

Abstract

Atmospheric trace gas measurements can be used to independently assess national greenhouse gas inventories through inverse modelling. Here, atmospheric nitrous oxide (N_2O) measurements made in the United Kingdom (U.K.) and Republic of Ireland are used to derive monthly N_2O emissions for 2013–2022 using two different inverse methods. We find mean U.K. emissions of 90.5 ± 23.0 (1σ) and 111.7 ± 32.1 (1σ) $\text{Gg N}_2\text{O yr}^{-1}$ for 2013–2022, and corresponding trends of -0.68 ± 0.48 (1σ) $\text{Gg N}_2\text{O yr}^{-2}$ and -2.10 ± 0.72 (1σ) $\text{Gg N}_2\text{O yr}^{-2}$, respectively for the two inverse methods. The U.K. National Atmospheric Emissions Inventory (NAEI) reported mean N_2O emissions of $73.9 \text{ Gg N}_2\text{O yr}^{-1}$ across this period, which is 14–33% smaller than the emissions derived from atmospheric data. We infer a pronounced seasonal cycle in N_2O emissions, with a peak occurring in the spring and a second smaller peak in the late summer for certain years. The springtime peak has a long seasonal decline that contrasts with the sharp rise and fall of N_2O emissions estimated from the bottom-up U.K. Emissions Model (UKEM). Bayesian inference is used to minimize the seasonal cycle mismatch between the average top-down (atmospheric data-based) and bottom-up (process model and inventory-based) seasonal emissions at a sub-sector level. Increasing agricultural manure management and decreasing synthetic fertilizer N_2O emissions reduces some of the discrepancy between the average top-down and bottom-up seasonal cycles. Other possibilities could also explain these discrepancies, such as missing emissions from NH_3 deposition, but these require further investigation.

1 Introduction

Atmospheric nitrous oxide (N_2O) is an important, long-lived greenhouse gas (GHG) that also contributes to the depletion of stratospheric ozone. Whilst global emissions of N_2O are well-constrained at around $17 \text{ Tg N}_2\text{O yr}^{-1}$ (Stell et al., 2022; Tian et al., 2020; Thompson et al., 2019; Wells et al., 2018) there are significant regional-scale differences between top-down (based on atmospheric data) and bottom-up (based on process models and/or inventories) N_2O emissions estimates (e.g. Wells et al., 2018; Jeong et al., 2018; Thompson et al., 2014).

The United Kingdom’s bottom-up N_2O emissions are reported in the U.K. National Atmospheric Emissions Inventory (NAEI; Ricardo Energy & Environment, 2019) each year which inform the National Inventory Reports (NIRs) submitted annually to the United Nations Framework Convention on Climate Change (UNFCCC). Anthropogenic emissions of N_2O are reported to be $72.6 \text{ Gg N}_2\text{O yr}^{-1}$ for the year 2021 (Brown et al., 2023), which are around 4% of the U.K.’s total carbon dioxide equivalent ($\text{CO}_{2,eq}$) GHG emissions. Agricultural N_2O emissions in the reported inventory account for 70% of 2021 U.K. N_2O emissions, with $\sim 10\%$ of U.K. N_2O emissions attributed to fossil fuel combustion and fugitive emissions, and the remaining anthropogenic emissions coming from the waste and industrial sectors (Fig. 1; Brown et al., 2023). Agricultural emissions are predominately from agricultural soils (57% of U.K. N_2O emissions) with the remaining agricultural N_2O emissions mostly from livestock manure management as well as indirect emissions from nitrogen runoff and leaching. A feature of the NAEI is that GHG fluxes are spatially distributed at 1 km^2 resolution for each SNAP (Selected Nomenclature for reporting of Air Pollutants) sector.

Under the 2008 Climate Change Act (UK Government, 2008) the U.K. has ambitious goals to achieve net-zero GHG emissions by 2050. Interim targets (“Carbon Budgets”) have been enacted, with the Fifth U.K. Carbon Budget requiring GHG emissions to be reduced to 57% of 1990 $\text{CO}_{2,eq}$ levels by 2030 (Committee on Climate Change, 2015). Current GHG emissions mitigation policy suggest U.K. N_2O emissions will remain approximately constant until 2030 (U.K. Department for Energy Security and Net Zero,

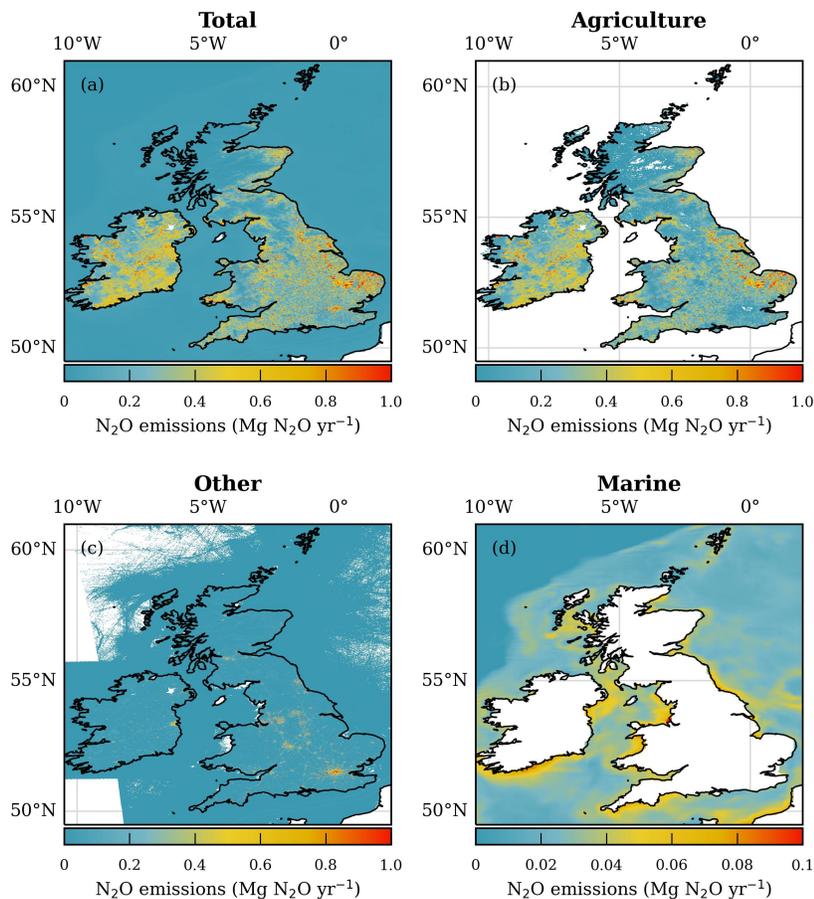


Figure 1. Bottom-up UKEM (a) total; (b) agricultural; (c) other anthropogenic (including shipping); (d) marine N₂O emissions for the year 2020 at 1 km² spatial resolution. Note that marine emissions are on a smaller scale than the other sectors and are not part of the NAEI. White areas indicate where emissions are zero.

2022). Given the potential uncertainties in inventories, which are used to guide these policy scenarios, there is a need to independently monitor progress towards such targets.

Atmospheric mole fraction measurements of trace gases can be used to evaluate emissions reported in national greenhouse gas inventories through inverse modelling. Atmospheric measurements are used in the inverse model to adjust bottom-up emissions through Bayesian inference. This approach has previously been used for deriving top-down N₂O emissions estimates for the U.K. (Manning et al., 2011; Ganesan et al., 2015) and elsewhere (e.g. Stell et al., 2022; Nevison et al., 2018; Thompson et al., 2014). Previous studies found top-down emissions were approximately 13–22% lower than those reported in the U.K. inventory at that time (Manning et al., 2011; Ganesan et al., 2015). Top-down estimates for earlier than 2012 had relatively large uncertainties as atmospheric observations were only available from the Mace Head station on the west coast of the Republic of Ireland (Fig. 2). The establishment of the U.K. Deriving Emissions linked to Climate Change (DECC) network in 2012 for making continuous measurements of atmospheric trace gases across the U.K. (Stanley et al., 2018; Stavert et al., 2019) has improved top-down emissions estimates, providing a higher sensitivity to emissions across

91 the U.K. and Republic of Ireland, leading to smaller top-down uncertainties (e.g. Lunt
92 et al., 2021).

93 Ganesan et al. (2015), using atmospheric N₂O measurements from the U.K. DECC
94 network with a hierarchical Bayesian inverse model (Ganesan et al., 2014), found aver-
95 age U.K. emissions of 101 (68-150) Gg N₂O yr⁻¹ for 2012-2014 and identified a pronounced
96 seasonal cycle with an amplitude of ~ 50 Gg N₂O yr⁻¹ peaking during the early sum-
97 mer. The observed seasonality was hypothesized to be due to agricultural application
98 of nitrogen-based fertilizers.

99 Seasonal changes in U.K. N₂O emissions are not captured in the U.K. NAEI in-
100 ventory, which only produces annual estimates. To facilitate comparison of top-down and
101 bottom-up seasonal cycles, temporal profiles derived from direct flux measurements and
102 other data sources are used to downscale the annual NAEI N₂O emissions at a source
103 sub-sector level to a monthly time resolution and combined with marine emissions from
104 Lessin et al. (2020) for the surrounding seas in the U.K. Emissions Model (UKEM, pre-
105 viously referred to as “UKGHG”) (P. Levy, 2020).

106 Here, we present results of top-down U.K. N₂O emissions from 2013-2022 derived
107 using two different Bayesian inverse models and compare these estimates to the U.K. NAEI
108 and the downscaled monthly estimates from the UKEM. Compared to previous U.K. N₂O
109 publications, we use observations from five stations in the U.K. DECC network (only three
110 stations were available for Ganesan et al. (2015)) and over a longer time period, result-
111 ing in better resolved emission trends and seasonal patterns.

112 2 Data and Methods

113 2.1 Atmospheric measurements

114 We use atmospheric N₂O mole fraction measurements from five U.K. DECC net-
115 work stations (Stanley et al., 2018; Stavert et al., 2019) between 2013-2022. Tall-tower
116 measurement stations in the U.K. DECC network sample ambient air from inlets mounted
117 > 90 m above ground level (agl) on telecommunications towers (Fig. 2), and the coastal
118 Mace Head (MHD) station measures closer to the surface at ~ 10 m agl. Atmospheric
119 N₂O measurements from MHD, Tacolneston (TAC) and Ridge Hill (RGL) stations sam-
120 pled across the entire 2013-2022 period, with Bilsdale (BSD) and Heathfield (HFD) be-
121 coming operational in 2014.

122 Both gas chromatograph coupled to electron capture detectors (GC-ECD) and optical-
123 based instruments (Picarro G5310, LGR; Text S1) are used for measuring atmospheric
124 N₂O mole fractions in the U.K. DECC network. Optical instruments are calibrated to
125 the World Meteorological Organization (WMO)-X2006A scale (Hall et al., 2007) with
126 samples averaged into ~ 1 min intervals. The GC-ECD instruments are calibrated to
127 the Scripps Institution of Oceanography (SIO-16) scale and sample approximately ev-
128 ery 10 min. A correction of -0.43 ppb is applied to N₂O measurements made on the SIO-
129 16 scale to adjust these measurements to the WMO-X2006A calibration scale (Prinn et
130 al., 2018).

131 At certain measurement stations (e.g. BSD) there are periods when concurrent GC-
132 ECD and optical-based atmospheric N₂O measurements are available. In such cases, we
133 use measurements from optical-based instruments in preference to those from the GC-
134 ECD instruments, due to their greater precision and higher frequency. Measurements
135 are always used from the highest air inlet available. Information about the instrumen-
136 tation and inlet heights used for each of the stations are summarized in Fig. 2, with the
137 U.K. DECC network further described in Stanley et al. (2018) and Stavert et al. (2019).

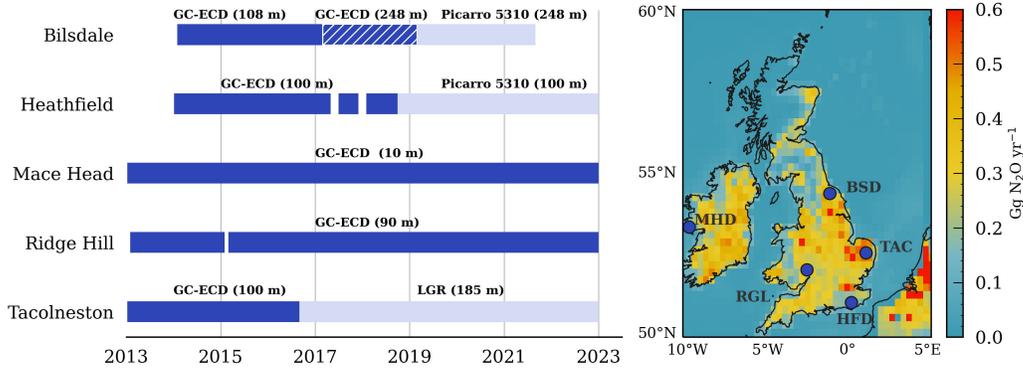


Figure 2. Summary of atmospheric N_2O mole fraction data used from the U.K. DECC network for 2013-2022. Instrument names and inlet heights (in brackets with “m” referring to metres above ground level) used for different periods are shown on the left. Gaps (white space) denote when measurements were unavailable. Shaded blue sections represent when gas chromatographs coupled to electron capture detectors (GC-ECD) were used, whereas sections shaded in lighter blue are when optical spectrometers were used. Hashed areas represent periods using GC-ECD measurements after a change in inlet height. Station locations are shown on the right in blue along with the May 2020 UKEM N_2O emissions regrided to match the NAME footprint resolution.

138 Measurements were averaged over 4 h periods and filtered to remove measurements
 139 that are more likely to be affected by local processes during times of stagnant air (e.g.
 140 Ganesan et al., 2015). Such meteorological conditions are unlikely to be accurately cap-
 141 tured at the spatiotemporal resolution of the atmospheric transport model (Section 2.2)
 142 used in this work. Each inverse model uses a different approach for filtering atmospheric
 143 measurements during stagnant air conditions. Further details about the filtering approaches
 144 are provided in Text S2. A comparison of top-down emissions derived using the same
 145 inverse method but with the different filtering approaches, and without any data filter-
 146 ing, is presented in Text S3.

147 Uncertainty in the atmospheric N_2O observations is quantified as the sum in quadra-
 148 ture of the instrument precision and observation variability in the 4 h averaging period.
 149 Observation uncertainties were on average ~ 0.28 ppb for the GC-ECD instruments and
 150 ~ 0.21 ppb for the optical instruments across the respective periods shown in Fig. 2.

151 2.2 Atmospheric transport model

152 The U.K. Met Office Lagrangian dispersion model: NAME (Numerical Atmospheric
 153 dispersion Modelling Environment; Jones et al., 2007) v7.2 was used to quantify the re-
 154 lationship between surface emissions and atmospheric mole fractions observed at each
 155 measurement station in both inverse models (Section 2.3). “Footprints” of surface emis-
 156 sion sensitivities were calculated from ensembles of particle back-trajectories in NAME.
 157 Each grid cell of the footprint describes the influence of emissions from that grid cell on
 158 the mole fraction measured at the receptor site at a certain time (Manning et al., 2011;
 159 Rigby et al., 2012).

160 Here, NAME footprints use the same specification as Manning et al. (2021). Hourly
 161 footprints were calculated with a ~ 25 km ($0.352^\circ \times 0.234^\circ$) spatial resolution over a
 162 model domain spanning approximately 98° W to 40° E and 11° N to 79° N using a 30-

163 day back-trajectory duration and particle release rate of $2 \times 10^4 \text{ h}^{-1}$. Meteorological
 164 fields from the Met Office Unified Model (UM) underlie the footprints with hourly, high-
 165 resolution up to $\sim 1.5 \text{ km}$ ($0.0135^\circ \times 0.0135^\circ$) with 57 vertical levels up to $\sim 12 \text{ km}$
 166 UKV meteorological fields used for over the British Isles and 3-hourly UM $0.1406^\circ \times 0.0938^\circ$
 167 with 59 vertical levels up to $\sim 30 \text{ km}$ global meteorological fields used for the rest of
 168 the modelling domain.

169 Footprints were combined with gridded a priori emissions (Section 2.4) to simu-
 170 late atmospheric N_2O mole fractions at each measurement station.

171 2.3 Inversion methodology

172 Two inverse models - InTEM (Inversion Technique for Emissions Modelling) and
 173 RHIME (Regional Hierarchical Inverse Modelling Environment) - are used for deriving
 174 top-down U.K. N_2O fluxes and are described below. Whilst both inverse models follow
 175 a Bayesian framework the inverse methods differ in several aspects: in the calculation
 176 of boundary condition/baseline mole fractions; the treatment of model-data uncertain-
 177 ties; and the approach for calculating the posterior emissions. Both inverse models in-
 178 fer emissions for each calendar month assuming N_2O emissions are constant in each 1-
 179 month period of inference.

180 2.3.1 InTEM inverse model

181 InTEM is an established Bayesian inverse model (Arnold et al., 2018; Manning et
 182 al., 2011, 2021) developed by the U.K. Met Office that has been widely used for trace
 183 gas inversions of different species across different regions (e.g. Rigby et al., 2019; Man-
 184 ning et al., 2021; Ganesan et al., 2020). InTEM is also used by the U.K. government for
 185 evaluating its nationally reported greenhouse gas emissions (Brown et al., 2022, 2023).

186 The InTEM framework minimizes the model-data mismatch constrained by obser-
 187 vation uncertainties, model uncertainties, and a priori information and its associated un-
 188 certainties, which are assumed to be Gaussian. To prevent non-physical solutions, In-
 189 TEM uses a non-negative least squares solver. The model-data uncertainty is calculated
 190 as the sum in quadrature of the observational uncertainty, baseline uncertainty and es-
 191 timated model uncertainty. The model uncertainty forms the largest contribution of the
 192 model-data uncertainty in InTEM and is calculated (for all trace gas species) by using
 193 the larger of: the concentration of the median pollution event over the year, or 10% of
 194 the concentration of the individual pollution event.

195 InTEM uses gridded N_2O emissions estimates (Section 2.4) and a time-varying mole
 196 fraction baseline as a priori constraints. The mole fraction baseline is derived using at-
 197 mospheric N_2O measurements from Mace Head that are representative of the well-mixed
 198 northern hemispheric baseline. This baseline is derived using filters to minimize influ-
 199 ences from populated regions, local sources, high altitudes and southerly latitudes (Manning
 200 et al., 2021). The mole fraction baseline is subsequently adjusted in the inversion by 11
 201 values depending on the geographical direction and altitude from which the air enters
 202 the model domain (Arnold et al., 2018), along with the spatial distribution of N_2O emis-
 203 sion values (Manning et al., 2021; Arnold et al., 2018). InTEM solves for posterior emis-
 204 sions in 100 scaling regions across Europe and imposes an a priori emissions uncertainty
 205 of 80% of NAEI N_2O emissions for the U.K. To account for small systematic differences
 206 in the modelling and between measurements, a prior offset uncertainty with 1σ uncer-
 207 tainty of 0.6 ppb is included for each station and subsequently solved for in the inver-
 208 sion. InTEM is further described in Manning et al. (2021); Arnold et al. (2018); Red-
 209 ington et al. (2023).

2.3.2 RHIME inverse model

RHIME (also previously referred to as “Bristol-MCMC”; Ganesan et al., 2014) has been frequently used for trace gas inversions of various atmospheric species across the globe (e.g. Say et al., 2021; Ganesan et al., 2015; Western et al., 2022). Here, Bayesian inference using MCMC (Markov Chain Monte Carlo) is used to quantify a mean scaling (with confidence intervals) of a priori emissions across 193 scaling regions in the model domain. A scaling of boundary conditions on the four cardinal boundaries of the NAME domain is also derived in the inversion.

Boundary condition fields representing mole fractions along the edges of the model domain are specified using CAMS (Copernicus Atmosphere Monitoring Service) v20r1 global inversion-optimized fluxes for each month. A scaling is calculated for each boundary in each 1-month period of inference to derive posterior boundary condition mole fractions. This ensures the CAMS data are not systematically underestimating or overestimating the baseline mole fractions at each of the measurement stations.

RHIME uses hyperparameters that characterize probability density functions (PDFs) of: the a priori emissions, boundary condition mole fractions, model-data covariances and offset terms. The RHIME framework allows uncertainties in the scaling parameters to be included in the model, with the a priori scaling sampled from a lognormal distribution $\sim LN(\mu = 0.346, \sigma = 0.693)$ and the model-data uncertainty from $\sim LN(\mu = -0.987, \sigma = 0.588)$. The model-data uncertainty is calculated as the sum in quadrature of the observational uncertainty and the absolute model uncertainty - derived from the model-data hyperparameter. The offset term is normally distributed with $\mu = 0$ ppb, $\sigma = 0.60$ ppb. Like Western et al. (2022) and Say et al. (2021), we use a No-U-Turn (NUTS) sampler (Hoffman & Gelman, 2014) for the a priori emissions and a slice sampler for estimating the model-data uncertainty. The samplers used a total 250,000 iterations (discarding the first 50,000) with 2 chains running in parallel. A Gelman-Rubin diagnostic is used to check for parameter convergence in both chains.

2.4 A priori emissions

The UKEM model (P. Levy, 2020) takes spatial data from the U.K. and Republic of Ireland inventories and processes them in a number of steps. These steps include: reprojection to a latitude-longitude grid, combining point- and area-based emissions, reconciling data from different sources into a single consistent classification scheme, rescaling to match national totals, back-projecting a time series of maps as inventories are updated annually, and adding in marine biogenic fluxes. For N_2O , marine biogenic fluxes from NEMO-ERSEM (the Nucleus for European Modelling of the Ocean model coupled with the European Regional Seas Ecosystem Model; Lessin et al., 2020) were used as the best available data on fluxes for the coastal sea around the U.K.

Inventory data are only produced annually. However, emissions may vary over much shorter timescales. The annual-scale emissions are therefore disaggregated in time in UKEM, to give the appropriate seasonal, day-of-week, and diurnal patterns. For some sectors, the variation in time is largely negligible (e.g. industry) or poorly known (e.g. LULUCF, waste), and these are represented as constant in time. For other sectors (agriculture, transport, energy), strong temporal patterns exist and can be characterized with activity data at higher temporal resolution from a variety of sources (P. Levy, 2020).

For N_2O , the key sector is agriculture, and the UKEM model uses activity data from the same process as the agricultural GHG inventory (Brown et al., 2023) but at monthly resolution when available. For example, the timing of synthetic fertilizer application is estimated from the British Survey of Fertiliser Practice (<https://www.gov.uk/government/collections/fertiliser-usage>). The timing of N_2O emission after application is estimated from existing data where fluxes have been measured in the field, generally us-

ing the closed static chamber technique. Typically, most of the emission occurs within only a few weeks of application (P. E. Levy et al., 2017). This produces a pronounced peak in the late spring, closely following the pattern in the application of synthetic fertilizer. The spatial distribution follows the distribution of cropland, where fertilizer inputs are highest (Fig. 1). An additional source is the mineralization of degrading peat soils in the Fenland areas in south-east England, though the magnitude of this is very uncertain.

Emissions from the marine sector are generally highest in the coastal zone, where inputs from rivers produce higher nitrogen concentrations. However, even near the coast, marine emissions are generally around ten times smaller than land emissions (Fig. 1), and although they peak during summer, their influence on the overall pattern is relatively small. At times, the modelled marine fluxes show net N_2O uptake, but only in a small region in the north-west of the domain.

As UKEM N_2O monthly emissions are only available for the U.K. and Republic of Ireland, $0.1^\circ \times 0.1^\circ$ monthly anthropogenic emissions from EDGAR (Emissions Database for Global Atmospheric Research; Crippa et al. (2021)) v6.0 are used for the rest of the modelling domain. Version 6.0 is used because monthly N_2O emissions are not currently available in the latest version (v7.0) of EDGAR. We “regrid” emissions data to match the NAME footprint spatial resolution and domain using a mass-conservative approach. As gridded emissions are only available for each month in 2013-2020, 2020 emissions are used as the a priori estimate for the 2021 and 2022 inversions.

Two sets of a priori emissions are used in the inversions. The first uses the monthly-resolved emissions and the second uses annual emissions, which are constant across each year. This provides an additional means of diagnosing whether any seasonal trends in the posterior emissions are driven by atmospheric measurements or by the a priori emissions.

2.5 Optimization of sectoral seasonal profiles

We are interested in comparing the U.K. top-down (derived using the inverse models; Section 2.3) and bottom-up (modelled in UKEM; Section 2.4) seasonal emissions profiles.

To investigate which N_2O emission sources could be driving seasonal differences we identified key agricultural sub-sectors that exhibit seasonal cycles in the U.K. These are: synthetic fertilizer usage, and manure management practices relating to: cattle during times they are housed (“cattle housing”), cattle during times of grazing (“cattle grazing”), the spreading of cattle manure, the spreading of poultry manure, and the spreading of digestate. Bottom-up emissions from all remaining N_2O sources are aggregated together. Each of these sub-sectors has a distinct temporal seasonal profile in the bottom-up UKEM emissions.

Here, we optimize the 2013-2022 averaged UKEM sub-sector emissions profiles to the averaged top-down seasonal emissions cycles - derived using InTEM and RHIME - to investigate which bottom-up emissions sources could be leading to differences with the top-down seasonal cycles.

We use a Bayesian approach to calculate the scale factors for each bottom-up sub-sector that minimizes the mismatch with the average top-down seasonal cycles. The posterior scale factors, \mathbf{X}_{post} , for each of the seven bottom-up sub-sector seasonal profiles are calculated using:

$$\mathbf{X}_{post} = \mathbf{X}_{prior} + (\mathbf{Q}_{prior} \mathbf{H}^T)(\mathbf{H} \mathbf{Q}_{prior} \mathbf{H}^T + \mathbf{R})^{-1}(\mathbf{Y} - \mathbf{H} \mathbf{X}_{prior}), \quad (1)$$

with associated covariance:

$$\mathbf{Q}_{post} = \mathbf{Q}_{prior} - (\mathbf{Q}_{prior}\mathbf{H}^T)(\mathbf{H}\mathbf{Q}_{prior}\mathbf{H}^T + \mathbf{R})^{-1}\mathbf{H}\mathbf{Q}_{prior}. \quad (2)$$

Here, \mathbf{X}_{prior} is a 7×1 matrix with elements of value 1 denoting the a priori scale factors for each UKEM emissions sub-sector. \mathbf{H} is a 12×7 matrix representing the monthly sub-sector UKEM seasonal emissions profiles (in $\text{Gg N}_2\text{O yr}^{-1}$). \mathbf{Y} is a 12×1 matrix representing the average top-down seasonal emissions (in $\text{Gg N}_2\text{O yr}^{-1}$) with associated uncertainties (in $\text{Gg N}_2\text{O yr}^{-1}$) captured in \mathbf{R} . Matrix \mathbf{R} is a 12×12 diagonal matrix where each element is the corresponding monthly 1σ value of the top-down seasonal emissions cycle. The uncertainty of the scale factors for each of the UKEM seasonal profiles are captured in \mathbf{Q} which is a 7×7 diagonal matrix with elements of value 1 denoting uncertainties of 100% on each of the bottom-up sub-sectors.

3 Results

3.1 Emissions and trends in N_2O (2013-2022)

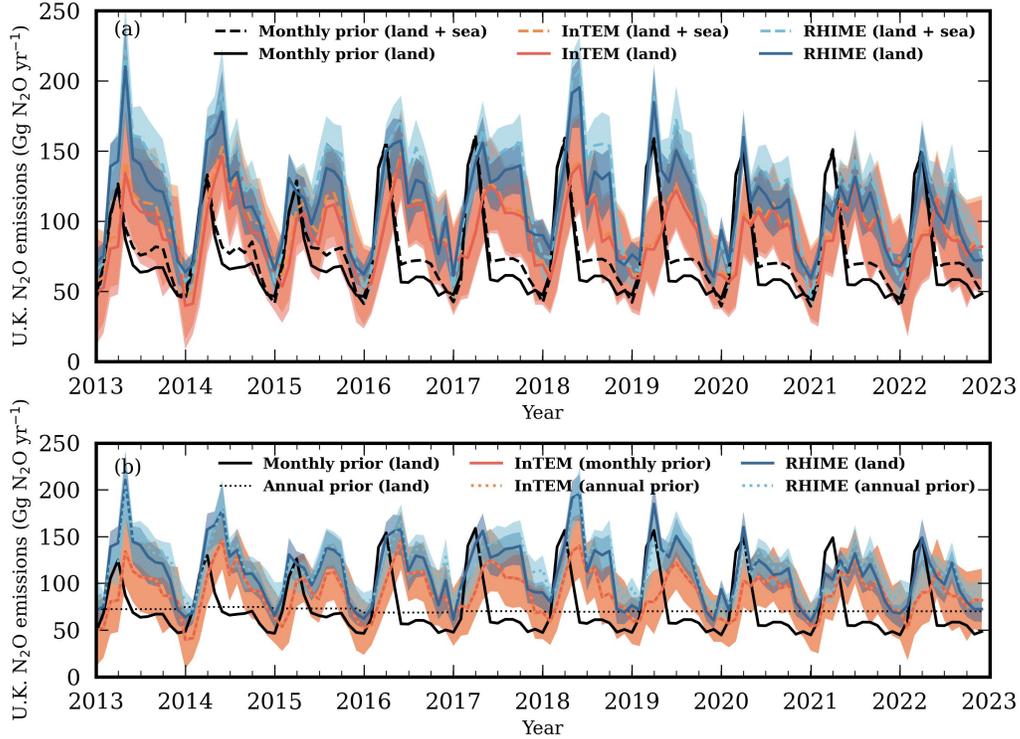


Figure 3. Time series of a priori and posterior monthly N_2O emissions for the U.K. (a) Mean monthly posterior emissions from InTEM (red) and RHIME (blue) for the land component of the U.K. (solid line) and the land plus surrounding seas of the U.K. (dashed line) with corresponding monthly a priori emissions in black. (b) InTEM and RHIME posterior emissions derived using monthly varying a priori emissions (solid line) and annually constant a priori emissions (dotted line). Shaded regions correspond to the 68% confidence interval ranges in both panels.

Top-down U.K. N_2O emissions were generally higher than their corresponding bottom-up UKEM/NAEI emissions across the 10-year period spanning January 2013 to December 2022 (Fig. 3, Table 1). On average, top-down emissions from RHIME were $38.0 \pm$

321 32.6(1 σ) Gg N₂O yr⁻¹ (32.7%) higher than the bottom-up estimates, and InTEM was
 322 16.8 ± 37.4(1 σ) Gg N₂O yr⁻¹ (14.0%) higher. Whilst RHIME posterior emissions are
 323 on average 21.2 ± 19.9(1 σ) Gg N₂O yr⁻¹ higher than the InTEM posterior emissions,
 324 there are overlapping 68% confidence intervals (CI) for most of the inversion period. The
 325 InTEM and RHIME top-down emissions are statistically well-correlated ($R^2 = 0.86$,
 326 $p < 0.01$). We later discuss potential reasons for differences between the InTEM and
 327 RHIME top-down emissions (Section 4).

328 We find an average of 5% of U.K. top-down emissions originate from the surround-
 329 ing seas (Fig. 3a; Table 1), whereas a priori emissions estimate 8% of U.K. N₂O emis-
 330 sions arise from the marine sector. InTEM infers that marine emissions should on av-
 331 erage be lower than the bottom-up estimates whereas RHIME infers the opposite (Ta-
 332 ble 1). Previous studies have highlighted that U.K. DECC stations are less sensitive to
 333 offshore emissions (Lunt et al., 2021) and as marine emissions form a small proportion
 334 of U.K. N₂O emissions we therefore limit our analysis to land-based emissions.

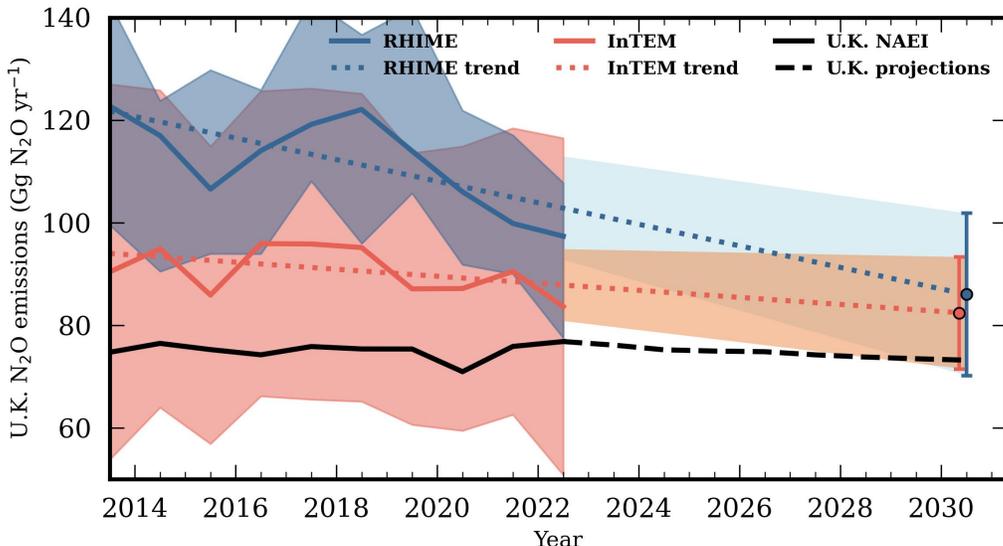


Figure 4. Annual top-down U.K. N₂O emissions from InTEM (red) and RHIME (blue) along with the linear trends for 2013-2022 that are extrapolated to 2030, with 2030 values indicated by circular markers. Shaded areas and error bars denote the 1 σ range. Current U.K. NAEI emissions and U.K. DESNZ projected emissions are shown in black.

335 We determine the RHIME and InTEM emissions trend over 2013-2022 by apply-
 336 ing a linear regression to the annual totals. Across this period, a mean negative trend
 337 of $-2.10 \pm 0.72(1\sigma)$ Gg N₂O yr⁻² ($p < 0.01$) was calculated for RHIME top-down emis-
 338 sions and $-0.68 \pm 0.48(1\sigma)$ Gg N₂O yr⁻² ($p < 0.01$) for InTEM top-down emissions
 339 (Fig. 4). Figure 4 shows that, comparatively, the NAEI N₂O emissions remain relatively
 340 constant across this period.

341 The U.K. Government Department for Energy Security and Net Zero (DESNZ) pro-
 342 duces GHG emissions projections based on current U.K. NAEI emissions and existing/near-
 343 finalised U.K. emissions mitigation policies (UK Department for Business Energy and
 344 Industrial Strategy, 2022). Projected U.K. N₂O emissions from 2022-2030 are shown in
 345 Fig. 4 and remain relatively constant at around 70 Gg N₂O yr⁻¹. Extrapolating the In-
 346 TEM and RHIME 2013-2022 emissions trends finds projected emissions of $82.3 \pm 10.9(1\sigma)$

Table 1. Summary of 2013-2022 mean U.K. posterior and a priori N₂O emissions and their differences reported with 1 σ uncertainty. “Monthly” and “Annual” reference posterior emissions derived using the monthly and annual UKEM emissions, respectively.

Inversion setup	InTEM		RHIME		UKEM	
	Monthly	Annual	Monthly	Annual	Monthly	Annual
Land emissions (Gg N ₂ O yr ⁻¹)	90.5 ± 23.0	90.5 ± 23.0	111.7 ± 32.1	116.1 ± 26.9	73.7 ± 31.9	73.7 ± 31.9
Marine emissions (Gg N ₂ O yr ⁻¹)	4.2 ± 2.1	4.2 ± 2.2	8.1 ± 10.5	17.9 ± 29.9	6.4 ± 6.9	6.4 ± 6.9
Land plus marine emissions (Gg N ₂ O yr ⁻¹)	94.7 ± 23.1	94.7 ± 23.0	119.8 ± 37.1	129.7 ± 36.0	80.1 ± 30.8	80.1 ± 30.8
Mean Posterior - A Priori Emissions						
Land						
Percentage difference	14.0 ± 41.0%	9.1 ± 41.0%	32.7 ± 23.2%	30.7 ± 15.5%	-	-
Gg N ₂ O yr ⁻¹ difference	16.8 ± 37.4	14.1 ± 22.9	38.0 ± 32.6	39.7 ± 26.5	-	-
Marine						
Percentage difference	-87.3 ± 230.4%	-120.0 ± 74.4%	29.6 ± 18.9%	23.7 ± 16.1%	-	-
Gg N ₂ O yr ⁻¹ difference	-2.2 ± 7.2	-3.4 ± 1.9	1.7 ± 4.3	5.9 ± 27.8	-	-
Land plus marine						
Percentage difference	12.7 ± 35.4%	5.3 ± 25.8%	30.7 ± 22.5%	31.5 ± 14.9%	-	-
Gg N ₂ O yr ⁻¹ difference	14.6 ± 33.4	10.6 ± 22.9	39.7 ± 34.4	31.5 ± 14.9	-	-

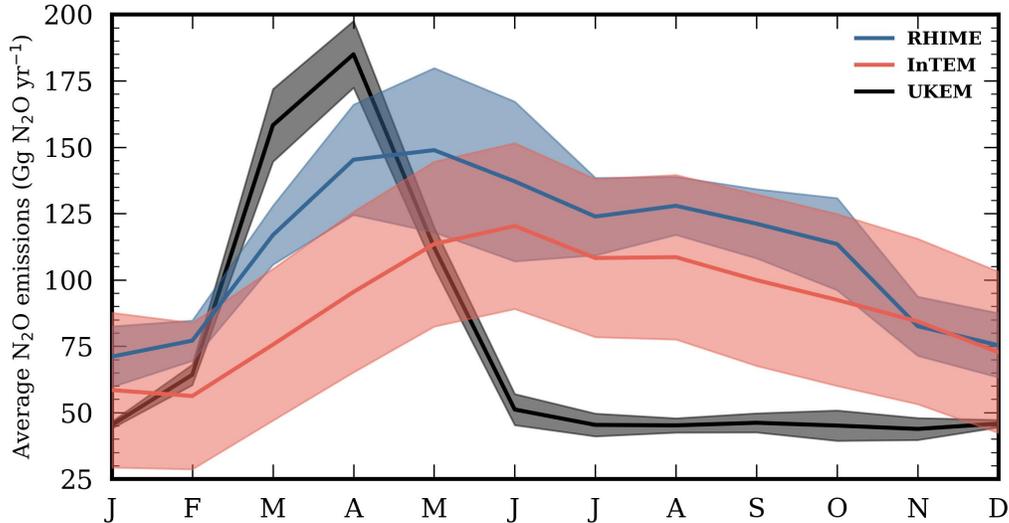


Figure 5. Average seasonal emissions profile across 2013-2022 with shaded regions denoting the 1σ standard deviation in emissions for each month.

347 Gg $\text{N}_2\text{O yr}^{-1}$ and $86.5 \pm 16.7 (1\sigma)$ Gg $\text{N}_2\text{O yr}^{-1}$, respectively, for the year 2030 (Fig.
 348 4).

349 3.2 Seasonal cycles

350 Pronounced seasonal cycles are observed in the top-down emissions with seasonal
 351 highs occurring during the summer and seasonal lows during the winter (Fig. 5). We find
 352 these seasonal cycles are insensitive to the seasonality in the prior. Fig. 3b shows a com-
 353 parison between posterior emissions that used a priori monthly UKEM emissions that
 354 either included or excluded a seasonal cycle. A pronounced seasonal peak is always de-
 355 rived in the top-down emissions, and there is strong agreement between the two sets of
 356 posterior emissions (Table 1), indicating that the derived seasonal cycle is primarily ob-
 357 servation driven.

358 In certain years there is a clearly identifiable springtime peak, as modelled in the
 359 bottom-up UKEM emissions, and a secondary, smaller peak occurring later in the year.
 360 From 2020 onwards these two seasonal peaks are less distinct in the top-down emissions
 361 compared to earlier years (Fig. 3).

362 The average seasonal emissions from 2013-2022 for top-down and bottom-up emis-
 363 sions are shown in Fig. 5. In UKEM, the modelled springtime seasonal peak rapidly de-
 364 clines. This is not seen in the top-down emissions, which instead fall more slowly. Ad-
 365 ditionally, the secondary peak that is sometimes seen in the top-down emissions seasonal
 366 cycles (Fig. 3) is not seen in UKEM. We find there are differences between the average
 367 InTEM and RHIME seasonal cycles. Notably, the occurrence of a late-summer/autumnal
 368 peak is less distinctive in InTEM than in RHIME, with the average InTEM seasonal cy-
 369 cle appearing more prolonged. The shaded regions in Fig. 5, which denote the 1σ stan-
 370 dard deviation in emissions for each month across 2013-2022, indicate variability in the
 371 RHIME seasonal cycle is generally lower than in InTEM for most months. Where the
 372 variability in the InTEM cycle is similar each month (Fig. 5), the variability in the RHIME
 373 seasonal cycle between April-July, when the seasonal maximum occurs, is much larger
 374 than in other months.

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3.3 Uncertainty analysis

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Median RHIME and InTEM model-data uncertainties with their 68% confidence intervals for each station are shown in Fig. 6a. Whilst median model-data uncertainties are less than 1 ppb, InTEM uncertainties are nearly four times larger than those from RHIME. These differences are reflected in the 68% CI uncertainty ranges of top-down emissions which are generally larger in InTEM. Whilst differences in model-data uncertainties could be partially attributed to the different inversion frameworks, it highlights the need for using different inverse methods that describe model-data uncertainties in different ways to ensure top-down emissions are more fully understood.

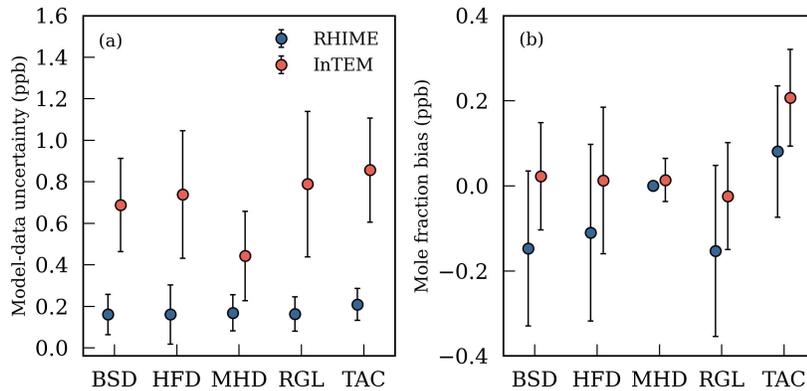


Figure 6. (a) Mean model-data uncertainties for InTEM (red) and RHIME (blue) for each of the measurement stations with error bars denoting the 68% confidence intervals. (b) Inverse model posterior offset bias median and 68% confidence intervals for each of the stations calculated in InTEM (red) and RHIME (blue).

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The InTEM and RHIME inversions account for any small instrumental (or model-related) differences that might occur in the atmospheric measurements across the network by solving for a mole fraction bias at each measurement station in each month. The median values with their 68% confidence intervals for each measurement station in InTEM and in RHIME are shown in Fig. 6b. The median bias values from each inversion are typically within ± 0.2 ppb at each station; similar to the measurement precision of the instruments (Section 2.1). However, median values differ by around 0.1 ppb between InTEM and RHIME with the largest discrepancies occurring at Bilsdale and Ridge Hill stations. The large, overlapping 68% confidence intervals indicate considerable variability in the biases that are solved in the monthly inversions. The posterior biases underscore that instrument or model-related differences exist across the U.K. DECC network at a magnitude comparable to the measurement uncertainty for N_2O . The high-variability in the calculated bias values suggests that these instrument-related differences are not constant in time and should be accounted for in N_2O inverse modelling studies using measurements from the U.K. DECC network (Thompson et al., 2011, 2014).

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3.4 Posterior UKEM sub-sector emissions seasonal profiles

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Figure 7a shows the 2013-2022 averaged UKEM seasonal cycles of six N_2O agricultural emissions sub-sectors (Section 2.5) with the remaining land-based N_2O emissions aggregated under “Other”. The pronounced springtime peak is driven by synthetic fertilizer emissions with some contribution from the manure management sub-sectors.

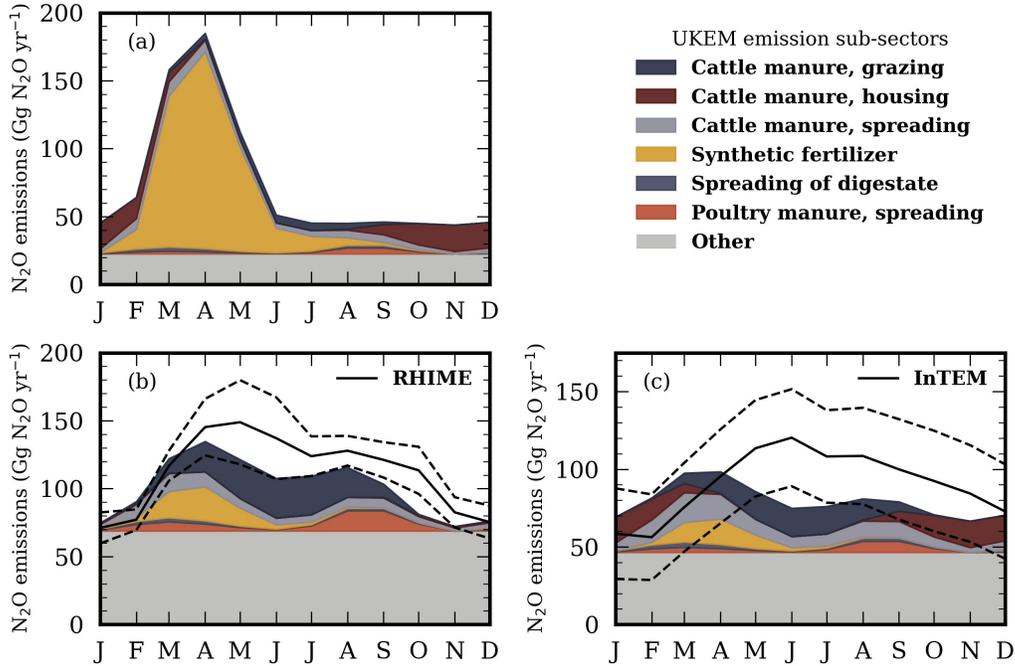


Figure 7. 2013-2022 average N₂O seasonal emissions profiles. (a) Disaggregated sub-sector emissions profiles as modelled in UKEM. Average RHIME and InTEM top-down seasonal emissions (solid lines) and 68% CI (dashed lines) with corresponding mean posterior sub-sector emissions are shown in (b) and (c), respectively.

404 As shown in Fig. 5, there are clear differences between the average bottom-up and
 405 top-down seasonal cycles for U.K. N₂O emissions across 2013-2022. To investigate what
 406 could be driving these differences, we use Bayesian inference (Section 2.5) to calculate
 407 scale factors for each UKEM sub-sector that optimize the fit to the observed average sea-
 408 sonal cycle. The seasonal profiles of the individual sectors from the UKEM are preserved.

409 Figure 7b and c show the optimized (posterior) sub-sectors that produce the best
 410 match with the derived seasonal cycles from RHIME and InTEM, respectively. We find
 411 the optimized fits do not generally have good agreement with the RHIME ($R^2 = 0.86$,
 412 $p < 0.01$) and InTEM ($R^2 = 0.11$, $p = 0.7$) seasonal cycles. Despite the large uncer-
 413 tainties in RHIME and InTEM there are certain months where the optimized bottom-
 414 up seasonal cycle does not fall within the RHIME/InTEM uncertainty. The fit to InTEM
 415 is generally worse than for RHIME, which could be due to the absence of two peaks in
 416 its average seasonal cycle and larger uncertainties than RHIME.

417 However for both RHIME and InTEM, a decrease in synthetic fertilizer emissions
 418 and an increase in “Other” emissions are inferred in the posterior seasonal profiles (Fig.
 419 7, Table 2). An increase in emissions related to cattle manure during times of grazing
 420 and the spreading of poultry manure could explain some of the seasonal differences with
 421 RHIME-inferred emissions in the latter half of the year. Whereas an increase in emis-
 422 sions related to the spreading of digestate, cattle manure during times of grazing and hous-
 423 ing could explain seasonal differences with InTEM-inferred emissions for the latter half
 424 of the year.

425 For January to April, there are only small differences between the RHIME/InTEM
 426 and optimized UKEM sub-sector emissions. After April, there is a larger mismatch be-

Table 2. Mean and 1σ uncertainty in the posterior scale factors for each UKEM emissions sub-sector when optimised to the average InTEM and RHIME seasonal cycles.

UKEM emissions sub-sector	InTEM-optimized scale factor	RHIME-optimized scale factor
Cattle manure, grazing	3.19 ± 1.61	5.12 ± 1.51
Cattle manure, housing	0.87 ± 0.87	0.05 ± 0.56
Cattle manure, spreading	1.83 ± 1.59	1.26 ± 1.28
Synthetic fertilizer	0.11 ± 0.20	0.17 ± 0.12
Spreading of digestate	1.11 ± 1.73	1.00 ± 1.70
Poultry manure, spreading	1.64 ± 1.73	3.41 ± 1.53
Other	2.05 ± 0.49	3.05 ± 0.43

427 tween these seasonal cycles. This mismatch could be attributed to retaining the seasonal
 428 profiles of each UKEM emissions sub-sectors, which could have time-correlated uncer-
 429 tainties.

430 4 Discussion

431 Continuous atmospheric N_2O measurements from the U.K. and Republic of Ire-
 432 land were used in InTEM and RHIME inverse models to derive top-down U.K. N_2O emis-
 433 sions estimates and their trends over 2013-2022. Mean posterior U.K. terrestrial emis-
 434 sions from InTEM and RHIME over 2013-2022 were $90.5 \pm 23.0(1\sigma)$ Gg N_2O yr $^{-1}$ and
 435 $111.7 \pm 32.1(1\sigma)$ Gg N_2O yr $^{-1}$, respectively. The two inversion systems inferred nega-
 436 tive trends of -0.68 ± 0.48 Gg N_2O yr $^{-2}$ and -2.10 ± 0.72 Gg N_2O yr $^{-2}$, respectively.
 437 However, it is unclear what could be driving a negative trend in the top-down emissions
 438 as the NAEI and UKEM sub-sector estimates do not have statistically significant trends
 439 over this period.

440 We find that the top-down emissions are close to the 2012-2014 U.K. emissions es-
 441 timates of 101 (68-150) Gg N_2O yr $^{-1}$ reported by Ganesan et al. (2015) and are approx-
 442 imately 14 – 33% higher than the bottom-up N_2O emissions reported in NAEI. This
 443 discrepancy is similar to the 13–22% differences between previously reported top-down
 444 and bottom-up estimates for the U.K. (Ganesan et al., 2015; Manning et al., 2011). How-
 445 ever, we find the inventory emissions estimates are lower than the top-down estimates
 446 whereas previous U.K. N_2O studies found inventory emissions were higher. The U.K.
 447 emissions inventory updates the entire record each year. Revisions to the U.K.’s N_2O
 448 emissions are reflected in the U.K.’s NIR submitted to the UNFCCC, which, for 2022,
 449 reported smaller N_2O emissions than were used as the a priori estimates in Ganesan et
 450 al. (2015) and Manning et al. (2011) at the time. Average differences of 14 – 33% be-
 451 tween top-down and bottom-up U.K. N_2O emissions underscore that there are still large
 452 uncertainties in U.K. N_2O emissions.

453 Similar to previous regional N_2O studies (e.g. Ganesan et al., 2015; Nevison et al.,
 454 2018; Wagner-Riddle et al., 2017; Jeong et al., 2012) we observe a pronounced seasonal
 455 cycle in the top-down N_2O emissions with a large springtime peak. We also observe a
 456 secondary peak in the late summer and/or a prolonged decay in the spring peak, which
 457 have not been previously observed in the U.K. The observed seasonal patterns are likely
 458 to be driven by agricultural sources since the timings of these peaks are broadly consis-
 459 tent each year but climatic patterns could also have an influence. Top-down N_2O emis-
 460 sions estimates in California, USA suggest the higher N_2O emissions observed during spring
 461 are linked to the application of fertilizers (Jeong et al., 2018, 2012; Xiang et al., 2013).

462 Nevison et al. (2018) observed a dual seasonal maxima in top-down emissions from 2008-
 463 2014 in the USA with a spring peak attributed to fertilizer applications and a late winter
 464 winter peak attributed to freeze-thaw effects (e.g. Wagner-Riddle et al., 2017). Higher aver-
 465 erage winter temperatures in the U.K. and the timing of the second peak make it un-
 466 likely for freeze-thaw effects to be a driver of seasonal variations of N_2O emissions in the
 467 U.K. Recent N_2O flux measurements over different terrestrial ecosystems in the U.K. found
 468 inconsistent seasonal patterns of total N_2O emissions across different land use types (Sgouridis
 469 & Ullah, 2017). This further suggests that seasonal variations with a springtime peak
 470 previously observed in the U.K. are driven by agricultural sources. However, environ-
 471 mental drivers such as precipitation and surface temperature also influence the magni-
 472 tude and timing of agricultural N_2O emissions (Skiba et al., 2012).

473 We find the agreement between the average top-down and bottom-up seasonal cycles
 474 could be improved by reducing UKEM synthetic fertilizer N_2O emissions by an aver-
 475 age factor of 0.15 ± 0.12 and increasing N_2O emissions from certain types of manure
 476 management - spreading of digestate and cattle manure during times of grazing and cat-
 477 tle housing - by an average factor of 4.2 ± 1.5 . Modelling the temporal changes of ma-
 478 nure management N_2O emissions is challenging as livestock waste management systems
 479 vary between sites and over time with little information available about their manage-
 480 ment practices (Chang et al., 2004). However, it is also possible the top-down and bottom-
 481 up seasonal cycle mismatch could be attributed to several other sources - such as miss-
 482 ing emissions from NH_3 deposition - but this requires further investigation.

483 Whilst we find top-down emissions from InTEM and RHIME are well-correlated
 484 ($R^2 = 0.86$, $p < 0.01$) there are differences in the magnitude and seasonal variations
 485 of N_2O emissions and uncertainties. Discrepancies between InTEM and RHIME poten-
 486 tially arise from differences in the calculation of boundary condition mole fractions, the
 487 treatment of model-data uncertainties and the approach for calculating posterior emis-
 488 sions in each inverse method. We find an average difference of $0.18 \pm 0.22(1\sigma)$ ppb be-
 489 tween the boundary condition mole fractions (Text S4). A lower mole fraction baseline
 490 would lead to higher emissions being derived in the inversion.

491 The results from these long-term U.K. atmospheric N_2O measurements from a dense
 492 measurement network demonstrate that they can be used for effective evaluation of re-
 493 gional N_2O emissions by using inverse modelling. Understanding of differences between
 494 top-down and bottom-up N_2O emissions could be enhanced by comparing top-down emis-
 495 sions derived using a different atmospheric transport model with different underlying me-
 496 teorological fields. Further long-term flux measurements of N_2O emissions sources could
 497 also provide additional constraints on the seasonal discrepancies observed between the
 498 average top-down and bottom-up seasonal cycles.

499 5 Conclusions

500 This study presents 10 years of top-down N_2O emissions derived using atmospheric
 501 measurements from across the U.K. and Republic of Ireland. Posterior emissions from
 502 both InTEM and RHIME inverse models find average U.K. emissions of 90.5 ± 23.0 Gg
 503 $\text{N}_2\text{O yr}^{-1}$ and 111.7 ± 32.1 Gg $\text{N}_2\text{O yr}^{-1}$, respectively across 2013-2020 which are \sim
 504 $14.0 \pm 41.0\%$ and $\sim 32.7 \pm 23.2\%$ higher than estimated in the U.K. national atmo-
 505 spheric emissions inventory, respectively. Average differences between the top-down and
 506 UKEM bottom-up seasonal patterns could be explained by decreasing synthetic fertil-
 507 izer emissions by an average factor of $\sim 0.15 \pm 0.12$, and increasing N_2O manure man-
 508 agement emissions (from cattle manure during grazing and housing) by an average fac-
 509 tor of 4.2 ± 1.5 . However, we find large uncertainties associated with the posterior scal-
 510 ing factors for the agricultural seasonal emissions profiles.

6 Open Research

Atmospheric measurements of N₂O used in this work are available from the CEDA Archive (O’Doherty et al., 2020): <https://data.ceda.ac.uk/badc/uk-decc-network/data/n2o/v23.08>. A doi for this data is currently being assigned. Python scripts used for data analysis are available for use in https://github.com/EricSaboya/uk_n2o

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References

- Arnold, T., Manning, A. J., Kim, J., Li, S., Webster, H., Thomson, D., . . . O’Doherty, S. (2018). Inverse modelling of CF₄ and NF₃ emissions in East Asia. *Atmospheric Chemistry and Physics*, 18(18), 13305–13320. doi: 10.5194/acp-18-13305-2018
- Brown, P., Cardenas, L., Choudrie, S., Del Vento, S., Karagianni, E., MacCarthy, J., . . . Wakeling, D. (2022). *UK Greenhouse Gas Inventory 1990 to 2020: Annual Report for submission under the Framework Convention on Climate Change* (Tech. Rep.).
- Brown, P., Cardenas, L., Del Vento, S., Karagianni, E., MacCarthy, J., Mullen, P., . . . Willis, D. (2023). *UK Greenhouse Gas Inventory, 1990 to 2021 Annual Report for Submission under the Framework Convention on Climate Change* (Tech. Rep.). UK Department for Energy Security and Net Zero.
- Chang, Y. H., Scrimshaw, M. D., & Lester, J. N. (2004). *IPCC Good Practice Guidance and Uncertainty Management in National Greenhouse Gas Inventories* (Vol. 25; Tech. Rep. No. 2).
- Committee on Climate Change. (2015). *The Fifth Carbon Budget: The next step towards a low-carbon economy* (Tech. Rep.). Retrieved from <https://www.theccc.org.uk/publication/the-fifth-carbon-budget-the-next-step-towards-a-low-carbon-economy/>
- Crippa, M., Guizzardi, D., Solazzo, E., Muntean, M., Schaaf, E., Monforti-Ferrario, F., . . . Vignati, E. (2021). *GHG emissions of all world countries - 2021 Report* (No. October).
- Ganesan, A. L., Manizza, M., Morgan, E. J., Harth, C. M., Kozlova, E., Lueker, T., . . . Rigby, M. (2020). Marine Nitrous Oxide Emissions From Three Eastern Boundary Upwelling Systems Inferred From Atmospheric Observations. *Geophysical Research Letters*, 47(14). doi: 10.1029/2020GL087822
- Ganesan, A. L., Manning, A. J., Grant, A., Young, D., Oram, D. E., Sturges, W. T., . . . O’Doherty, S. (2015, 6). Quantifying methane and nitrous oxide emissions from the UK and Ireland using a national-scale monitoring network. *Atmospheric Chemistry and Physics*, 15(11), 6393–6406. Retrieved from <https://acp.copernicus.org/articles/15/6393/2015/> doi: 10.5194/acp-15-6393-2015
- Ganesan, A. L., Rigby, M., Zammit-Mangion, A., Manning, A. J., Prinn, R. G., Fraser, P. J., . . . Weiss, R. F. (2014). Characterization of uncertainties in atmospheric trace gas inversions using hierarchical Bayesian methods. *Atmospheric Chemistry and Physics*, 14(8). doi: 10.5194/acp-14-3855-2014
- Hall, B. D., Dutton, G. S., & Elkins, J. W. (2007). The NOAA nitrous oxide standard scale for atmospheric observations. *Journal of Geophysical Research Atmospheres*, 112(9). doi: 10.1029/2006JD007954
- Hoffman, M. D., & Gelman, A. (2014). The no-U-turn sampler: Adaptively set-

- 562 ting path lengths in Hamiltonian Monte Carlo. *Journal of Machine Learning*
563 *Research*, 15.
- 564 Jeong, S., Newman, S., Zhang, J., Andrews, A. E., Bianco, L., Dlugokencky, E., ...
565 Fischer, M. L. (2018). Inverse Estimation of an Annual Cycle of California's
566 Nitrous Oxide Emissions. *Journal of Geophysical Research: Atmospheres*,
567 123(9). doi: 10.1029/2017JD028166
- 568 Jeong, S., Zhao, C., Andrews, A. E., Dlugokencky, E. J., Sweeney, C., Bianco, L.,
569 ... Fischer, M. L. (2012). Seasonal variations in N₂O emissions from central
570 California. *Geophysical Research Letters*, 39(16). doi: 10.1029/2012GL052307
- 571 Jones, A., Thomson, D., Hort, M., & Devenish, B. (2007, 4). The U.K. Met Office's
572 Next-Generation Atmospheric Dispersion Model, NAME III. In *Air pollution*
573 *modeling and its application xvii* (pp. 580–589). Springer US. doi: 10.1007/978
574 -0-387-68854-1{_}62
- 575 Lessin, G., Polimene, L., Artioli, Y., Butenschön, M., Clark, D. R., Brown, I., &
576 Rees, A. P. (2020). Modeling the Seasonality and Controls of Nitrous Oxide
577 Emissions on the Northwest European Continental Shelf. *Journal of Geophysi-*
578 *cal Research: Biogeosciences*, 125(6). doi: 10.1029/2019JG005613
- 579 Levy, P. (2020). *Greenhouse Gas Fluxes from the UK (ukghg)*. Retrieved from
580 <https://github.com/NERC-CEH/ukghg/blob/>
- 581 Levy, P. E., Cowan, N., van Oijen, M., Famulari, D., Drewer, J., & Skiba, U. (2017).
582 Estimation of cumulative fluxes of nitrous oxide: uncertainty in temporal up-
583 scaling and emission factors. *European Journal of Soil Science*, 68(4). doi:
584 10.1111/ejss.12432
- 585 Lunt, M. F., Manning, A. J., Allen, G., Arnold, T., Bauguitte, S. J.-B., Boesch,
586 H., ... Rigby, M. (2021, 11). Atmospheric observations consistent with
587 reported decline in the UK's methane emissions (2013–2020). *Atmo-*
588 *spheric Chemistry and Physics*, 21(21), 16257–16276. Retrieved from
589 <https://acp.copernicus.org/articles/21/16257/2021/> doi: 10.5194/
590 acp-21-16257-2021
- 591 Manning, A. J., O'Doherty, S., Jones, A. R., Simmonds, P. G., & Derwent, R. G.
592 (2011). Estimating UK methane and nitrous oxide emissions from 1990 to
593 2007 using an inversion modeling approach. *Journal of Geophysical Research*
594 *Atmospheres*, 116(2), 1–19. doi: 10.1029/2010JD014763
- 595 Manning, A. J., Redington, A. L., Say, D., O'Doherty, S., Young, D., Simmonds,
596 P. G., ... Arnold, T. (2021). Evidence of a recent decline in UK emissions of
597 hydrofluorocarbons determined by the InTEM inverse model and atmospheric
598 measurements. *Atmospheric Chemistry and Physics*, 21(16), 12739–12755. doi:
599 10.5194/acp-21-12739-2021
- 600 Nevison, C., Andrews, A., Thoning, K., Dlugokencky, E., Sweeney, C., Miller, S.,
601 ... Nehrkorn, T. (2018). Nitrous Oxide Emissions Estimated With the
602 CarbonTracker-Lagrange North American Regional Inversion Framework.
603 *Global Biogeochemical Cycles*, 32(3). doi: 10.1002/2017GB005759
- 604 O'Doherty, S., Say, D., Stanley, K., Spain, G., Arnold, T., Rennick, C., ...
605 Grant, A. (2020). *UK DECC (Deriving Emissions linked to Climate*
606 *Change) Network*. Centre for Environmental Data Analysis. Retrieved from
607 <http://catalogue.ceda.ac.uk/uuid/f5b38d1654d84b03ba79060746541e4f>
- 608 Prinn, R. G., Weiss, R. F., Arduini, J., Arnold, T., Langley Dewitt, H., Fraser,
609 P. J., ... Zhou, L. (2018). History of chemically and radiatively important
610 atmospheric gases from the Advanced Global Atmospheric Gases Experiment
611 (AGAGE). *Earth System Science Data*, 10(2). doi: 10.5194/essd-10-985-2018
- 612 Redington, A. L., Manning, A. J., Henne, S., Graziosi, F., Western, L. M., Ar-
613 duini, J., ... Young, D. (2023, 7). Western European emission estimates
614 of CFC-11, CFC-12 and CCl₄ derived from atmospheric measurements from
615 2008 to 2021. *Atmospheric Chemistry and Physics*, 23(13), 7383–7398. doi:
616 10.5194/acp-23-7383-2023

- 617 Ricardo Energy & Environment. (2019). UK Emission Mapping Methodology 2017.
618 (1), 49.
- 619 Rigby, M., Manning, A. J., & Prinn, R. G. (2012). The value of high-frequency,
620 high-precision methane isotopologue measurements for source and sink esti-
621 mation. *Journal of Geophysical Research Atmospheres*, *117*(12), 1–14. doi:
622 10.1029/2011JD017384
- 623 Rigby, M., Park, S., Saito, T., Western, L. M., Redington, A. L., Fang, X., ...
624 Young, D. (2019). Increase in CFC-11 emissions from eastern China based on
625 atmospheric observations. *Nature*, *569*(7757). doi: 10.1038/s41586-019-1193-4
- 626 Say, D., Manning, A. J., Western, L. M., Young, D., Wisher, A., Rigby, M., ...
627 O’Doherty, S. (2021). Global trends and European emissions of tetrafluoro-
628 romethane (CF₄), hexafluoroethane (C₂F₆) and octafluoropropane (C₃F₈).
629 *Atmospheric Chemistry and Physics*, *21*(3). doi: 10.5194/acp-21-2149-2021
- 630 Sgouridis, F., & Ullah, S. (2017). Soil Greenhouse Gas Fluxes, Environmental Con-
631 trols, and the Partitioning of N₂O Sources in UK Natural and Seminat-
632 ural Land Use Types. *Journal of Geophysical Research: Biogeosciences*, *122*(10).
633 doi: 10.1002/2017JG003783
- 634 Skiba, U., Jones, S. K., Dragosits, U., Drewer, J., Fowler, D., Rees, R. M., ...
635 Manning, A. J. (2012). *UK emissions of the greenhouse gas nitrous oxide*
636 (Vol. 367) (No. 1593). doi: 10.1098/rstb.2011.0356
- 637 Stanley, K. M., Grant, A., O’Doherty, S., Young, D., Manning, A. J., Stavert, A. R.,
638 ... Derwent, R. G. (2018). Greenhouse gas measurements from a UK network
639 of tall towers: Technical description and first results. *Atmospheric Measure-*
640 *ment Techniques*, *11*(3), 1437–1458. doi: 10.5194/amt-11-1437-2018
- 641 Stavert, A. R., O’Doherty, S., Stanley, K., Young, D., Manning, A. J., Lunt, M. F.,
642 ... Arnold, T. (2019). UK greenhouse gas measurements at two new tall
643 towers for aiding emissions verification. *Atmospheric Measurement Techniques*,
644 *12*(8). doi: 10.5194/amt-12-4495-2019
- 645 Stell, A. C., Bertolacci, M., Zammit-Mangion, A., Rigby, M., Fraser, P. J., Harth,
646 C. M., ... Ganesan, A. L. (2022). Modelling the growth of atmospheric ni-
647 trous oxide using a global hierarchical inversion. *Atmospheric Chemistry and*
648 *Physics*, *22*(19). doi: 10.5194/acp-22-12945-2022
- 649 Thompson, R. L., Chevallier, F., Crotwell, A. M., Dutton, G., Langenfelds, R. L.,
650 Prinn, R. G., ... Aoki, S. (2014). Nitrous oxide emissions 1999 to 2009 from a
651 global atmospheric inversion. *Atmospheric Chemistry and Physics*, *14*(4). doi:
652 10.5194/acp-14-1801-2014
- 653 Thompson, R. L., Gerbig, C., & Rödenbeck, C. (2011). A Bayesian inversion
654 estimate of N₂O emissions for western and central Europe and the assess-
655 ment of aggregation errors. *Atmospheric Chemistry and Physics*, *11*(7). doi:
656 10.5194/acp-11-3443-2011
- 657 Thompson, R. L., Lassaletta, L., Patra, P. K., Wilson, C., Wells, K. C., Gressent,
658 A., ... Canadell, J. G. (2019). Acceleration of global N₂O emissions seen from
659 two decades of atmospheric inversion. *Nature Climate Change*, *9*(12). doi:
660 10.1038/s41558-019-0613-7
- 661 Tian, H., Xu, R., Canadell, J. G., Thompson, R. L., Winiwarter, W., Sunthar-
662 alingam, P., ... Yao, Y. (2020). A comprehensive quantification of
663 global nitrous oxide sources and sinks. *Nature*, *586*(7828). doi: 10.1038/
664 s41586-020-2780-0
- 665 UK Department for Business Energy and Industrial Strategy. (2022). *Updated en-*
666 *ergy and emissions projections 2021 to 2040* (Tech. Rep.). Retrieved from
667 [https://www.gov.uk/government/publications/energy-and-emissions](https://www.gov.uk/government/publications/energy-and-emissions-projections-2021-to-2040)
668 [-projections-2021-to-2040](https://www.gov.uk/government/publications/energy-and-emissions-projections-2021-to-2040)
- 669 U.K. Department for Energy Security and Net Zero. (2022). *Energy and emissions*
670 *projections: 2021 to 2040* (Tech. Rep.). Retrieved from [https://www.gov.uk/](https://www.gov.uk/government/publications/energy-and-emissions-projections-2021-to-2040)
671 [government/publications/energy-and-emissions-projections-2021-to](https://www.gov.uk/government/publications/energy-and-emissions-projections-2021-to-2040)

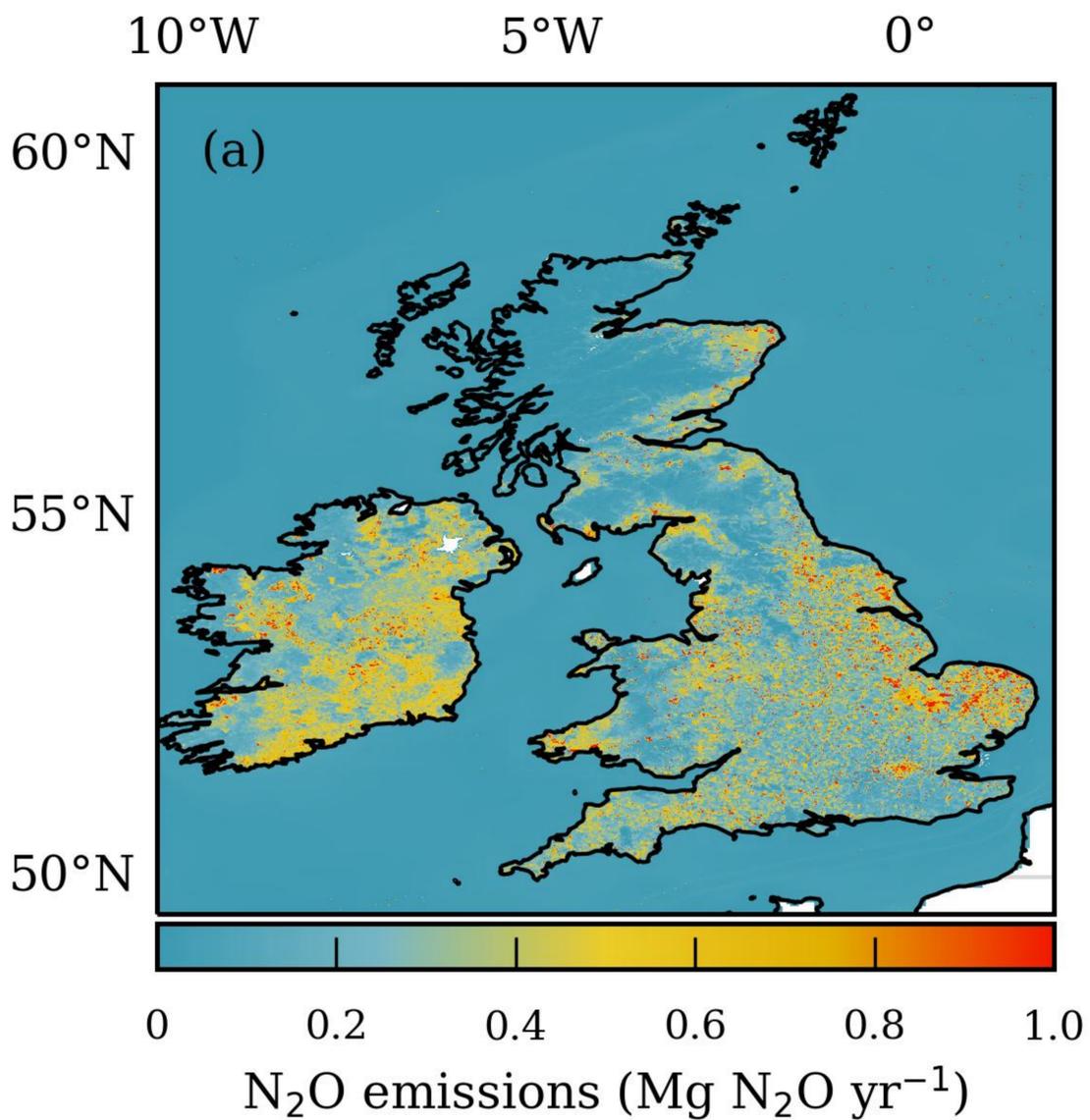
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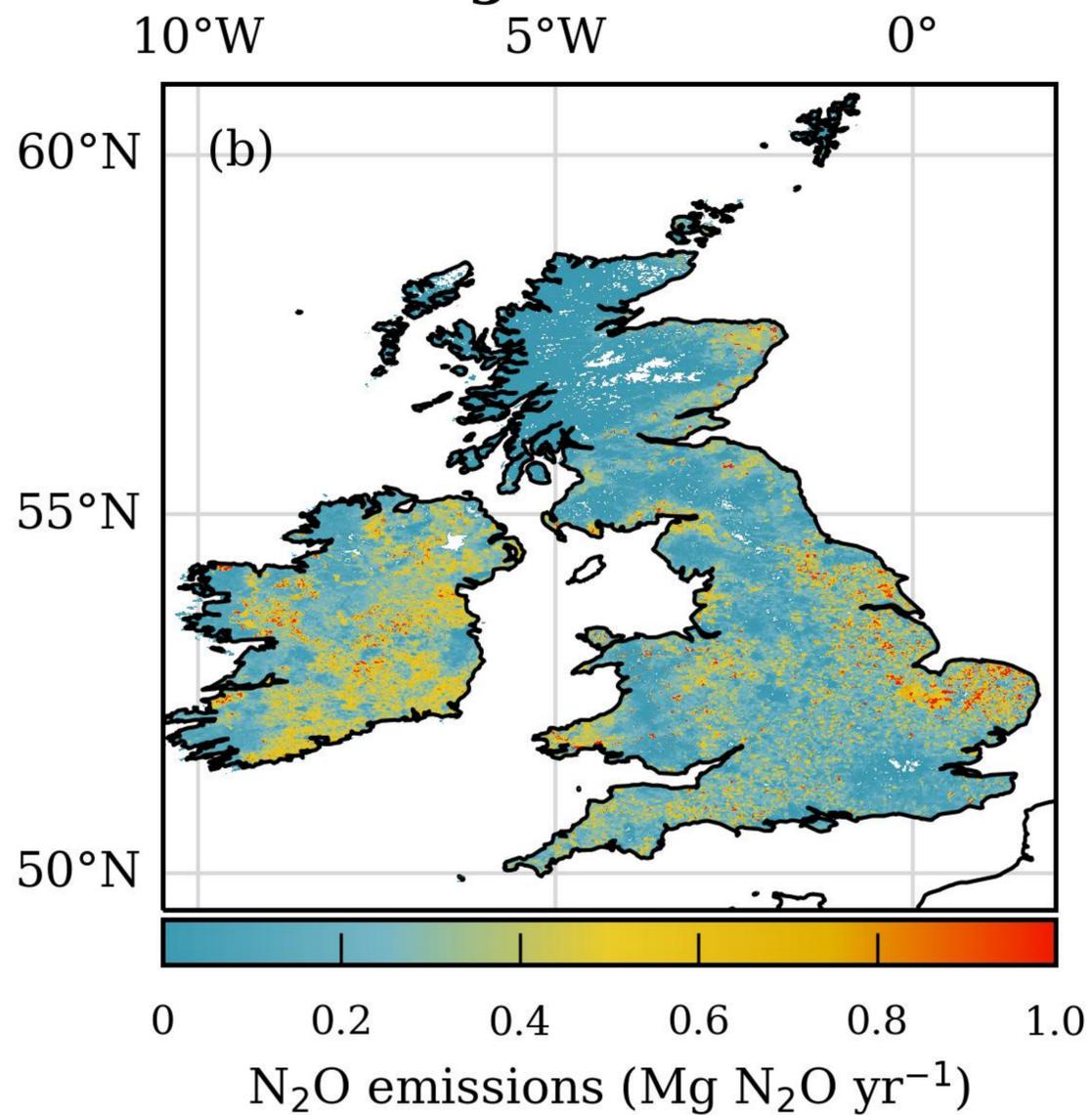
- UK Government. (2008). *Climate Change Act 2008: Elizabeth II. Chapter 27*.
- Wagner-Riddle, C., Congreves, K. A., Abalos, D., Berg, A. A., Brown, S. E., Am-
badan, J. T., ... Tenuta, M. (2017). Globally important nitrous oxide emis-
sions from croplands induced by freeze-thaw cycles. *Nature Geoscience*, *10*(4).
doi: 10.1038/ngeo2907
- Wells, K. C., Millet, D. B., Bousserez, N., Henze, D. K., Griffis, T. J., Chaliyakun-
nel, S., ... Paul Steele, L. (2018). Top-down constraints on global
N₂O emissions at optimal resolution: Application of a new dimension re-
duction technique. *Atmospheric Chemistry and Physics*, *18*(2). doi:
10.5194/acp-18-735-2018
- Western, L. M., Redington, A. L., Manning, A. J., Trudinger, C. M., Hu, L., Henne,
S., ... Rigby, M. (2022, 7). A renewed rise in global HCFC-141b emissions
between 2017-2021. *Atmospheric Chemistry and Physics*, *22*(14), 9601–9616.
doi: 10.5194/acp-22-9601-2022
- Xiang, B., Miller, S. M., Kort, E. A., Santoni, G. W., Daube, B. C., Commane, R.,
... Wofsy, S. C. (2013). Nitrous oxide (N₂O) emissions from California based
on 2010 CalNex airborne measurements. *Journal of Geophysical Research*
Atmospheres, *118*(7). doi: 10.1002/jgrd.50189

Figure 1.

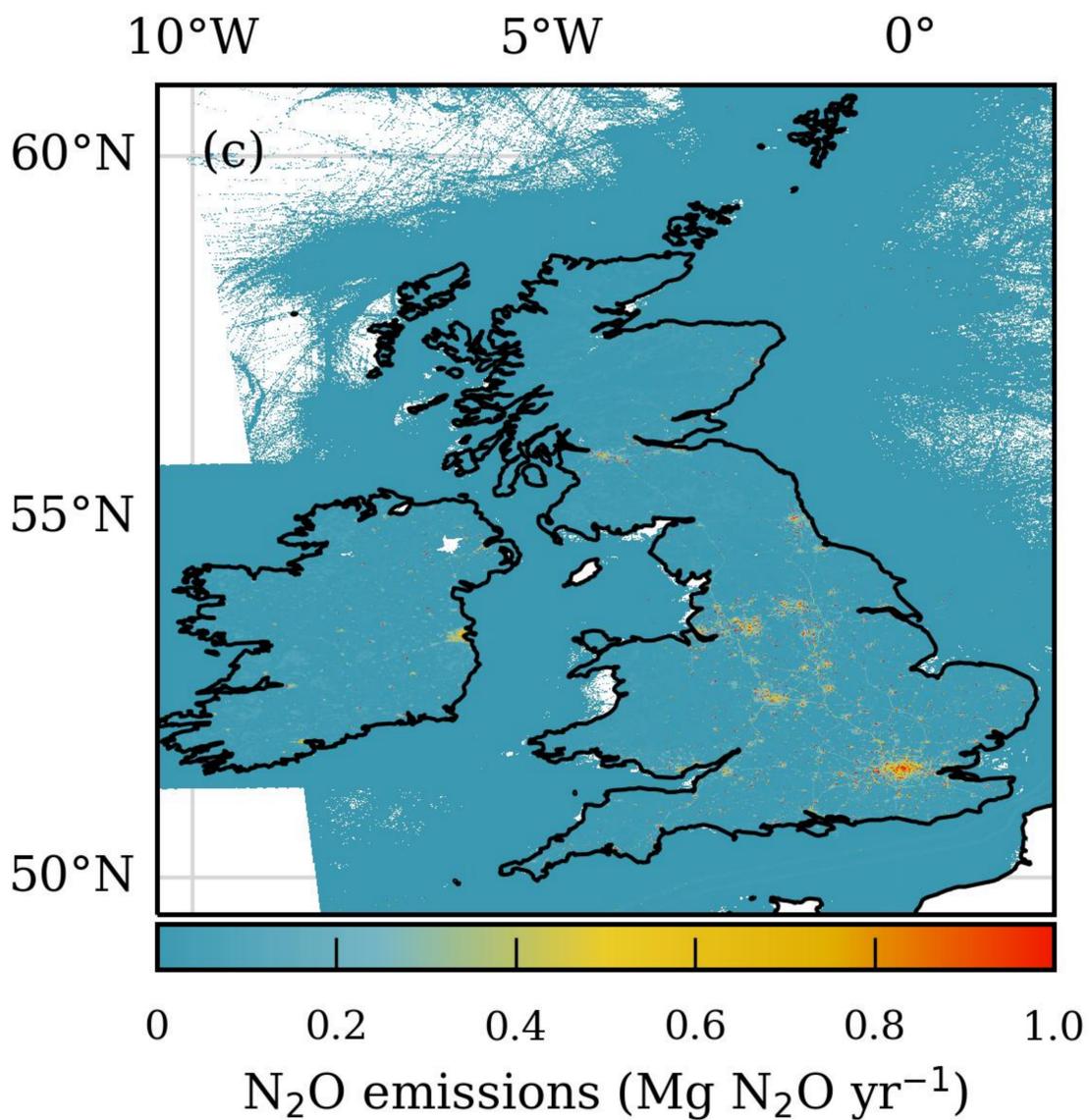
Total



Agriculture



Other



Marine

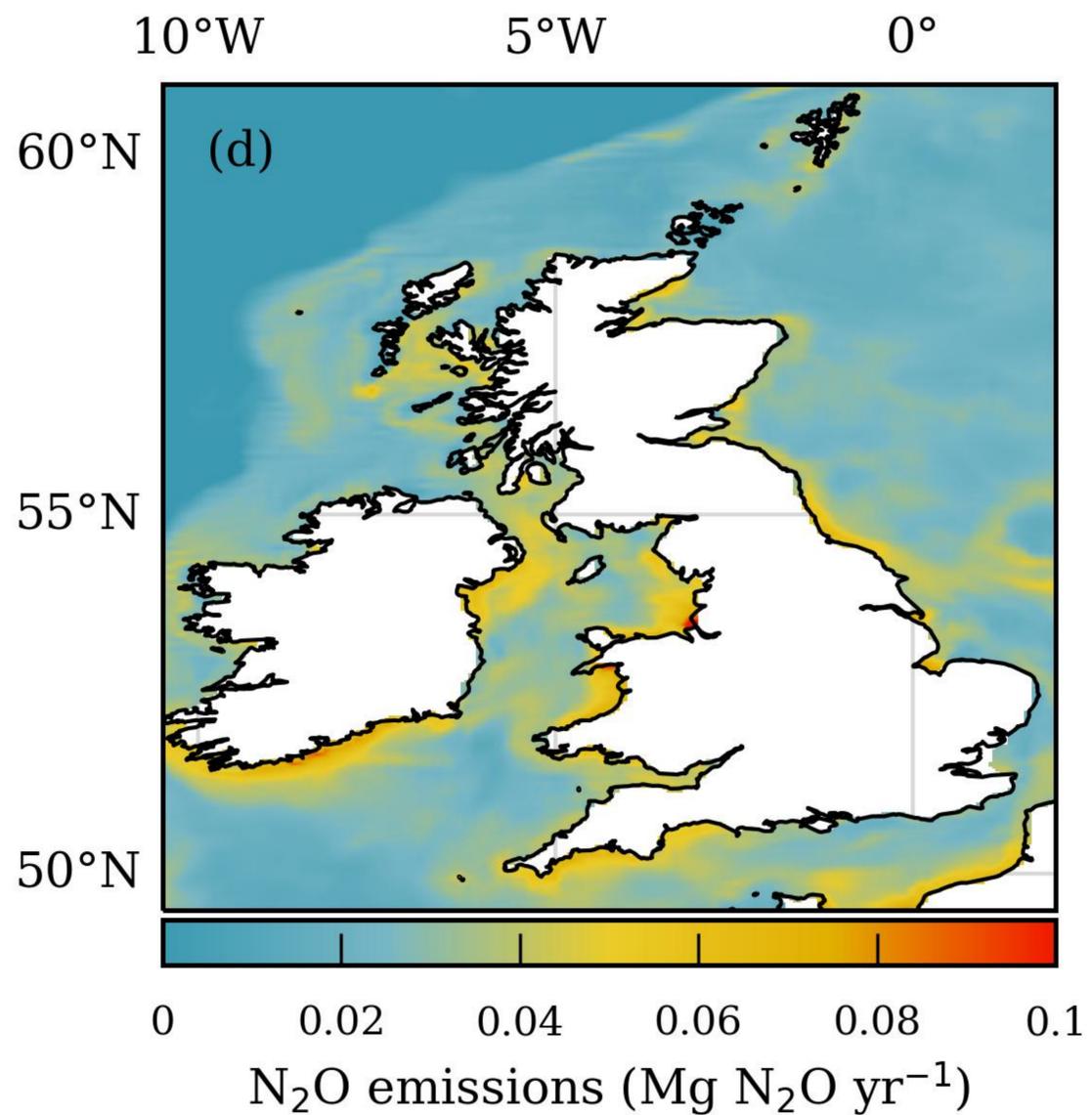


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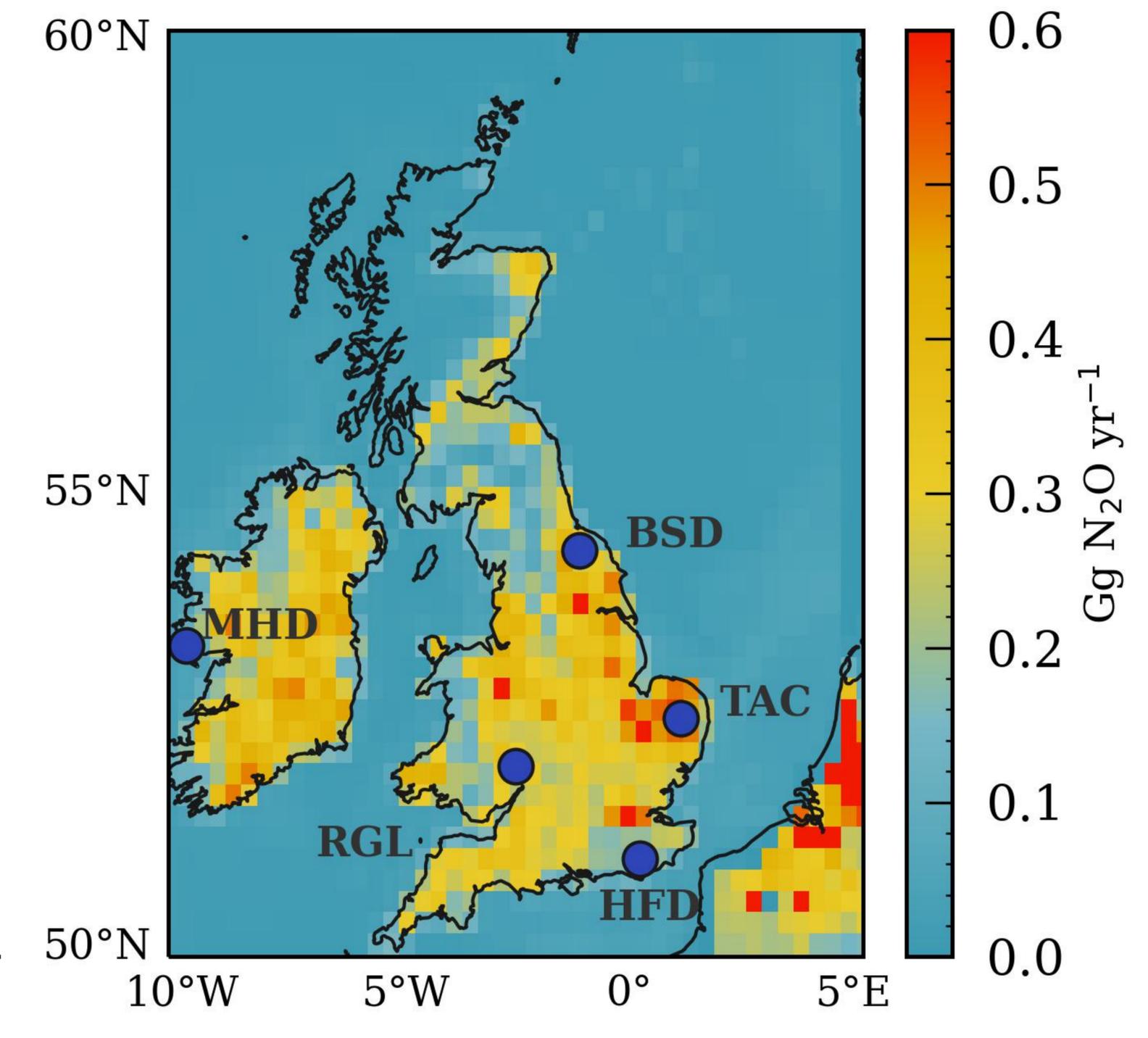
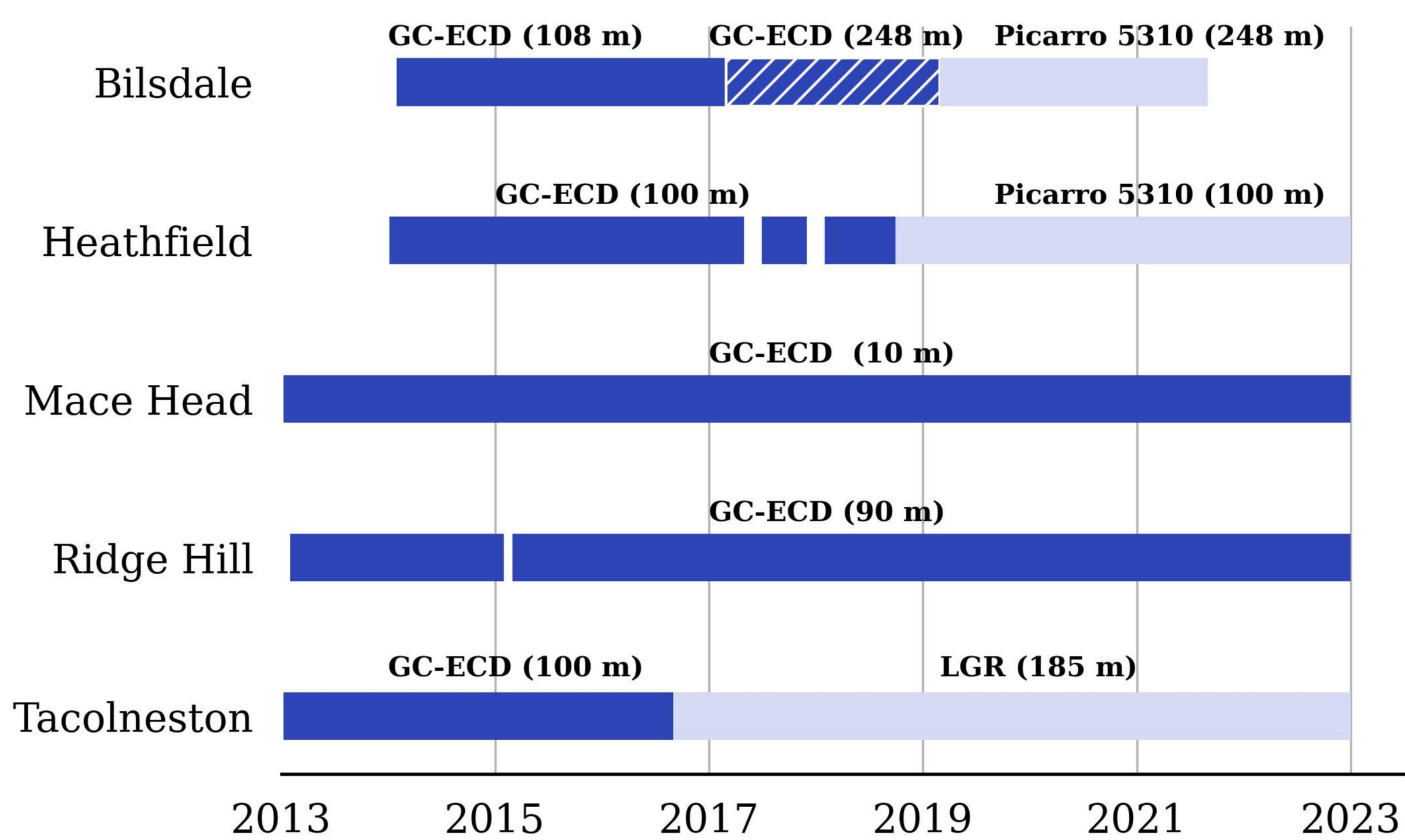


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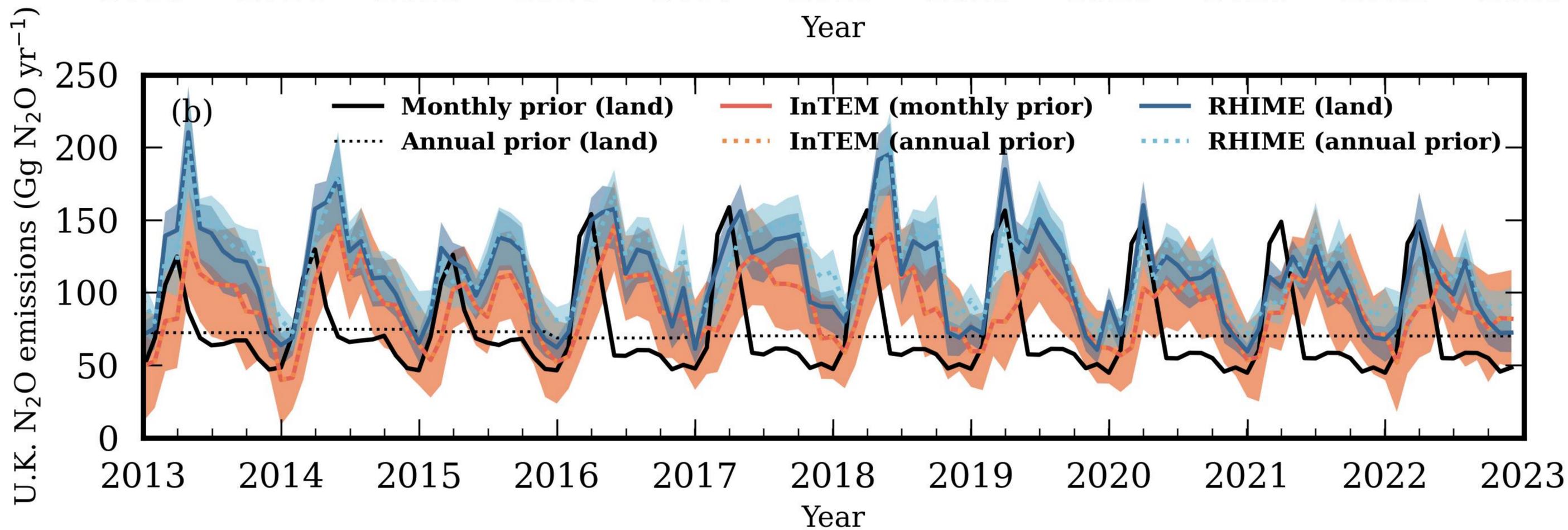
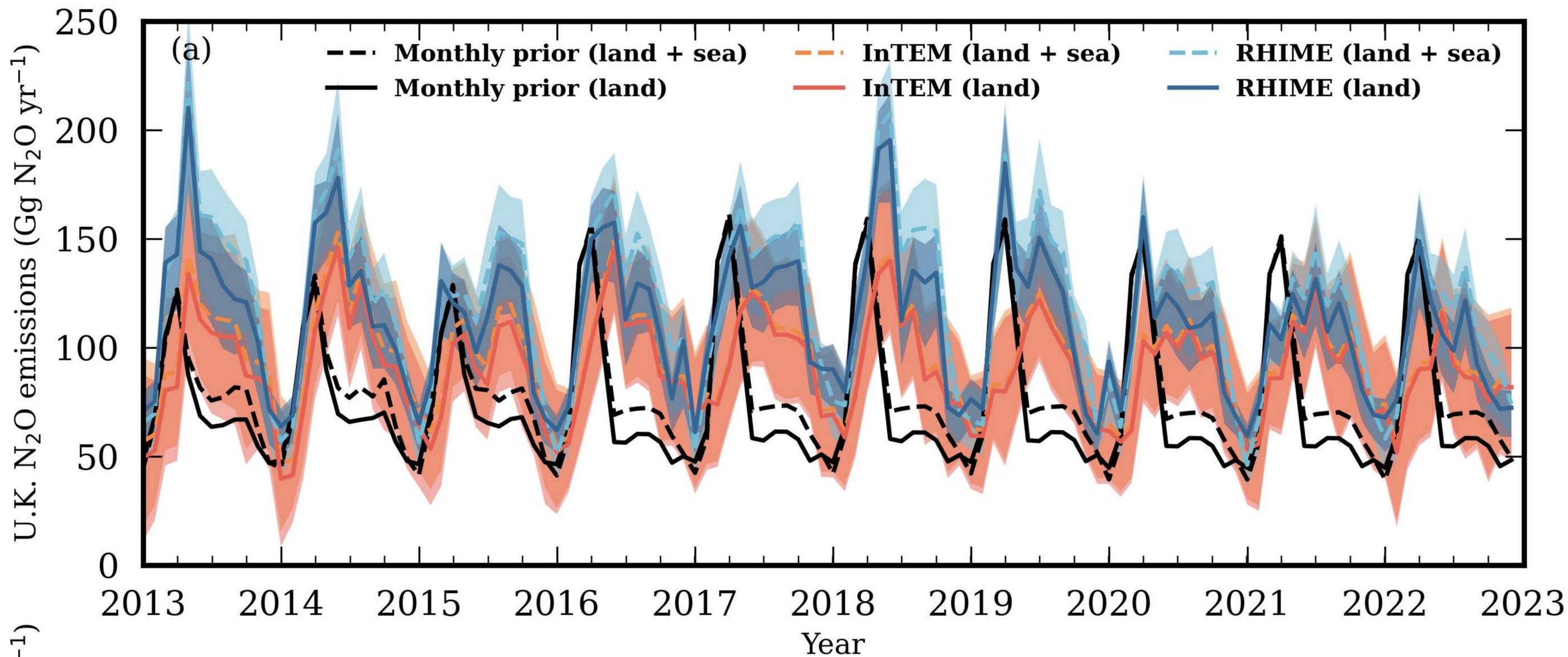


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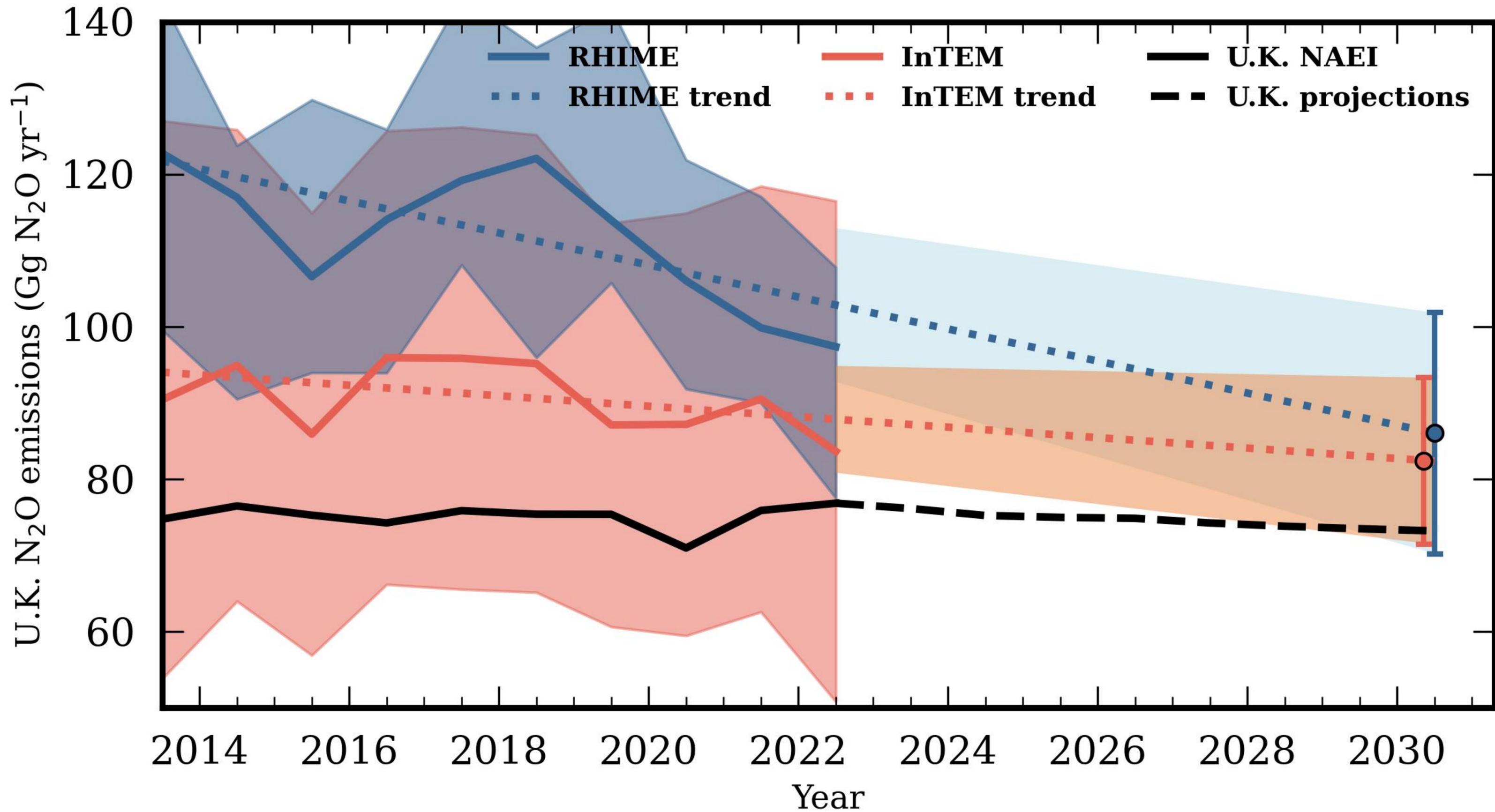


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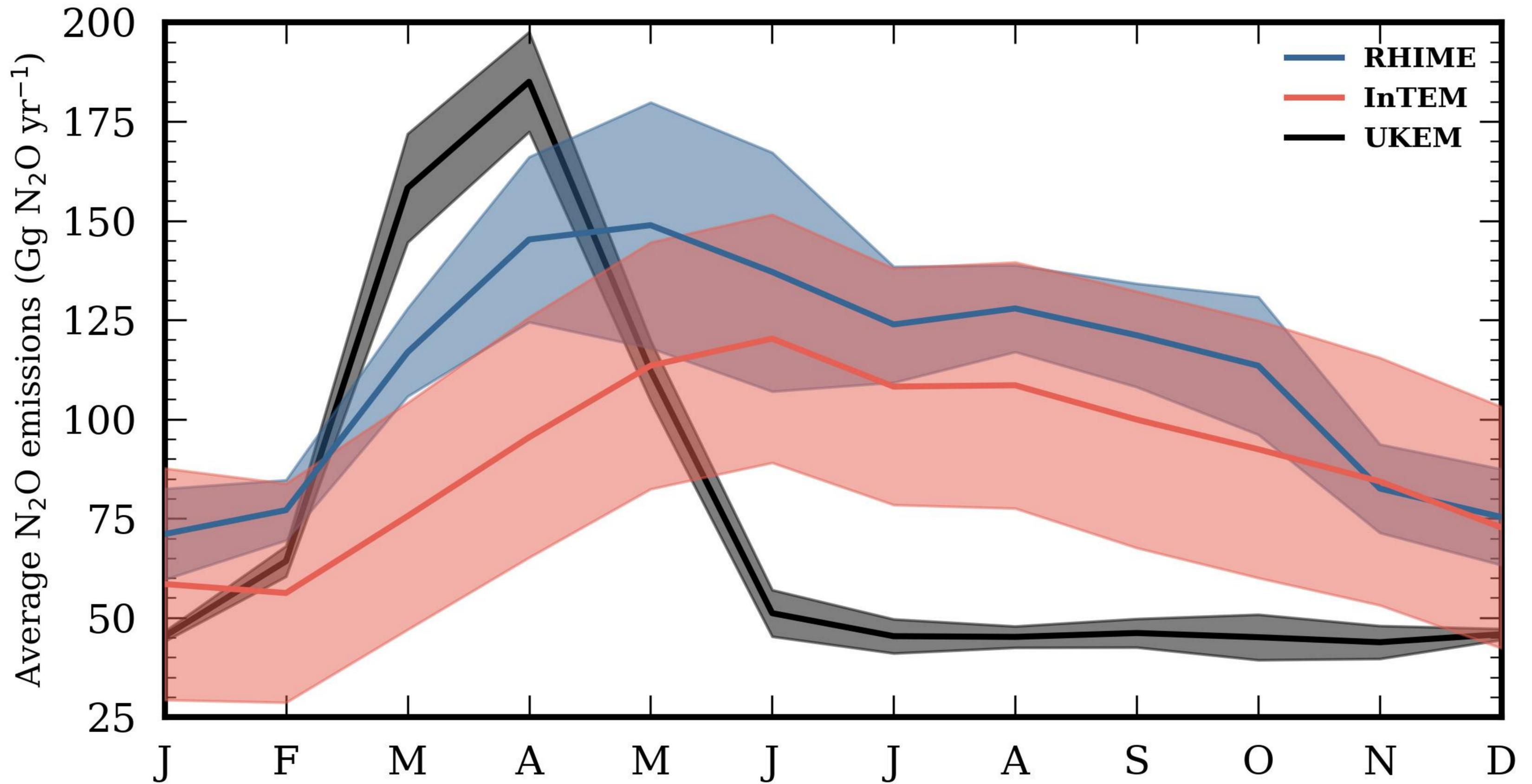


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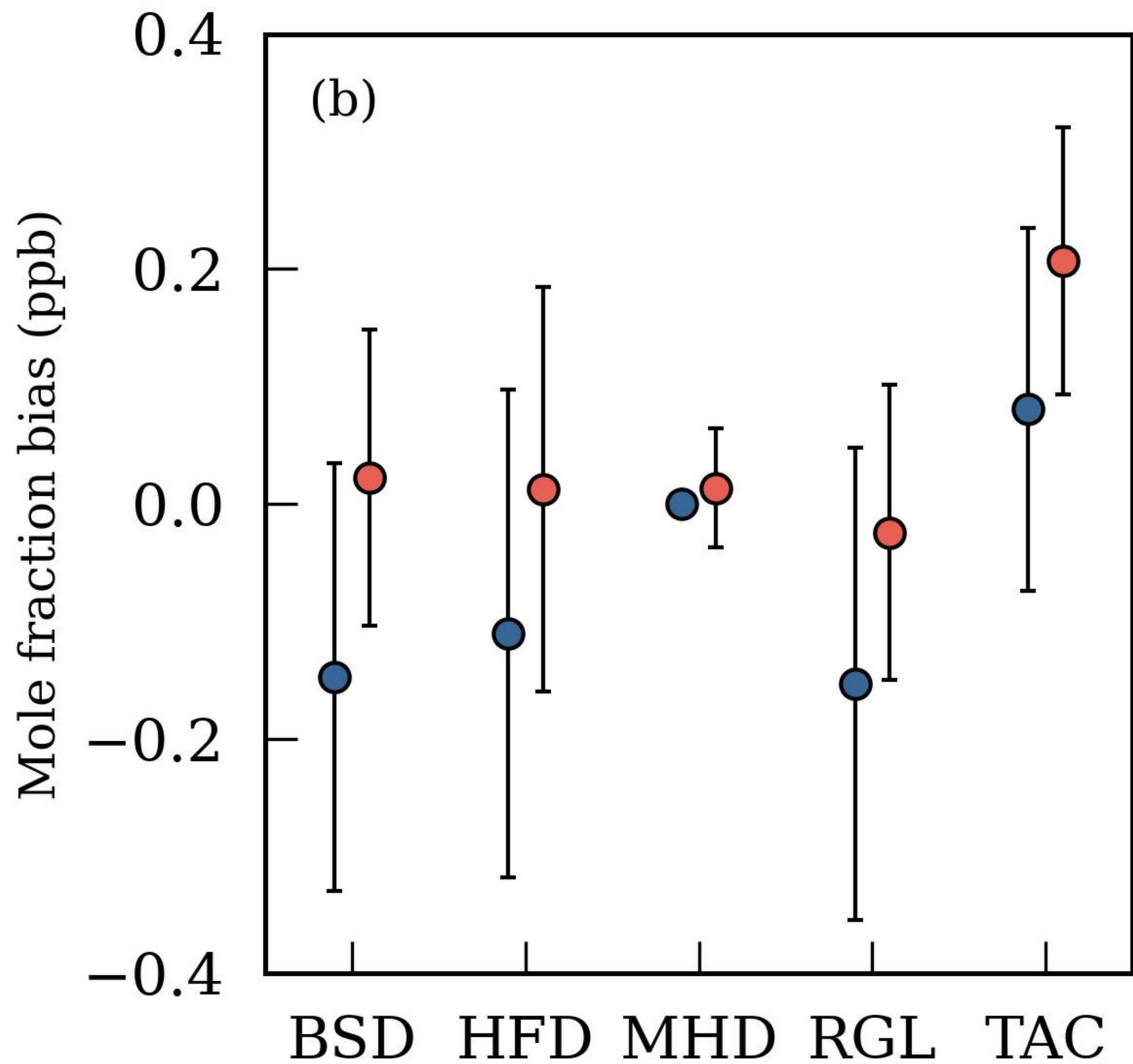
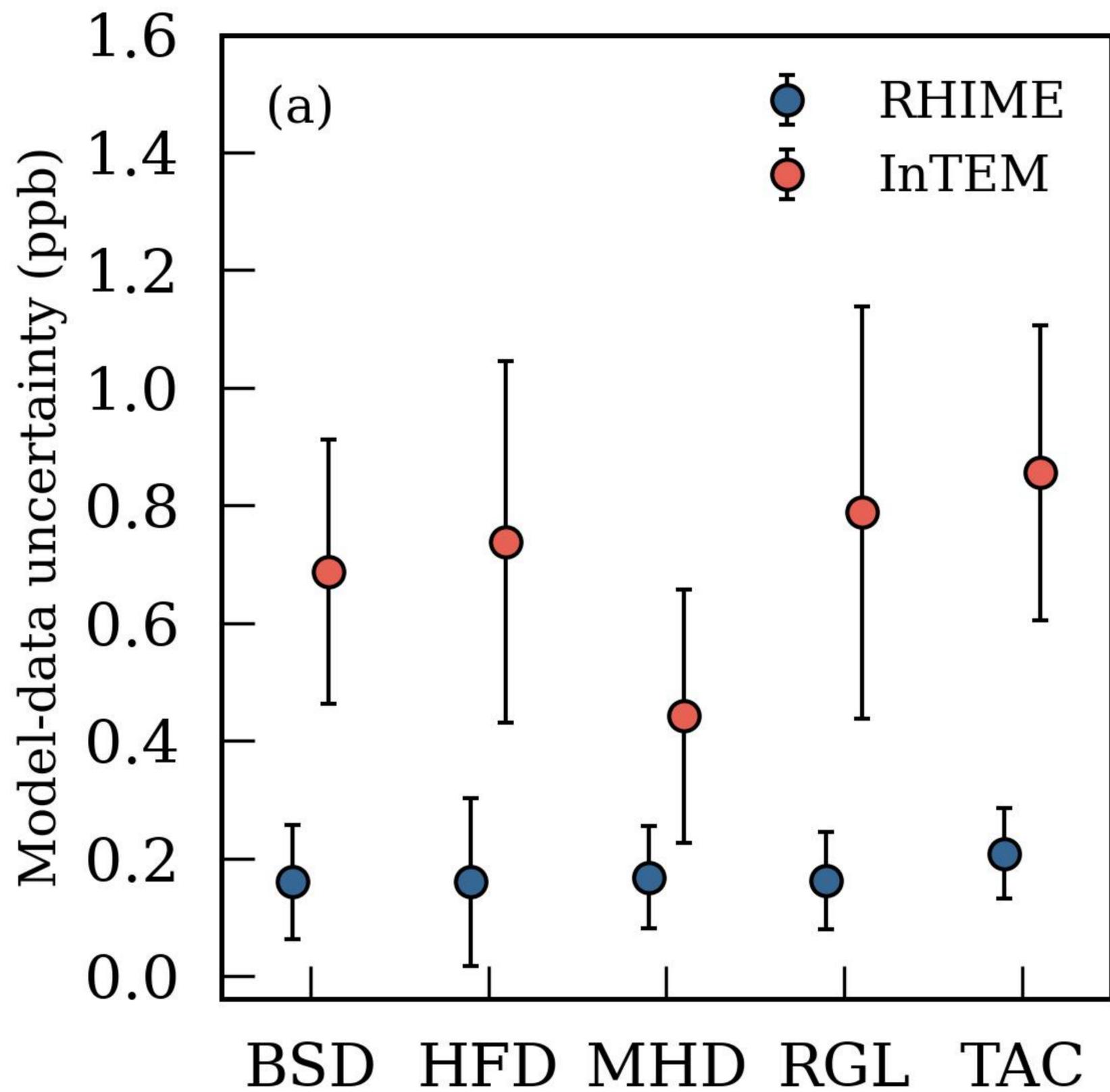


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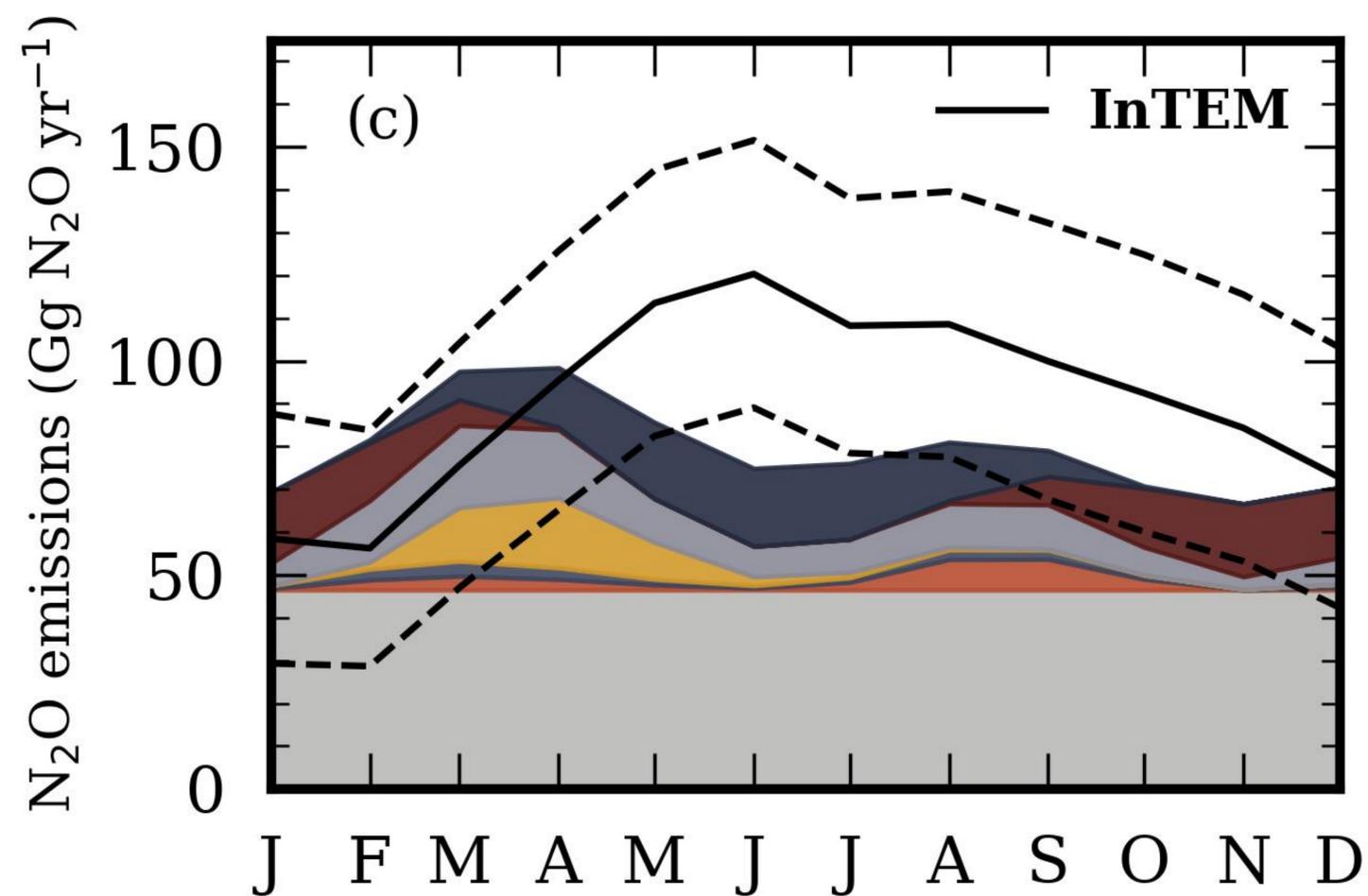
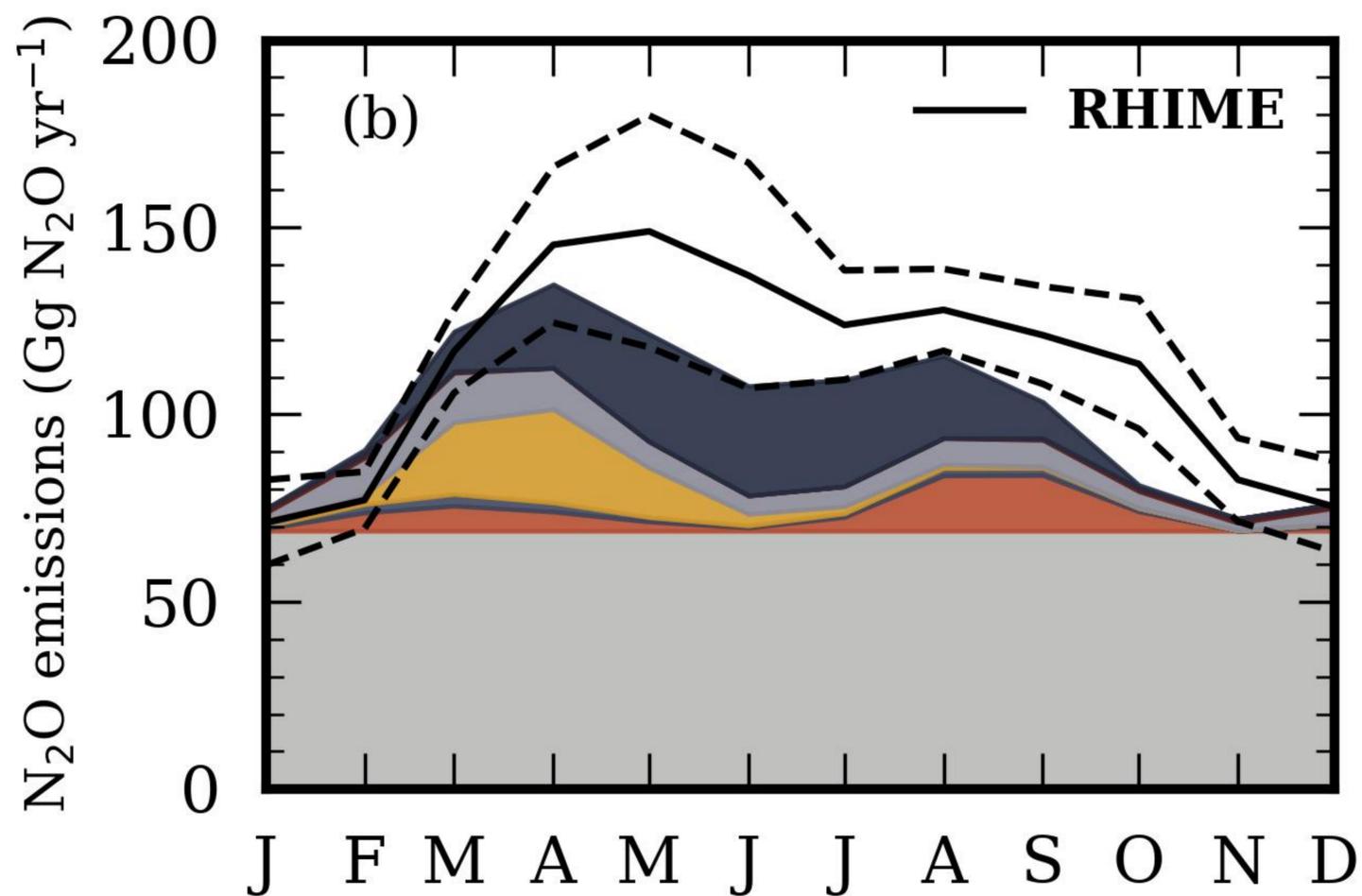
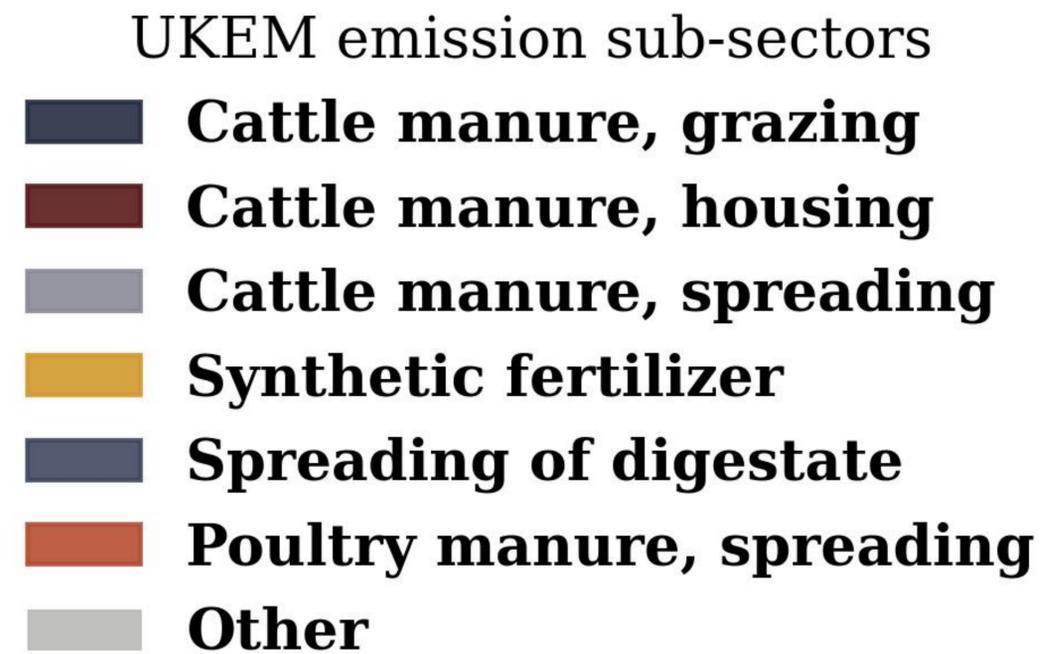
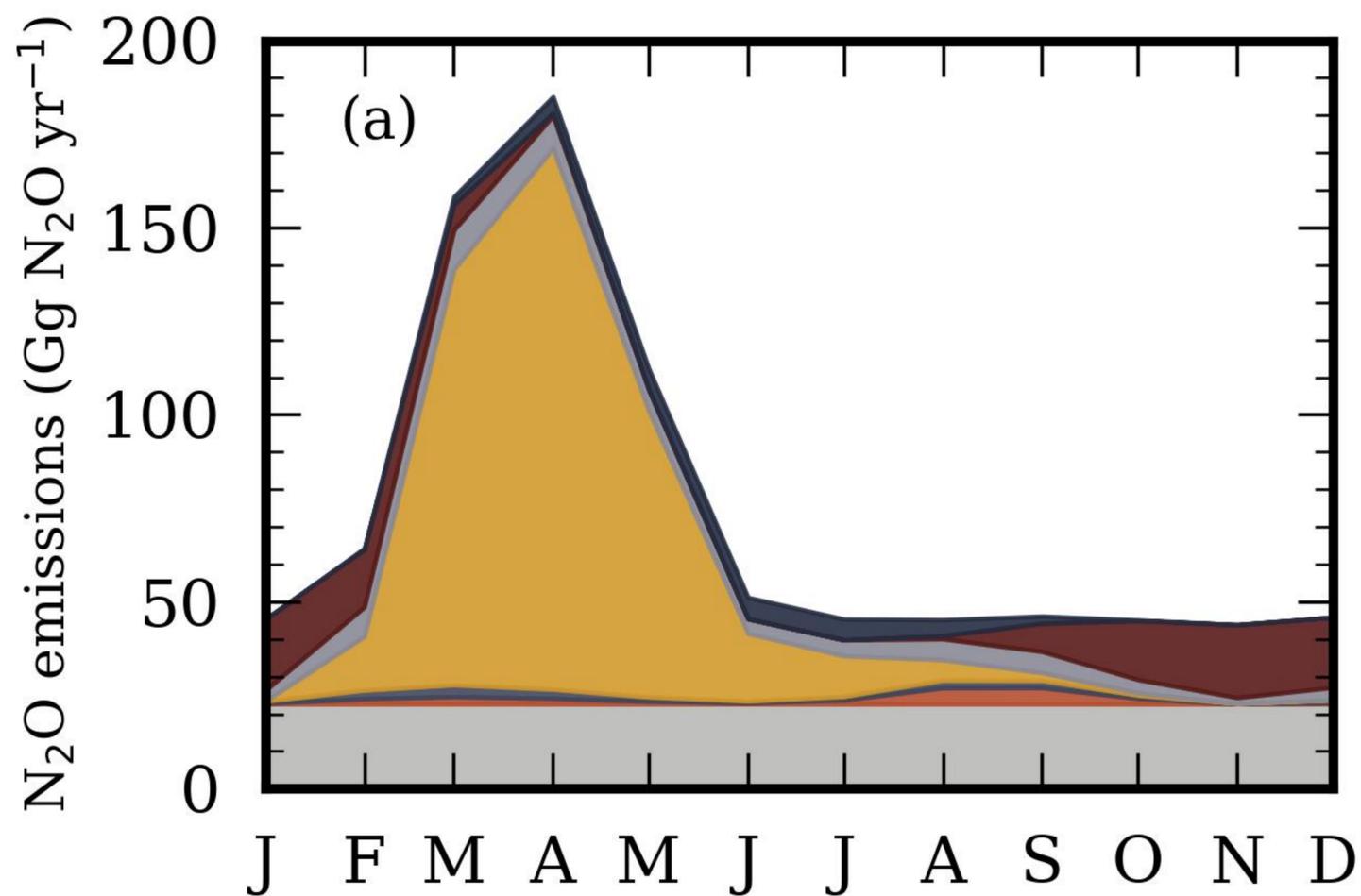


Figure S1.

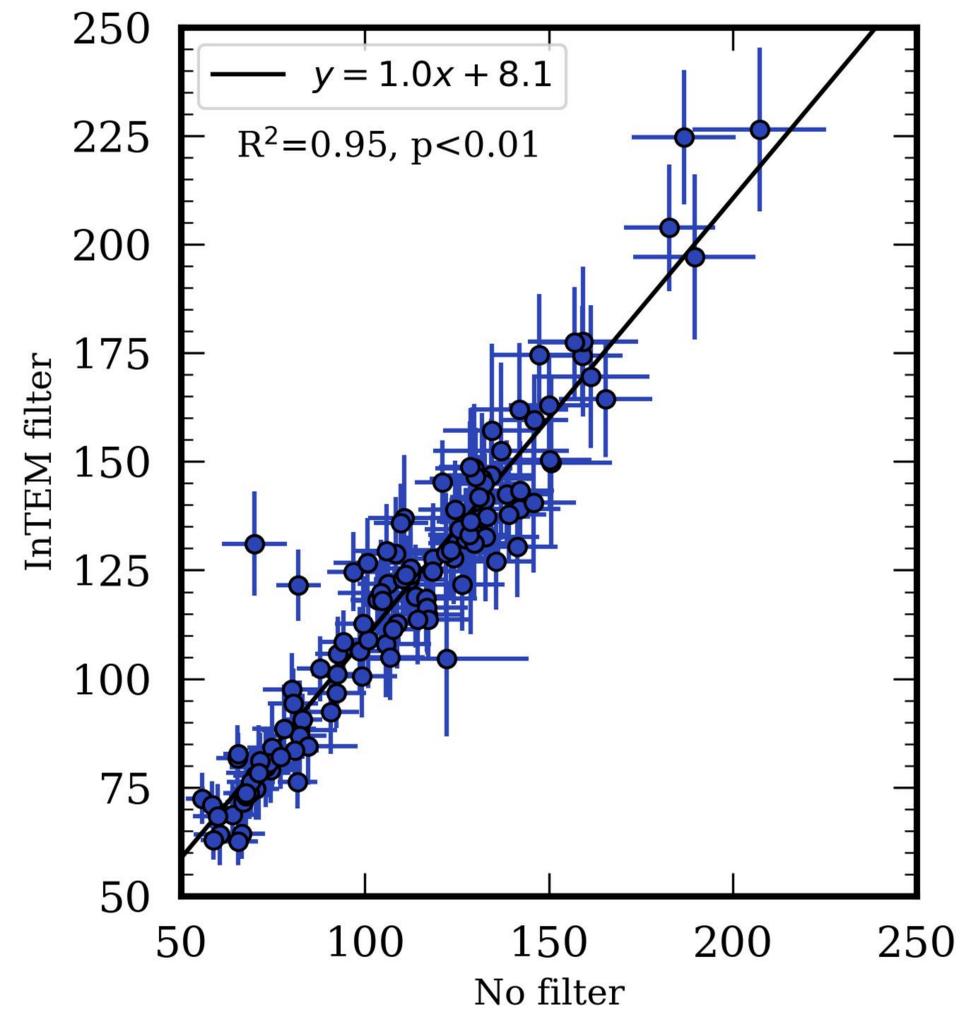
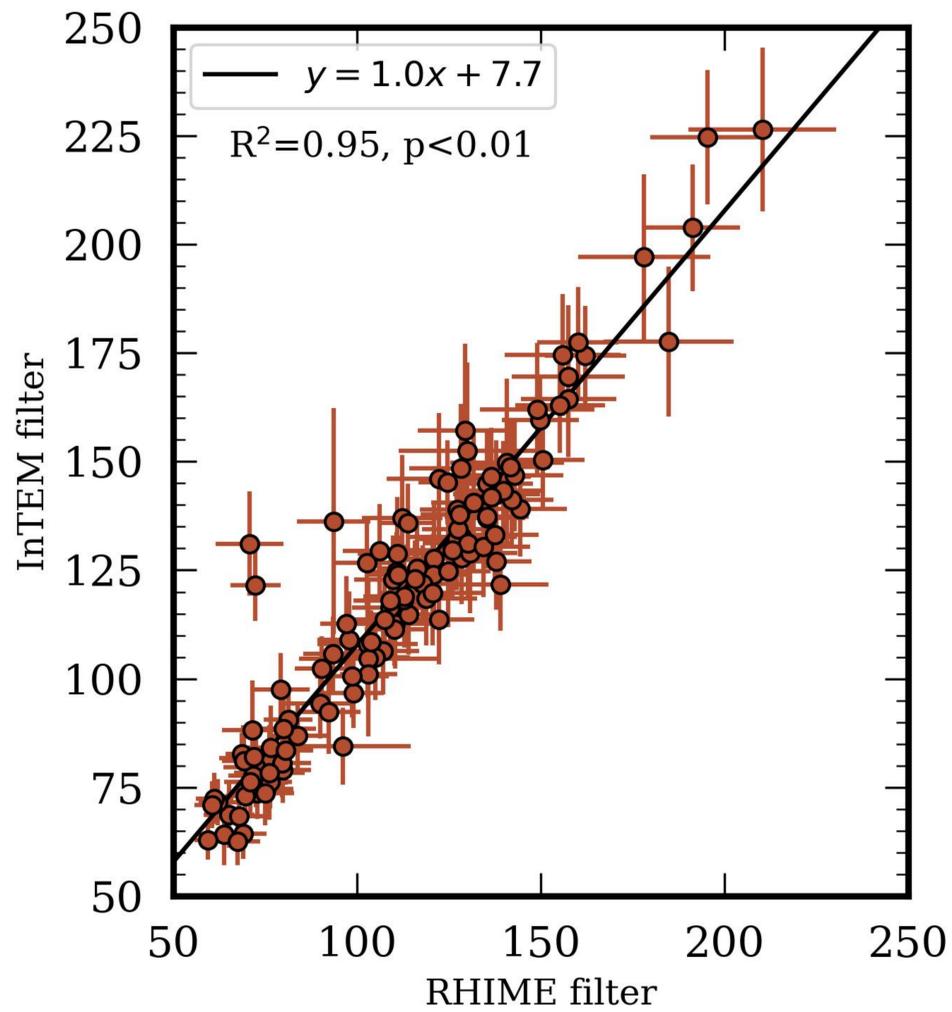
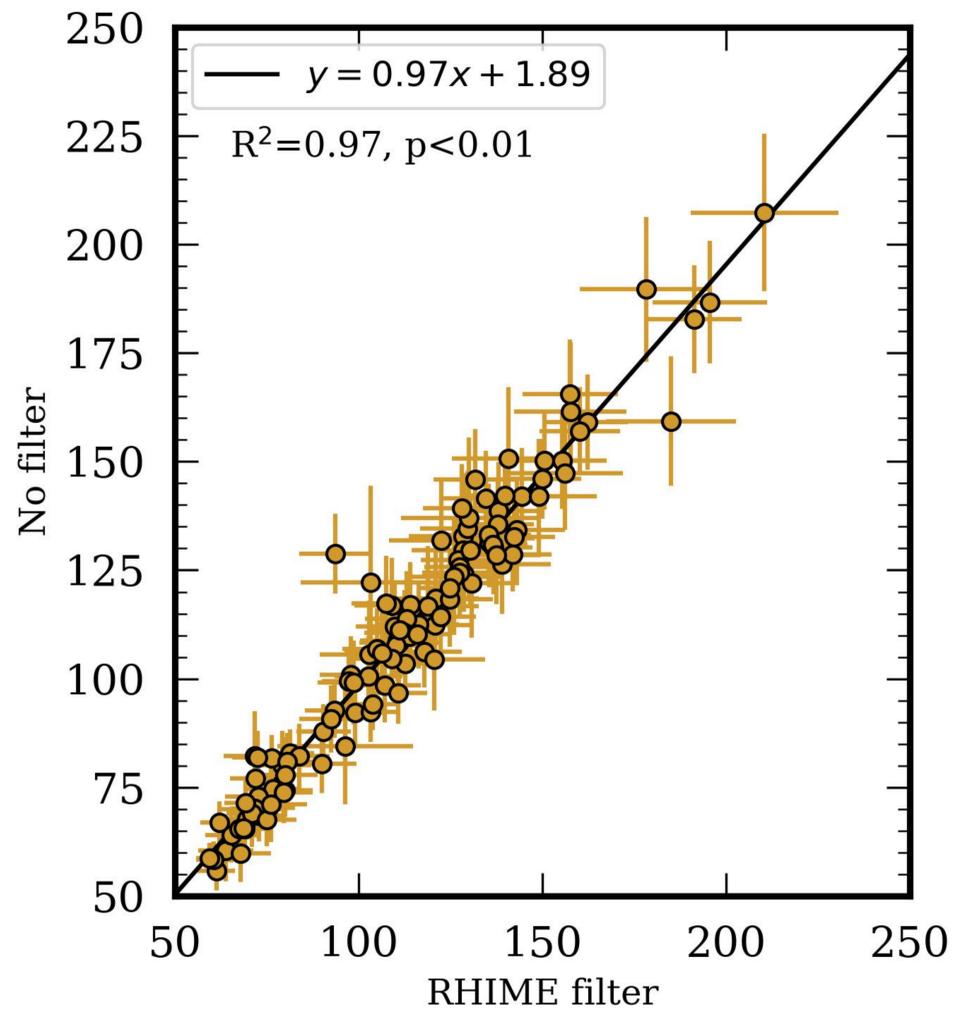
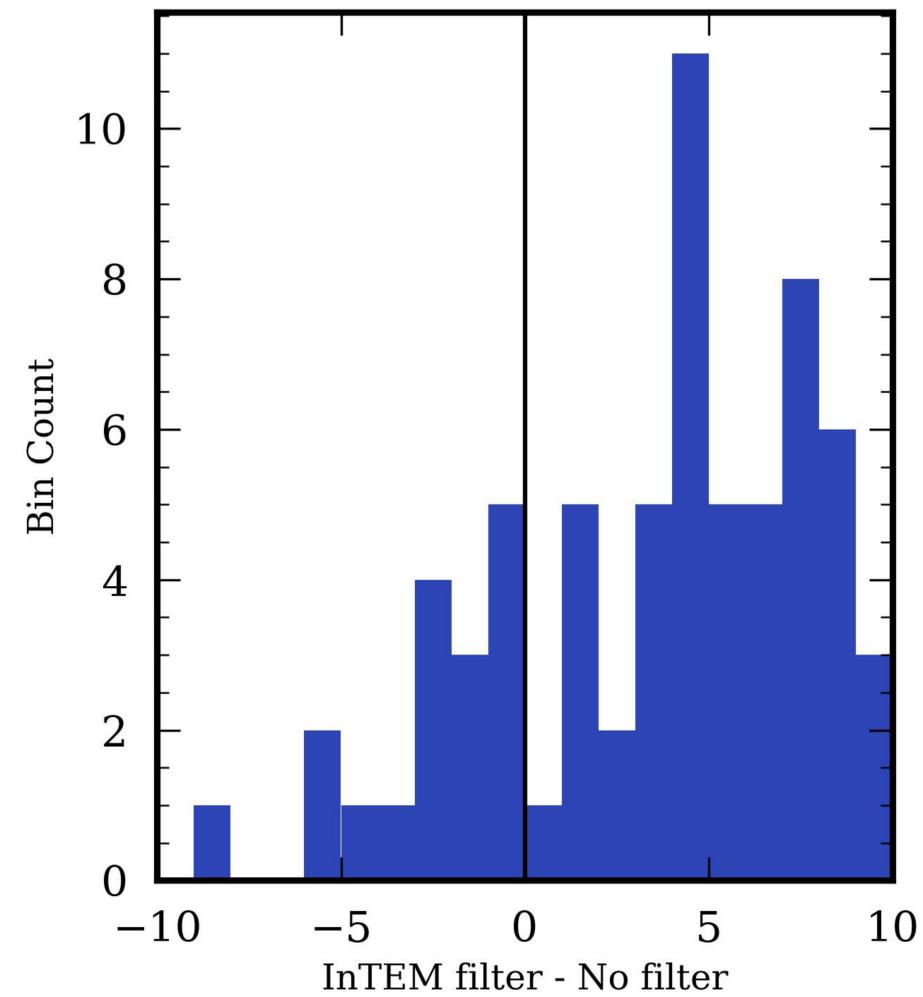
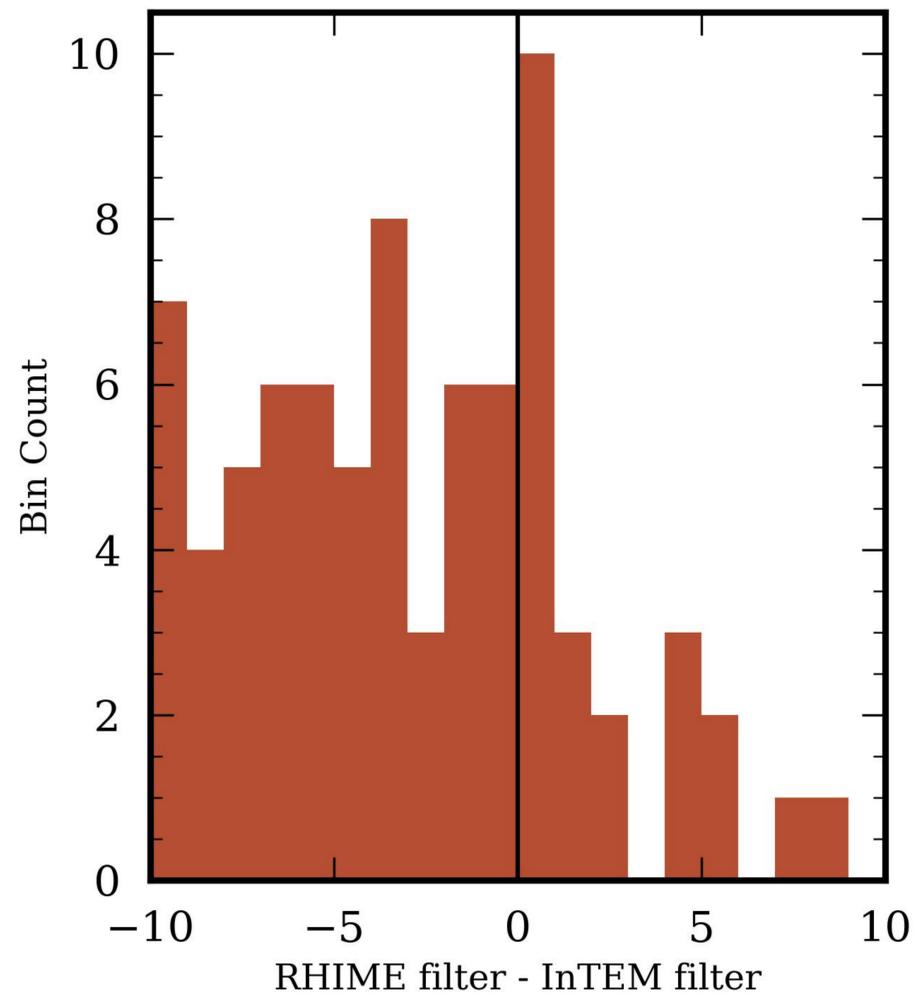
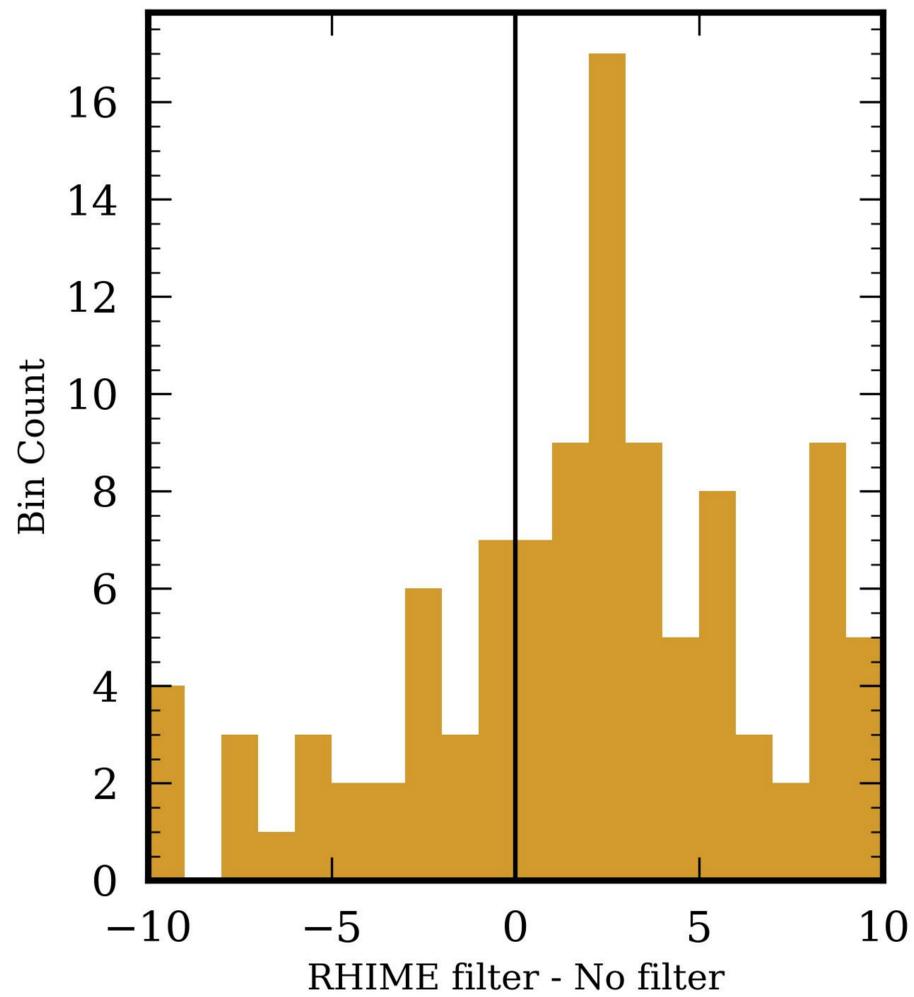


Figure S2.

