

# Neutral Composition Information in ICON EUV Dayglow Observations

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## Key Points:

- A comprehensive introduction to the ICON EUV dayglow observations is presented.
- Some dim emission features are identified as originating from  $O^+$  from similarity to known features.
- Emission near 87.8 nm follows  $N_2$  and, combined with 61.6 nm data, contains information about  $\Sigma O/N_2$ .

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## Abstract

Since the earliest space-based observations of Earth’s atmosphere, ultraviolet (UV) airglow has proven a useful resource for remote sensing of the ionosphere and thermosphere. The NASA Ionospheric Connection Explorer (ICON) spacecraft, whose mission is to explore the connections between ionosphere and thermosphere utilizes UV airglow in the typical way: an extreme-UV (EUV) spectrometer uses dayglow between 54 nm and 88 nm to measure the density of  $O^+$ , and a far-UV spectrograph uses the  $O$  135.6 nm doublet and  $N_2$  Lyman-Birge-Hopfield band dayglow to measure the column ratio of  $O$  to  $N_2$  in the upper thermosphere. Two EUV emission features,  $O^+$  61.6 nm and 83.4 nm, are used for the  $O^+$  retrieval; however, many other features are captured along the EUV instrument’s spectral dimension. In this study, we examine the other dayglow features observed by ICON EUV and demonstrate that it measures a nitrogen feature around 87.8 nm which can be used to observe the neutral thermosphere.

## Plain Language Summary

The ionosphere is a region of near-Earth space made up of plasma. NASA’s ICON mission seeks explore the factors influencing formation of the ionosphere and how it interacts with Earth and its atmosphere. One of the ways ICON does this is by measuring airglow: light released by the air in the upper atmosphere. This occurs with visible light, with the same shades seen in the aurora; it also occurs in the ultraviolet range, invisible to the human eye but visible to ICON instruments.

An imager is included on ICON to measure extreme-ultraviolet light, almost as energetic as X-rays. Certain atoms and molecules in the atmosphere are known to glow at specific wavelengths. By measuring the brightness of airglow at certain wavelengths, ICON is able see the structure of ionospheric oxygen. The instrument also measures dimmer emissions at other wavelengths, some of which are known to come from certain atmospheric species and others which are unknown or uncertain. Here we look at the other wavelengths and attempt to find their origins. We find that most are likely coming from oxygen. Interestingly, we find one that we think comes from nitrogen. This could be useful for measuring the abundance of molecular nitrogen in the upper atmosphere, a task currently performed by another instrument on the ICON spacecraft. We make a case for the practicality of this approach.

## 1 Introduction

Extreme ultraviolet (EUV) dayglow has long been a phenomenon of interest in the study of the terrestrial upper atmosphere and ionosphere. Beginning with the earliest rocket measurements in the late 1960s and early 1970s (e.g. (Young et al., 1968), (Donahue & Kumer, 1971)) which include the identification of the bright 83.4 nm  $O^+$  line produced by photoionization excitation of  $O$  (Carlson & Judge, 1973), EUV observations immediately provided a means of studying the ionosphere. Study of the terrestrial EUV dayglow continued in 1972 when the first spectral observations were taken from the lunar surface on the Apollo 16 mission (Carruthers & Page, 1976) and were followed by higher resolution spectra from rocket experiments (Christensen, 1976) (Gentieu et al., 1979) (Gentieu et al., 1984). The first satellite measurements, taken from the USAF STP 78-1 spacecraft (Chakrabarti et al., 1983), combined with the rocket measurements, form the much of the basis of the modern understanding of Earth’s EUV dayglow between 30 and 91 nm. EUV and far-UV (FUV) emissions longward of 91 nm have received much scrutiny, often due to the numerous atomic oxygen features found in those wavelength ranges (e.g. (Cotton et al., 1993)).

In the time intervening, some individual EUV features have seen extended study. One such example is the aforementioned 83.4 nm  $O^+$  line, which has been observed by multiple spacecraft (notably by SSULI and RAIDS prior to ICON) and has been used to retrieve altitude profiles of ionospheric  $O^+$  (Stephan, 2016). More recently, the 61.6 nm (sometimes referred to as 61.7 nm)  $O^+$  feature has been proven useful for ionospheric retrieval, in part due to being optically thin in the thermosphere, where the 83.4 nm line undergoes significant scattering and absorption by  $O^+$  in the ionosphere (Stephan et al., 2012). Additionally, the 58.4 nm  $He$  feature has been studied with some detail (e.g. (Bush & Chakrabarti, 1995), (Anderson et al., 1979)). However, outside of these few features, much of the remaining EUV spectrum has not seen much study since the initial spectroscopy experiments in the 1980s. As a result, although other lines have been observed and documented (Meier, 1991), little is known about the production mechanisms of several EUV dayglow features.

Thermospheric neutral composition has significant impact on the composition and density of the ionosphere. In particular, high  $O$  density enhances  $O^+$  production via photoionization of  $O$  (increasing total ionospheric density) and charge exchange with  $N_2^+$ ;

meanwhile high  $N_2$  density enhances one of the primary loss processes for  $O^+$ : the reaction  $O^+ + N_2 \rightarrow NO^+ + N$  (Prölss & Bird, 2004). As  $O^+$  is the dominant ion in the F-region, the column  $O/N_2$  ratio,  $\Sigma O/N_2$  (Strickland et al., 1995), is an important quantity in the study of ionosphere-thermosphere coupling (e.g. (Lean et al., 2011) (Mendillo et al., 2005)).  $\Sigma O/N_2$  is defined by (Strickland et al., 1995) as the ratio of the vertical column densities of  $O$  and  $N_2$  above a base where the  $N_2$  column density is  $10^{-17} \text{ cm}^{-2}$ . (Meier, 2021) discusses the column density ratio in some detail and provides the prescription for its computation.  $\Sigma O/N_2$  has long been retrieved using far UV (FUV) dayglow observations of the 135.6 nm  $O$  doublet and Lyman-Birge-Hopfield (LBH) bands, starting with the AIRS instrument on the Polar BEAR spacecraft (Evans et al., 1995) through TIMED-GUVI (Meier et al., 2015), for example.

Launched in late 2019, the NASA Ionospheric Connection Explorer (ICON) seeks to further characterize Earth’s low-mid latitude ionosphere and the factors that influence its formation and evolution (Immel et al., 2018). ICON hosts a suite of four instruments, three of which remotely sense atmospheric parameters using airglow in various bands. These include an EUV spectrometer, which remotely senses  $O^+$  profiles using the known 61.6 nm and 83.4 nm lines (Stephan et al., 2017), and an FUV imager, which retrieves daytime  $\Sigma O/N_2$  using the 135.6 nm oxygen and a band of LBH wavelengths (Stephan et al., 2018) (Meier, 2021).

In this work, we examine features other than 61.6 nm and 83.4 nm that are captured by the ICON EUV instrument and demonstrate that the emissions near 87.8 nm have potential utility for observing the neutral thermosphere. In Section 2.1, we delve into greater detail about the function of and data reported by the ICON EUV instrument. Section 2.2 follows with a brief description of the ICON FUV instrument with an emphasis on contrasting to the EUV instrument. In Section 3, we examine the brightness profiles of the 12 emission features reported in the Level 1 (L1) EUV data and identify the source of some of the emission features by comparing them to known lines in the dataset. In Section 4, we show that the ratio of the brightness of the 61.6 nm and 87.8 nm bins from the L1 EUV contains information about  $\Sigma O/N_2$ . In the concluding discussion, we state our proposed identification of the emission features in the EUV L1 data and make the case for retrieval of  $\Sigma O/N_2$  from EUV observations.

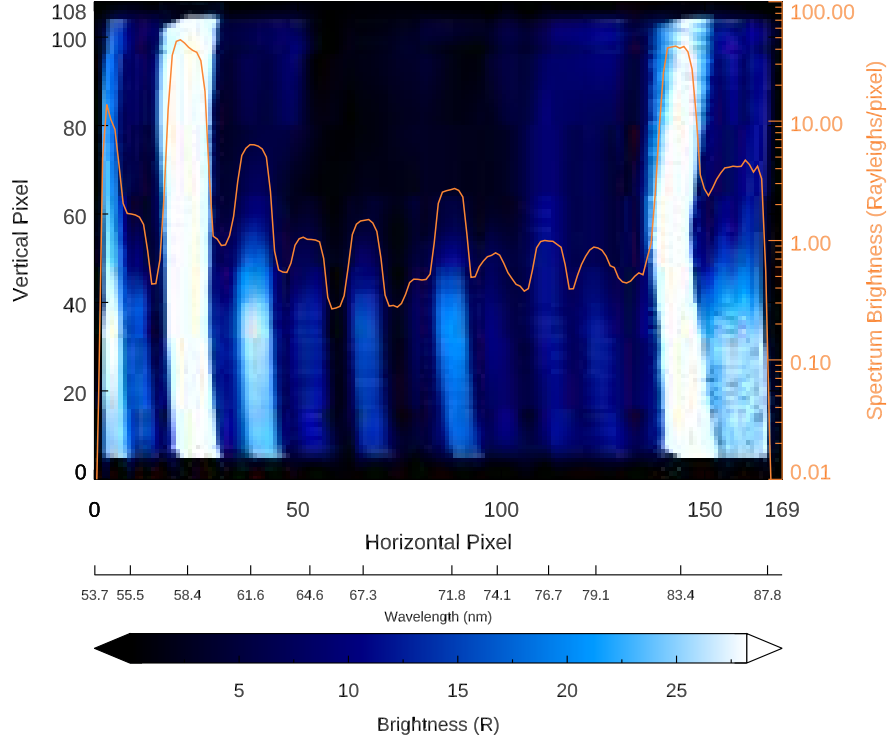
## 2 Data

### 2.1 ICON EUV

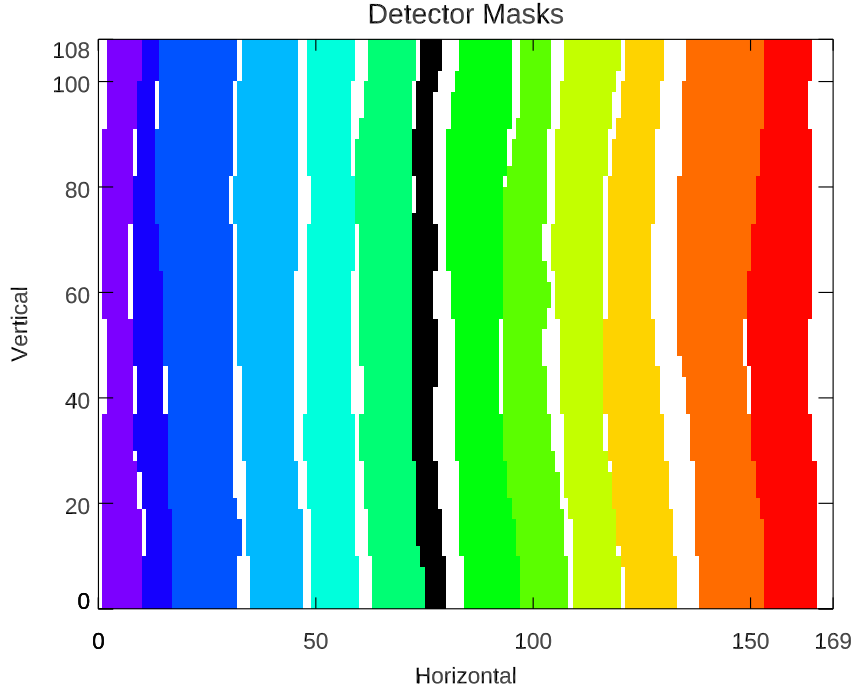
The ICON EUV instrument is a 1D imaging spectrometer, included on the ICON mission to produce  $O^+$  density profiles in the daytime ionosphere (Sirk et al., 2017) (Stephan et al., 2017). It does so by measuring altitude profiles of  $O^+$  airglow at 83.4 nm and 61.6 nm on the limb. During operation, EUV radiation from the atmosphere enters the instrument through a slit, illuminates a toroidal diffraction grating, and is focused onto a micro channel plate detector (MCP). The instrument was designed to have the necessary resolution to resolve the 61.6 nm  $O^+$  line from the bright 58.4 nm  $He$  line (Sirk et al., 2017). Low level output of the instrument is a 2D array of photon counts. One dimension is 108 pixels across and spatial, oriented along the local vertical on the ICON spacecraft. In normal science operations, the field of view spans tangent altitudes approximately 20-550 km. The second dimension is 169 pixels across and measures wavelength. The image is flat-field corrected and the counts are converted to brightness in Rayleighs resulting in an image similar to that in Figure 1, which is the sum of all images over the course of a single day. The tall and bright features on the left and right (around 25 and 150 pixels) are the 58.4 nm and 83.4 nm lines, respectively. The 61.6 nm line is seen near pixel 45, just to the right of the 58.4 nm line.

In order to convert the brightness of each pixel to brightness of the desired emission features, masks are applied to the detector image, as depicted in Figure 2. Each mask represents a wavelength bin, intended to capture a single emission feature on the 2.1 nm resolution of the instrument. The mask boundaries were set at the minimum brightness between features, based on the first week of instrument exposure. For each row, the mask determines which pixels belong to which of the 12 wavelength bins. The variation by row, seen as distortion of emission features in pixel space (easily seen in the 83.4 nm line in Figure 1, which would otherwise be a straight vertical feature) is caused by nonuniformity in the MCP electric field with some contribution from optical aberration. After a 12 second integration for each exposure, the pixels belonging to the same wavelength bin are summed, yielding the brightness measurement for that wavelength at a given vertical pixel.

In addition to being calibrated on the ground (Sirk et al., 2017), the instrument undergoes a monthly in-flight calibration. The absolute calibration is performed by com-



**Figure 1.** EUV detector image summed over the course of 10 June 2020. Overplotted, temporally averaged spectral brightness measured by ICON throughout 2020 for vertical pixels 35-40 (around 250-280 km tangent altitude, near peak brightness for most features) and 10-14 LT. The approximate wavelengths are indicated and are most accurate for the central row. Nonuniformity in the MCP electric field combined with optical aberrations cause the emission features to appear curved in pixel space.



**Figure 2.** A visualization of the mask used to generate the L1 EUV data from the detector image. The 12 line regions are shown in color. The black region around pixel 75 is the background region, which is used to estimate the background signal.

paring instrument response when pointed at the full moon to the spectrum brightness predicted by the combination of lunar geometric albedo and measurements of the solar ultraviolet spectrum from the Solar Dynamics Observatory’s Extreme Ultraviolet Variability Experiment (EVE). A monthly flat-fielding calibration is performed by pointing the instrument at the terrestrial disk with the spatial dimension aligned in the in-track direction. The estimated systematic error in the L1 data from uncertainties in the ground calibration, lunar albedo measurements, EVE solar EUV measurements, and flat-fielding is 21%.

The ICON Level 1 (L1) EUV data product reports the calibrated airglow brightness along the vertical axis over the 12 second integration in 12 wavelength bins, which are tabulated in Table 1. The bins were named according to the wavelength of the airglow feature which was expected to dominate. On the edges of the detector, the epony-

156 mous wavelengths lie outside of the nominal wavelength range of the bin because the spec-  
 157 tral features chosen, while not centered on the detector, were believed to be detected and  
 158 to dominate the bin through instrumental broadening. Among these bins are three in-  
 159 tended to capture the  $O^+$  features at 61.6 nm and 83.4 nm and the  $He$  line centered at  
 160 58.4 nm. These lines are bright and well-known, so we will refer to them as the primary  
 161 features. In higher level data products, an inversion is performed on the 61.6 nm and  
 162 83.4 nm brightness to obtain  $O^+$  density (Stephan et al., 2017). The remaining nine bins  
 163 are likewise centered around observed wavelengths and most bins contain known emis-  
 164 sions, as indicated in Table 1. We will refer to these as secondary features. They are not  
 165 part of the ICON mission objectives, but the data are there as a consequence of the res-  
 166 olution and passband required to measure the primary features. Nonetheless, ICON has  
 167 collected a wealth of data about these features, which we examine in this study. We use  
 168 the Version 03 L1 EUV Data collected over the year 2020.

169 In our discussion of these data, we will use the tangent altitudes corresponding to  
 170 each vertical pixel of the detector as the independent variable. This is a geometrical def-  
 171 inition and does not define a unique altitude for the emitting volume elements. Indeed,  
 172 the altitude of unit optical depth along the ICON line-of-sight is typically around 250  
 173 km (due to  $O$ ,  $O_2$ , and  $N_2$  absorption, see Section 3.2), so emissions at lower tangent  
 174 altitudes come from foreground emitters above 250 km or so.

175 In order to protect the instrument from damaging particle radiation, the EUV in-  
 176 strument powers down when ICON passes through the South Atlantic Anomaly (SAA).  
 177 As a result, our dataset includes limited observations in and around the SAA. This case  
 178 excepted, ICON’s orbit allows for full latitude-LT (local time) coverage over its 41 day  
 179 coverage cycle. To reduce latitude-LT coverage effects, which are expected to be dom-  
 180 inant over a slight seasonal bias, this analysis is restricted to 8 full coverage cycles from  
 181 2020 DOY 1-328 (1 January to 23 November).

## 182 2.2 ICON FUV

183 The ICON FUV instrument is a spectrographic imager, which observes airglow in  
 184 two wavelength bands (Immel et al., 2018)(Mende et al., 2017). The first passband, dubbed  
 185 the long wave (LW), captures a portion of the  $N_2$  LBH emission (emitted by excited  $N_2$   
 186 during the daytime) centered at 157 nm. The second, dubbed the short wave (SW), cap-



**Table 1.** Overview of the ICON EUV Bins and Known Airglow Emissions

EUV Bin Name	Dominant Source (Notes)	ICON Limb Peak Brightness (R)	Nominal Wavelength Range	Documented Airglow Lines	Limb Brightness (R)
53.7 nm	$O^+$ ( <i>He</i> contamination at high alt.)	51	54.0-55.1 nm	53.7 nm <i>He</i> 53.8 nm $O^+$ 53.9 nm $O^+$	86 94
55.5 nm	$O^+$	9.2	55.3-56.5 nm	55.5 nm $O^+$	75
<b>58.4 nm</b>	<i>He</i>	<b>450</b>	<b>56.7-59.9 nm</b>	<b>58.4 nm <i>He</i></b>	<b>58</b>
<b>61.6 nm</b>	$O^+$	<b>59</b>	<b>60.6-62.8 nm</b>	<b>61.6 nm <math>O^+</math></b>	<b>103</b>
64.6 nm		9.0	63.6-65.6 nm	64.4 nm $O^+$ 64.5 nm $N^+$	
67.3 nm	$O^+$	12	66.1-68.4 nm	67.1 nm $N^+$ 67.3 nm $O^+$	18
71.8 nm	$O^+$	23	70.6-72.5 nm	71.8 nm $O^+$	42
74.1 nm		6.0	72.9-74.8 nm	74.5 nm $O$ , $N^+$	20
76.7 nm		7.5	75.6-77.5 nm		
79.1 nm	$O^+$	8.4	77.7-80.0 nm	79.7 nm $O^+$ , $N_2$	28
<b>83.4 nm</b>	$O^+$ (Optically thick)	<b>630</b>	<b>81.2-84.1 nm</b>	<b>83.4 nm <math>O^+</math></b>	<b>580</b>
87.8 nm	$N$ and/or $N^+$ (low altitudes)	89	84.5-87.2 nm	87.5 nm $N$ 87.8 nm $O$	40 33

Primary lines are shown in **bold**. The second column contains the dominant source of airglow, identified in Section 3. The third column contains mean peak brightness on the limb over 382,404 ICON EUV exposures between 10-14 LT from 1 January 2020 to 23 November 2020. The 53.7 nm and 87.8 nm bins are on the edge of the detector and are known to capture emissions outside of the nominal wavelength range, so relevant known emissions are included even if out of range. Additionally, the edge bins are known to suffer from vignetting of 50-75%. The fifth and sixth columns contain previously published intensities from limb measurements made at 200-250 km, if available. (Gentieu et al., 1984) (Chakrabarti et al., 1983) These measurements were made in 1980 near solar maximum, while the 2020 ICON measurements were made at solar minimum, so the brightness of the former tends to be higher.

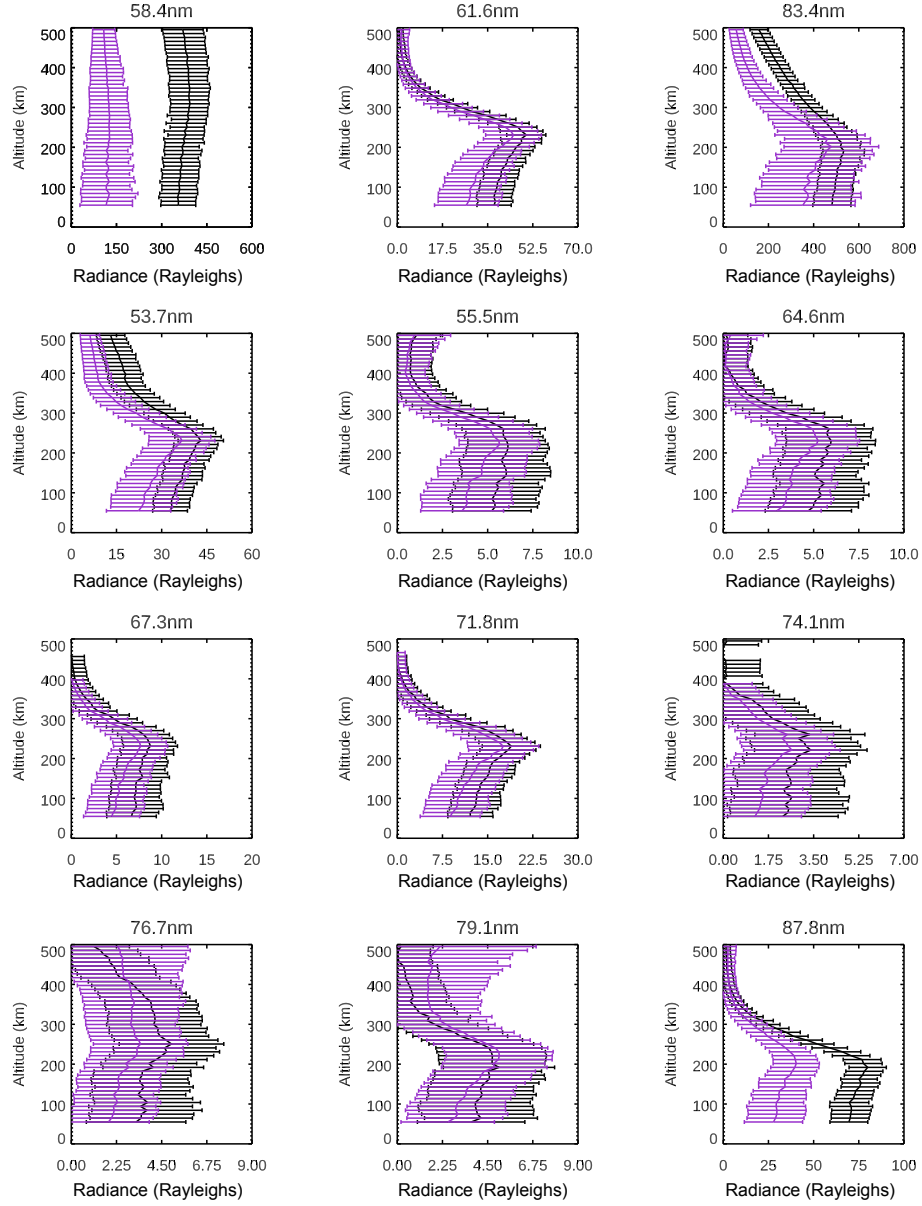
tures the 135.6 nm oxygen doublet, primarily emitted by excited  $O$  during the daytime and by  $O^+$  recombination at night, as well a significant (20%) contribution from another portion of the  $N_2$  LBH band. The LBH contamination in the SW channel is removed in the process of retrieving the daytime Level 2 FUV product,  $\Sigma O/N_2$  ratio (Stephan et al., 2018).

While FUV observations are not the subject of this study, they are introduced as means of testing the utility of EUV observations for retrieval of  $\Sigma O/N_2$ . We do this by comparing the EUV observations to the brightness ratio SW/LW. The disk  $\Sigma O/N_2$  data reported in the Level 2 FUV product is not suitable for comparison because the disk pixels used for the retrieval are displaced from the EUV limb tangent points by about 15 degrees of latitude. However, the EUV and FUV fields of view overlap on the limb in the 50-140 km tangent altitude range. We make use of this by examining the L1 FUV data, which reports the SW and LW brightness profiles before the removal of the LBH contribution to the SW channel and the inversion to  $\Sigma O/N_2$ . These measurements are taken at a 12 second cadence, the same as that for the EUV instrument, although the two are not synchronized. During daytime science operations, the FUV and EUV fields of view are aligned horizontally.

### 3 EUV Dayglow Characteristics

#### 3.1 Altitude Profiles and Primary features

Figure 3 contains altitude-binned average profiles for each wavelength reported in the L1 EUV data for high and low solar zenith angles (SZA). We use altitude bins rather than pixel number since the tangent altitude of an individual pixel experiences some small variation due in part to the non-spherical shape of the Earth. We chose a bin size of 10 km, which is small compared to the scale height in the thermosphere (around 50 km, 30 km, and 30 km for  $O$ ,  $O_2$ , and  $N_2$ , respectively) and larger than the tangent altitude spacing between pixels (2-9 km, with closer spacing at higher tangent altitudes). The former condition is important for capturing emission structure while the latter ensures that each bin is populated for each exposure. Additionally, average midday peak brightness for each emission feature on the limb and average brightness on the disk from EUV nadir-pointing calibrations are included in Table 1. First, we discuss the observations of the



**Figure 3.** Averaged altitude profiles of ICON L1 EUV wavelength bin for low (0-20, black) and high (75-90, purple) SZA, from 2020 DOY 1-328. Error bars represent the standard deviation in each altitude bin, which exceeds the Poisson error and the systematic error reported in the L1 EUV data. The top row contains the primary lines. The EUV L1 background subtraction allows for (small) negative values to be reported in the L1 product, which is why the 74.1 nm line, for example, trails off of the plot scale at high altitudes.

primary features at 58.4 nm, 61.6 nm, and 83.4 nm, since their behavior is well understood.

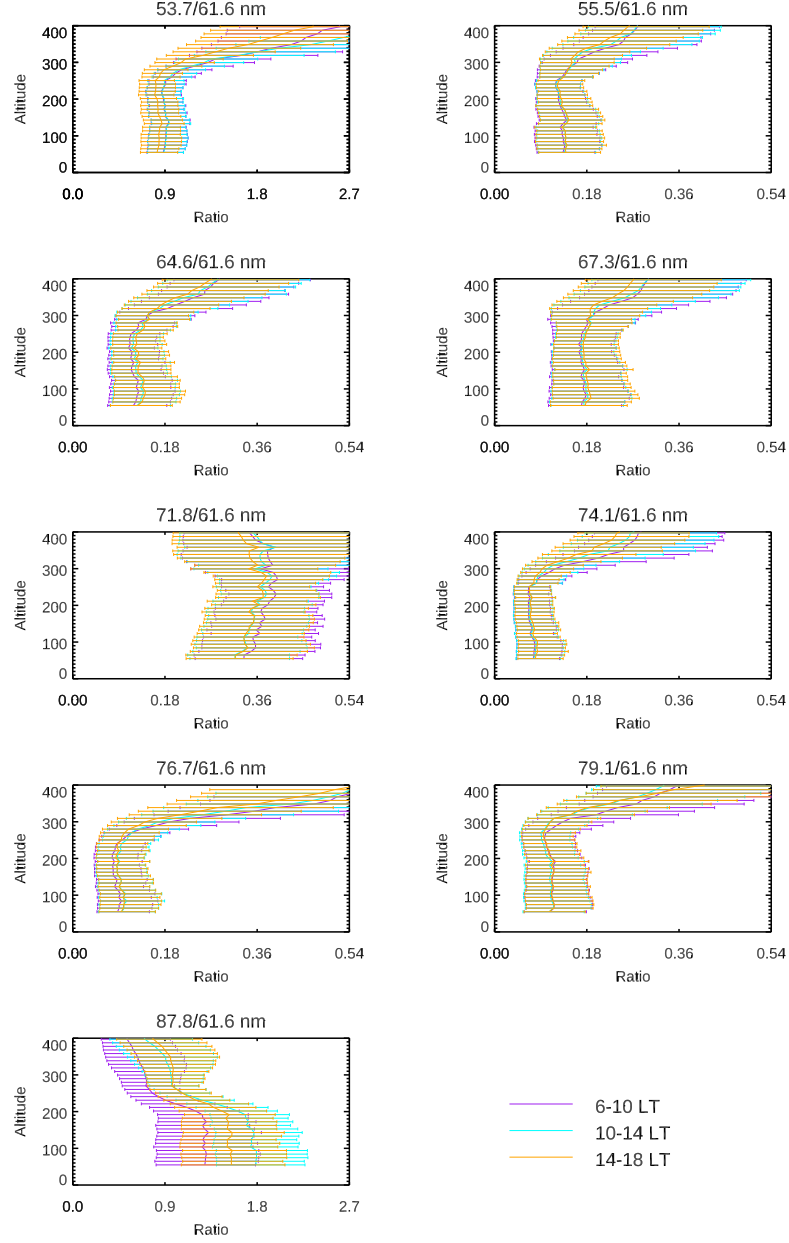
Emissions at 61.6 nm and 83.4 nm are each emitted by  $O^+$  after photoionization excitation of  $O$  (Meier, 1991)(Chakrabarti et al., 1983), but their profiles differ in shape and in magnitude (83.4 nm emissions are about 10 times brighter than 61.6 nm) due to their respective production and transport mechanisms. The most significant differences are that 83.4 nm is produced more efficiently and undergoes multiple scattering by ionospheric  $O^+$ , while 61.6 nm is optically thin to  $O^+$  in the ionosphere (Stephan et al., 2012).

The bright 58.4 nm line comes from resonant scattering of solar photons by neutral  $He$  (Meier, 1991).  $He$  has a very large scale height and is optically thick at 58.4 nm, so resonant scattering results in a far more uniform radiation field than if it were optically thin. The profile is essentially uniform across all ICON lines of sight and has a strong dependence on solar zenith angle, as one might expect given that solar photon scattering is a source.

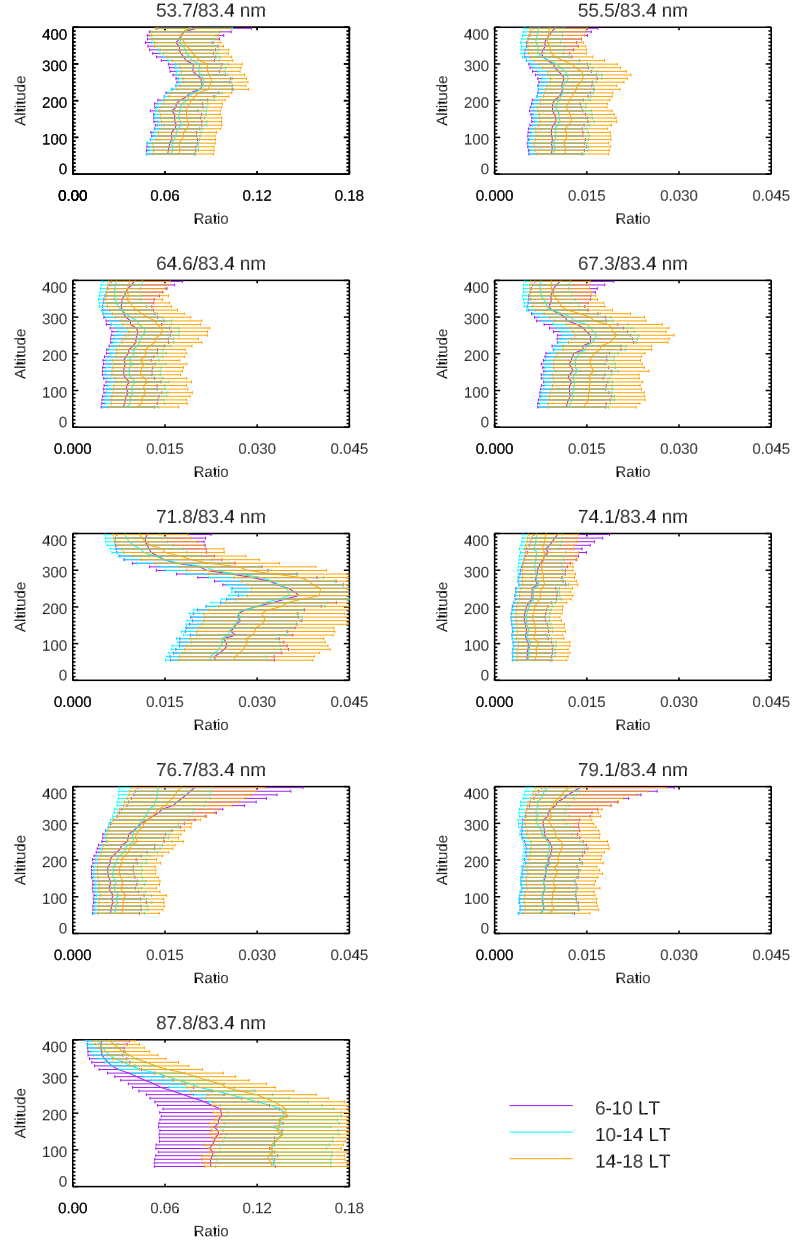
We now turn our attention to the secondary wavelengths. These emissions have been less studied, and some EUV features remain unidentified.

### 3.2 Secondary features: Brightness Ratios and Profile Shapes

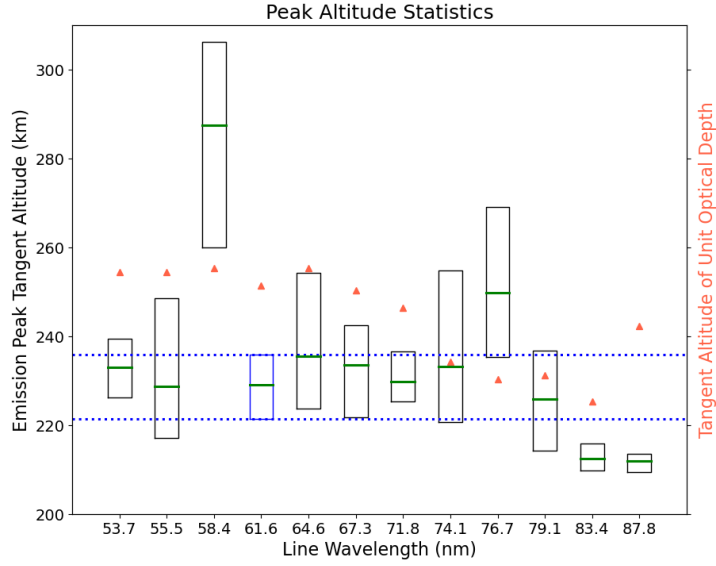
To compare one wavelength to another, it is useful to look at the brightness ratio at the time of measurement as a function of altitude. By using the ratio of individual measurements rather than the ratio of the average brightness, the variation due to differing production processes can be isolated from variation due to other factors, such as solar flux, local time, solar zenith angle, and physical location. In a further effort to isolate the sources of variation, we break down the ratios by local time and solar zenith angle. Measurements (on a single vertical pixel basis) are excluded if the brightness of the primary feature is below 5 Rayleighs or if the brightness of the secondary feature does not exceed 1 Rayleigh. We use the same scheme for altitude bins that we used for the altitude profiles. Since the brightness ratio measurements are not normally distributed and are heavily right skewed, we take the median value in each bin and represent the error as the 15.866 and 84.134 percentiles, which are equivalent to the  $\pm 1\sigma$  range for normally distributed data.



**Figure 4.** Brightness relative to 61.6 nm brightness for each secondary line, broken down by LT. Many of these features track well with 61.6 nm at low altitude but diverge at higher altitudes.

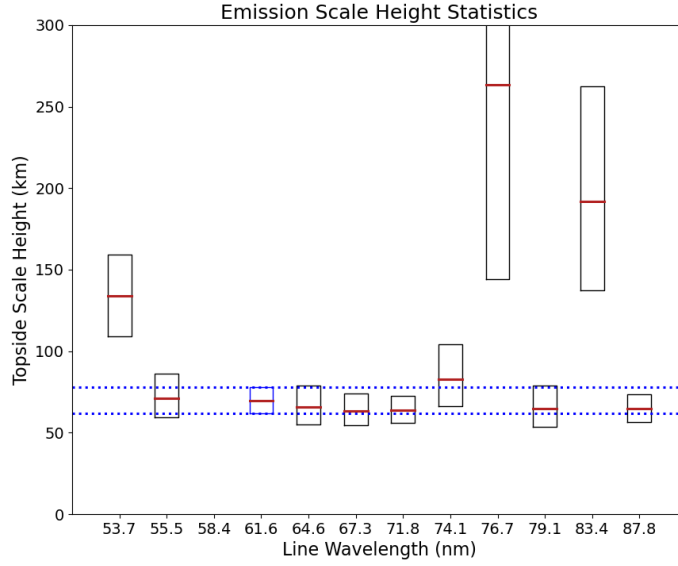


**Figure 5.** Brightness relative to 83.4 nm brightness for each secondary line, broken down by LT.



**Figure 6.** Green lines represent the median altitude of the emission peak for each L1 line. The upper and lower quartiles are indicated by boxes for all emissions and by the dotted blue line for the 61.6 nm line. Red triangles indicate the tangent altitude at which the tangent point reaches unit optical depth from neutral absorption along the ICON line of sight.

The brightness ratios relative to the 61.6 nm and 83.4 nm lines and broken down by local time are shown in Figures 4 and 5. The remaining ratio plots, broken down by solar zenith angle and month and including the ratio to 58.4 brightness are contained in Appendix A. The most immediately noticeable trend in Figure 4 is that these ratios increase significantly at high altitude for most wavelengths. This is surprising because we would expect the brightness of features other than 58.4 nm and 83.4 nm to be essentially zero at and above around 400 km. The increase of the ratio of all of these features to 61.6 nm is likely the result of a small error in the absolute calibration of the instrument or in the background subtraction, which could result in a bias that would disproportionately affect the dimmer features. Since we are dividing by the small 61.6 nm brightness at high altitudes, an unnoticeable error in the flux could create a large error in the ratio. Another explanation would be that these features all scatter more efficiently than 61.6 nm or have a *H* or *He* source. It is unlikely that this holds for all features, as we will see that most of these have an *O* parent.



**Figure 7.** Red lines represent the median topside emission scale height (see Section 3.2) for each L1 line. The upper and lower quartiles are indicated by boxes for all emissions and by the dotted blue line for the 61.6 nm line. Given the nature of the line and the large *He* scale height, the lower quartile, median, and upper quartile of the 58.4 nm topside emission scale height far exceed the plot range with values of 700, 1600, and 4000 km, respectively. The upper quartile of the 76.7 nm scale height has exceeded the plot range with a value of 480 km.

A vertical line in which the ratio is constant with little coherent variation between local times means that the brightness ratio of the primary and secondary feature is fixed and suggests that the emissions have the same or a similar source. Most of the secondary features resemble 61.6 nm in this way at low altitudes, suggesting that *O* could be their source. However, the altitude profile of the emission is also affected by atmospheric absorption and the scattering properties of each feature.

Approaching the peak from above, brightness increases as the line of sight begins to pass through regions with higher density of the emission source. As the line of sight pierces lower altitudes, three factors begin to determine the shape of the profile. First, the increase in overall density leads to increases in absorption/scattering between the source and the satellite. Second, production ceases to increase at the same rate as the density as the solar EUV flux is attenuated. Third, lines of sight with tangent altitudes below the production peak will see decreased brightness as the slant path through the



production layer shortens. If two features have the same source, the first factor will shape the profiles differently, as scattering and absorption cross sections are wavelength dependent; and the second can also have a differing effect, solar flux is attenuated unevenly across all wavelengths, and photoionization cross sections are likewise wavelength dependent.

Figure 6 shows the statistical location of the peak of each L1 feature, as determined by a three altitude pixel rolling average. The altitude of the peak indicates where the three factors discussed above begin to dominate over the increasing density along the line of sight. We focus on comparison to 61.6 nm since Figure 4 show that many features have similar altitude structure. We use an MSIS atmosphere and lab-measured neutral absorption cross sections (Conway, 1988) to determine the optical depth of the tangent point, and Figure 6 shows the tangent altitude at which the slant path between the tangent point and ICON has unit optical depth from absorption by  $O$ ,  $O_2$ , and  $N_2$ . This allows for a simple comparison of the expected neutral absorption affects between wavelength bins; although it should be noted that the cross sections used are of low resolution (0.1-0.3 nm), and small scale structure is known to exist in parts of the EUV instrument's spectral range (see (Fennelly & Torr, 1992), for example).

We are also interested in the shape of the emission profile above the peak, where we expect the brightness to correspond most strongly to production and, thus, number density of the parent species. To parameterize the shape, we treat the brightness similarly to a density, fitting to an exponential  $I = I_0 e^{\Delta z/H} + b$  to find the scale height  $H$ , allowing for some bias  $b$  to account for bias in the background subtraction. We fit only between the peak and 350 km, as the signal to noise ratio is very low above this altitude. Figure 7 contains the scale height statistics.

Looking at Figures 4, 5, 6, and 7, together, we can comprehensively relate the behavior of the secondary features to that of the primary features. In Figure 4, we see that the only feature which clearly behaves differently from 61.6 nm is the 87.8 nm feature which undergoes a large shift around the 61.6 nm peak. In Figure 6, we see that the 53.7, 55.5, 67.3, and 71.8 features have a peak in family with 61.6 nm while undergoing similar absorption. Figure 7 shows that the scale height of these emissions is also in family with 61.6 nm, except for 53.7 nm, which is significantly higher. This confirms that

the 55.5, 67.3, and 71.8 nm L1 bins are dominated by the known  $O^+$  emissions (see Table 1) and that 53.7 nm is mostly  $O^+$  with  $He$  contamination at high altitudes.

The 64.6 nm bin peaks noticeably higher than 61.6 nm while undergoing similar absorption. This feature is visible in the rocket limb spectrum of Gentieu et al. (1984), although the authors do not identify the source. Rocket measurements of the disk identify this feature as a blend of  $O^+$  and  $N^+$ . (Chakrabarti et al., 1983) The scale height from ICON is similar to that of 61.6 nm, so it is likely that the instrument is mostly seeing the  $O^+$  line at 64.4 nm. However, contamination by the nearby known  $N^+$  emission would not explain the raised peak, as we would expect  $N_2$  rather than atomic nitrogen to be the effective source of such an emission. The very similar 74.1 nm bin contains a previously identified blend of  $O$  and  $N^+$  lines, peaks slightly above 61.6 nm while being optically thin to a lower altitude, and has a larger scale height than 61.6 nm. The 64.6 nm and 74.1 nm bins are clearly measuring the respective  $O^+$  and  $O$  features, but there is a significant additional contribution in each case.

The 79.1 nm bin peaks slightly lower than 61.6 nm, but is also optically thin to a lower altitude (see Figure 6), which would have the effect of lowering the peak. Given the constant ratio to 61.6 nm at low altitudes and that the scale height is similar to that of the  $O^+$  features, this seems to be the previously identified  $O^+$  feature at 79.7 nm.

In Figure 5, the suspected  $O/O^+$  features (53.7, 55.5, 67.3, 71.8, and 79.1 nm) all have a bump around 250 km tangent altitude that is most pronounced in the afternoon. The multiple scattering of 83.4 nm photons by  $O^+$  has the effect of broadening the altitude profile. For this reason, the 83.4 nm profile is smoother in the afternoon than in the morning (see Figure A1) when the  $O^+$  concentration is higher. The aforementioned wavelengths are optically thin to  $O^+$ , resulting in a sharper peak and an enhancement relative to 83.4 nm near the peak.

The 76.7 nm feature is a clear outlier. It peaks well above 61.6 nm, despite having a lower tangent altitude of unit optical depth. As seen in Figure 7, the scale height of this emission is very large. Figure 3 indicates that this feature has a strong SZA dependence. It is possible that this feature is instrumental, a re-imaging of light reflected from the MCP to the diffraction grating and back. This effect was observed on the Emirates Mars Ultraviolet Spectrometer which employs a similar toroidal diffraction grating for UV-spectroscopy. (Holsclaw et al., 2021) The re-imaging of reflected light would

result in the loss of spatial information, which could explain the relatively flat profile seen in Figure 3.

Another outlier is the 87.8 nm feature. This (relatively) bright feature (see Figure 3) peaks lower than any other at around 210 km, while undergoing similar absorption to 61.6 nm (see Figure 6). This is most similar to the 83.4 nm line, but Figure 7 indicates that 87.8 nm has a much smaller scale height than 83.4 nm and a slightly smaller one than 61.6 nm. These indicators are consistent with an emission coming from a heavier species, like  $N_2$ . In fact,  $N$  features have previously been observed at 87.5 nm, 87.8 nm, and 88.7 nm (Gentieu et al., 1984); it is likely that these are produced by excited  $N^*$  or  $N^{+*}$  from photodissociation  $N_2$ , as has been observed in nearby emission features (Meier et al., 1991). In Section 4, we will show that this bin is picking up nitrogen emissions by establishing that the brightness ratio of 61.6 nm and 87.8 nm emissions acts as a proxy for  $\Sigma O/N_2$ .

#### 4 Relationship of 61.6/87.8 to $\Sigma O/N_2$

A common metric for tracking upper-atmospheric composition is  $\Sigma O/N_2$  ratio, obtained by a column integral of number density ( $n$ ) down until a threshold  $N_2$  column density ( $\mathbf{N}$ ) of  $10^{17} \text{ cm}^{-2}$  is reached, at altitude  $z_{17}$ .

$$\Sigma O/N_2 = \frac{\int_{z_{17}}^{\infty} n_O dz}{\int_{z_{17}}^{\infty} n_{N_2} dz} = \frac{\int_0^{\mathbf{N}_O} d\mathbf{N}'_O}{\int_0^{\mathbf{N}_{N_2}} d\mathbf{N}'_{N_2}} = \frac{\mathbf{N}_O}{10^{17} \text{ cm}^{-2}} \quad (1)$$

The altitude  $z_{17}$  is typically the 130-140 km range but is not considered to be of inherent physical significance (Meier, 2021).

As stated in Section 2.2, one of the motivations for the inclusion of the FUV instrument on the ICON mission is to measure  $\Sigma O/N_2$  using airglow emissions of  $O$  at 135.6 nm (short wave, SW) and the  $N_2$  LBH band near 157 nm (long wave, LW). The inversion process is complex, depending on external factors such as solar zenith angle, solar flux, and viewing angle. However, for a given set of observation conditions,  $\Sigma O/N_2$  on the disk is established to be determinable by the brightness ratio  $I_{SW}/I_{LW}$  (Meier, 2021) (Strickland et al., 1995).

Because  $O$  is the parent source of the 61.6 nm emission and the 87.8 nm feature is likely from one or both known  $N$  lines with an  $N_2$  parent, we expect the brightness

ratio of the 61.6 nm and 87.8 nm bins to likewise provide a measurement of  $\Sigma O/N_2$ . If the instrument observes low enough tangent altitudes, below the point of extinction, our observation will share a key property with disk observations: the line of sight will pass through all relevant altitudes once without contributions from beyond the foreground (that is, beyond the tangent point). This will also mitigate any concerns of out-of-band Lyman- $\alpha$  contamination, which we consider to be a possibility at high tangent altitudes near the edge of the detector.

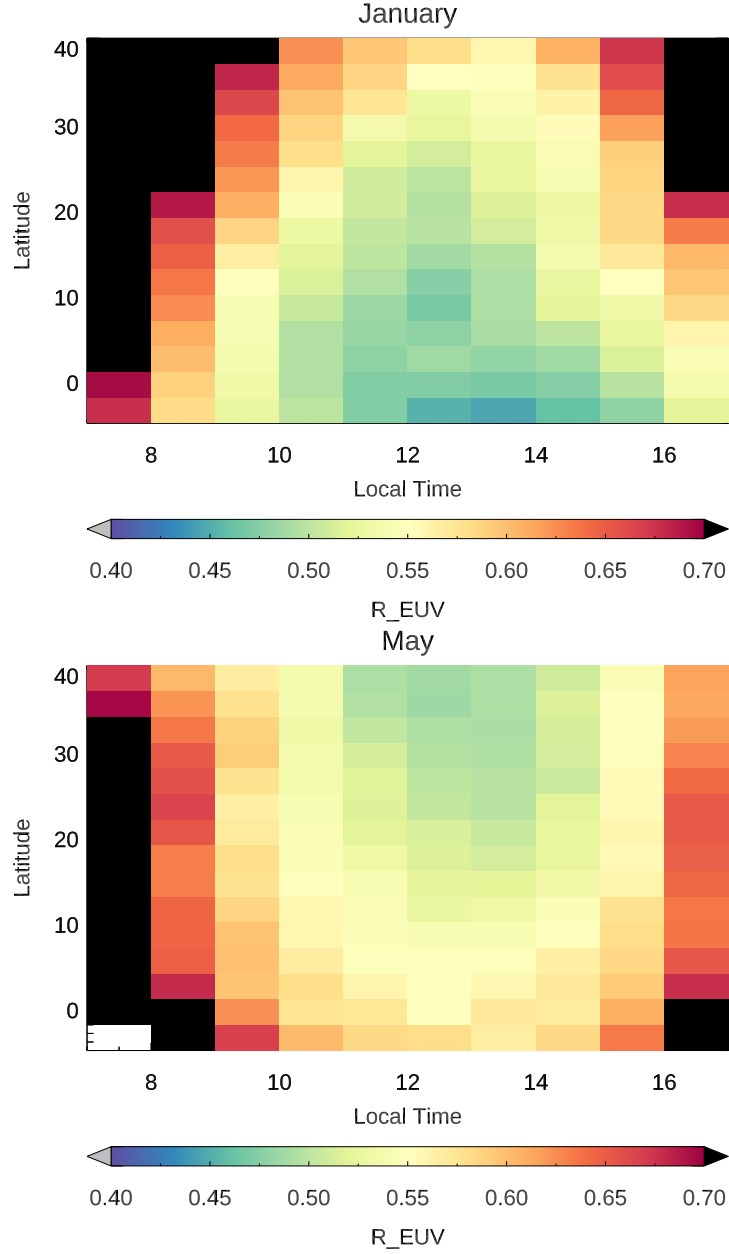
The FUV features are far brighter than either of these EUV emissions. As seen in Table 1, midday 61.6 nm emissions typically peak around 60 Rayleighs and 87.8 nm emissions around 90 Rayleighs. In contrast, the FUV SW and LW measurements typically peak around 8000 and 2000 Rayleighs, respectively. Any attempt to retrieve  $\Sigma O/N_2$  using a single vertical pixel on the EUV detector will be dominated by shot noise. Therefore, we must not only look low enough to have a disk-like measurement but also sum over a range of vertical pixels so that we have sufficient signal. To do this, we define a sub-limb region, which contains lines of sight with tangent point altitudes between 50 and 140 km. The lower boundary is approximately the lower limit of the EUV field of view. With ICON’s viewing geometry, we expect the tangent points to have unit optical depth from neutral absorption at a tangent altitude of around 240 km for both of these wavelengths, so the upper boundary of 140 km tangent altitude is well within the region where emissions beyond the tangent point do not significantly contribute to the measurement (as in a disk measurement). In analogy with SW/LW, we define a simple metric to compare to  $\Sigma O/N_2$ :

$$R_{EUV} = \frac{\sum_i I_{616,i}}{\sum_i I_{878,i}} \text{ for } i \text{ such that } z_i \in [50, 140] \text{ km} \quad (2)$$

Here  $I_{616,i}$  and  $I_{878,i}$  are the 61.6 nm and 87.8 nm brightnesses and  $z_i$  the tangent altitude of the  $i^{\text{th}}$  pixel of the EUV detector. We will show that  $R_{EUV}$  is correlated to  $\Sigma O/N_2$ , supporting our claim that the 87.8 nm feature comes from one or more  $N$  emissions and suggesting the possibility of observing  $\Sigma O/N_2$  using EUV observations.

#### 4.1 Latitude Structure of $R_{EUV}$ and Comparison to MSIS $\Sigma O/N_2$

One of the key characteristics of  $N_2$  in the upper atmosphere is the seasonal variation of its spatial structure. In the summer hemisphere, increased temperature leads



**Figure 8.** A comparison of the latitude-LT structure of  $R_{EUV}$  for January and May 2020. The color of each pixel represents the median  $R_{EUV}$  value in the month, LT, and latitude indicated, and white pixels indicate missing data. The figures show a reversal of the latitude structure around local noon, qualitatively similar to what we would expect from  $\Sigma O/N_2$ .

to a larger scale height, resulting in enhanced density of heavier species at high altitude. Figure 8 shows the latitude-LT structure of  $R_{EUV}$  in January and May 2020. Both plots show an increased ratio near sunrise and sunset. However, the latitudinal pattern dif-

fers between the two. In January,  $R_{EUV}$  is lowest around the equator and increases with latitude. In May, the trend is reversed, with low  $R_{EUV}$  at mid-latitudes and higher ones near the equator. We would expect to see the same qualitative trends from  $\Sigma O/N_2$ , and such behavior has indeed been observed in  $\Sigma O/N_2$  retrieved from FUV observations (Strickland et al., 2004).

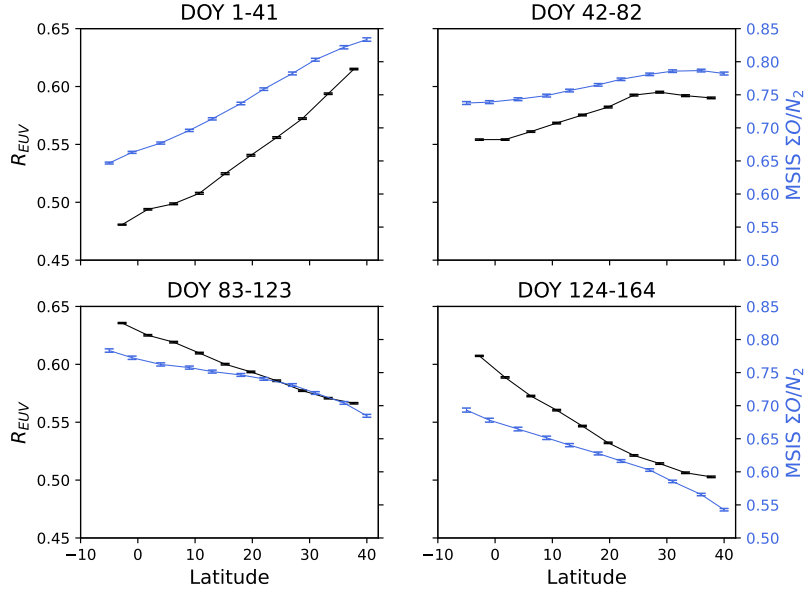
To provide a more convincing resemblance to  $\Sigma O/N_2$ , we now compare  $R_{EUV}$  to NRLMSISE-00 (MSIS) (Picone et al., 2002)  $\Sigma O/N_2$  for some consecutive 41 day coverage cycles of ICON. Since we are mostly interested in the evolution of the latitudinal structure, we reduce the EUV data to one dimension by averaging into latitude bins over 10-14 LT. We run MSIS at five UTCs each day at 11 equally spaced latitudes from -5 to 40 degrees, to match the latitudes covered by ICON. At each UTC and latitude, we run MSIS for longitudes corresponding to 10, 12, and 14 LT. We use a constant F10.7 of 70 sfu, since the variation in F10.7 during the first half of 2020 was very small (ranging from 68-75 sfu over the dates considered). Finally, we use a constant magnetic index ( $A_p$ ) of 4 as this adequately characterizes the solar-quiet conditions from the first half of 2020.  $\Sigma O/N_2$  is averaged at each latitude, resulting, in combination with the ICON data, in the plots in Figure 9.

There is a clear resemblance between the EUV and MSIS data in the evolution of structure from winter to summer. We conclude that  $R_{EUV}$  does vary with  $\Sigma O/N_2$  on seasonal timescales.

## 4.2 Comparison of $R_{EUV}$ to ICON FUV Data

The ICON FUV data provides simultaneous measurements to compare to  $R_{EUV}$ . However, due to differences in viewing geometry, the location of the EUV sub-limb and the disk  $\Sigma O/N_2$  observations are too far apart, separated by about 14 degrees of latitude. Fortunately, the EUV and FUV detectors overlap in the 50-140 km tangent altitude range so we can investigate the correlation between the two. We define  $R_{FUV}$  as the ratio of mean SW and LW brightness:

$$R_{FUV} = \frac{\sum_i I_{SW,i}/N_{SW}}{\sum_j I_{LW,j}/N_{LW}} \begin{cases} \text{for } i \text{ such that } H_{SW,i} \in [50, 140] \text{ km} \\ \text{for } j \text{ such that } H_{LW,j} \in [50, 140] \text{ km} \end{cases} \quad (3)$$

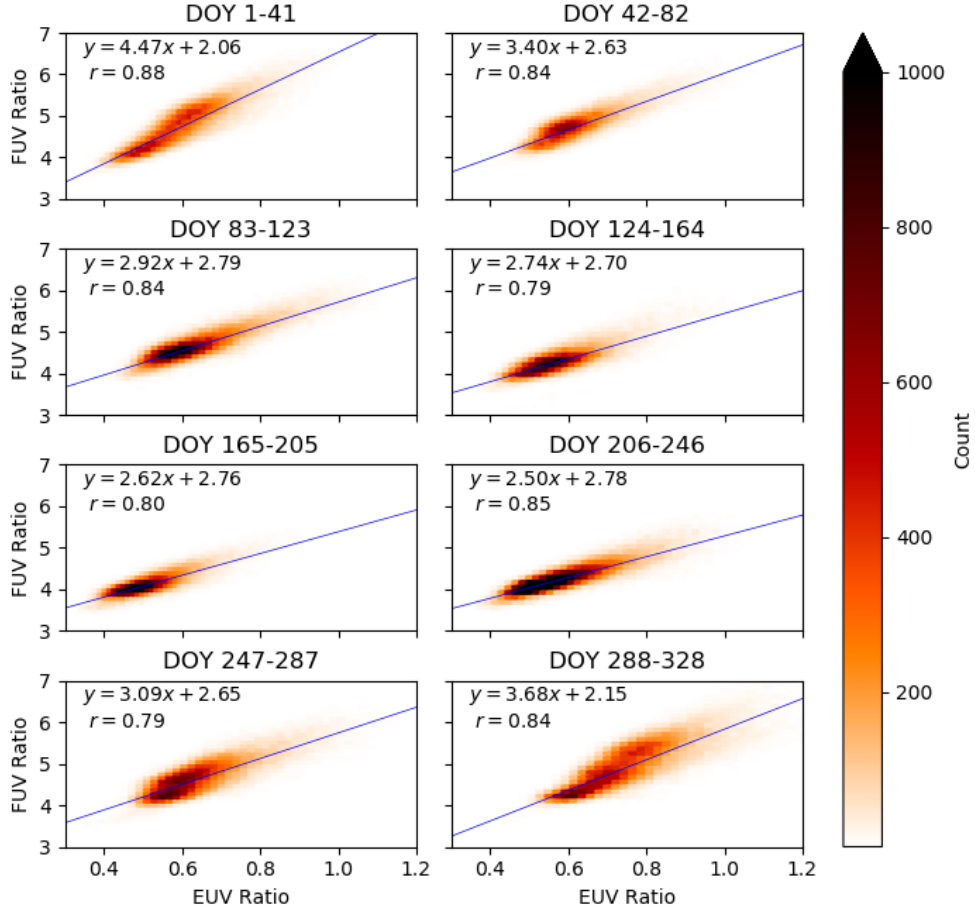


**Figure 9.**  $R_{EUV}$  is plotted along with MSIS column  $O/N_2$  versus latitude for 10-14 LT and date ranges from January to June. Error bars represent uncertainty in the mean.  $R_{EUV}$  is shown to undergo the same seasonal reversal seen in  $\Sigma O/N_2$ .

Here,  $H_{SW,i}$  and  $H_{LW,i}$  are the tangent altitudes of the  $i^{\text{th}}$  pixel in the SW and LW channels, respectively, which differ slightly due to the FUV instrument optics.  $N_{SW}$  and  $N_{LW}$  are the number of pixels that fall within the desired altitude range in the SW and LW channels. We are taking the ratio of the mean brightness here, since  $N_{SW}$  and  $N_{LW}$  can differ.

Since the fields of view are aligned between the two instruments (nominally and upon inspection), we need only look for coincidences in time. For each EUV exposure, we find the time of the nearest FUV exposure and consider the measurements coincident if the time difference is less than 12 seconds, which is the cadence of both detectors. The coincident measurements for ICON coverage cycles from 2020 DOY 1-328 are plotted in Figure 10.

Each EUV and FUV measurement has a reported uncertainty, which propagates to  $R_{EUV}$  and  $R_{FUV}$ . When measuring the correlation between the datasets, we would like to give greater consideration to measurements with lesser uncertainty. Since there is uncertainty in the x and y variables, a typical chi-squared minimization will not do.



**Figure 10.** Density plots of correlation between  $R_{EUV}$  and  $R_{FUV}$  during the eight full coverage cycles of 2020. The range in  $R_{EUV}$  varies between cycles: for example DOY 124-164 in the spring covers a small range, while DOY 1-41 in the winter covers a larger range.

Instead, we compute a fit to each curve using orthogonal distance regression (Boggs & Donaldson, 1989), which minimizes the orthogonal distance to the fit line and weights the points as the inverse square of the uncertainty. We then use these same weights to calculate the weighted Pearson correlation coefficient. The resulting fit is shown for the data in Figure 10, and the correlation coefficients are tabulated in Table 2, broken down by ICON coverage cycle and local time.

The correlation between  $R_{EUV}$  and  $R_{FUV}$  is generally quite strong. The full dataset shows a correlation coefficient of 0.83. Smaller subsets restricted to a two hour local time



**Table 2.** Weighted Pearson correlation coefficients for  $R_{EUV}$  and  $R_{FUV}$ 

	6-8 LT	8-10 LT	10-12 LT	12-14 LT	14-16 LT	16-18 LT	6-18 LT
DOY 1-41	0.75	0.92	0.92	0.90	0.87	0.73	0.88
DOY 42-82	0.73	0.78	0.64	0.66	0.74	0.74	0.84
DOY 83-123	0.71	0.75	0.62	0.60	0.66	0.70	0.84
DOY 124-164	0.81	0.78	0.67	0.62	0.72	0.76	0.79
DOY 165-205	0.88	0.85	0.76	0.72	0.74	0.76	0.80
DOY 206-246	0.80	0.78	0.66	0.70	0.80	0.83	0.85
DOY 247-287	0.71	0.80	0.72	0.63	0.71	0.61	0.79
DOY 288-328	0.78	0.89	0.88	0.82	0.84	0.77	0.84
DOY 1-328	0.87	0.86	0.81	0.79	0.81	0.80	0.83

The color of each cell scales with the value, included for easy comparison with Table 3.

**Table 3.** 5<sup>th</sup>-95<sup>th</sup> percentile range of  $R_{EUV}$ 

	6-8 LT	8-10 LT	10-12 LT	12-14 LT	14-16 LT	16-18 LT	6-18 LT
DOY 1-41	0.27	0.41	0.30	0.25	0.25	0.23	0.35
DOY 42-82	0.30	0.25	0.16	0.14	0.19	0.22	0.34
DOY 83-123	0.35	0.26	0.19	0.17	0.18	0.25	0.37
DOY 124-164	0.40	0.32	0.21	0.18	0.20	0.24	0.31
DOY 165-205	0.36	0.26	0.17	0.16	0.21	0.25	0.27
DOY 206-246	0.36	0.25	0.18	0.17	0.22	0.30	0.35
DOY 247-287	0.37	0.32	0.21	0.17	0.20	0.23	0.36
DOY 288-328	0.38	0.47	0.35	0.27	0.30	0.30	0.41
DOY 1-328	0.49	0.41	0.32	0.29	0.31	0.32	0.41

Here, the color of each cell scales with the value but clips at 0.4, so that lower values can more easily be distinguished. We can see that regions of Table 2 with low correlation coefficients generally correspond to regions of low  $R_{EUV}$  spread in Table 3.

range for a given coverage cycle show correlation coefficients as high as 0.92 and no lower than 0.60. Possible origins of variance between  $R_{EUV}$  and  $R_{FUV}$  include ionospheric contamination of the SW (135.6 nm photons produced by radiative recombination), differing atmospheric extinction between the correspondent EUV and FUV features, and noise in the EUV measurements.

There is a pattern in the correlation strength where the correlation coefficients are lowest around local noon and near equinox. The second and third panels of Figure 9, indicate that there is little variation in  $R_{EUV}$  with latitude during the spring, one of the periods with low correlation coefficients. Table 3 captures the spread of  $R_{EUV}$  within each subset of the data, and the color of each cell scales with the value. Tables 2 and 3 match up quite well, indicating that the correlation between  $R_{EUV}$  and  $R_{FUV}$  is weaker when only a small range of  $R_{EUV}$  values are considered. The exception is a dawn, when we would expect the EUV measurements very dim and, therefore, noisy.

## 5 Summary and Discussion

### 5.1 Identity of Emission Features

In Section 3.2 we showed that several emission features in the ICON EUV data behave similarly to 61.6 nm  $O^+$ . Specifically, the 53.7, 55.5, 67.3, 71.8, and 79.1 nm bins are dominated by optically thin  $O^+$  airglow.

The 53.7 nm bin is likely picking up one of the previously identified lines at 53.8 nm and/or 53.9 nm. The dimness of the combined features relative to the 61.6 nm brightness compared to what was seen in the rocket measurements (Gentieu et al. (1984), see Table 1) is evidence of vignetting on this edge of the detector. The 79.1 nm bin seems to be dominated by the  $O^+$  doublet at 79.7 nm. The 55.5 nm, 67.3 nm, and 71.8 nm features observed by ICON are likely the known  $O^+$  lines at the same wavelengths with the 67.3 nm bin likely containing a negligible amount of the  $N^+$  line at 67.1 nm and the 79.1 nm bin containing a negligible amount of the  $N_2$  79.7 nm emission. In these four cases, the rocket measurements of Gentieu et al. (1984) measure brighter dayglow on the limb, which is to be expected from the difference in solar irradiance. However, the magnitude of this difference varies greatly between wavelengths, with low solar 2020 55.5 nm, 67.3 nm, 71.8, and 79.1 nm brightness being 12%, 67%, 55%, and 30% of their high solar 1980 values, respectively. This discrepancy likely comes from a combination of cross-instrument comparison and a physical difference in the response of each emission feature to changes in solar irradiance; if the ICON observations continue into solar maximum, a future study could identify the degree to which solar factors are responsible for this difference.

The 64.6 nm and 74.1 nm bins do not behave as pure  $O/O^+$  features, so we conclude that these bins have significant blending from the respectively contained  $N^+$  lines. The source of the 76.7 nm feature remains unknown, and it is possibly an instrumental artifact.

Our analysis in Section 4 has linked the 87.8 nm bin to  $N_2$ . It is likely that the bin is dominated by some combination of the known  $N$  multiplets between 86.5 nm and 88.9 nm or even the  $N^+$  doublet near 86.0 nm (Kramida et al., 2021). These emissions are not well modeled, but the connection to  $N_2$  suggests that one or both of these features are produced in the thermosphere by photodissociation of  $N_2$  resulting in excited  $N^*$  (or

$N^{+*}$ ), as has been observed in the lab and in the dayglow (Meier et al., 1991). As with other  $N$  and  $N^+$  EUV emissions, photoelectron impact excitation may also contribute to the excitation rate.

It must be noted that this region on the edge of the detector is subject to an unknown amount of vignetting, which makes it difficult to interpret the absolute magnitude of the ICON measurements. Additionally, there is concern that this region of the detector may suffer from Ly- $\alpha$  scattering within the instrument, which would affect the shape of the profile since it would disproportionately brighten pixels at higher tangent altitudes.

Our identification of the 55.5, 67.3, 71.8, and 79.1 nm bins as  $O^+$  dominated is supported by presence of the atomic features in the oxygen spectra from the ground calibration of the EUV instrument. (Figure 9 of Sirk et al. (2017)) The only feature seen in the oxygen calibration spectra that we do not identify in the in-flight data is found in the 87.8 nm band, which contains a previously observed neutral O feature. It is possible that this feature is present in the 87.8 nm data and that the profile is really a blend of the  $N_2$  and  $O$  parent features. This would still be consistent with our findings in Section 4 if the  $O$  contribution is only relevant at tangent altitudes higher than 140 km, which is plausible given the shape of the  $O^+$  profiles.

## 5.2 Utility of $R_{EUV}$ for Retrieving $\Sigma O/N_2$

The correlation of  $R_{EUV}$  with  $R_{FUV}$  indicates that the EUV and FUV sub-limb observations contain much of the same information. In principle, this is sufficient to suggest that the EUV instrument is capable of  $\Sigma O/N_2$  retrieval, since FUV sub-limb observations are known to be sufficient for inversion to  $\Sigma O/N_2$  (Meier, 2021). However, the lack of an EUV dayglow model for the observed 87.8 nm feature presents an obstacle to the development of an inversion algorithm. It is likely that the ICON EUV limb observations could also provide measurements of neutral density altitude profiles, as has been done using FUV observations (Meier et al., 2015).

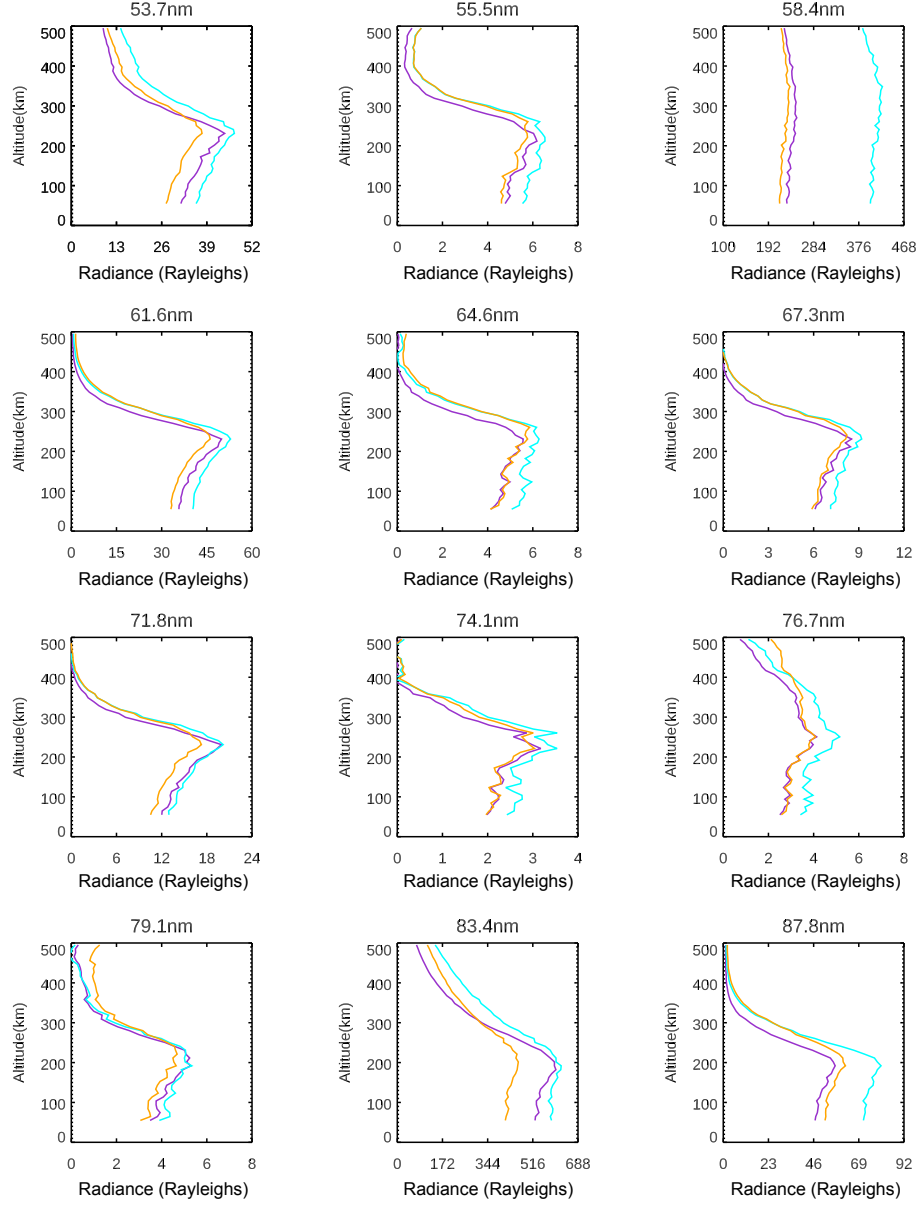
An EUV  $\Sigma O/N_2$  retrieval algorithm offers a potential advantage over current FUV methods for retrieving  $\Sigma O/N_2$ . The 135.6 nm oxygen doublet used for FUV retrieval is dominated by photoionization and electron impact ionization excitation of  $O$  with minor, but often non-negligible, contributions from  $O^+$  radiative recombination, especially

at high altitudes. In certain regions and atmospheric conditions, this can lead to ionospheric contamination of the retrieved  $\Sigma O/N_2$  (Kil et al., 2013). In contrast, the 61.6 nm line is only produced in significant amounts by photoionization excitation of  $O$ . This advantage could potentially offset the difficulty and uncertainties that arise from the dimness of the EUV features and from the development of an EUV retrieval, which would include a more detailed characterization of the ICON 87.8 nm band. An effective EUV retrieval of neutral composition would be practically advantageous since it would allow for the measurement of the neutral thermosphere and the ionosphere with a single instrument.

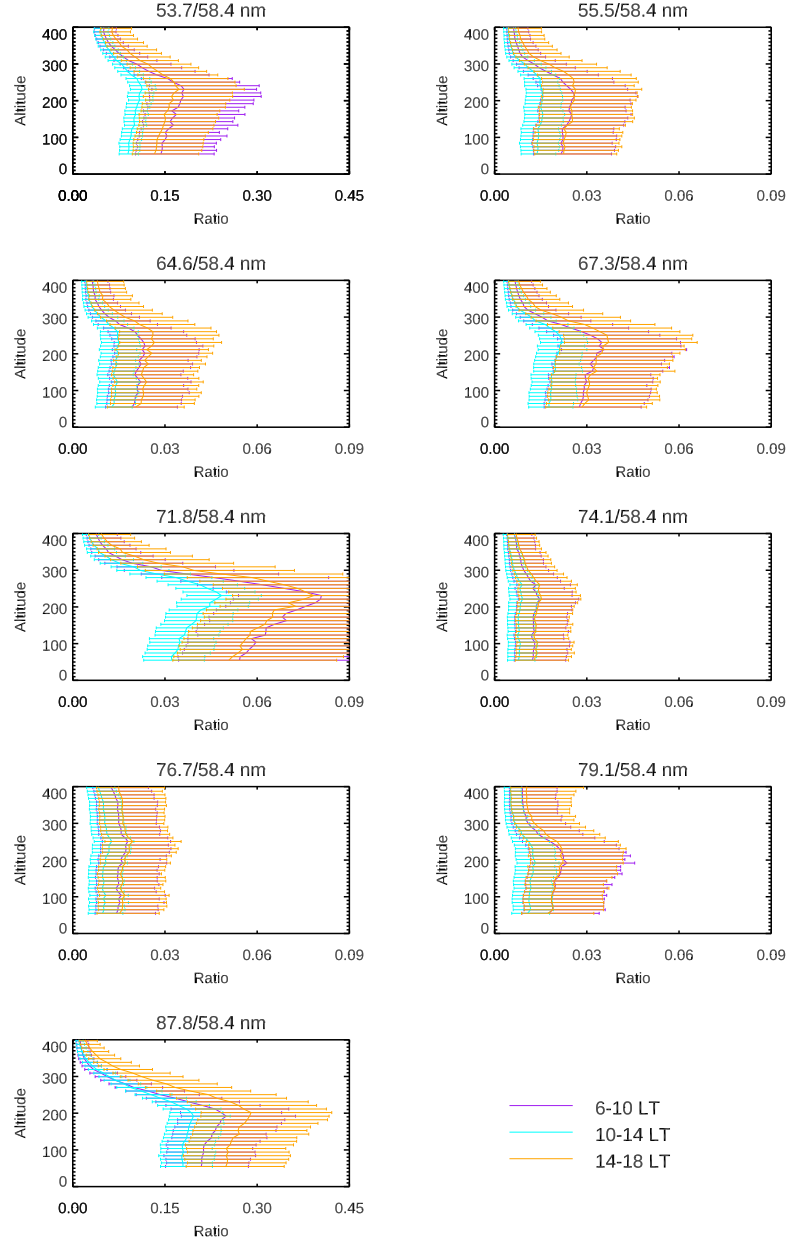
## 6 Data Availability Statement

This analysis used version 03 of the Level 1.5 ICON-EUV data, version 03 of the Level 1.3 ICON-FUV data, and version 03 of the ICON-Ancillary data which are available from the ICON website ( <https://icon.ssl.berkeley.edu/Data> ) and NASA’s Space Physics Data Facility ( <https://cdaweb.gsfc.nasa.gov/pub/data/icon/> ). The NRLMSISE-00 model is available from NASA’s Community Coordinated Modeling Center ( <https://kauai.ccmc.gsfc.nasa.gov/instantrun/msis> ) and is available for download from the US Naval Research Laboratory ( <https://map.nrl.navy.mil/map/pub/nrl/NRLMSIS/NRLMSISE-00/> ).

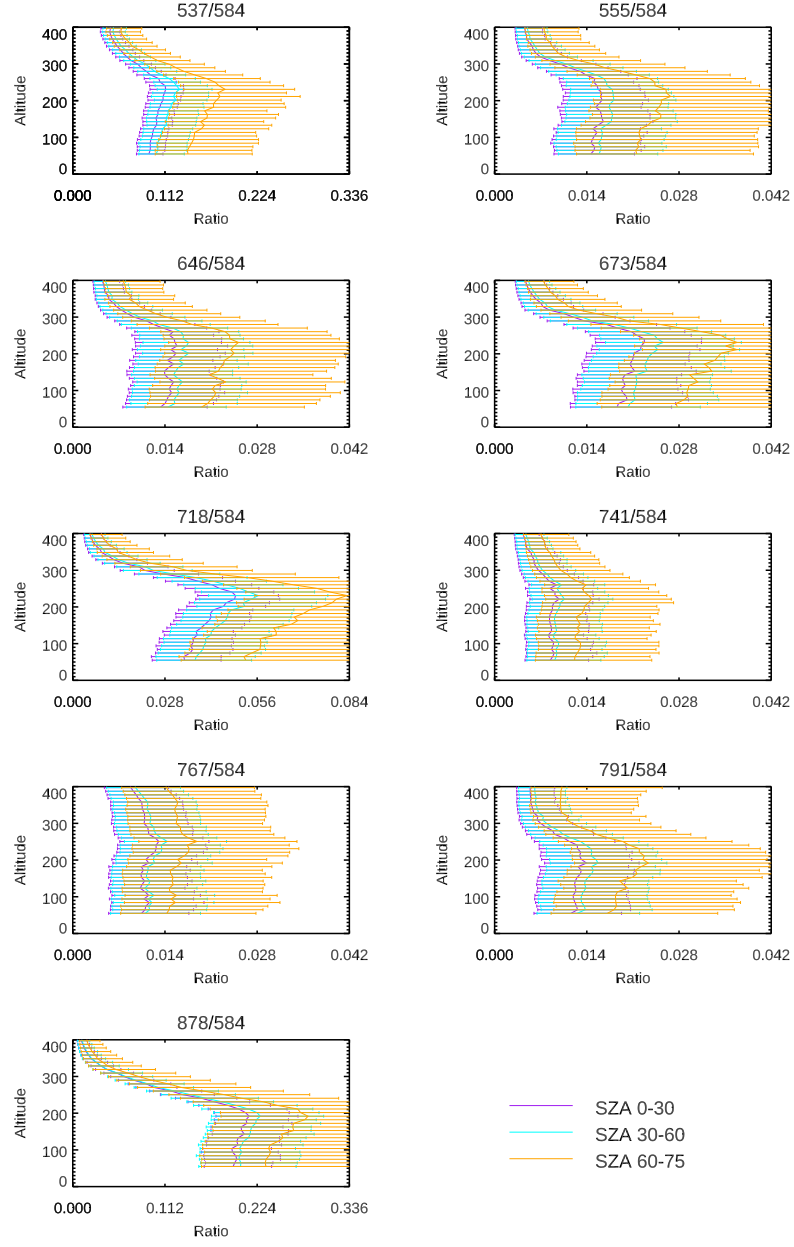
## **Appendix A Additional Figures**



**Figure A1.** Averaged altitude profiles of ICON L1 EUV wavelength bin for 6-10 LT (purple), 10-14 LT (cyan), and 14-18 LT (orange) from 2020 DOY 1-328.

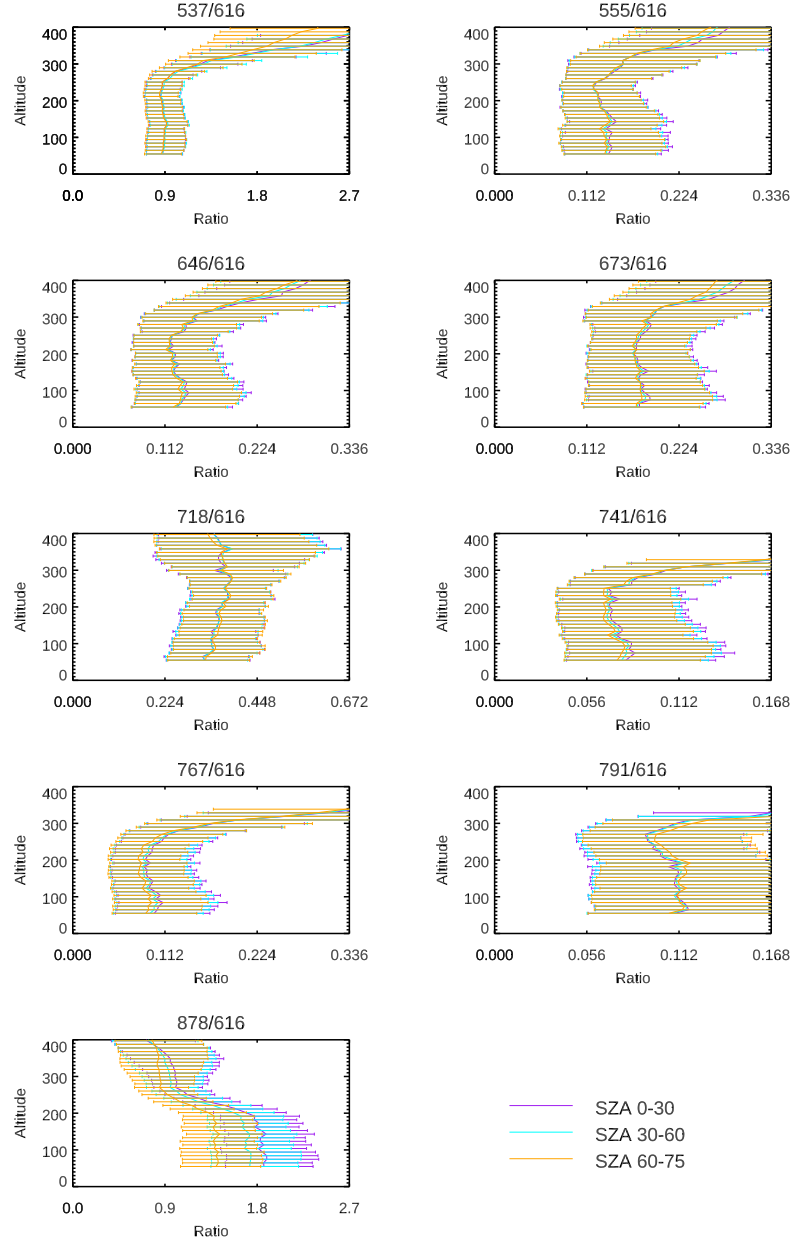


**Figure A2.** Brightness relative to 58.4 nm brightness for each secondary line, broken down by LT.

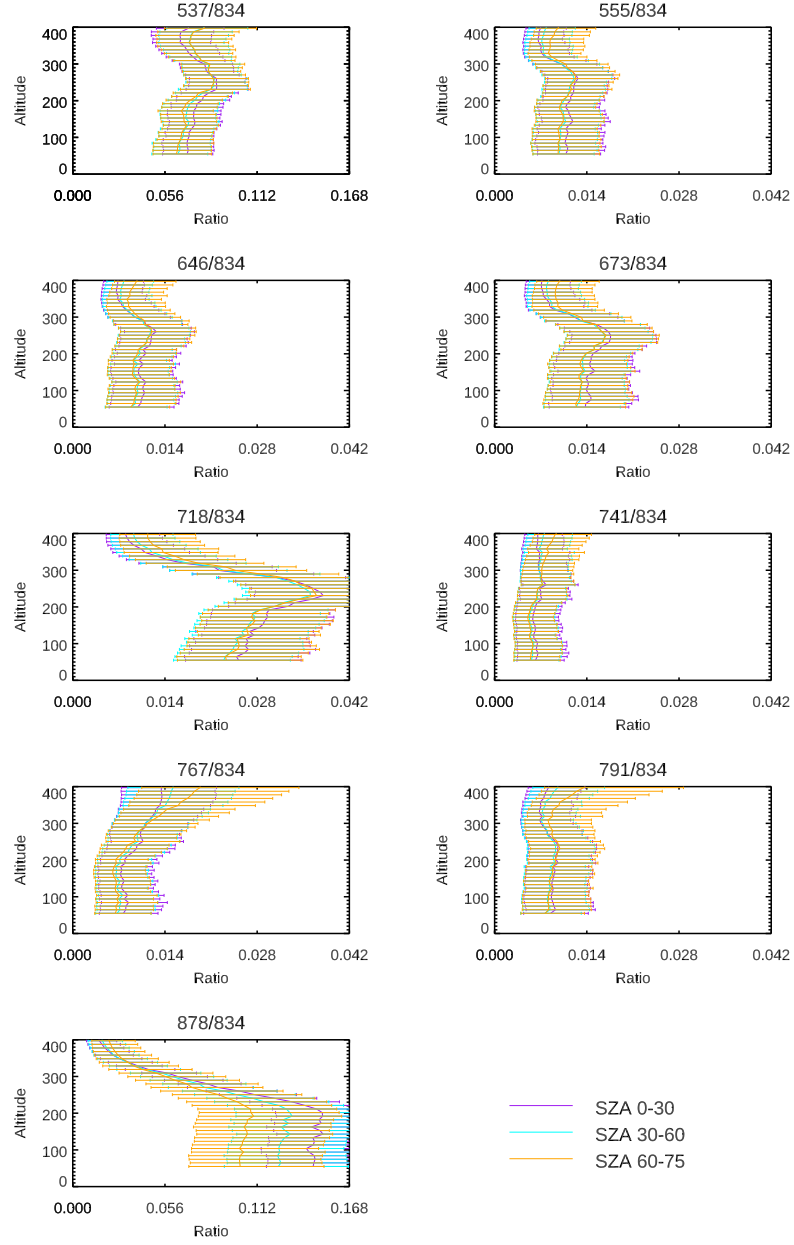


**Figure A3.** Brightness relative to 58.4 nm brightness for each secondary line, broken down by SZA.





**Figure A4.** Brightness relative to 61.6 nm brightness for each secondary line, broken down by SZA.



**Figure A5.** Brightness relative to 83.4 nm brightness for each secondary line, broken down by SZA.

## Acknowledgments

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## References

- Anderson, J., D. E., Meier, R. R., & Weller, C. S. (1979, May). The seasonal-latitudinal variation of exospheric helium from He 584-Å dayglow emissions. *Journal of Geophysical Research*, *84*(A5), 1914-1920. doi: 10.1029/JA084iA05p01914
- Boggs, P., & Donaldson, J. R. (1989). Orthogonal distance regression..
- Bush, B. C., & Chakrabarti, S. (1995, October). Analysis of Lyman  $\alpha$  and He I 584-Å airglow measurements using a spherical radiative transfer model. *Journal of Geophysical Research*, *100*(A10), 19609-19626. doi: 10.1029/95JA01210
- Carlson, R. W., & Judge, D. L. (1973, May). Rocket observations of the extreme ultraviolet dayglow. *Planetary and Space Science*, *21*(5), 879-880. doi: 10.1016/0032-0633(73)90105-0
- Carruthers, G. R., & Page, T. (1976, April). Apollo 16 Far ultraviolet spectra of the terrestrial airglow. *Journal of Geophysical Research*, *81*(10), 1683. doi: 10.1029/JA081i010p01683
- Chakrabarti, S., Paresce, F., Bowyer, S., Kimble, R., & Kumar, S. (1983, June). The extreme ultraviolet day airglow. *Journal of Geophysical Research*, *88*(A6), 4898-4904. doi: 10.1029/JA088iA06p04898
- Christensen, A. B. (1976, April). A rocket measurement of the extreme ultraviolet dayglow. *Geophysical Research Letters*, *3*(4), 221-224. doi: 10.1029/GL003i004p00221
- Conway, R. (1988, March). Photoabsorption and Photoionization Cross Sections of O, O<sub>2</sub>, and N<sub>2</sub> for Photoelectron Production Calculations: A Compilation of Recent Laboratory Measurements, NRL Memorandum Report 6155 [Computer software manual].
- Cotton, D. M., Chakrabarti, S., & Gladstone, G. R. (1993, December). Preliminary results from the Berkeley EUV airglow rocket spectrometer: O I and N<sub>2</sub> FUV/EUV dayglow in the thermosphere and lower exosphere. *Journal of Geophysical Research*, *98*(A12), 21627-21642. doi: 10.1029/93JA02267

- Donahue, T. M., & Kumer, J. B. (1971, January). An observation of the helium I 584-A dayglow radiation between 400 and 1000 km. *Journal of Geophysical Research*, 76(1), 145. doi: 10.1029/JA076i001p00145
- Evans, J. S., Strickland, D. J., & Huffman, R. E. (1995, July). Satellite remote sensing of thermospheric O/N<sub>2</sub> and solar EUV. 2. Data analysis. *Journal of Geophysical Research*, 100(A7), 12227-12234. doi: 10.1029/95JA00573
- Fennelly, J. A., & Torr, D. G. (1992, January). Photoionization and Photoabsorption Cross Sections of O, N<sub>2</sub>, O<sub>2</sub>, and N for Aeronomic Calculations. *Atomic Data and Nuclear Data Tables*, 51, 321. doi: 10.1016/0092-640X(92)90004-2
- Gentieu, E. P., Feldman, P. D., Eastes, R. W., & Christensen, A. B. (1984, December). EUV airglow during active solar conditions 2. Emission between 530 and 930 Å. *Journal of Geophysical Research*, 89(A12), 11053-11058. doi: 10.1029/JA089iA12p11053
- Gentieu, E. P., Feldman, P. D., & Meier, R. R. (1979, April). Spectroscopy of the extreme ultraviolet dayglow at 6.5Å resolution: Atomic and ionic emissions between 530 and 1240Å. *Geophysical Research Letters*, 6(4), 325-328. doi: 10.1029/GL006i004p00325
- Holsclaw, G. M., Deighan, J., Almatroushi, H., Chaffin, M., Correia, J., Evans, J. S., ... Tyagi, K. (2021, December). The Emirates Mars Ultraviolet Spectrometer (EMUS) for the EMM Mission. *Space Science Reviews*, 217(8), 79. doi: 10.1007/s11214-021-00854-3
- Immel, T. J., England, S. L., Mende, S. B., Heelis, R. A., Englert, C. R., Edelstein, J., ... Sirk, M. M. (2018, February). The Ionospheric Connection Explorer Mission: Mission Goals and Design. *Space Science Reviews*, 214(1), 13. doi: 10.1007/s11214-017-0449-2
- Kil, H., Lee, W. K., Shim, J., Paxton, L. J., & Zhang, Y. (2013). The effect of the 135.6 nm emission originated from the ionosphere on the timed/guvi o/n<sub>2</sub> ratio. *Journal of Geophysical Research*, 118(2), 859-865. Retrieved from <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2012JA018112> doi: <https://doi.org/10.1029/2012JA018112>
- Kramida, A., Yu. Ralchenko, Reader, J., & and NIST ASD Team. (2021). NIST Atomic Spectra Database (ver. 5.9), [Online]. Available: <https://physics.nist.gov/asd> [2022, March 30]. National Institute of

- Standards and Technology, Gaithersburg, MD.
- Lean, J. L., Meier, R. R., Picone, J. M., & Emmert, J. T. (2011, October). Ionospheric total electron content: Global and hemispheric climatology. *Journal of Geophysical Research*, *116*(A10), A10318. doi: 10.1029/2011JA016567
- Meier, R. R. (1991, December). Ultraviolet spectroscopy and remote sensing of the upper atmosphere. *Space Science Reviews*, *58*(1), 1-185. doi: 10.1007/BF01206000
- Meier, R. R. (2021, March). The Thermospheric Column O/N<sub>2</sub> Ratio. *Journal of Geophysical Research*, *126*(3), e29059. doi: 10.1029/2020JA029059
- Meier, R. R., Picone, J. M., Drob, D., Bishop, J., Emmert, J. T., Lean, J. L., ... Gibson, S. T. (2015). Remote sensing of earth's limb by timed/guvi: Retrieval of thermospheric composition and temperature. *Earth and Space Science*, *2*(1), 1-37. Retrieved from <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/2014EA000035> doi: <https://doi.org/10.1002/2014EA000035>
- Meier, R. R., Samson, J. A. R., Chung, Y., Lee, E. M., & He, Z. X. (1991, August). Production of N<sup>+</sup> from N<sub>2</sub> + hv: Effective EUV emission yields from laboratory and dayglow data. *Planetary and Space Science*, *39*(8), 1197-1207. doi: 10.1016/0032-0633(91)90171-6
- Mende, S. B., Frey, H. U., Rider, K., Chou, C., Harris, S. E., Siegmund, O. H. W., ... Ellis, S. (2017, October). The Far Ultra-Violet Imager on the Icon Mission. *Space Science Reviews*, *212*(1-2), 655-696. doi: 10.1007/s11214-017-0386-0
- Mendillo, M., Huang, C.-L., Pi, X., Rishbeth, H., & Meier, R. (2005, October). The global ionospheric asymmetry in total electron content. *Journal of Atmospheric and Solar-Terrestrial Physics*, *67*(15), 1377-1387. doi: 10.1016/j.jastp.2005.06.021
- Picone, J. M., Hedin, A. E., Drob, D. P., & Aikin, A. C. (2002). Nrlmsise-00 empirical model of the atmosphere: Statistical comparisons and scientific issues. *Journal of Geophysical Research*, *107*(A12), SIA 15-1-SIA 15-16. Retrieved from <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2002JA009430> doi: <https://doi.org/10.1029/2002JA009430>
- Prölss, G. W., & Bird, M. K. (2004). *Physics of the Earth's Space Environment: an introduction*.
- Sirk, M. M., Korpela, E. J., Ishikawa, Y., Edelstein, J., Wishnow, E. H., Smith,

- C., ... Immel, T. J. (2017, October). Design and Performance of the  
 ICON EUV Spectrograph. *Space Science Reviews*, 212(1-2), 631-643. doi:  
 10.1007/s11214-017-0384-2
- Stephan, A. W. (2016, September). Advances in remote sensing of the daytime iono-  
 sphere with EUV airglow. *Journal of Geophysical Research*, 121(9), 9284-9292.  
 doi: 10.1002/2016JA022629
- Stephan, A. W., Korpela, E. J., Sirk, M. M., England, S. L., & Immel, T. J. (2017,  
 October). Daytime Ionosphere Retrieval Algorithm for the Ionospheric Connec-  
 tion Explorer (ICON). *Space Science Reviews*, 212(1-2), 645-654. doi: 10  
 .1007/s11214-017-0385-1
- Stephan, A. W., Meier, R. R., England, S. L., Mende, S. B., Frey, H. U., & Immel,  
 T. J. (2018, February). Daytime O/N<sub>2</sub> Retrieval Algorithm for the Iono-  
 spheric Connection Explorer (ICON). *Space Science Reviews*, 214(1), 42. doi:  
 10.1007/s11214-018-0477-6
- Stephan, A. W., Picone, J. M., Budzien, S. A., Bishop, R. L., Christensen, A. B.,  
 & Hecht, J. H. (2012, January). Measurement and application of the O II  
 61.7 nm dayglow. *Journal of Geophysical Research*, 117(A1), A01316. doi:  
 10.1029/2011JA016897
- Strickland, D. J., Evans, J. S., & Paxton, L. J. (1995, July). Satellite remote sensing  
 of thermospheric O/N<sub>2</sub> and solar EUV. 1. Theory. *Journal of Geophysical Re-  
 search*, 100(A7), 12217-12226. doi: 10.1029/95JA00574
- Strickland, D. J., Meier, R. R., Walterscheid, R. L., Craven, J. D., Christensen,  
 A. B., Paxton, L. J., ... Crowley, G. (2004, January). Quiet-time seasonal  
 behavior of the thermosphere seen in the far ultraviolet dayglow. *Journal of  
 Geophysical Research*, 109(A1), A01302. doi: 10.1029/2003JA010220
- Young, J. M., Carruthers, G. R., Holmes, J. C., Johnson, C. Y., & Patterson, N. P.  
 (1968, May). Detection of Lyman- $\beta$  and Helium Resonance Radiation in the  
 Night Sky. *Science*, 160(3831), 990-991. doi: 10.1126/science.160.3831.990