kim, 2017) and parameterization schemes (Hasan & Islam, 2018), particularly when it comes to predicting precipitation. To address this challenge, a deep learning model for semantic segmentation using a Swin-Transformer backbone and a hierarchical Unet structure is proposed in this study. This model leverages basic meteorological variables such as air temperature, pressure, wind speed, and humidity to significantly improve the performance of the baseline WRF model in simulating precipitation, particularly for extreme events induced by strong convection. The overall effectiveness of this deep learning post-processing framework is demonstrated through a comprehensive performance evaluation, including an analysis of its spatial and quantile distributions, and a detailed discussion of two case studies.

To evaluate the model's performance, we assessed hourly precipitation amounts across intensity thresholds ranging from 0.1mm to 20mm during the period of June 2021 to September 2021. The results demonstrated that our deep learning model outperformed the baseline WRF simulation for all precipitation intensities. Specifically, the model improved the baseline WRF simulation by 21.7% for light rainfall and drizzle (precipitation amount less than  $1mmh^{-1}$ ), and by 60% for moderate rainfall events with precipitation thresholds of  $3mmh^{-1}$  and  $5mmh^{-1}$ . For heavy rainfall events with hourly precipitation intensity exceeding 10mm, the improvements reflected by the TS and ETS scores reached as high as 50% compared to the baseline WRF. The overall quantile distribution of baseline WRF and the proposed deep learning framework are also compared, with results showing that the prediction of rainfall intensity across all the quantiles received various degrees of improvement. Additionally, the spatial distribution of the 95<sup>th</sup> percentile rainfall intensity and its zonal and meridional averages were also revealing a significantly better alignment with observational data. However, minor challenges were noted in regions with possible orographic precipitation trigger mechanisms, particularly in the mountainous southwest part of China. Future exploratory efforts could be directed towards amplifying the model's proficiency in recognizing and integrating finer-scale terrain and land surface effects. Such advancements could potentially elevate the forecast skill in these currently less confident areas.

434 In addition to the overall evaluation, we presented two case studies of precipitation  
 435 events triggered by different synoptic conditions to demonstrate the model's ability  
 436 to capture complex weather phenomena. For both events, we investigated the 6-hourly  
 437 accumulated rainfall of four intervals and showed how our deep learning model can pro-  
 438 vide more accurate precipitation forecasts by learning from meteorological datasets and  
 439 extracting relevant features.

440 The first precipitation event is caused by the large-scale movement of frontal rain  
 441 bands, while the second event is induced by strong convection during the monsoon sea-  
 442 son. In both cases, we observed rapid degradation of model performance as precipita-  
 443 tion thresholds increased in the baseline WRF model. On the contrary, our deep learn-  
 444 ing model was able to compensate for the insufficient predictability of the baseline WRF  
 445 model simulation and achieve improved ETS scores over each temporal interval at var-  
 446 ious precipitation thresholds. Notably, our model demonstrated particular success in cap-  
 447 turing extreme precipitation amounts exceeding 30mm or 50mm, which are often diffi-  
 448 cult to predict using traditional modelling approaches. These findings demonstrate the  
 449 effectiveness of our deep learning model in capturing precipitation characteristics from  
 450 basic meteorological variables and further quantitatively estimating precipitation based on  
 451 extracted features.

452 As we are feeding the WRF model simulated meteorological fields into our deep  
 453 learning model, the quality of the precipitation estimation results is ultimately depen-  
 454 dent on the quality of our baseline WRF simulation, correspondingly, it should also be  
 455 related to the initial forcing data used to drive the WRF simulation. Therefore, the re-  
 456 sults presented in this study is showing substantial relative improvements against the  
 457 baseline WRF simulation, demonstrating the ability of this neural network in captur-  
 458 ing triggering processes which are not currently described in the existing precipitation  
 459 parameterizations of the WRF model. Moreover, due to the model abstraction, the model  
 460 grid states may not fully match the CMORPH precipitation observations used as labels  
 461 for training, making it challenging to accurately estimate precipitation amounts in terms  
 462 of intensity and location. As a consequence, this spatial and temporal inconsistency of  
 463 prediction and observation was reflected in the increased false alarm rate (FAR) at higher  
 464 thresholds. Similar results were also noted by Hess and Boers (2022), they attributed  
 465 this issue to the localized intermittent nature of the heavy rainfall events. We believe  
 466 that this limitation can be mitigated by accumulating hourly prediction results over sev-  
 467 eral hours, or by adjusting the hyperparameter  $\alpha$  and  $\beta$  in the Tversky loss which con-  
 468 trols the trade-off between precision and recall.

469 Compared to solely optimizing the model with a global MSEloss, several studies  
 470 have attempted to manage the extremely skewed precipitation data by performing log-  
 471 transformation scaling (Shi et al., 2017; Pathak et al., 2022), modifying traditional re-  
 472 gression loss functions such as MAE loss and MSE loss by binning precipitation data and  
 473 assigning different weights to each category (Shi et al., 2017; Franch et al., 2020), or com-  
 474 bining structural similarity measure (SSIM) loss function during optimization (Tran &  
 475 Song, 2019; Hess & Boers, 2022). The usage of the Tversky loss function has also demon-  
 476 strated its superior ability in dealing with strongly imbalanced distributed precipitation  
 477 data in our study. The current optimization process involves pixel-wise optimization of  
 478 the Tversky loss function, followed by using MSE loss for global refinement. Exploring  
 479 custom loss functions that can focus on local features and the overall distribution might  
 480 be beneficial for further improving the temporal and spatial accuracy of the deep learn-  
 481 ing model.

482 Additionally, the data augmentation technique is also worth exploring. The exist-  
 483 ing permutation study by Li et al. (2022) has shown that the moisture-related predic-  
 484 tor dominates the precipitation estimation in this deep learning framework. By using  
 485 feature patch masking, mixing, and shuffling techniques, it may be possible to further  
 486 improve the model's generalizing ability by increasing the difficulty of the original task

487 and enriching the dataset's diversity. This approach can help reduce the model's reliance  
 488 on specific predictors, leading to more robust and accurate predictions.

## 489 5 Open Research

490 The experiments are conducted using Pytorch software v2.0 (<https://pytorch.org/>).  
 491 The precipitation data used for training as references are from CMORPH (<https://www.ncdc.noaa.gov/products/climate-data-records/precipitation-cmorph>). The me-  
 492 teorological feature map used for training is generated from the WRF model (<https://github.com/wrf-model/WRF>), while due to the large size of the training dataset (~300  
 493 G), it is currently archived in HKUST ENVF database (<http://envf.uswt.hk/itf-si/>),  
 494 which can be provided upon request. The source code for preprocessing WRF data, train-  
 495 ing the Swin-Transformer-Unet model and post-processing to generate plots are avail-  
 496 able at (<https://doi.org/10.5281/zenodo.8210356>).  
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