

1           **Frictional Properties of Feldspar-chlorite Altered Gouges and**  
2           **Implications for Fault Reactivation in Hydrothermal Systems**

3   **Zijuan Hu<sup>1,2</sup>, Chongyuan Zhang<sup>2,3,4</sup>, Lei Zhang<sup>5</sup>, Derek Elsworth<sup>6,7</sup>, Quan Gan<sup>8</sup>,**  
4   **Huiru Lei<sup>5</sup>, Manchao He<sup>3</sup>, Leihua Yao<sup>1</sup>**

5   <sup>1</sup>School of Engineering and Technology, China University of Geosciences (Beijing),  
6   Beijing 100083, China

7   <sup>2</sup>Institute of Geomechanics, Chinese Academy of Geological Sciences, Beijing  
8   100081, China

9   <sup>3</sup>School of Mechanics and Civil Engineering, China University of Mining and  
10   Technology, Beijing 100083, China

11   <sup>4</sup>Technology Innovation Center for In-situ Stress, Ministry of Natural Resources,  
12   Beijing 100081, China

13   <sup>5</sup>State Key Laboratory of Earthquake Dynamics, Institute of Geology, China  
14   Earthquake Administration, Beijing, China

15   <sup>6</sup>Department of Energy and Mineral Engineering, EMS Energy Institute and G3  
16   Center, the Pennsylvania State University, PA, 16802, USA

17   <sup>7</sup>Department of Geosciences, the Pennsylvania State University, University Park, PA,  
18   16802, USA

19   <sup>8</sup>School of Resources and Safety Engineering, Chongqing University, Chongqing  
20   400044, China

21   Corresponding author: Chongyuan Zhang ([zhchongyuan@126.com](mailto:zhchongyuan@126.com), [zhangchy@ca](mailto:zhangchy@ca)  
22   [gs.ac.cn](mailto:zhangchy@ca.gs.ac.cn))

23   **Key Points:**

- 24       ● Pure-feldspar gouges change from v-strengthening to v-weakening with  
25       elevated temperature and remain v-weakening at pore pressures
- 26       ● Conversely, feldspar-chlorite altered have reduced strength but change  
27       response from v-weakening to v-strengthening
- 28       ● Frictionally-unstable slip is likely on feldspar gouges but less-likely where  
29       chlorite altered, under P-T conditions of representative of deep geothermal  
30       reservoirs

31

32 **Abstract**

33 As a particularly common minerals in granites, the presence of feldspar and altered  
34 feldspar-chlorite gouges at hydrothermal conditions have important implications in  
35 fault strength and reactivation. We present laboratory observations of frictional  
36 strength and stability of feldspar (K-feldspar and albite) and altered feldspar-chlorite  
37 gouges under conditions representative of deep geothermal reservoirs to evaluate the  
38 impact on fault stability. Velocity- stepping experiments are performed at a confining  
39 stress of 95 MPa, pore pressures of 35-90 MPa and temperatures of 120-400°C  
40 representative of in situ conditions for such reservoirs. Our experiment results show  
41 that the feldspar gouge is frictionally strong ( $\mu\sim 0.71$ ) at all experimental temperatures  
42 ( $\sim 120\text{-}400^\circ\text{C}$ ) but transits from velocity-strengthening to velocity-weakening at  
43  $T>120^\circ\text{C}$ . Increasing the pore pressure increases the friction coefficient ( $\sim 0.70\text{-}0.87$ )  
44 and the gouge remains velocity weakening, but this weakening decreases as pore  
45 pressures increase. The presence of alteration-sourced chlorite leads to a transition  
46 from velocity weakening to velocity strengthening in the mixed gouge at experimental  
47 temperatures and pore pressures. As a ubiquitous mineral in reservoir rocks, feldspar  
48 is shown to potentially contribute to unstable sliding over ranges in temperature and  
49 pressure typical in deep hydrothermal reservoirs. These findings emphasize that  
50 feldspar minerals may increase the potential for injection-induced seismicity on  
51 pre-existing faults if devoid of chlorite alteration.

52 **Plain Language Summary**

53 Granites are an important habitat for deep geothermal reservoirs where the advantages  
54 of low-carbon energy are offset by the potential hazard of injection-triggered  
55 seismicity. Feldspars in these granites are highly susceptible to water–rock  
56 interactions and produce chlorite under appropriate hydrothermal conditions. Many  
57 studies have shown that chlorite itself, as well as its coexistence with other minerals  
58 (not yet including feldspar), can generate earthquakes during fluid-injection into such  
59 reservoirs. We conduct laboratory measurements of frictional properties of simulated  
60 feldspar and chlorite gouges to determine the likelihood of spawning earthquakes. We  
61 show that high temperatures and pore pressures favor unstable slide on feldspar  
62 gouges – and thus the potential to generate earthquakes. However, chlorite as an  
63 alteration product stabilizes faults and reduces the potential for earthquakes. Our  
64 results highlight the importance of feldspar and its alteration in controlling fault  
65 strength and the potential for triggered earthquakes in geothermal reservoirs.

66 **1 Introduction**

67 Fluid injection triggering fault reactivation and thus induced seismicity during  
68 unconventional resource extraction such as for shale gas recovery and geothermal  
69 energy has received widespread attention (Majer et al., 2007; Ellsworth et al., 2013;  
70 Schultz et al., 2020). Large volume and high-rate injections may elevate fluid  
71 pressures and reactivate pre-existing fractures and faults and enhance the potential for  
72 induced seismicity (Hubbert & Rubey, 1959; Bao & Eaton., 2016; Faulkner et al.,  
73 2018; Schultz et al., 2020; Eyre et al., 2019). Thus, understanding mechanisms of

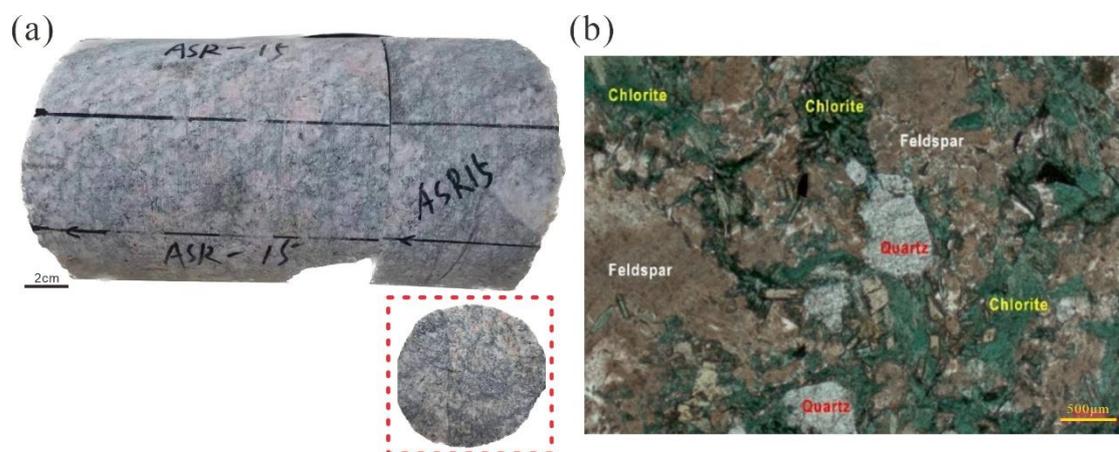
74 injection-induced earthquakes is vital to mitigate the hazard (Hunfeld et al., 2017).  
75 Mature faults are typically gouge filled and thus slip nucleates on this weak  
76 component. The mineralogy of fault gouge exerts a significant control on the  
77 frictional properties (including both strength and sliding stability) and thus on the  
78 potential for induced earthquakes (Scholz, 1998, 2019; Niemeijer & Spiers, 2007;  
79 Ikari et al., 2011). Therefore, frictional properties of fault gouge, primarily the  
80 frictional strength, strength evolution with slip velocity and history (velocity and state  
81 dependence) are of great significance in defining the seismic behavior of faults.  
82 Variation in reservoir temperature as well as fluid pressures impact frictional  
83 properties and are important to define (van der Elst et al., 2013; Faulkner et al., 2018;  
84 Andrés et al., 2019; An et al., 2020).

85 Many previous experimental studies have explored the frictional stability of  
86 faults using both simulated and natural fault gouges spanning broad ranges of  
87 temperature and pressure, including for granites. A common feature of granite gouges  
88 is that frictional properties are strongly influenced by temperature - increasing  
89 temperature results in a transition from velocity weakening to velocity strengthening  
90 (Lockner et al., 1986; Blanpied et al., 1991, 1995). Specifically, the frictional  
91 properties of quartz and feldspar, the main constituent minerals in granites, have also  
92 been the subject of extensive experimental studies. Quartz-rich fault gouge commonly  
93 exhibits high frictional strength ( $\mu \sim 0.7$ ) and promotes velocity strengthening behavior  
94 (Chester & Higgs, 1992; Tembe et al., 2010; Lu & He, 2018; Masuda et al., 2019;  
95 Bedford et al., 2022). The major feldspar-group minerals in the crust are albite  
96 ( $\text{NaAlSi}_3\text{O}_8$ ), anorthite ( $\text{CaAl}_2\text{Si}_2\text{O}_8$ ) and K-feldspar ( $\text{KAlSi}_3\text{O}_8$ ). Among them, the  
97 anorthite- albite series (plagioclase) are the most abundant minerals and may  
98 modulate the response of quartz in granites. To date, few experiments define the  
99 frictional properties of feldspar under hydrothermal conditions typical of the shallow  
100 crust – representative of triggering in the recovery of deep geothermal energy. A few  
101 experimental studies define the frictional strength of plagioclase as similar to quartz,  
102 with a coefficient of friction of  $\sim 0.7$  (Masuda, 2019, 2020). However, compared with  
103 quartz, plagioclase exhibits different frictional stability response. Over the  
104 temperature range 200–500°C, plagioclase is velocity weakening with negative ( $a-b$ )  
105 values (Masuda, 2019, 2020; He et al., 2013) suggesting the potential for seismic  
106 reactivation.

107 In addition to the comminution products from granites, as quartz and feldspar,  
108 metamorphic transformations under long-term hydrothermal conditions produce  
109 important alteration products – again with potentially key controls on frictional  
110 stability (Brown et al., 2003; Okamoto et al., 2019). For example, chlorite is a  
111 widespread product of hydrothermal alteration within subducting oceanic crust and  
112 geothermal reservoirs (Elders et al., 1979; Schiffman & Fridleifsson, 1991; Morrow et  
113 al., 2000; Okamoto et al., 2019), resulting from the alteration of micas or as products  
114 of water–rock interactions (Lucie, 2016; Yuguchi et al., 2015, 2021). Chlorite is  
115 produced by the dissolution of K-feldspar in the presence of weakly alkaline pore  
116 fluids in the temperature range of 210–350 °C (Yuguchi et al., 2021), representative of  
117 deep geothermal reservoirs. Previous laboratory studies of simulated chlorite-rich

118 gouges sheared at elevated temperatures and pore pressures indicate relatively low  
 119 shear strengths (frictional coefficient  $\mu < 0.5$ ) although velocity strengthening  
 120 behavior (Shimamoto & Logan, 1981; Ikari et al., 2009; Okamoto et al., 2019;  
 121 Fagereng & Ikari, 2020; An et al., 2021). Of particular concern is that the alteration  
 122 pathway from feldspar to chlorite is rapid, with the potential to evolve over  
 123 engineering timescales and in fracture systems newly accessed in geothermal  
 124 reservoirs. Therefore, a careful understanding of the frictional properties of pervasive  
 125 feldspar chloritization is necessary in assessing and mitigating potential  
 126 injection-induced seismic risks.

127 Typical geological hosts for deep geothermal reservoirs are granitic and with  
 128 abundant low-grade metamorphic minerals (such as chlorite and epidote) developed  
 129 with in the reservoirs (Okamoto et al., 2019; An et al., 2021). In addition, typical  
 130 enhanced geothermal systems (EGS) reservoirs often host pyrite, indicative as a  
 131 geothermometer of hydrothermal intrusion  $> 250^{\circ}\text{C}$  (Zhang et al., 2022) - with  $250^{\circ}\text{C}$   
 132 present within the stability field for chlorite. In addition, K-feldspar is hydrolyzed in a  
 133 weakly alkaline environment at this temperature and with the potential to form  
 134 chlorite. Such strong feldspar-chlorite alteration is apparent in the Matouying EGS  
 135 reservoir of North China (Figure 1). To constrain the impact of such alteration  
 136 pathways on the stability of faults in such generic granitic reservoirs, we conduct  
 137 friction-stability experiments on feldspar (K-feldspar and albite) and chlorite mixtures  
 138 at elevated temperatures and pore pressures typifying geothermal reservoirs.



139  
 140 **Figure 1.** Photographs of (a) granite core recovered from  $\sim 4$  km depth of the  
 141 Matouying EGS reservoir in Tangshan, North China and (b) photomicrograph  
 142 showing strong alteration of feldspar-chlorite.

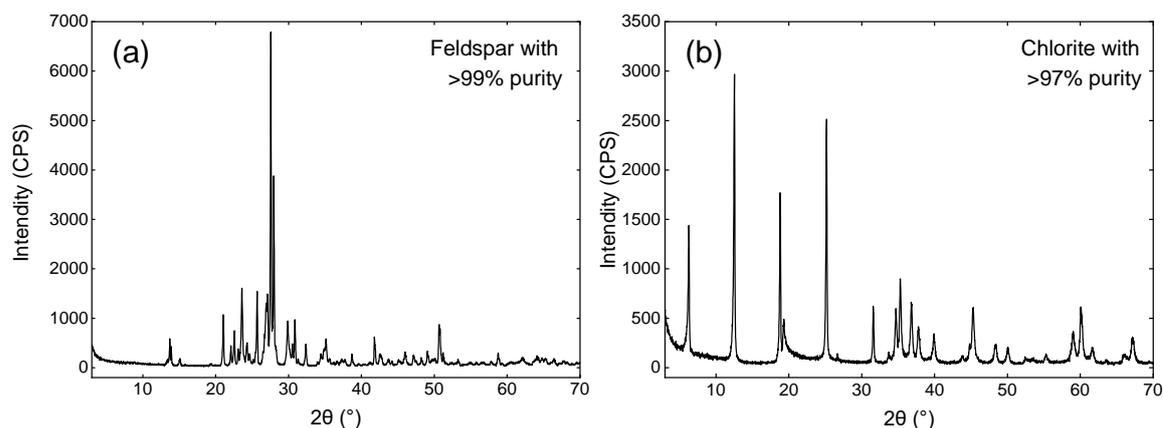
## 143 2 Experimental Methods

144 We complete friction stability measurements on synthetic feldspar gouges  
 145 representing chlorite alteration. Experiments are conducted under recreated  
 146 hydrothermal conditions representative of deep geothermal systems to examine the  
 147 influence of temperature, stress and pore pressures on friction and stability.

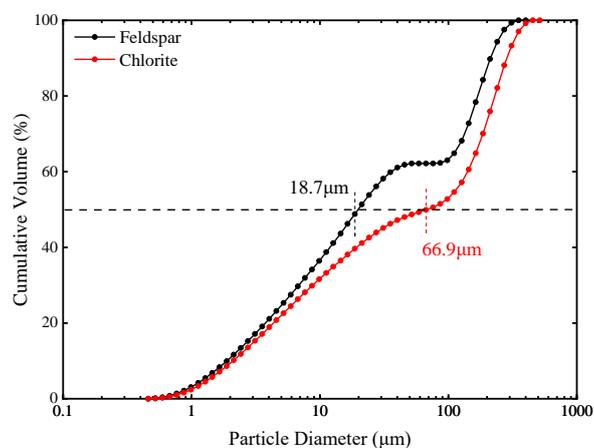
### 148 2.1 Gouge Preparation

149 The feldspar and chlorite minerals used to simulate fault gouges in our

150 experiments were obtained commercially. The mineral compositions of the two  
 151 mineral powders were analyzed by X-ray diffraction (XRD) and shown to be of >99  
 152 wt.% (feldspar) and 97 wt.% (chlorite) purity (Figure 2). The mineral composition of  
 153 feldspar is 53 % K-feldspar and 47 % albite. To simulate fault gouge, both mineral  
 154 powders were crushed and sieved to pass through a #200-mesh sieve. Median particle  
 155 sizes of the feldspar (18.7  $\mu\text{m}$ ) and chlorite (66.9  $\mu\text{m}$ ) (Figure 3) are defined by laser  
 156 classifier. Mixed gouges were then prepared from the feldspar and chlorite minerals in  
 157 the same proportions by weight.



158 **Figure 2.** XRD results of feldspar and chlorite mineral powders used in the simulated  
 159 fault gouges for the experiments. (a) feldspar purity at >99% and (b) chlorite purity  
 160 at >97%.  
 161



162 **Figure 3.** Particle size distributions of feldspar and chlorite gouges for the  
 163 friction-stability experiments. The median particle sizes of the feldspar and chlorite  
 164 gouges are 18.7  $\mu\text{m}$  and 66.9  $\mu\text{m}$ , respectively.  
 165

## 166 2.2 Apparatus and Experimental Procedure

167 Shear experiments were performed using an argon gas-confined triaxial testing  
 168 apparatus (Figure 4) located at the Institute of Geology, China Earthquake  
 169 Administration, Beijing, China (He et al., 2006, 2013). An electro-hydraulic  
 170 servocontrol system drives the axial displacement in this apparatus that can apply a  
 171 confining pressure up to 420 MPa, a temperature up to 600 °C, a fluid pressure up to  
 172 200 MPa, and an axial displacement rate down to  $10^{-2}$   $\mu\text{m/s}$ . We conducted shear  
 173 experiments at temperature ( $T=180$  °C) and pressure ( $P_c=95$  MPa,  $P_f=35$  MPa)

174 conditions relevant to the depth of fluid injection for the 3.7 km Gonghe geothermal  
 175 reservoir. We add deionized water to the mineral powders and bring them to chemical  
 176 equilibrium by stirring for ~10 min. A 1-mm-thick layer of fault gouge was  
 177 sandwiched between gabbro driving blocks 20 mm in diameter and 40 mm in height.  
 178 The sawcut surface of the driving blocks is inclined at 35° to the loading axis. Two  
 179 interconnected 2-mm-microboreholes in the upper driving block provide fluid access  
 180 to the layer of fault gouge, and thus access to pore pressures. The assembled sample  
 181 was jacketed in a 0.35-mm-thick annealed copper sleeve. Double O-rings seal both  
 182 ends of the assembled sample to guard against incursion of the argon gas into the  
 183 gouge layer. A thermocouple monitors the changes in temperature local to the fault  
 184 gouge with the temperature maintained constant within ±1°C by an independent  
 185 controller throughout each experiment. At the initiation of each experiment, we  
 186 increase the confining pressure (argon gas), inject water to increase the pore pressure  
 187 to 2/3 of the target pressure, heat to the target temperature and then fully pressurize to  
 188 the desired value. Fluctuation in confining pressure and pore pressure were  
 189 maintained constant within ±0.3 and ±0.1 MPa, respectively, through two servo-  
 190 controlled intensifiers. All stress and displacement data were recorded at a sampling  
 191 frequency of 1 Hz, except for a sampling frequency of 10 Hz at the maximum loading  
 192 rate.

193 A total of 14 shear experiments were conducted at a constant confining  
 194 pressure ( $P_c = 95$  MPa), different pore pressures ( $P_f = 35, 50, 70$  and  $90$  MPa) and  
 195 different temperatures ( $T = 120, 180, 300,$  and  $400^\circ\text{C}$ ) for monomineralic gouges  
 196 (feldspar) and uniformly mixed feldspar: chlorite gouges (1:1). Experimental details  
 197 are listed in Table 1. The results are representative of EGS development with the  
 198 confining pressure  $P_c = 95$  MPa and pore pressure  $P_f = 35$  MPa corresponding to the  
 199 lithostatic and hydrostatic pressures at ~3.7-km depth (assuming a rock density of  
 200  $2,630 \text{ kg/m}^3$ ) in the EGS reservoir at Gonghe. In addition, the elevated pore pressures  
 201  $P_f = 50\text{--}90$  MPa are consistent with fluid injection during reservoir stimulation. In the  
 202 initial stage of the velocity stepping experiments, the gouge was sheared at a constant  
 203 axial loading velocity of  $0.5 \mu\text{m/s}$  until steady state friction was achieved. Then, the  
 204 axial loading velocity was stepped between  $5, 0.5$  and  $0.05 \mu\text{m/s}$ , corresponding to  
 205 shear velocities of  $6.1, 0.61$  and  $0.061 \mu\text{m/s}$ , to explore the frictional properties of the  
 206 feldspar and mixed feldspar-chlorite gouges.

### 207 2.3 Data Analysis

208 The raw data of the shear experiments are corrected for the decrease in gouge  
 209 contact area with shearing and the shear resistance from the copper jacket. The  
 210 corrected data were then processed to obtain the corrected shear ( $\tau$ ) and normal  
 211 stresses ( $\sigma_n$ ). Frictional strength of the simulated fault gouge is defined by the  
 212 frictional coefficient  $\mu$  as

$$\mu = \frac{\tau}{\sigma_{neff}} = \frac{\tau}{(\sigma_n - P_f)} \quad (1)$$

213 where the  $\sigma_{neff}$  is the effective normal stress and  $P_f$  is the pore pressure.

214 The velocity dependence parameter ( $a-b$ ) was estimated based on rate- and

215 -state-friction (RSF) theory (Dieterich, 1978; Ruina, 1983; Scholz, 1998). In the  
 216 framework of RSF friction, the frictional coefficient  $\mu$  is expressed as

$$\mu = \mu^* + a \ln\left(\frac{V}{V^*}\right) + b \ln\left(\frac{V^* \theta}{D_c}\right) \quad (2)$$

$$\frac{d\theta}{dt} = 1 - \frac{V\theta}{D_c} \quad (3)$$

217 where  $\mu^*$  is the friction coefficient at the reference shear velocity  $V^*$ ,  $a$  is the friction  
 218 parameters reflecting the direct effect,  $b$  is the evolutionary effects of the shear  
 219 velocity transition, with  $D_c$  denoting the critical slip distance over which frictional  
 220 strength evolves to a new steady state. Equation (3) is the common evolution equation  
 221 for the state variable  $\theta$ , namely, the slowness law.

222 At a steady state friction, the state variable  $\theta$  does not change with time  $t$  and  
 223 thus  $d\theta/dt=0$ . Then, combining Equations (2) and (3), yields the frictional stability  
 224 parameter ( $a-b$ ) as

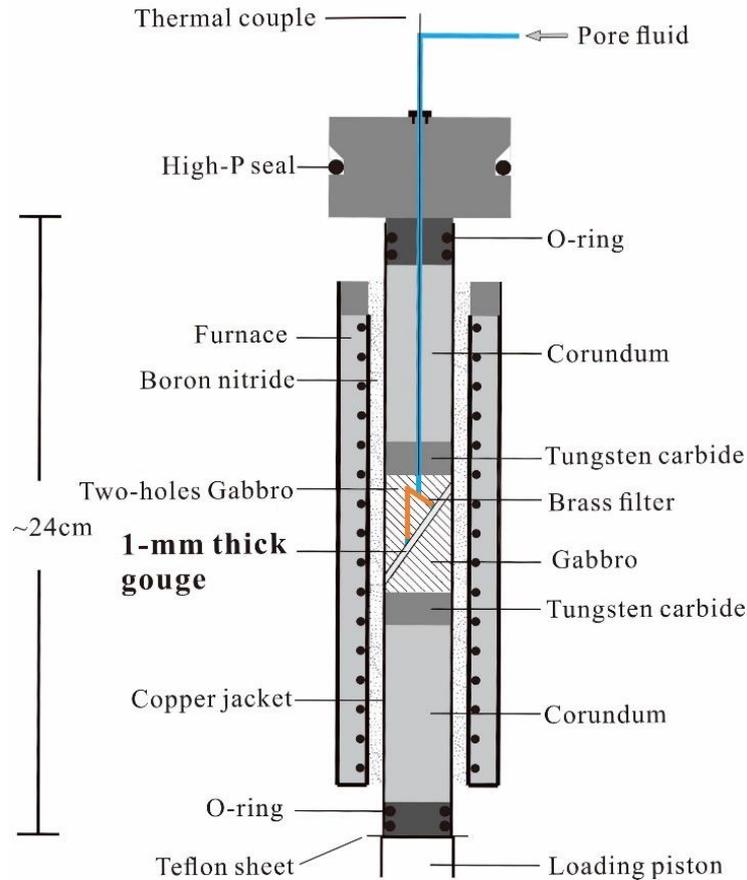
$$a - b = \frac{\mu - \mu^*}{\ln(V/V^*)} \quad (4)$$

225 Positive values of ( $a-b$ ) indicate that frictional coefficient increases with increasing  
 226 velocity, namely, velocity strengthening behavior. Fault gouges with positive ( $a-b$ )  
 227 promote inherently stable sliding and inhibit seismic rupture. Conversely, negative  
 228 values of ( $a-b$ ) indicate velocity weakening behavior and may host unstable fault slip  
 229 if the fault system stiffness falls below the critical stiffness. Velocity-weakening  
 230 behavior offers the potential for stick-slip and may result in earthquake nucleation  
 231 (Marone, 1998; Scholz, 1998). The values of ( $a-b$ ) under stable sliding can be  
 232 recovered directly from the friction- displacement curves. Conversely, the values of  
 233 ( $a-b$ ) for quasi-static oscillations or stick-slip may be obtained from data fitting  
 234 (described in detail in He et al., 2013), as completed here.

235 The potential for induced earthquakes requires that velocity- weakening  
 236 conditions and a fault stiffness below the critical stiffness ( $K_{cr}$ ) are both met. Fault  
 237 instability results when ( $a-b$ )  $< 0$  and the loading stiffness  $K$  is smaller than the critical  
 238 stiffness  $K_{cr}$ , that is,  $K \leq K_{cr}$  (Gu et al., 1984). The critical stiffness of the fault (Ruina,  
 239 1983; He et al., 2013) is defined as

$$K_{cr} = \frac{-(a - b) \cdot \sigma_{neff}}{D_c} \quad (5)$$

240 where  $D_c$  is the characteristic slip distance and  $\sigma_{neff}$  is the effective normal stress.  
 241 Due to the necessary condition of velocity weakening for the unstable sliding of faults,  
 242 ( $a-b$ ) is a topical parameter in shear experiments. Thus, we use the values of ( $a-b$ ) as  
 243 indices to explore the sliding behavior of faults in our study.



244

245 **Figure 4.** Schematic of the high-temperature and -pressure triaxial test apparatus.  
246 Both ends of the copper jacket are sealed with O-rings to prevent argon gas  
247 penetrating into the fault gouge sample. A high-pressure seal on the upper piston was  
248 used to prevent the leakage of the argon gas and retain the seal on the entire assembly.

#### 249 2.4 Microstructural Methods

250 Post-shear, the corundum and tungsten carbide blocks were carefully removed  
251 with the gouge-filled gabbro driving blocks retained in the copper jacket. The samples  
252 were then placed into rubber molds-impregnated with epoxy resin in a vacuum  
253 chamber then cured in an oven at 65°C for ~24 h until the epoxy had completely  
254 hardened. Thin sections were then prepared by slicing the hardened samples along  
255 shear and across the thin axis. All thin sections were polished and coated with carbon.  
256 With the microstructure observed using scanning electron microscopy (SEM).  
257

258 **Table 1.** Experimental matrix and key data

Testing ID	Gouge	$T$ (°C)	$P_c$ (MPa)	$P_f$ (MPa)	$\sigma_{neff}$ (MPa)	$l_{final}$ (mm)
Data set 1: Pure gouge at constant $P_c=95$ MPa						
CK-08	Fsp	120	95	35	118	3.49
CK-07	Fsp	180	95	35	117	3.71
CK-12	Fsp	180	95	50	89	4.04
CK-14	Fsp	180	95	70	52	3.53
CK-10	Fsp	180	95	90	13	3.20
CK-06	Fsp	300	95	35	120	3.32
CK-05	Fsp	400	95	35	120	3.64
Data set 2: Mixed gouges at constant $P_c=95$ MPa						
CK-09	50% Fsp+50% Chl	120	95	35	95	3.83
CK-01	50% Fsp+50% Chl	180	95	35	96	2.28
CK-11	50% Fsp+50% Chl	180	95	50	74	3.58
CK-13	50% Fsp+50% Chl	180	95	70	43	3.14
CK-02	50% Fsp+50% Chl	180	95	90	10	3.36
CK-04	50% Fsp+50% Chl	300	95	35	95	2.93
CK-03	50% Fsp+50% Chl	400	95	35	105	2.86

259 *Note.* Fsp=Feldspar, Chl=Chlorite,  $P_c$ = confining pressure,  $P_f$  = pore fluid pressure,  
 260  $\sigma_{neff}$ = effective normal stress,  $T$  = temperature,  $l_{final}$  = final shearing displacement.

### 261 3 Results

262 We report experiments on both pure-feldspar and feldspar-chlorite altered gouges  
 263 to examine the impact of alteration of friction and frictional stability, and link  
 264 response to microstructural observations.

#### 265 3.1 Feldspar Gouge

266 The friction-displacement curves for the feldspar and mixed gouges initially  
 267 exhibit a linear increase in friction (shear stress to normal stress ratio) with  
 268 displacement followed by inelastic yield point then steady state friction at a shearing  
 269 displacement of  $\sim 1.83$  mm. This is in turn followed by slight strain hardening  
 270 response until the final shear displacements of 3–4 mm (Table 1 and Figure S1 in  
 271 Supporting Information). The simulated feldspar gouge shows stable sliding at  
 272  $T=120^\circ\text{C}$  and  $P_f=35$  MPa (Figure S1a in Supporting Information), while stick-slips  
 273 were observed at a shear velocity of  $0.061 \mu\text{m/s}$  at  $T = 180^\circ\text{C}$  (Figure S1b in  
 274 Supporting Information). As shown in Figure 4, the feldspar gouge exhibits a stable  
 275 sliding at  $T>180^\circ\text{C}$ . A small oscillation lasting for  $\sim 0.1$  mm of displacement occurs  
 276 when the axial shear rate is switched to  $5 \mu\text{m/s}$  at  $T=300^\circ\text{C}$  (Figure S1f in Supporting  
 277 Information). The friction-displacement curves of the feldspar fault gouge show stable  
 278 sliding and a strong dependence of friction coefficient on velocity (Figure S1g in  
 279 Supporting Information). We measured the steady-state-friction coefficient at a shear  
 280 displacement of  $\sim 1.83$  mm for each experiment. Our results show that the friction

281 coefficient of feldspar gouge is  $\sim 0.71$  over the range of experimental temperature  
 282 (120–400°C), with no significant change with increasing temperature (Figure 5a). In  
 283 contrast, the coefficients of friction  $\mu$  for the simulated feldspar gouge at varied pore  
 284 pressure are in the range  $\sim 0.70$ – $0.87$  and increase with pore pressure (Figure 5b).

285 Frictional stabilities ( $a-b$ ) for the simulated feldspar gouge were obtained over  
 286 the range of experimental temperature and pore pressure conditions (Table 2). A  
 287 transition from velocity strengthening behavior ( $a-b = 0.00016$  to  $0.00163$ ) to velocity  
 288 weakening behavior ( $a-b = -0.00098$ ) can be identified at  $T = 120$ – $180$  °C and  $P_f = 35$   
 289 MPa. This is then followed by a transition to velocity strengthening at  $T = 300$  °C ( $a-b$   
 290  $= 0.00108$ ) and finally to velocity weakening at  $T = 400$  °C ( $a-b = 0.00101$ ). This  
 291 indicates that, temperature has a significant impact on stability of the feldspar gouge  
 292 (Figure S2c in Supporting Information). At  $T = 180$  °C and  $P_f = 35$ – $90$  MPa, the feldspar  
 293 gouge exhibits velocity weakening behavior ( $a-b = -0.00098$  to  $-0.00013$ ). The  
 294 velocity-weakening behavior of feldspar gouge decreases with increasing pore  
 295 pressure at  $P_f \geq 35$  MPa (Table 2 and Figure 5d). Moreover, the effective normal stress  
 296 also significantly affects the velocity dependence of frictional stability. A higher  
 297 effective normal stress returns a higher absolute value of ( $a-b$ ) that increase nearly  
 298 linearly with increasing effective normal stresses at  $T = 180$  °C (Figure 5e). In  
 299 summary, higher effective normal stresses promote velocity-weakening behavior at  
 300 constant temperature. In addition, we analyzed the variation of friction stability with  
 301 shear velocity- the feldspar gouge exhibits a velocity-weakening behavior ( $a-b =$   
 302  $-0.00270$  to  $-0.00090$ ) at lower shear velocities in the range of the experimental  
 303 temperature ( $T = 180$ – $400$  °C) and pressures (Figure 5a-b).

### 304 3.2 Mixed Feldspar-Chlorite Gouges

305 We evaluate the competing influences of feldspar and the alteration mineral  
 306 chlorite on the frictional stability of the gouge. The shear experiments were conducted  
 307 on the feldspar/chlorite mixed gouges at constant confining pressure  $P_c = 95$  MPa,  
 308 pore pressure  $P_f = 35$ – $90$  MPa, and temperature  $T = 120$ – $400$  °C. After  $\sim 1.83$  mm of  
 309 shear displacement, the friction- displacement curves exhibit a slight strain  
 310 strengthening and stable sliding to a final shear displacement of 3–4 mm (Figure S2 in  
 311 Supporting Information). The mixed gouges slide stably over the range of  
 312 experimental temperature  $T = 120$ – $300$  °C (Figure S2a, S2b and S2f in Supporting  
 313 Information). At  $T = 400$  °C and for an axial loading rate  $V = 0.5$   $\mu\text{m/s}$ , small  
 314 oscillation in friction result, lasting  $\sim 0.2$  mm, followed by stable sliding (Figure S2g  
 315 in Supporting Information). The method for determining the frictional coefficient of  
 316 the mixed gouges is identical to that for the feldspar gouge. The results show that the  
 317 mean frictional coefficient for the mixed gouges is  $\sim 0.55$ , with the coefficient of  
 318 friction  $\mu$  at  $T = 400$  °C deviating significantly from that at  $T \leq 300$  °C (Figure 5a).  
 319 The frictional strength increases with increasing pore pressure and the friction  
 320 coefficient is less than that for the feldspar gouge over the range of experimental pore  
 321 pressures (Figure 5b).

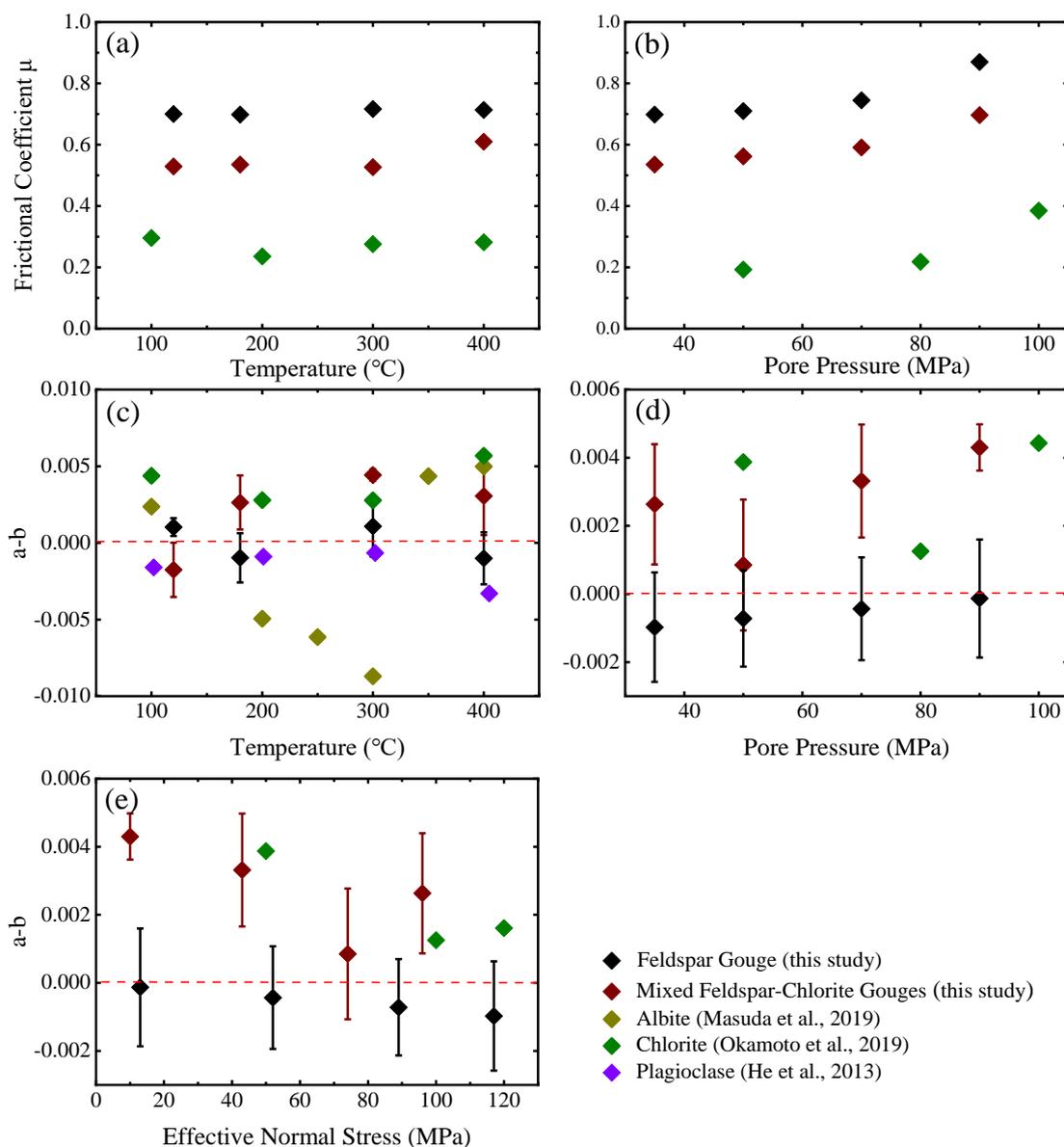
322 Trends in frictional stability ( $a-b$ ) with applied temperature and pore pressure  
 323 for the feldspar-chlorite mixed gouges are illustrated in Figure 5 and Table 2. This

324 sample exhibits a transition from velocity weakening at  $T = 120^\circ\text{C}$  and  $P_f = 35$  MPa  
 325 ( $a-b = -0.00176$ ) to slight velocity strengthening at  $T \geq 180^\circ\text{C}$  ( $a-b = 0.00263$  to  
 326  $0.00443$ ). Frictional stability ( $a-b$ ) increases with increasing temperature when the  
 327 axial displacement rates were stepped between  $0.5$  and  $0.05 \mu\text{m/s}$ . Moreover, the  
 328 value of ( $a-b$ ) shows a variety of trends as the axial displacement rates were stepped  
 329 between  $0.5$  and  $5 \mu\text{m/s}$  (Figure 6). We compared velocity- dependent coefficient of  
 330 friction for both the feldspar and mixed gouges, with the mixed gouges stronger than  
 331 the feldspar at  $T > 120^\circ\text{C}$  (Figure 5c). The friction and stability of mixed gouges vary  
 332 little with porosity at  $T = 180^\circ\text{C}$  and  $P_f = 35\text{--}50$  MPa - however, ( $a-b$ ) values increase  
 333 with increasing pore pressure at  $P_f \geq 50$  MPa. The frictional properties of the  
 334 uniformly mixed gouges are significantly different from those of feldspar alone and  
 335 change the trend as a function of pore pressure (Figure 5d). The velocity dependence  
 336 parameter ( $a-b$ ) of the mixed gouges first decreases and then increases with an  
 337 increase in effective normal stress, with  $74$  MPa as the transition stress (Figure 5e).  
 338 The values of ( $a-b$ ) decreases with increasing effective normal stress ( $\sigma_{neff}$ ) at  $T =$   
 339  $180^\circ\text{C}$  and  $\sigma_{neff} < 74$  MPa. This can be interpreted as the lower-bound effective  
 340 stress promoting velocity-strengthening behavior of the mixed fault gouge at constant  
 341 temperature. Moreover, the sliding rate influences the frictional stability of the mixed  
 342 gouges. The mixed feldspar-chlorite gouges exhibit velocity-strengthening behavior  
 343 within the range of experimental temperature and pressures, and ( $a-b$ ) reduces with an  
 344 increase in shearing velocity (Figure 6c-d).

345 **Table 2.** Results of Shear Experiments

Test ID	$\mu_{ss}$	$a-b$ values at ( $\mu\text{m/s}$ )			
		6.1–0.61	0.61–0.061	0.061–0.61	0.61–6.1
CK-08	0.6998	0.00137±0.00046	0.00034±0.00016	0.00078	0.00163
CK-07	0.698	-0.00059±0.0003	-0.00145±0.00077	-0.00285	0.00099
CK-12	0.71	0.00051±0.00082	-0.0009±0.00014	-0.00264±0.00015	0.00016
CK-14	0.7451	0.00085±0.00039	-0.00095±0.00007	-0.00234±0.00057	0.0007
CK-10	0.8696	0.00045±0.00003	0.00165±0.00013	-0.00247±0.00021	-0.00016
CK-06	0.7161	0.00395	-0.00136	0.00022	0.00309
CK-05	0.7132	-0.00155±0.0005	-0.00112±0.00005	-0.0027	0.00134
CK-09	0.5291	0.00075±0.0007	-0.00336±0.00101	-0.0002±0.00063	-0.00242
CK-01	0.5352	0.00062	0.00391	0.00337	-
CK-11	0.5616	-0.00108±0.00241	0.00132±0.00059	0.00332±0.00084	-0.00016
CK-13	0.5909	0.00187±0.00105	0.00493±0.00021	0.00456±0.00061	0.0019
CK-02	0.6961	0.00385±0.00273	0.00531±0.0017	0.00407±0.00209	0.00397
CK-04	0.5264	0.00465±0.00043	0.00478	0.00434	0.00394
CK-03	0.6094	0.00311±0.0002	0.00565	0.00391	-0.00043

346 *Note.* The ( $a-b$ ) values are the average values from the same velocity steps in each  
 347 test with the error calculated from the standard deviation. Values of ( $a-b$ ) at each  
 348 shear velocity are also shown in Table S1 in the Supporting Information S1. The first  
 349 velocity step ( $0.61\text{--}6.1 \mu\text{m/s}$ ) was excluded in the calculation.  $\mu_{ss}$  is the coefficient of  
 350 friction determined at  $\sim 1.83$  mm shear displacement.  
 351



352

353

354

355

356

357

358

359

360

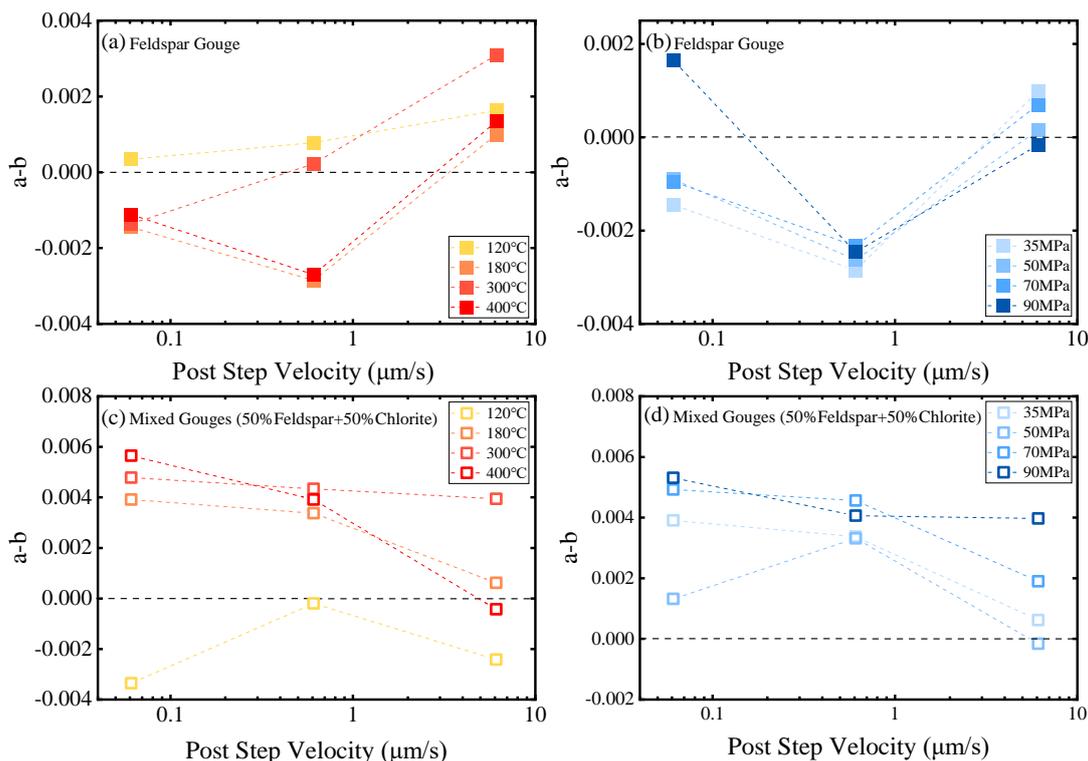
361

362

363

364

**Figure 5.** Coefficient of friction  $\mu$  and frictional stability ( $a-b$ ) for feldspar and feldspar-chlorite mixed gouges at varied conditions. (a) Values of  $\mu$  versus temperature of pure feldspar gouge and mixtures at  $P_c = 95$  MPa and  $P_f = 35$  MPa. (b) Relationship between friction coefficient and pore pressure of pure gouge and mixtures at  $P_c = 95$  MPa and  $T = 180$   $^{\circ}\text{C}$ . (c) ( $a-b$ ) as a function of temperatures for pure feldspar gouge and mixtures at  $P_c = 95$  MPa and  $P_f = 35$  MPa. The gray-green, green, and purple diamonds represent the results from Masuda et al. (2019), Okamoto et al. (2019) and He et al. (2013), respectively. (d) ( $a-b$ ) as a function of pore pressure of pure gouge and mixtures at  $P_c = 95$  MPa and  $T = 180$   $^{\circ}\text{C}$ . (e) ( $a-b$ ) as a function of effective normal stress of pure gouge and mixtures at  $P_c = 95$  MPa and  $T = 180$   $^{\circ}\text{C}$ .



365

366 **Figure 6.** Frictional stability ( $a-b$ ) for the feldspar and feldspar/chlorite mixed gouges  
 367 at various post step velocities.

368

### 3.3 Microstructural Observations

369

370

371

372

373

374

375

376

377

378

379

380

381

382

383

384

385

386

387

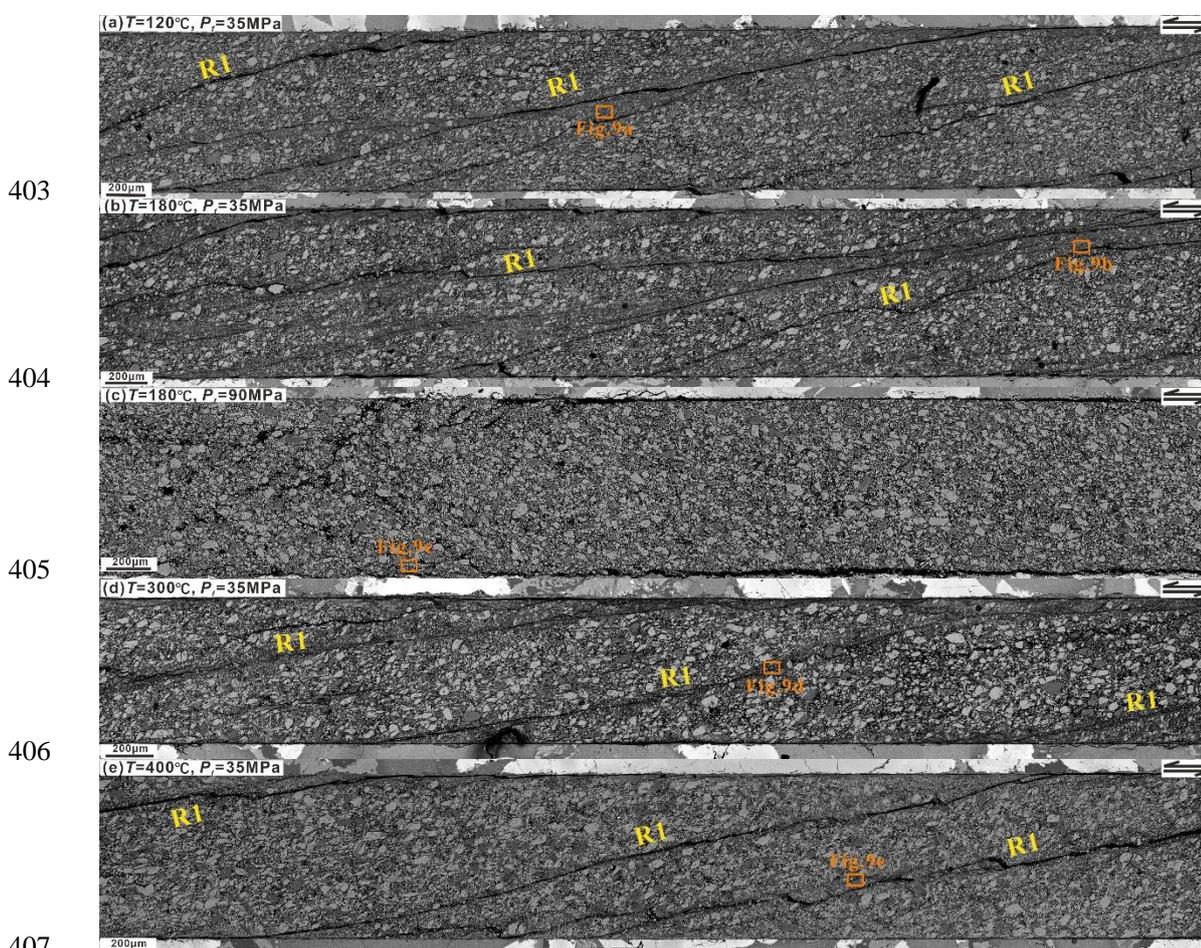
388

389

Following the method described in Section 2.4, we performed microstructural observations on the deformed samples using scanning electron microscopy (SEM), with the methods of Logan et al. (1992) adopted to define the resulting fabric. Among the post-experiment samples, disassembly destroyed the feldspar samples at  $P_f = 50$  and 70 MPa and the mixed gouges at  $P_f = 70$  MPa and  $T = 300^\circ\text{C}$ . The fabrics of the gouge samples differ characteristically for the various gouge frictional responses (velocity-strengthening / velocity-weakening). Among them, grain size reduction appears to control the velocity-weakening behavior, and instability (Zhang et al., 2016; Casas et al., 2023). The microstructures of the deformed gouges are illustrated in Figures 9–10. At  $P_f = 35$  MPa and  $T = 120$ – $400^\circ\text{C}$ ,  $R_1$  angle shears commonly develop in the feldspar gouge (Figure 7). The fault gouge exhibits velocity-weakening behavior with a narrow-localized shear zone with significant local grain crushing at  $T = 180^\circ\text{C}$ . Cracks tortuous penetrate throughout the fault zone (Figure 7b), although the number density of  $R_1$  shear diminish with increasing temperature. At  $T = 300$ – $400^\circ\text{C}$ , there are few  $R_1$  fractures and local shear zones are inconspicuous that penetrate the fault zone in the gouge layer. When the temperature is increased to  $400^\circ\text{C}$ , the fault gouge exhibits a velocity weakening response and the extent of shear crack deflection is significantly higher than at  $180^\circ\text{C}$  (Figure 7e). The number of  $R_1$  shears in the feldspar gouge decreases with increasing pore pressure at  $T = 180^\circ\text{C}$  and  $P_f = 35$ – $90$  MPa. At  $P_f = 90$  MPa, uniformly shears evolve with pervasive particle crushing ( $< 50 \mu\text{m}$ ) for the deformed feldspar gouges (Figure 9), although no obvious

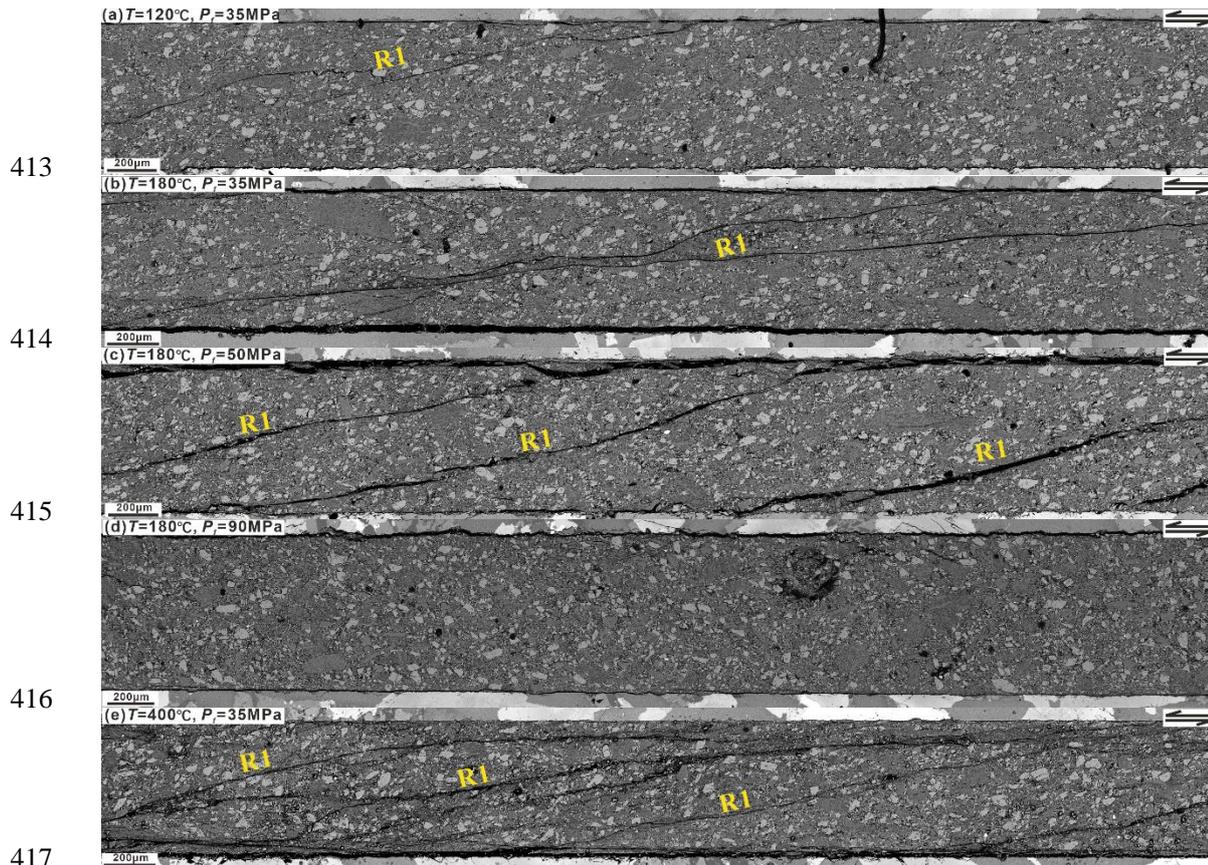
390  $R_1$  shear zone penetrates the fault gouge layer (Figure 7c). In addition, the gouge  
 391 particles are sheared and uniformly crushed, with particle sizes ranging from a few  
 392 micrometers to tens of micrometers at high pore pressure ( $P_f=90$  MPa), i.e., at low  
 393 effective stresses.

394 The chlorite particles are more fully comminuted than the feldspar particles in  
 395 the mixed gouge (Figure S3 in Supporting Information). The number of  $R_1$  shears in  
 396 the mixed gouges increases with temperature at  $P_f=35$  MPa and  $T=120$ – $400^\circ\text{C}$  (Figure  
 397 8), and the fault exhibits velocity-weakening behavior at  $T=120^\circ\text{C}$ . At  $T=180$ – $400^\circ\text{C}$ ,  
 398 the gouge layer exhibits multiple cross-distributed groups of  $R_1$  shears. The number of  
 399  $R_1$  shear decreases with increasing pore pressure and a few parallel  $R_1$  shears develop  
 400 in the fault zone at  $T=180^\circ\text{C}$  and  $P_f=50$  MPa (Figure 8c). However, the particles in  
 401 the mixed fault gouge are sheared and crushed, with particle sizes ranging from few  
 402 micrometers to nearly a hundred micrometers at  $P_f=90$  MPa (Figure 8d).

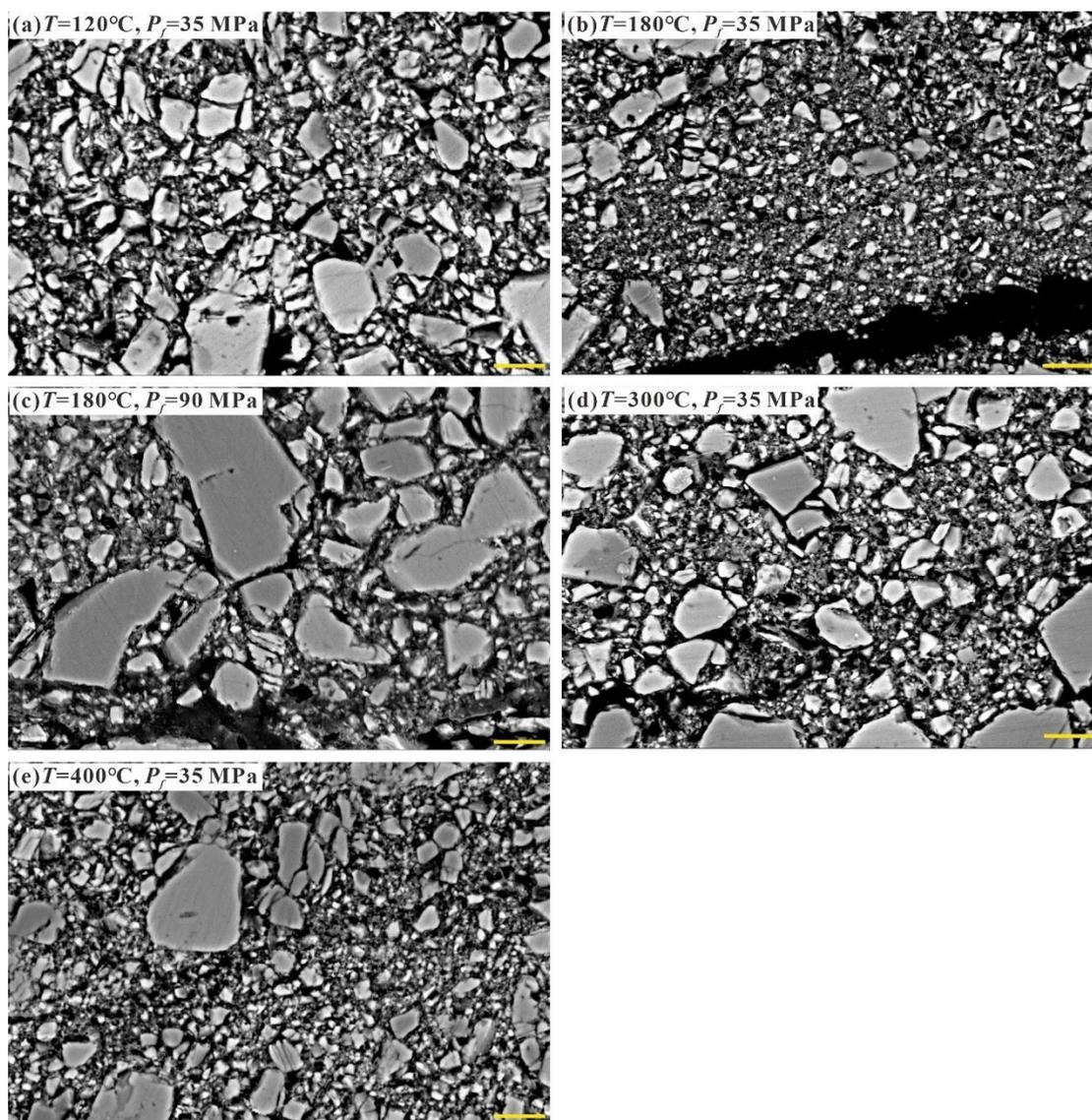


407 **Figure 7.** Microstructure (backscattered images) of the deformed feldspar gouge post-  
 408 experiment. Refer to Logan et al. (1992) for relevant terms used to describe  
 409 microscopic features. We observed local shear zones (LSZ) and  $R_1$  shear (Riedel shear)  
 410 in partial samples.  
 411

412



418 **Figure 8.** Microstructure (backscattered images) of the deformed feldspar/chlorite  
419 mixed gouges post-experiment. Refer to Logan et al. (1992) for relevant terms used to  
420 describe microscopic features. Experimental conditions: (a)  $T=120\text{ }^{\circ}\text{C}$ ,  $P_f=35\text{ MPa}$ , (b)  
421  $T=180\text{ }^{\circ}\text{C}$ ,  $P_f=35\text{ MPa}$ , (c)  $T=180\text{ }^{\circ}\text{C}$ ,  $P_f=50\text{ MPa}$ , (d)  $T=180\text{ }^{\circ}\text{C}$ ,  $P_f=90\text{ MPa}$ , (e)  
422  $T=400\text{ }^{\circ}\text{C}$ ,  $P_f=35\text{ MPa}$



423

424 **Figure 9.** Particle fragmentation characteristics near the shear zone in the feldspar  
425 gouge after shearing. Among them, (b), (c) and (e) exhibit unstable sliding in velocity  
426 weakening. Scale bar: 5  $\mu\text{m}$ .

#### 427 **4 Discussion**

428 We explore impacts and mechanisms of temperature and pore pressure on both  
429 pure-feldspar (Section 4.1) and mixed feldspar-chlorite gouges (Section 4.2) in  
430 controlling friction and frictional stability. We use these observations and projected  
431 mechanisms to define implications for injection-induced seismicity (Section 4.3) in  
432 such common feldspar-chlorite altered systems.

##### 433 4.1 Feldspar Stability Dependence on Temperature and Pore Pressure

434 The magnitudes of temperature and pore pressure are both important factors  
435 influencing fault instability (Blanpied et al., 1995). There is no significant variation in  
436 the frictional strength of the feldspar gouge with increasing temperature in our

437 experiments. The average value of frictional coefficient for the feldspar gouge is  $\sim 0.7$ ,  
438 consistent with previous experimental studies (Masuda, 2019, 2020; He et al, 2013).  
439 At  $P_f=35$  MPa and  $T=120\text{--}400$  °C, the feldspar gouge exhibits an evolving velocity  
440 dependence- velocity-strengthening behavior at  $T=120$  °C and  $300$  °C and conversely  
441 velocity-weakening behavior at  $T=180$  °C and  $400$  °C (Figure 5c). We also observe  
442 the apparent localized shear zones at this transition. The grain sizes in the localized  
443 shear zones significantly decreases increasing in the surface areas of the gouge and  
444 promoting unstable sliding (Bedford & Faulkner, 2021). Moreover, the feldspar gouge  
445 also shows a different transition in velocity dependence with the shear velocity  
446 stepping. For axial loading rates switched between  $0.5$   $\mu\text{m/s}$  and  $0.05$   $\mu\text{m/s}$ , the  
447 frictional stability coefficients ( $a-b$ ) of the gouge decrease with increasing  
448 temperature. The transition from velocity-strengthening to velocity-weakening in our  
449 shear experiments can commonly be explained by a microphysical model (Niemeijer  
450 & Spiers, 2007; den Hartog et al., 2012; Chen et al., 2015, 2016). The model assumes  
451 that the strength and velocity dependence of the gouge is regulated by the competition  
452 between an expansion velocity (due to intergranular particle flow) and a compaction  
453 velocity (due to thermally activated pressure solution). The velocity of pressure  
454 solution increases with an increase in temperature and the gouge is further compacted.  
455 When the compaction velocity of the gouge particles is approximately equal to the  
456 expansion velocity then velocity-weakening behavior results (Verberne et al., 2020;  
457 Chen et al., 2020). Although we did not observe remarkable features indicating  
458 pressure solution for our range of experimental temperatures, previous experimental  
459 studies have indeed indicated the presence of intergranular pressure solution in  
460 feldspar minerals at higher temperatures (He et al., 2013). Thus, our experimental  
461 results are consistent with such a model. Our experimental results show a  
462 velocity-strengthening behavior at  $T=300$ °C which transforms to velocity weakening  
463 at  $T=400$ °C (Figure 5c), which is inconsistent with the above microphysical model.  
464 However, we note that our experimental conditions of  $T=300$  and  $400$  °C and  $P_f=35$   
465 MPa approach those for supercritical water ( $T=374$  °C,  $P_f=22.1$  MPa; Maxim et al.,  
466 2021). Solubility of solutes and reaction rates in the supercritical state are  
467 significantly elevated (Weingärtner & Franck, 2005) and thus may impact  
468 intergranular pressure solution. Therefore, when the temperature exceeds the critical  
469 point between  $300\text{--}400$ °C, the density of the water will change drastically (Sakuma &  
470 Ichiki, 2016) and may promote an elevated rate of pressure solution of feldspar in the  
471 range of  $300$ °C to  $400$ °C. This, combined with the microstructural observation that  
472 localized shear zones disappear at  $T=300$  °C – suggests that feldspar undergoes a  
473 decomposition reaction and produces other alteration minerals such as chlorite at  
474  $T=300$  °C, and thus changes the chemical composition of the gouge (Yuguchi et al.,  
475 2021), which then affects the frictional behavior. However, this conjecture requires  
476 further analysis of the post-experimental sample composition.

477 Our experimental results show that the effect of pore pressure on the frictional  
478 stability of gouge exhibits specific trends at constant temperature (Figure 5d). The  
479 frictional strength of feldspar gouge increases with increasing pore pressure, showing  
480 velocity weakening behavior, but the velocity dependence coefficient ( $a-b$ ) increases

481 linearly with increasing pore pressure. Previous systematic studies of quartz solubility  
 482 have shown that quartz solubility increases with increasing pore pressure at the  
 483 identical temperature (Anderson & Burnham, 1965), and feldspar follows a similar  
 484 pro-grade solubility trend. Thus, the rate of feldspar pressure solution increases with  
 485 increasing pore pressure. This is consistent with the above microphysical model. That  
 486 is, that the pressure solution of feldspar becomes more vigorous and effective as a  
 487 dissolution-diffusion process with an increase in pore pressure at the same  
 488 experimental conditions- thus increasing the intergranular compaction velocity and  
 489 decreasing porosity. As the intergranular contact area decreases, the gouge frictional  
 490 strength increases and the value of  $(a-b)$  also gradually increases. Few distinct  $R_1$   
 491 shear zones are observed in the microstructures of our experiments at  $P_f = 90$  MPa,  
 492 rather, exhibiting uniform shear fragmentation. The deformation of the gouge is  
 493 dominated by rolling at high pore pressure and low effective stresses, and thus brittle  
 494 rupture processes may be masked by the disappearance of localized shear zones in the  
 495 gouge. Thus, the feldspar gouge exhibits velocity-weakening behavior at high pore  
 496 pressure, but with no obvious local shear zones observed in the microstructure.

#### 497 4.2 Influence of Mineral Composition on Fault Friction

498 Fault may transition from strengthen to weaken over long fluid-injection  
 499 timescales at elevated temperatures, due to secondary mineral precipitation of  
 500 hydrothermal alteration (Jeppson et al., 2023). Chlorite, as a low frictional strength  
 501 altered mineral (Ikari et al., 2009; Okamoto et al., 2020; An et al., 2021) reduces the  
 502 overall frictional strength of the mixed gouges. The frictional coefficients of the  
 503 feldspar-chlorite mixed gouge were significantly lower than those of the feldspar  
 504 gouge over the range of experimental temperatures. There is no clear evolution of the  
 505 frictional coefficient of the mixed gouge with increasing temperature at  $T \leq 300$  °C  
 506 (Figure 5a). Chlorite has a low coefficient of friction ( $\mu = 0.25-0.30$ ) over our range  
 507 of experimental temperature ( $\leq 400$  °C) and does not change significantly with  
 508 temperature (Okamoto et al., 2019, 2020). Therefore, we speculate that the sudden  
 509 increase in the frictional coefficient of the mixed gouges from 0.55 to 0.61 at  $T =$   
 510  $400^\circ\text{C}$  is due to the impact of K-feldspar chloritization (Yuguchi et al., 2021). The  
 511 frictional strengths of the feldspar-chlorite mixed gouges increase with increasing  
 512 pore pressure. Moreover, the frictional strength of the feldspar gouge varies  
 513 consistently for both contents at different temperatures and pore pressures. Our results  
 514 show that the presence of chlorite in the gouge overall reduces the strength but does  
 515 not change the trend in frictional strength. The velocity dependence of the mixed  
 516 gouges over the range of experimental temperatures reveals that the mixed gouges  
 517 exhibit both velocity-weakening and velocity-strengthening behavior, and that the  
 518 velocity dependence parameter  $(a-b)$  increases with increasing temperature.

519 Our results indicate that the effect of different pore pressures on the frictional  
 520 properties of feldspar-chlorite mixed gouges is similar to that of temperature. The  
 521 frictional strength of the mixed gouges is significantly lower than that of the feldspar  
 522 gouge and increases with increasing pore pressure. There is a consistent trend of  
 523 frictional strength of gouges with pore pressure for the two feldspar contents (1:0 and

524 1:1). The feldspar-chlorite mixed gouges show velocity-strengthening behavior that  
525 increases with increasing pore pressure. The variation of (*a-b*) values with pore  
526 pressure is nearly linear at  $P_f \geq 50$  MPa (Figure 5d). The frictional stability of the  
527 mixed gouges is significantly enhanced compared to the feldspar gouge due to the  
528 presence of chlorite. Combining with the observations of microstructural  
529 characteristics of the deformed samples, we find that the number of  $R_1$  shear decreases  
530 with increasing pore pressure. Moreover, compared with the frictional stability of  
531 chlorite (Okamoto et al., 2019), albite (Masuda et al., 2019) and plagioclase (He et al.,  
532 2013) from previous experiments, the presence of feldspar minerals reduces the  
533 frictional stability of the mixed gouges (Figure 5c-e). The stability of the feldspar  
534 gouge increases with an increase in pore pressure.

### 535 4.3 Implication for Injection-Induced Seismicity

536 These results have important implications in understanding the mitigation of  
537 induced seismicity at different temperatures and pore pressures in granitic EGS  
538 (enhanced geothermal system) reservoirs. EGS requires the injection of cold fluids  
539 (typically water) into the subsurface which typically results in a significant increase in  
540 pore pressure and a reduction in effective stress. Fluid transport relies on a connected  
541 network of natural and artificial fractures. Unexpectedly, large amounts of fluid may  
542 access undetected critical stress faults (Grigoli et al., 2018), with alteration minerals  
543 conditioning the strength and stability of such faults (An et al., 2021, 2022). In  
544 high-temperature fractured reservoirs, the enhanced artificial circulation fluid volume  
545 will further accelerate the consumption of feldspar and the occurrence of chlorite.  
546 This effect will lead to a decrease in the strength of exposed faults, as confirmed by  
547 laboratory friction experiments and microstructural observations (Jeppson et al.,  
548 2023). Our experimental results show that the feldspar gouge presents  
549 velocity-weakening behavior at  $T \geq 180$  °C. This indicates that feldspar-rich faults  
550 may undergo unstable sliding as the depth of the modified reservoir increases, but the  
551 feldspar/chlorite mixed gouges may promote stable sliding of faults, which implies  
552 that feldspar chloritization in EGS reservoirs of granite rock is a process with  
553 dynamic competitive effects. Furthermore, our experimental results suggest that the  
554 process of feldspar to chlorite transformation has an influential role in the unstable  
555 sliding of faults. The results of our shear experiments demonstrate that temperature  
556 alone may not be the main determinative factor contributing to seismicity in granitic  
557 EGS reservoirs and highlights the importance of pore pressures. Meanwhile, high  
558 temperature may promote the fluid-assisted processes and modulate the local stresses  
559 in reservoirs (den Hartog et al., 2012; Martínez- Garzón et al., 2014; Chen et al., 2015;  
560 Westaway & Burnside, 2019). Hence, the combined effect of temperature and pore  
561 pressure may be important in triggering seismicity in feldspar-rich EGS reservoirs.

562 In addition, other compounding factors cannot be neglected in the triggering of  
563 earthquakes in EGS reservoirs, such as in situ driving slip rates. Our experiments  
564 show that the velocity dependence of feldspar gouge weakens with decreasing  
565 shearing rate. Feldspar gouge exhibits velocity weakening behavior at axial shear  
566 velocities of  $\sim 0.5$   $\mu\text{m/s}$  and  $\sim 0.05$   $\mu\text{m/s}$  (Figure 6), which suggests that unstable

567 sliding of the fault occurs at low rates of shear sliding and thus may induce seismicity.  
568 The natural driving slip rates of the faults are much lower than those utilized in the  
569 laboratory (Chen et al., 2015). Thus, lower tectonic driving velocities may further  
570 reduce frictional stability, that is, induce smaller and more negative values of  $(a-b)$   
571 and then cause unstable sliding of the faults. Further experimental studies are needed  
572 to analyze the effects of different mineral contents and shear velocity on fault stability.  
573 Our results have implications for natural faults that may undergo feldspar-chlorite  
574 alteration processes and are present at elevated temperatures and pressures in the  
575 subsurface.

## 576 **5 Conclusions**

577 Based on the hydrothermal conditions at the depth of the granitic EGS  
578 reservoir, we conducted shear experiments to explore the frictional stability of  
579 simulated feldspar (K-feldspar and albite) gouge and feldspar/chlorite mixed gouge.  
580 Meanwhile, the effects of different temperatures and pore pressures on fault sliding  
581 were analyzed. The experimental results show that the frictional strength of  
582 chlorite-rich mixed gouges is lower than that of feldspar gouge. The microstructural  
583 observation reveals that the chlorite particles are more strongly sheared. Thus, the  
584 decrease in frictional strength with increasing chlorite content may be explained by  
585 this phenomenon. The velocity dependence enhances with increasing chlorite content,  
586 which suggests that the content of clay minerals in the gouge may be one of the  
587 controlling factors for the frictional behavior of the fault. Additionally, the velocity  
588 dependence of feldspar gouge enhances and the values of the velocity dependence  
589 coefficients  $(a-b)$  increase with increasing pore pressure under hydrothermal  
590 conditions. Combined with the analysis of the effects of temperature and pore  
591 pressure, we hypothesize that fluid injection operations within the depth range of the  
592 EGS modification may promote unstable sliding of potentially feldspar-rich faults,  
593 which suggests that feldspar-rich faults may be reactivated in granitic EGS reservoirs.  
594 Hence, the presence of feldspar chloritization sequences and the range of injection  
595 pressures needs to be considered in the development of EGS to minimize the risk of  
596 injection-induced seismicity.

## 597 **Acknowledgments**

598 The research was funded by National Natural Science Foundation of China (No.  
599 42177175), Central Public-interest Scientific Institution Basal Research Fund (No.  
600 DZLXJK202204), and China Geological Survey (Nos. DD20190138, DD20221660).  
601 We thank Prof. Changrong He for useful discussions of the raw test data. We  
602 appreciate the assistance of Wenming Yao and Xi Ma in running the original  
603 experiments. DE gratefully acknowledges support from the G. Albert Shoemaker  
604 endowment.

## 605 **Data Availability Statement**

606 The experimental data presented in this study are available at

607 <https://doi.org/10.5061/dryad.4b8gthtkc>

608 **References**

- 609 Anderson, G. M., & Burnham, C. W. (1965). The solubility of quartz in super-critical  
610 water. *American Journal of Science*, 263(6), 494–511.  
611 <https://doi.org/10.2475/ajs.263.6.494>.
- 612 Andrés, S., Santillán, D., Mosquera, J. C., & Cueto- Felgueroso, L. (2019). Delayed  
613 weakening and reactivation of rate- and- state faults driven by pressure changes due  
614 to fluid injection. *Journal of Geophysical Research: Solid Earth*, 124, 11,917–  
615 11,937. <https://doi.org/10.1029/2019JB018109>.
- 616 An, M., Zhang, F., Chen, Z., Elsworth, D., & Zhang, L. (2020). Temperature and fluid  
617 pressurization effects on frictional stability of shale faults reactivated by hydraulic  
618 fracturing in the Changning block, southwest China. *Journal of Geophysical  
619 Research: Solid Earth*, 125, e2020JB019584.  
620 <https://doi.org/10.1029/2020JB019584>.
- 621 An, M., Zhang, F., Min, K. B., Elsworth, D., Marone, C., & He, C. (2021). T  
622 he potential for low - grade metamorphism to facilitate fault instability in a  
623 geothermal reservoir. *Geophysical Research Letters*, 48(11). <https://doi.org/10.1029/2021GL093552>.
- 624
- 625 An, M., Zhang, F., Min, K.-B., Elsworth, D., He, C., & Zhao, L. (2022). Frictional  
626 stability of metamorphic epidote in granitoid faults under hydrothermal conditions  
627 and implications for injection-induced seismicity. *Journal of Geophysical Research:  
628 Solid Earth*, 127, e2021JB023136. <https://doi.org/10.1029/2021JB023136>.
- 629 Bao, X., & Eaton, D. W. (2016). Fault activation by hydraulic fracturing in western  
630 Canada. *Science*, 354(6318), 1406-1409. <https://doi.org/10.1126/science.aag2583>.
- 631 Bedford, J. D., & Faulkner, D. R. (2021). The role of grain size and effective normal  
632 stress on localization and the frictional stability of simulated quartz gouge.  
633 *Geophysical Research Letters*, 48, e2020GL092023.  
634 <https://doi.org/10.1029/2020GL092023>.
- 635 Brown, K. M., Kopf, A., Underwood, M. B., & Weinberger, J. L. (2003). Com  
636 positional and fluid pressure controls on the state of stress on the Nankai su  
637 bduction thrust: A weak plate boundary. *Earth and Planetary Science Letters*,  
638 214(3–4), 589–603. [https://doi.org/10.1016/S0012-821X\(03\)00388-1](https://doi.org/10.1016/S0012-821X(03)00388-1).
- 639 Blanpied, M. L., Lockner, D. A., & Byerlee, J. D. (1991). Fault stability inferr  
640 ed from granite sliding experiments at hydrothermal conditions. *Geophysical  
641 Research Letters*, 18(4), 609–612. <https://doi.org/10.1029/91GL00469>.
- 642 Blanpied, M. L., Lockner, D. A., & Byerlee, J. D. (1995). Frictional slip of granite at  
643 hydrothermal conditions. *Journal of Geophysical Research: Solid Earth*, 100(B7),  
644 13045–13064. <https://doi.org/10.1029/95JB00862>.
- 645 Casas, N., Mollon, G., & Daouadji, A. (2023). Influence of grain-scale properties on  
646 localization patterns and slip weakening within dense granular fault gouges.  
647 *Journal of Geophysical Research: Solid Earth*, 128, e2022JB025666.  
648 <https://doi.org/10.1029/2022JB025666>.
- 649 Chen, J., Verberne, B. A., & Spiers, C. J. (2015). Effects of healing on the se  
650 ismogenic potential of carbonate fault rocks: Experiments on samples from th  
651 e Longmen Shan Fault, Sichuan, China. *Journal of Geophysical Research: So*

- 652 *lid Earth*, 120(8), 5479–5506. <https://doi.org/10.1002/2015JB012051>.
- 653 Chen, J., & Spiers, C. J. (2016). Rate and state frictional and healing behavior  
654 of carbonate fault gouge explained using microphysical model. *Journal of G*  
655 *eophysical Research: Solid Earth*, 121(12), 8642–8665. <https://doi.org/10.1002/2016JB013470>.
- 656
- 657 Chen, J., Verberne, B. A., & Niemeijer, A. R. (2020). Flow - to - friction transi  
658 tion in simulated calcite gouge: Experiments and microphysical modeling. *Jo*  
659 *urnal of Geophysical Research: Solid Earth*, 125(11). <https://doi.org/10.1029/2020JB019970>.
- 660
- 661 den Hartog, S. A., Peach, C. J., de Winter, D. M., Spiers, C. J., & Shimamot  
662 o, T. (2012). Frictional properties of megathrust fault gouges at low sliding  
663 velocities: New data on effects o-f normal stress and temperature. *Journal of*  
664 *Structural Geology*, 38, 156–171. <https://doi.org/10.1016/j.jsg.2011.12.001>.
- 665 Dieterich, J. H. (1978). Preseismic fault slip and earthquake prediction. *Journal of*  
666 *Geophysical Research: Solid Earth*, 83(B8), 3940–3948.  
667 <https://doi.org/10.1029/JB083iB08p03940>.
- 668 Elders, W. A., Hoagland, J. R., Mcdowell, S. D., & Cobo, J. (1979). Hydrothermal  
669 mineral zones in the geothermal reservoir of Cerro Prieto. *Geothermics*, 8(3–4),  
670 201–209. [https://doi.org/10.1016/0375-6505\(79\)90042-7](https://doi.org/10.1016/0375-6505(79)90042-7).
- 671 Ellsworth, W. L. (2013). Injection-induced earthquakes. *Science*, 341(6142), 1225942.  
672 <https://doi.org/10.1126/science.1225942>.
- 673 Er-yong Zhang; Dong-guang Wen; Gui-ling Wang et al. (2022). The first powe  
674 r generation test of hot dry rock resources exploration and production demon  
675 stration project in the Gonghe Basin, Qinghai Province, China. *China Geolog*  
676 *y*, Vol.5(3): 372–382. <https://doi.org/10.31035/cg2022038>.
- 677 Eyre, T. S., Eaton, D. W., Garagash, D. I., Zecevic, M., Venieri, M., Weir, R., &  
678 Lawton, D. C. (2019). The role of aseismic slip in hydraulic fracturing-induced  
679 seismicity. *Science Advances*, 5(8), 1–11. <https://doi.org/10.1126/sciadv.aav7172>.
- 680 Fagereng, Å., & Ikari, M. J. (2020). Low - temperature frictional characteristics  
681 of chlorite - epidote - amphibole assemblages: Implications for strength and s  
682 eismic style of retrograde fault zone-s. *Journal of Geophysical Research: Soli*  
683 *d Earth*, 125(4), e2020JB019487. <https://doi.org/10.1029/2020JB019487>.
- 684 Faulkner, D. R., Sanchez-Roa, C., Boulton, C., & den Hartog, S. A. M. (2018). Pore  
685 fluid pressure development in compacting fault gouge in theory, experiments, and  
686 nature. *Journal of Geophysical Research: Solid Earth*, 123, 226–241.  
687 <https://doi.org/10.1002/2017JB015130>.
- 688 Fengshou Zhang, Mengke An, Lianyang Zhang, Yi Fang, Derek Elsworth. (2020).  
689 Effect of mineralogy on friction-dilation relationships for simulated faults:  
690 Implications for permeability evolution in caprock faults. *Geoscience Frontiers*,  
691 (11): 439–450. <https://doi.org/10.1016/j.gsf.2019.05.014>.
- 692 F. M. Chester; N. G. Higgs. (1992). Multimechanism Friction Constitutive Model for  
693 Ultrafine Quartz Gouge at Hypocentral Conditions. *Journal of Geophysical*  
694 *Research. Part B: Solid Earth*, 97: 1859–1870. <https://doi.org/10.1029/91JB02349>.
- 695 Grigoli, F., Cesca, S., Rinaldi, A. P., Manconi, A., López-Comino, J. A., Clinton, J. F.,  
696 Wiemer, S. (2018). The November 2017 Mw5.5 Pohang earthquake: A possible

- 697 case of induced seismicity in South Korea. *Science*, 360(6392), 1003–1006.  
698 <https://doi.org/10.1126/science.aat2010>.
- 699 Gu, J. C., Rice, J. R., Ruina, A. L., & Tse, S. T. (1984). Slip motion and stability of a  
700 single degree of freedom elastic system with rate and state dependent friction.  
701 *Journal of the Mechanics and Physics of Solids*, 32(3), 167–196.  
702 [https://doi.org/10.1016/0022-5096\(84\)90007-3](https://doi.org/10.1016/0022-5096(84)90007-3)
- 703 He, C., Yao, W., Wang, Z., & Zhou, Y. (2006). Strength and stability of frictional  
704 sliding of gabbro gouge at elevated temperatures. *Tectonophysics*, 427(1–4), 217–  
705 229. <https://doi.org/10.1016/j.tecto.2006.05.023>.
- 706 He, C., Luo, L., Hao, Q. M., & Zhou, Y. (2013). Velocity - weakening behavior of  
707 plagioclase and pyroxene gouges and stabilizing effect of small amounts of quartz  
708 under hydrothermal conditions. *Journal of Geophysical Research: Solid Earth*,  
709 118(7), 3408–3430. <https://doi.org/10.1002/jgrb.50280>.
- 710 Hubbert, M. K., & Rubey, W. W. (1959). Role of fluid Pressure in mechanics of  
711 overthrust faulting. *Geological Society of America Bulletin*, 70(2), 115–166.  
712 [https://doi.org/10.1130/0016-7606\(1959\)70\[115:ROFPIM\]2.0.CO;2](https://doi.org/10.1130/0016-7606(1959)70[115:ROFPIM]2.0.CO;2).
- 713 Hunfeld, L. B., Niemeijer, A. R., & Spiers, C. J. (2017). Frictional properties of  
714 simulated fault gouges from the seismogenic Groningen gas field under in situ P–T  
715 -chemical conditions. *Journal of Geophysical Research: Solid Earth*, 122, 8969–  
716 8989. <https://doi.org/10.1002/2017JB014876>.
- 717 Ikari, M. J., Saffer, D. M., & Marone, C. (2009). Frictional and hydrologic pr  
718 operties of clay - rich fault gouge. *Journal of Geophysical Research: Solid E*  
719 *arth*, 114(B5). <https://doi.org/10.1029/2008JB006089>.
- 720 Ikari, M. J., Marone, C., & Saffer, D. M. (2011). On the relation between fault  
721 strength and frictional stability. *Geology*, 39(1), 83–86.  
722 <https://doi.org/10.1130/G31416.1>.
- 723 Jeppson, T. N., Lockner, D. A., Beeler, N. M., & Moore, D. E. (2023).  
724 Time-dependent weakening of granite at hydrothermal conditions. *Geophysical*  
725 *Research Letters*, 50, e2023GL105517. <https://doi.org/10.1029/2023GL105517>
- 726 John D. Bedford, Daniel R. Faulkner & Nadia Lapusta. (2022). Fault rock  
727 heterogeneity can produce fault weakness and reduce fault stability. *Nature*  
728 *communications*, Vol.13(1): 326, <https://doi.org/10.1038/s41467-022-27998-2>.
- 729 Lockner, D. A., Summers, R., & Byerlee, J. D. (1986). Effects of temperature and  
730 sliding rate on frictional strength of granite. *Pure and Applied Geophysics*, 124(3),  
731 445–469. <https://doi.org/10.1007/BF00877211>.
- 732 Logan, J. M., Dengo, C. A., Higgs, N. G., & Wang, Z. Z. (1992). Fabrics of  
733 experimental fault zones: Their development and relationship to mechanical  
734 behavior. In *International geophysics*, 51, 33–67. Academic Press.  
735 [https://doi.org/10.1016/S0074-6142\(08\)62814-4](https://doi.org/10.1016/S0074-6142(08)62814-4).
- 736 Lucie Mathieu. (2016). Quantifying hydrothermal alteration with normative minerals  
737 and other chemical tools at the Beattie Syenite, Abitibi greenstone belt, Canada.  
738 *EOCHEMISTRY-EXPLORATION ENVIRONMENT ANALYSIS*, Vol.16: 233–244.  
739 <https://doi.org/10.1144/geochem2016-410>.
- 740 Majer, E. L., Baria, R., Stark, M., Oates, S., Bommer, J., Smith, B., & Asanuma, H.  
741 (2007). Induced seismicity associated with enhanced geothermal systems.

- 742 *Geothermics*, 36(3), 185–222. <https://doi.org/10.1016/j.geothermics.2007.03.003>.
- 743 Marone, C. (1998). Laboratory-derived friction laws and their application to seismic  
744 faulting. *Annual Review of Earth and Planetary Sciences*, 26(1), 643–696.  
745 <https://doi.org/10.1146/annurev.earth.26.1.643>.
- 746 Martínez- Garzón, P., Kwiatak, G., Sone, H., Bohnhoff, M., Dresen, G., & Hartline, C.  
747 (2014). Spatiotemporal changes, faulting regimes, and source parameters of  
748 induced seismicity: A case study from the Geysers geothermal field. *Journal of*  
749 *Geophysical Research: Solid Earth*, 119, 8378–8396.  
750 <https://doi.org/10.1002/2014JB011385>.
- 751 Masuda, K.; Arai, T.; Takahashi, M. (2019) Effects of frictional properties of quartz  
752 and feldspar in the crust on the depth extent of the seismogenic zone. *Progress in*  
753 *Earth and Planetary Science*, 6(1): 1–8,  
754 <https://doi.org/10.1186/s40645-019-0299-5>.
- 755 Masuda, Koji. (2020). Frictional properties of anorthite (feldspar): implications for  
756 the lower boundary of the seismogenic zone. *Earth, Planets & Space*, Vol.72(1): 1–  
757 6, <https://doi.org/10.1186/s40623-020-01271-6>.
- 758 Maxim, F., Karalis, K., Boillat, P., Banuti, D. T., Marquez Damian, J. I., Niceno, B., &  
759 Ludwig, C. (2021). Thermodynamics and dynamics of supercritical water pseudo-  
760 boiling. *Advanced Science*, 8(3). <https://doi.org/10.1002/advs.202002312>.
- 761 Morrow, C. A., Moore, D. E., & Lockner, D. A. (2000). The effect of mineral bond  
762 strength and adsorbed water on fault gouge frictional strength. *Geophysical*  
763 *Research Letters*, 27(6), 815–818. <https://doi.org/10.1029/1999GL008401>.
- 764 Niemeijer, A. R., & Spiers, C. J. (2007). A microphysical model for strong velocity  
765 weakening in phyllosilicate-bearing fault gouges. *Journal of Geophysical Research*,  
766 112, B10405. <https://doi.org/10.1029/2007JB005008>.
- 767 Okamoto, A. S., Verberne, B. A., Niemeijer, A. R., Takahashi, M., Shimizu, I., Ueda,  
768 T., & Spiers, C. J. (2019). Frictional properties of simulated chlorite gouge at  
769 hydrothermal conditions: Implications for subduction megathrusts. *Journal of*  
770 *Geophysical Research: Solid Earth*, 124(5), 4545–4565.  
771 <https://doi.org/10.1029/2018JB017205>.
- 772 Okamoto, A. S., Niemeijer, A. R., Takeshita, T., Verberne, B. A., & Spiers, C. J.  
773 (2020). Frictional properties of actinolite-chlorite gouge at hydrothermal conditions.  
774 *Tectonophysics*, 779, 228377. <https://doi.org/10.1016/j.tecto.2020.228377>.
- 775 Ruina, A. (1983). Slip instability and state variable friction laws. *Journal of*  
776 *Geophysical Research: Solid Earth*, 88(B12), 10359–10370.  
777 <https://doi.org/10.1029/JB088iB12p10359>.
- 778 Sakuma, H., & Ichiki, M. (2016). Density and isothermal compressibility of  
779 supercritical H<sub>2</sub>O–NaCl fluid: molecular dynamics study from 673 to 2000 K, 0.2  
780 to 2 GPa, and 0 to 22 wt% NaCl concentrations. *Geofluids*, 16(1), 89–102.  
781 <https://doi.org/10.1111/gfl.12138>.
- 782 Schiffman, P., & Fridleifsson, G. O. (1991). The smectite-chlorite transition in  
783 drillhole NJ-15, Nesjavellir geothermal field, Iceland: XRD, BSE and electron  
784 microprobe investigations. *Journal of Metamorphic Geology*, 9(6), 679–696.  
785 <https://doi.org/10.1111/j.1525-1314.1991.tb00558.x>.
- 786 Scholz, C. (1998). Earthquakes and friction laws. *Nature*, 391, 37–42.

- 787 <https://doi.org/10.1038/34097>.
- 788 Scholz, C. (2019). *The Mechanics of Earthquakes and Faulting* (3rd ed.). Cambridge:  
789 Cambridge University Press. <https://doi.org/10.1017/9781316681473>.
- 790 Schultz, R., Skoumal, R. J., Brudzinski, M. R., Eaton, D., Baptie, B., & Ellsworth, W.  
791 (2020). Hydraulic fracturing-induced seismicity. *Reviews of Geophysics*, 58,  
792 e2019RG000695. <https://doi.org/10.1029/2019RG000695>.
- 793 Shimamoto, T., & Logan, J. M. (1981). Effects of simulated clay gouges on the  
794 sliding behavior of Tennessee sandstone. *Tectonophysics*, 75(3–4), 243–255.  
795 [https://doi.org/10.1016/0040-1951\(81\)90276-6](https://doi.org/10.1016/0040-1951(81)90276-6).
- 796 Yuguchi, T., Sasao, E., Ishibashi, M., & Nishiyama, T. (2015). Hydrothermal  
797 chloritization processes from biotite in the Toki granite, Central Japan: Temporal  
798 variations of the compositions of hydrothermal fluids associated with chloritization.  
799 *American Mineralogist*, 100(5–6), 1134–1152.  
800 <http://dx.doi.org/10.2138/am-2015-5126>.
- 801 Takashi Yuguchi, Takanobu Matsuki, Yuya Izumino, Eiji Sasao, Tadao Nishiya  
802 ma (2021), Mass transfer associated with chloritization in the hydrothermal a  
803 lteration process of granitic pluton. *American Mineralogist*; 106 (7): 1128–11  
804 42. <https://doi.org/10.2138/am-2020-7353>.
- 805 Tembe, S., D. A. Lockner, and T.-F. Wong (2010), Effect of clay content and  
806 mineralogy on frictional sliding behavior of simulated gouges: Binary and ternary  
807 mixtures of quartz, illite, and montmorillonite, *Journal of Geophysical Research:*  
808 *Solid Earth*, 115, B03416, <https://doi:10.1029/2009JB006383>.
- 809 van der Elst, N. J., Savage, H. M., Keranen, K. M., & Abers, G. A. (2013). Enhanced  
810 remote earthquake triggering at fluid- injection sites in the midwestern United  
811 States. *Science*, 341(6142), 164–167. <https://doi.org/10.1126/science.1238948>.
- 812 Verberne, B. A., van den Ende, M. P., Chen, J., Niemeijer, A. R., & Spiers, C. J.  
813 (2020). The physics of fault friction: Insights from experiments on simulated  
814 gouges at low shearing velocities. *Solid Earth Discussions*, 2020, 1–37.  
815 <https://doi.org/10.5194/se-11-2075-2020>.
- 816 Weingärtner, H., & Franck, E. U. (2005). Supercritical water as a solvent. *Angewandte*  
817 *Chemie International Edition*, 44(18), 2672–2692.  
818 <https://doi.org/10.1002/anie.200462468>.
- 819 Westaway, R., & Burnside, N. M. (2019). Fault “corrosion” by fluid injection: A  
820 potential cause of the November 2017 Mw 5.5 Korean earthquake. *Geofluids*, 2019,  
821 1–23. <https://doi.org/10.1155/2019/1280721>.
- 822 Zhang, F., Huang, R., An, M., Min, K. B., Elsworth, D., Hofmann, H., & Wa  
823 ng, X. (2022). Competing Controls of Effective Stress Variation and Chloritiz  
824 ation on Friction and Stability of Faults in Granite: Implications for Seismici  
825 ty Triggered by Fluid Injection. *Journal of Geophysical Research: Solid Eart*  
826 *h*, 127(8). <https://doi.org/10.1029/2022JB024310>.
- 827 Zhang, L., and C. He (2016), Frictional properties of phyllosilicate-rich mylonit  
828 e and conditions for the brittle-ductile transition, *J. Geophys. Res. Solid Eart*  
829 *h*, 121, 3017–3047. <https://doi.org/10.1002/2015JB012489>.
- 830 Zhen Lu, Changrong He. (2018). Friction of foliated fault gouge with a biotite

831 interlayer at hydrothermal conditions. *Tectonophysics*, 740, 72–92. [https://doi.](https://doi.org/10.1016/j.tecto.2018.05.003)  
832 [org/10.1016/j.tecto.2018.05.003.](https://doi.org/10.1016/j.tecto.2018.05.003)