

Minimal climate impacts from short-lived climate forcers following emission reductions related to the COVID-19 pandemic

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Key Points:

- Emission reductions are likely to have led to a global reduction in short-lived climate forcers and tropospheric oxidising capacity.
- Reductions in O₃ and aerosol from both lower emissions and decreased sulphate oxidation resulted in a net negative radiative forcing.
- The radiative impacts are small and short-lived. Longer term climate impacts must come through future sustained emission reductions.

Abstract

We present an assessment of the impacts on atmospheric composition and radiative forcing of short-lived pollutants following worldwide decrease in anthropogenic activity and emissions comparable to what has occurred in response to the COVID-19 pandemic, using the global composition-climate model UKCA. Changes in emissions reduce tropospheric hydroxyl radical and ozone burdens, increasing methane lifetime. Reduced SO₂ emissions and oxidising capacity lead to a decrease in the sulphate aerosol burden and increase in aerosol particle size, with accompanying reductions to cloud droplet number concentration. However, large reductions in black carbon emissions increase the albedo of aerosols. Overall, the changes in ozone and aerosol direct effects (neglecting aerosol-cloud interactions) result in an instantaneous radiative forcing of -31 to -74 mWm⁻². Upon cessation of emission reductions, the short-lived climate forcers rapidly return to pre-COVID levels, meaning these changes are unlikely to have lasting impacts on climate assuming emissions return to pre-intervention levels.

41 Plain Language Summary

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43 As a result of the global COVID-19 pandemic, unprecedented lockdown measures have been
44 imposed worldwide to reduce the spread of the disease, causing huge reductions in economic
45 activity and corresponding reductions in transport, industrial and aviation emissions. As well as
46 lowering emissions of long lived greenhouse gases, such as carbon dioxide, this has resulted in a
47 dramatic reduction in the emissions of components which affect climate in the short term. In this
48 study we have used state-of-the-art computer simulations to quantify how changes in these
49 components are likely to impact the chemical make-up of the atmosphere and the likely short-
50 term impacts on climate. Despite large decreases in nitrogen dioxide and atmospheric particles,
51 we find these changes result in a very small impact on the energy balance of the atmosphere but
52 one that would act to cool the planet, without considering the knock-on impacts on clouds.
53 However, these effects are all likely to be short-lived if emissions return to pre-lockdown levels.

54 1 Introduction

55 The outbreak of the COVID-19 coronavirus disease in China in December 2019 and its global
56 spread in early 2020 has led to the most deadly and disruptive pandemic in recent memory. As of
57 8 June, there have been 6.8 million confirmed cases and 395,000 deaths globally (WHO). In
58 response, governments around the world have implemented varying lockdown measures. The
59 resulting decreases in transport and economic activity have led to the unprecedented reduction of
60 anthropogenic emissions of carbon dioxide (CO₂) (Le Quéré, 2020) and short-lived climate
61 forcers (SLCF) (Zhang et al., 2020). The SLCFs include sulphur dioxide (SO₂), nitrogen oxides
62 (NO and NO₂, which together form NO_x), carbon monoxide (CO), and organic carbon and black
63 carbon (OC and BC respectively). Such species perturb the oxidant balance of the atmosphere
64 (O'Connor et al., 2020), the ozone budget (Young et al., 2018) and aerosol burden (Karset et al.,
65 2018), and thus the radiative balance of the atmosphere and climate (Myhre et al., 2013). This
66 paper aims to assess how the perturbations to atmospheric composition arising from changes to
67 emissions of SLCFs due the COVID-19 pandemic affect parameters important for climate.

68

69 There remains uncertainty in the temporal, spatial, and composition changes to emissions arising
70 from the restrictions imposed. Le Quéré et al (2020) calculated reductions in daily CO₂
71 emissions of between 11 and 25% by April 2020 relative to April 2019. Despite this uncertainty
72 there exists common themes to emissions changes on which this study focuses.

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75 2 Methods

76 2.1 Model description

77 Five experiments were performed using the United Kingdom Chemistry and Aerosols Model
78 (UKCA) run at a horizontal resolution of 1.25°×1.9° with 85 vertical levels up to 85 km (Walters
79 et al., 2019) with the fully interactive stratospheric and tropospheric chemistry (Archibald et al.,
80 2020), and GLOMAP-mode aerosol scheme which simulates sulphate, sea-salt, black carbon,
81 organic matter, and dust but not currently nitrate aerosol (Mulcahy et al., 2020). Emissions were
82 from the CMIP6 CEDS inventories (Hosely et al. 2018). Emissions of methane (CH₄) and carbon
83 dioxide (CO₂) were not simulated, rather a prescribed value is applied for CO₂ and a lower

84 boundary condition (fixed surface concentration) used for methane. The simulations were run
 85 using nudging (Telford et al., 2008) to atmospheric reanalyses from ECMWF (Dee et al., 2011)
 86 to constrain the simulations to consistent meteorology enabling a small ensemble of three
 87 different years: 2012, 2013, and 2014. The years chosen were the most recent CMIP6 emissions
 88 available at the time, and were averaged to filter out the influence of interannual meteorological
 89 variation. Nudging prevented temperatures and horizontal winds from responding to the forcings
 90 produced by the emissions changes, thus preventing changes in aerosols from affecting clouds
 91 and the subsequent impacts on the radiation budget (Zhang et al., 2014).

92 93 94 **2.2 Scenario descriptions**

95
96 Five scenarios were considered, each with different perturbations to emissions (Table 1). Emitted
 97 species are specified in Table S1. The perturbation scenarios A1-A4 were developed by reducing
 98 global anthropogenic emissions in the aviation, surface transport, and industrial sectors by a set
 99 factor. In all perturbation scenarios, emissions were held at the control run values until mid-
 100 February before declining linearly until mid-March to their minimum value. They remained at
 101 their minimum value until mid-May before increasing linearly to the control levels by mid-June
 102 (Fig. S1). We made the approximation of all countries in the world making parallel emission
 103 reductions. As these scenarios were developed early in the COVID-19 pandemic when
 104 information on the impact of the lockdown on all sectors was unknown, we drew on available
 105 information from a number of sources to compile emission reduction scenarios that span likely
 106 representative changes in emissions. See Text S1 for further details.

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109
110 **Table 1 - Scenarios and emission changes**

Scenario	Transport	Aircraft	Industry	% Global change in surface emissions during “lockdown period” (March-May)			
				NO	SO ₂	BC	OC
Control	No reduction	No reduction	No reduction	No reduction	No reduction	No reduction	No reduction
A1	-50%	-50%	-25%	-15.8	-8.84	-11.88	-3.66
A2	-50%	-25%	-25%	-15.8	-8.84	-11.88	-3.66
A3	-75%	-50%	-25%	-22.2	-9.48	-16.48	-4.52
A4	-50%	-50%	No reduction	-12.8	-1.27	-9.19	-1.73

111
112 The scenarios were designed to allow a comparison between the effects of decreasing different
 113 sectors on emissions. By comparing A1 with A3 and A4, we saw that global NO_x emissions were

114 approximately twice as sensitive to surface transport emissions than industrial emissions, while
 115 the majority of SO₂ emission decreases were due to industrial emissions. Comparing the primary
 116 aerosol emissions, BC was more sensitive to the surface transport sector, whilst OC was more
 117 sensitive to industry. While reducing aviation emissions resulted in a negligible decrease in the
 118 total mass of emissions, these emissions were injected directly into the free troposphere which is
 119 more sensitive to NO_x emissions (Stevenson et al., 2004). These reductions are in line with those
 120 in the recently published study by Le Quéré et al (2020) which estimates decreases in aviation of
 121 50-90%, in surface transport of 40-75% and various industrial emissions such as Chinese coal
 122 (40%) and US steel (35%).

123

124 3 Results

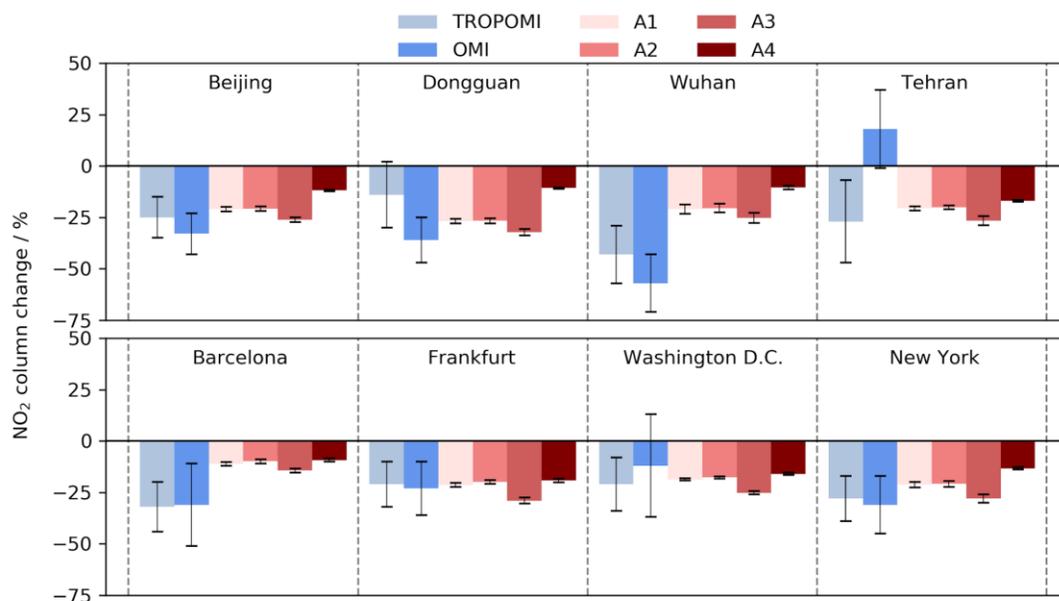
125 In all cases we combined the results from the simulations with different years of meteorology to
 126 generate an ensemble mean, and compared the results of the different scenarios (A1-A4) to the
 127 control case. In all the scenarios the effects of emissions changes were short lived and
 128 atmospheric composition returned to control levels within a couple months of the emissions
 129 reductions ceasing. In the following analyses we focus on the lockdown period (mid-March to
 130 mid-May), where emissions are prescribed to be at their lowest, and quantify changes in
 131 composition and average instantaneous radiative forcing (IRF) from O₃ and aerosol direct
 132 effects.

133

134 3.1 Evaluation of NO₂ Column

135

136 Observations of tropospheric NO₂ columns have exhibited significant reductions globally
 137 (Bauwens et al., 2020, Zhang et al., 2020) with decreases in excess of 20% over many major
 138 cities. Figures 1 and S2 show NO₂ column changes from observation (Bauwens et al., 2020) and
 139 model scenarios.



140

141 **Figure 1. Observed and modelled tropospheric NO₂ column changes. Observations are**
142 **from TROPOMI and OMI relative to 2019, see Bauwens et al. (2020) for more details.**
143 **Model results are from the 4 scenarios relative to the control averaged over the period of**
144 **lowest emissions (mid-March to mid-May).**

145
146 Figure 1 highlights that our model simulations are in good agreement with observed NO₂ column
147 decreases by Bauwens et al. (2020), with the A1 scenario being within error in most cases. This
148 increases confidence in the representativeness of our emissions scenarios for the COVID-19
149 changes. However, we note that the model simulations generally underestimate the magnitude of
150 NO₂ column changes, suggesting our emission perturbations may be at the lower end of what
151 happened during the pandemic. Shi and Brasseur (2020) showed through surface observation
152 analyses across China that the COVID-19 lockdowns resulted in significant decreases in NO₂,
153 but increases in ozone (O₃). These local increases in surface O₃ in polluted regions are also
154 captured in our simulations, although with a smaller magnitude (Fig. 2), and are driven by the
155 non-linear NO_x-VOC chemistry that produces O₃ in the troposphere (Monks et al., 2016).
156 However, all scenarios exhibited a general decrease in global tropospheric O₃, attributed to the
157 reduction in NO_x emissions.

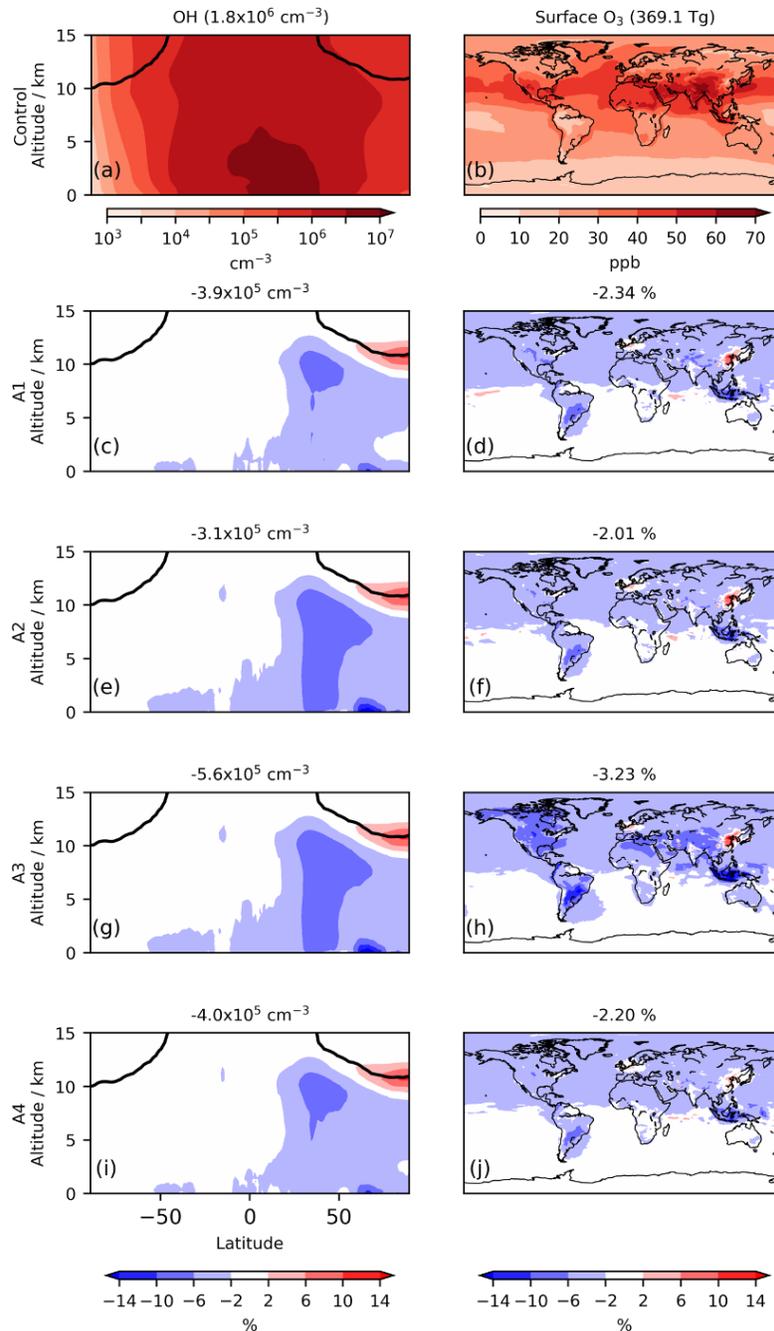
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159 **3.2 Reduction in Oxidant Burden**

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161 Globally averaged, the changes to emissions from transport, industry and aviation led to
162 decreases in the tropospheric O₃ burden of 2.0-3.2 % (Fig. 2, S3), which recovered quickly once
163 emissions increased. The OH concentration was also simulated to have decreased (Fig. 2). The
164 reduction in tropospheric O₃ was most pronounced in A3 where localised decreases exceeded
165 7%, illustrating the large impact of reducing surface transport emissions. The Northern
166 Hemisphere midlatitudes, the location of the largest absolute change in emissions, saw the
167 greatest reductions. Spatial heterogeneity in OH and O₃ depletion between scenarios revealed the
168 importance of emissions from surface transport and aviation (Fig. S4). The additional decreases
169 in low altitude O₃ and OH in Scenario A3 relative to A1 were attributed to the greater reduction
170 in surface transport emissions in A3, while smaller decreases in mid altitude O₃ and OH in A2
171 were due to the smaller reduction in aviation emissions. The similarity in O₃ and OH between
172 scenarios A1 and A4 signified the tropospheric oxidant budget is relatively insensitive to
173 industrial emissions.

174



175
 176 **Figure 2. Zonal mean OH and surface O₃ mixing ratios in control runs and respective**
 177 **changes (mid-March to mid-May). Model results are the ensemble mean for each scenario.**
 178 **Black lines in the OH plots show the tropopause. Titles in the left column show mean**
 179 **tropospheric air-mass-weighted [OH] in control (top) and change (lower panels). Titles in**
 180 **the right column show mean tropospheric O₃ burden in control (top) and change (lower**
 181 **panels).**

182
 183 The decrease in tropospheric OH did not affect model methane concentration due to the fixed
 184 methane surface boundary condition. However, the change in methane concentration, c , which

185 would have occurred can be calculated from the methane lifetime (Eq.1) (Thornhill et al., 2020),
 186 where $f=1.33$ is methane's feedback on its own lifetime (Fiore et al., 2009).

$$187 \quad \frac{\Delta c}{c} = \left(\frac{\Delta \tau}{\tau} + 1\right)^f - 1 \approx f \frac{\Delta \tau}{\tau}$$

188 **Equation 1**

189
 190 Methane lifetime increased by 2-2.5% (A1-2, A4) and 4% (A3) over the period of emissions
 191 reduction (Fig. S5). This would correspond to increases in methane concentration of ~20-40 ppb
 192 if steady state were reached. However, given the perturbations' brevity, much smaller increases
 193 of 1-2 ppb were calculated as an upper bound (Text S2). We therefore conclude the effect on
 194 methane concentration and the associated forcing are negligible.

195 **3.3 Reduction in Sulphate Aerosol Burden**

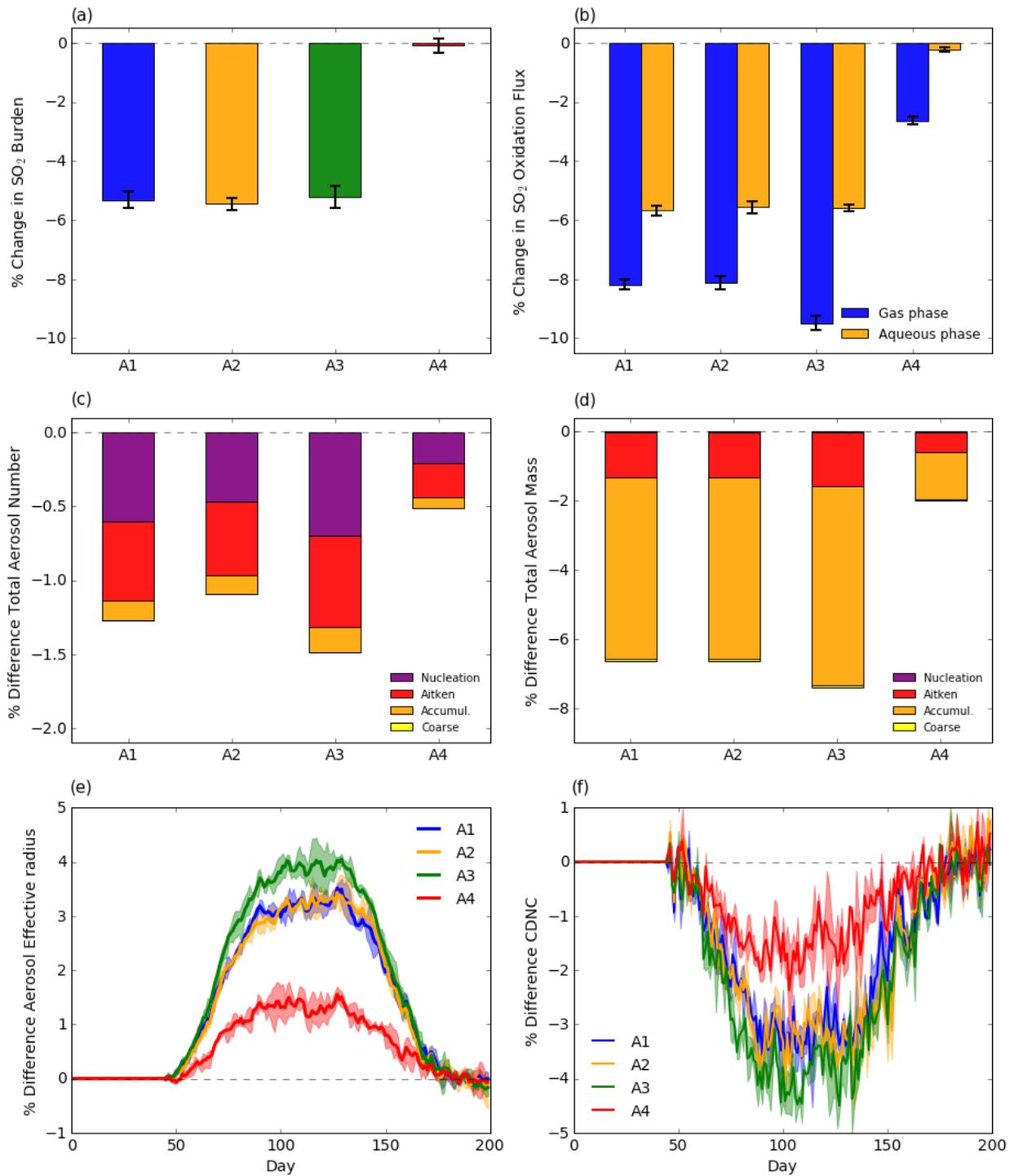
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 197
 198 The perturbation to oxidants reduced the oxidation flux of SO₂ beyond the change due to the
 199 reduction in SO₂ emissions alone, illustrating the coupling between emissions, oxidants and
 200 sulphate aerosol, an important climatic forcer. SO₂ production fluxes (emissions plus chemical
 201 production) decreased by around 8% (A1-A3) and 1.3% (A4), highlighting the sensitivity of SO₂
 202 to industrial emission reductions. However, the corresponding drop in SO₂ burden (5.4% (A1-
 203 A3) and 0.1% (A4)) (Fig. 3(a)) was smaller than the production flux decrease due to a reduction
 204 in chemical loss driven by oxidant decreases. This effect was most pronounced with the
 205 tropospheric gas phase OH + SO₂ flux which decreased by 8-9.5% (A1-A3) and 2.6% (A4) (Fig.
 206 3(b)) and showed significant spatial similarity to [OH] change and exceeded the changes in SO₂
 207 alone (Fig. S5).

208
 209 The other SO₂ oxidation pathway, aqueous oxidation by H₂O₂ and O₃, decreased by only 4%
 210 (A1-A3), meaning relatively more SO₂ was oxidised via aqueous phase chemistry. This is
 211 important because in UKCA, the H₂SO₄ produced via OH + SO₂ oxidation can nucleate new
 212 particles and thus affects aerosol number and size distribution. However, the aqueous phase
 213 pathway only adds mass to existing particles. The different reductions in gaseous and aqueous
 214 flux causes an additional perturbation to the aerosol size distribution resulting in fewer, larger
 215 aerosols (Fig. 3 (c, e)).

216
 217 We calculated a reduction in sulphate aerosol burden (with rapid post-lockdown recovery) with
 218 non-uniform reduction across the aerosol modes and largest changes in the mid latitude Northern
 219 Hemisphere (Figs. S6,7). The largest decrease in mass occurred in the accumulation mode (Fig.
 220 3(c)) and the largest decrease in number in the nucleation mode (Fig. 3(d)). This perturbation to
 221 the size distribution produced an increase in the mean aerosol effective radius (r_{eff}) of 1-4% (Fig.
 222 3(e)) and is attributed in part to the greater relative reduction of gas phase oxidation of SO₂ (and
 223 thus new particle nucleation) than aqueous phase oxidation: a further illustration of coupling
 224 between composition and climatically-relevant agents.

225
 226 The perturbation to the aerosol size and number distribution resulted in cloud droplet number
 227 concentration (CDNC) decreases of up to 4% globally (Fig. 3(f)), with localised decreases
 228 exceeding 10% (Fig. S8) and commensurate increase in effective cloud droplet radius of 0.25-
 229 0.4% (Figs. S9, S10). The drop in CDNC is likely to reduce cloud albedo (Twomey, 1977) and
 230 thus contribute a positive forcing but this has not been calculated here.

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Figure 3 - Mean change in (a) SO₂ burden, (b) SO₂ oxidation flux, (c) sulphate aerosol number, and (d) mass burden split by aerosol size (March to May). Mean change in (e) r_{eff} and (f) CDNC (error bars and shading show ensemble range).

3.4 Aerosol Optical Depth

239

240 The decrease in simulated sulphate aerosol burden and emissions of primary aerosol (BC, OC)
 241 results in decreases in aerosol optical depth (AOD) at 550 nm across most terrestrial regions in
 242 scenarios A1-A3 (Fig. S11) with rapid increase upon emission increase. Eastern China exhibits
 243 the largest absolute decreases (Fig. S12) while A4 showed much smaller decreases, highlighting
 244 the major contribution of industrial SO₂ emissions to AOD. Observed AOD changes between
 245 2017-2019 and 2020 from VIIRS (Sayer et al., 2018) were analysed (Fig. S13) but showed little
 246 significant sign due to considerable noise.

247

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249 4. Radiative Effects

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251 The impact of ozone reduction (Fig. S14) was estimated using the conversion factor of 0.042
 252 Wm⁻² DU⁻¹ (Stevenson et al., 2013). The IRF resulting from aerosol direct radiative effects,
 253 IRF_{DRE} , was calculated by comparing the total outgoing flux, F , and outgoing clean air flux,
 254 F_{clean} , between the perturbed and control runs (Eq. 2) following Ghan (2013).

255

$$IRF_{DRE} = \Delta(F - F_{clean})$$

256

Equation 2

257

258 The IRF_{DRE} was calculated to be significantly smaller than the O₃ forcings in A1 and A2 but
 259 comparable in A3 and A4. Despite the warming effect expected from the reduction in sulphate
 260 aerosol, the global aerosol IRF was simulated to be negative in all scenarios (Table 2).

261

262 **Table 2 - IRF relative to control runs averaged over period of lowest emissions (mid**
 263 **March - mid May).**

264

Values in parentheses show the ensemble range.

Instantaneous Radiative Forcing / mWm ⁻²	A1	A2	A3	A4
Ozone	-34 (-37 to -31)	-29 (-32 to -27)	-47 (-50 to -43)	-32 (-35 to -30)
Aerosol Direct Effect IRF	-4 (-9 to +3)	-2 (-8 to +6)	-27 (-34 to -18)	-44 (-47 to -40)
Ozone and Aerosol IRF	-38	-31	-74	-66

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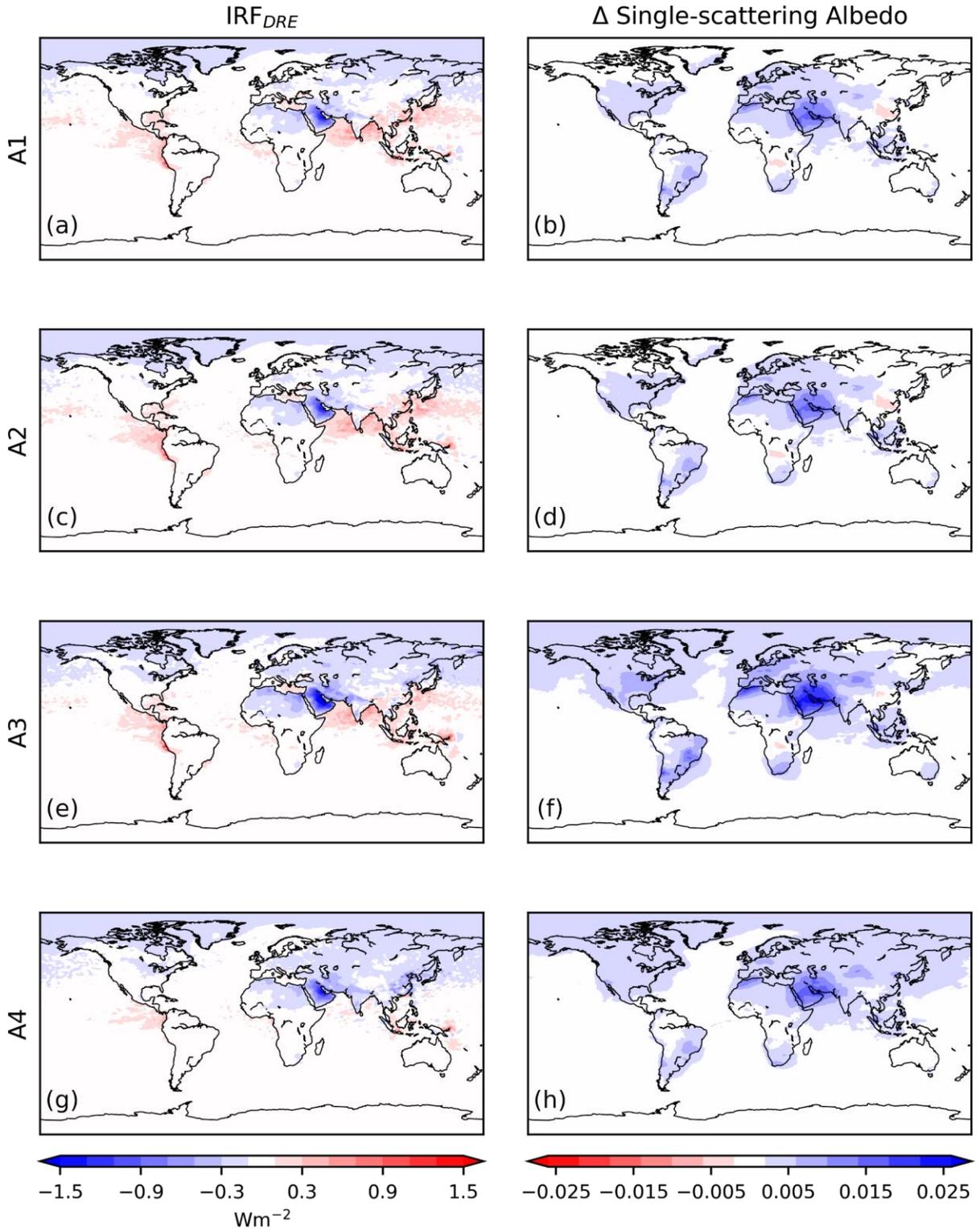


Figure 4. IRF from aerosol direct effects (IRF_{DRE}) and change in single-scattering albedo (March to May) averaged over the 3 years investigated.

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270 Spatial analysis of the IRF_{DRE} (Fig. 4) revealed a negative forcing over large Northern
271 Hemisphere continental regions except Eastern China, the location of the greatest reduction in
272 SO_2 emissions, which exhibited a warming in A1-A3. This warming was attributed to SO_2
273 emission reductions associated with industry as it was not simulated in A4.

274
275 The negative forcing was especially strong over the Arabian peninsula. This was attributed to the
276 fact that the reduction in aerosol exposed solar radiation to a surface with a higher albedo than
277 the original aerosol population, resulting in a greater fraction of insolation being reflected
278 (Haywood and Shine., 1995). This effect was compounded on the Arabian peninsula by the large
279 decreases in black carbon emissions from both surface transport and industry sectors (Fig. S16),
280 a strongly absorbing aerosol component with low single-scattering albedo (Bond et al., 2013).
281 Accordingly, the increase in single-scattering albedo (Fig. 4) is most pronounced over the
282 Arabian peninsula and correlates well with the negative IRF_{DRE} . In addition, the reduction in SO_2
283 emissions (Fig. S14) was much more modest in this region and therefore the associated warming
284 effects were smaller. Globally these competing aerosol forcing effects almost completely offset
285 in A1 and A2 while the greater reductions in black carbon from surface transport resulted in net
286 cooling in A3. The even larger cooling in A4 was attributed to the combination of BC emissions
287 reduction from transport with the much smaller reduction in SO_2 emissions without industry
288 mitigation, resulting in higher aerosol SSA and a negative forcing.

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292

293 **5. Conclusion**

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295 In this study we investigated the impacts on atmospheric composition and radiative forcing from
296 changes in anthropogenic road transport, aviation, and industrial emissions comparable to those
297 resulting from the response to the COVID-19 pandemic. Our model results have shown these
298 emission reductions led to significant changes in atmospheric composition, driven by the
299 changes in the oxidising capacity of the troposphere and oxidant-aerosol-precursor interactions.
300 Decreases in NO_x emissions reduced tropospheric O_3 and as a result the oxidising capacity, with
301 concomitant increases in methane lifetime although a negligible increase in methane forcing.
302 SO_2 emission reductions and the reduction in tropospheric oxidising capacity led to decreases in
303 sulphate burden. The reduction in sulphate aerosol number is predominantly manifest in the
304 nucleation mode; attributed in part to the greater relative reduction in gas phase SO_2 oxidation
305 compared to aqueous phase oxidation and supported by increases in aerosol effective radius and
306 decreases in CDNC. This highlights the influence of oxidant changes on the aerosol size
307 distribution (as well as aerosol burden), an important climatic parameter.

308

309 Despite reduction in the sulphate aerosol burden, decreases in BC emissions resulted in a
310 negative forcing from the aerosol direct effect which, when combined with the negative forcing
311 from tropospheric O_3 reduction, led to a small negative forcing of 31-74 mWm^{-2} . This change is
312 short-lived and comparable to a *temporary* decrease of 3-6 ppm of CO_2 . Due to model setup
313 limitations these estimates do not include impacts from aerosol-cloud interactions. However we
314 can speculate from the reductions in aerosol number and CDNC that cloud effects would have a

315 positive forcing, reducing the overall magnitude of the forcing we calculated and potentially
316 changing its sign.

317

318 Our results suggest that temporary changes to SLCF emissions due to the COVID-19 emergency
319 measures are not going to have a significant impact on near-term climate change, implying that
320 changes in CO₂ emissions during the lockdown period and following recovery will be more
321 important in determining the lasting impact of the pandemic on climate. Elucidating the full
322 effective radiative forcing, including aerosol-cloud interactions, and the climate response due to
323 emission changes warrants further investigation using longer free-running simulations.

324

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336 *All co-authors contributed to writing and reviewing the manuscript.*

337 *Data availability: All necessary data is in the process of being deposited on the CEDA archive*
338 *https://arrivals.ceda.ac.uk/uploader/browse/Covid-19_emiss_reduc_study/*

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