

# Analysis of Recent Anthropogenic Surface Emissions from Bottom-up Inventories and Top-down Estimates: Are Future Emission Scenarios Valid for the Recent Past?

N. Elguindi<sup>1</sup>, C. Granier<sup>1,2</sup>, T. Stavrou<sup>3</sup>, S. Darras<sup>4</sup>, M. Bauwens<sup>3</sup>, H. Cao<sup>5</sup>, C. Chen<sup>6,7</sup>, H.A.C. Denier van der Gon<sup>8</sup>, O. Dubovik<sup>6,7</sup>, T.M. Fu<sup>9</sup>, D. Henze<sup>10</sup>, Z. Jiang<sup>11</sup>, J.J.P. Kuenen<sup>8</sup>, J. Kurokawa<sup>12</sup>, C. Lioussé<sup>1</sup>, K. Miyazaki<sup>13</sup>, J.-F. Müller<sup>3</sup>, Z. Qu<sup>10,14</sup>, K. Sekou<sup>15,16</sup>, F. Solmon<sup>1</sup>, B. Zheng<sup>17</sup>

<sup>1</sup>Laboratoire d'Aérodynamique, CNRS, Université de Toulouse, France

<sup>2</sup>NOAA/ESRL/CSD- CIRES/University of Colorado, Boulder, CO, USA

<sup>3</sup>Royal Belgian Institute for Space Aeronomy, Brussels, Belgium

<sup>4</sup>Observatoire Midi-Pyrénées, Toulouse, France

<sup>5</sup>University of Colorado Boulder, USA

<sup>6</sup>Laboratoire d'Optique Atmosphérique (LOA), UMR8518 CNRS, Université de Lille, Villeneuve D'ASCQ, 59655, France

<sup>7</sup>GRASP-SAS, Remote Sensing Developments, Université de Lille, Villeneuve D'ASCQ, 59655, France

<sup>8</sup>TNO, dept. Climate, Air and Sustainability, Utrecht, The Netherlands

<sup>9</sup>School of Environmental Science and Engineering, Southern University of Science and Technology, Shenzhen, Guangdong Province, China

<sup>10</sup>Department of Mechanical Engineering, University of Colorado Boulder, Boulder, CO, USA

<sup>11</sup>School of Earth and Space Sciences, University of Science and Technology of China, Hefei, Anhui, 230026, China

<sup>12</sup>Asia Center for Air Pollution Research, 1182 Sowa, Nishi-ku, Niigata, Niigata, 950-2144, Japan

<sup>13</sup>Jet Propulsion Laboratory, California Institute of Technology

<sup>14</sup>John A. Paulson School of Engineering and Applied Science, Harvard University, Cambridge, MA, USA

<sup>15</sup>University PGC, UFR sciences Biologiques - Korhogo, Côte d'Ivoire

<sup>16</sup>Laboratoire de Physique de l'Atmosphère (LAPA-UFHB)

<sup>17</sup>Laboratoire des Sciences du Climat et de l'Environnement, CEA-CNRS-UVSQ, UMR8212, Gif-sur-Yvette, France

Corresponding author: N. Elguindi ([elgn@aero.obs-mip.fr](mailto:elgn@aero.obs-mip.fr))

## Key Points:

- Top-down emissions are generally within the range of bottom-up inventories and exhibit a similar level of uncertainty, or even less in regions such as China.
- In China, the U.S. and Europe emission trends in the last decade from SSP126 match most closely actual trends from bottom-up and top-down estimates.
- In Western Africa and India recent emission trends from low pollution control scenarios (SSP460 and SSP370, respectively) match most closely actual trends.

## 45 **Abstract**

46 This study compares recent CO, NO<sub>x</sub>, NMVOC, SO<sub>2</sub>, BC and OC anthropogenic emissions from  
47 several state-of-the-art top-down estimates to global and regional bottom-up inventories and  
48 projections from five SSPs in several regions. Results show that top-down emissions exhibit  
49 similar uncertainty as bottom-up inventories in most regions, and even less in some such as  
50 China. In general, for all species the largest discrepancies are found outside of regions such as  
51 the U.S., Europe and Japan where the most accurate and detailed information on emissions is  
52 available. In some regions such as China, which has undergone dynamical economic growth and  
53 changes in air quality regulations during the last several years, the top-down estimates better  
54 capture recent emission trends than global bottom-up inventories. These results show the  
55 potential of top-down estimates to complement bottom-up inventories and to aide in the  
56 development of emission scenarios, particularly in regions where global inventories lack the  
57 necessary up-to-date and accurate information regarding regional activity data and emission  
58 factors such as Africa and India. Areas of future work aimed at quantifying and reducing  
59 uncertainty are also highlighted. A regional comparison of recent CO and NO<sub>x</sub> trends in the five  
60 SSPs indicate that SSP126, a strong-pollution control scenario, best represents the trends from  
61 the from top-down and regional bottom-up inventories in the U.S., Europe and China, while  
62 SSP460, a low-pollution control scenario, lies closest to actual trends in West Africa. This  
63 analysis can be a useful guide for air quality forecasting and near-future pollution  
64 control/mitigation policy studies.

## 65 **1 Introduction**

66 Anthropogenic activities such as energy production, industrial processes, transportation,  
67 agriculture and waste management are responsible for the emissions of gaseous and particulate  
68 pollutants which can both modify the climate and reduce air quality, leading to adverse impacts  
69 on the environment and human health. Accurate and up-to-date emission inventories are essential  
70 to understand the contribution of various human activities, model and predict the related changes  
71 in atmospheric composition, and design cost-effective mitigation strategies. Despite their  
72 paramount importance, large uncertainties and limitations exist in current state-of-the-art global  
73 and regional emission inventories (Crippa et al., 2018). Based on emission estimates from  
74 inventories, along with information regarding socio-economic, environmental, and technological

75 trends, future emission scenarios such as the Shared Socioeconomic Pathways (SSPs) are created  
76 and used by atmospheric and chemistry models to generate future climate and pollutant  
77 concentration projections. As such, uncertainties associated with current emission estimates are  
78 directly propagated into future scenarios of emission trajectories, model climate projections and  
79 air quality forecasts.

80

81 For a given chemical compound, emission inventories rely on the definition of key socio-  
82 economic sectors (i.e. road traffic) involving certain technologies (i.e. car engine) characterized  
83 by specific emission factors (i.e. CO emissions per unit of fuel used per km). This information is  
84 then scaled up using geographically distributed information of the corresponding activities (e.g.  
85 car traffic intensity map) to create large-scale gridded inventories. The complexity of emission  
86 modelling lies in the diversity of chemical species, as well as in the characterization and  
87 quantification of emission factors and sector activities, all of which are highly variable and  
88 influenced by socio-economic and environmental factors. Country-level indicators used to build  
89 global emission inventories and projections often lack up-to-date regional specific information,  
90 especially in developing regions. These inventories are largely created using emission factors  
91 (EF) that are representative of conditions in developed countries, such as Europe and North  
92 America. Although often an EF is selected that represents a low technology level, the origin of  
93 the EF data is from developed countries, thereby introducing large errors and uncertainties into  
94 the emission estimates for developing countries. Furthermore, the collection of data for all  
95 countries throughout the world takes considerable time; by the time the inventories are updated  
96 there is thus often a significant lag from the present day.

97

98 Alternatively, inverse modelling techniques, which constrain atmospheric models by  
99 observations to estimate surface emissions, have been used to derive emissions for various air  
100 pollutants (e.g. Miyazaki et al. 2017; Stavrou et al. 2012; Müller et al. 2005; Arellano et al.  
101 2004). A main advantage of these estimation techniques is their high spatial coverage, in  
102 particular when spaceborne atmospheric data are used as constraints. Another benefit is that they  
103 can provide more timely emission estimates than traditional bottom-up inventories, which are  
104 generally delayed by a few years. Inverse modelling has the potential to reduce uncertainty in air  
105 quality and chemistry climate models by providing more constrained emission data, especially in

106 regions of the world where bottom-up estimates are believed to be deficient. Nevertheless,  
107 inverse modelling has its limitations and uncertainties that should be addressed and quantified.  
108 Notably, uncertainties are associated with the use of atmospheric models, in particular their  
109 representation of transport and chemical processes (e.g. Jiang et al., 2013; Stavrou et al.,  
110 2013). Another source of uncertainties is related to the observations used to constrain the models.  
111 For example, satellite data intercomparison studies revealed large differences between different  
112 retrievals of the same compound as well as significant biases against well-calibrated validation  
113 data, reaching up to a factor of two (Zhu et al., 2016; Zhu et al., 2019b). Therefore, before  
114 inverse modelling estimates can be used to supplement bottom-up emission inventories, an  
115 assessment of their reliability and associated uncertainty over diverse regions is needed.

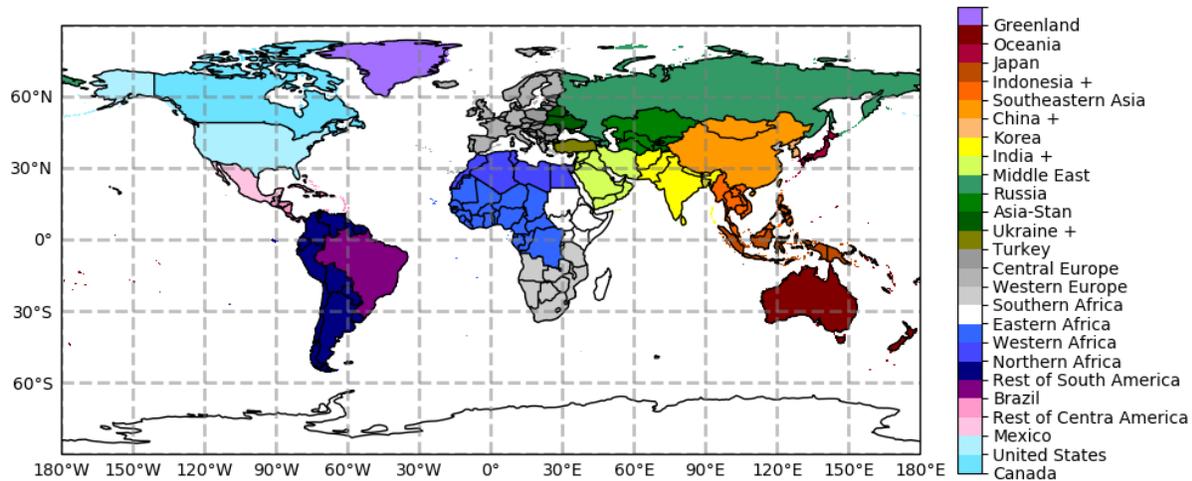
116 The main goal of this study is to provide a comprehensive, systematic comparison of emission  
117 estimates derived from inverse modelling techniques for various species (CO, NO<sub>x</sub>, NMVOC,  
118 SO<sub>2</sub>, BC, and OC) to current state-of-the-art emission inventories for several regions of the  
119 world. To this end, we compiled estimates of anthropogenic emissions from eleven different  
120 sources based on various inverse methods. These estimates are compared to the most widely  
121 used global inventories EDGARv4.3.2 (Crippa et al., 2018) and CEDS (Hoesly et al., 2018), as  
122 well as to recent regional inventories for Europe (CAMs-REG-AP, Granier et al. 2019), Africa  
123 (DACCIWA, <https://www.eccad3.sedoo.fr>) and China (MEICv1.3: <http://www.meicmodel.org>,  
124 [Zheng et al., 2018](#)).

125

126 Another objective is to compare the trends in the projected emissions of the Shared  
127 Socioeconomic Pathways (SSPs) developed for the sixth Climate Model Intercomparison Project  
128 (CMIP6) (Riahi et al. 2017; Kreigler et al. 2012) to the current best estimates of emission trends  
129 for the recent past in selected regions. We aim to evaluate which narrative describes best the  
130 direction taken by the different regions in terms of emissions. Assessing how well the scenarios  
131 capture recent emission trends will be useful for determining their suitability for studies that  
132 evaluate emissions in the recent past and near future, such as pollution control/mitigation impact  
133 studies and air quality forecasting. Furthermore, we compare CO and NO<sub>x</sub> emissions from the  
134 SSPs to those of the Representative Concentration Pathways (RCPs) that were used in the fifth  
135 Climate Model Intercomparison Project (CMIP5) (Moss et al. 2010; van Vuuren et al, 2011) for  
136 several world regions. While a comprehensive overview of the SSP emissions and comparison  
137 with the RCPs is presented in Gidden et al. (2019), we are focused on the regional scale, and  
138 intend to aid in the interpretation and analysis of regional climate change studies.

## 139 **2 Materials and Methods**

140 This study compares 10 different top-down emission estimates to several state-of-the-art global  
141 and regional bottom-up inventories for CO, NO<sub>x</sub>, NMVOC, SO<sub>2</sub>, OC and BC. A description of  
142 each of the top-down estimates is provided below. Details and references for each of the top-  
143 down and bottom-up inventories are also given in Tables 1 and 2, respectively. We focus on the  
144 following 12 regions as defined by the IMAGE2.4 26 regions (Bouwman et al., 2006): China,  
145 the Middle East, Western Africa, the United States, Western and Central Europe, South America,  
146 India+, Southeast Asia, Indonesia+, Oceania and Southern Africa. Note that India+ includes  
147 neighbouring countries such as Pakistan, Afghanistan, Nepal, and Bangladesh. A map of the  
148 regions is shown in Figure 1. These regions were selected based their economic and geographic  
149 diversity, as well as on the availability of top-down and bottom-up regional inventories. The  
150 regional averages for all datasets are calculated on their native grids.



151

152 Figure 1: This study focuses on the following 11 regions as defined by Bouwman et al. (2006),  
 153 the Middle East, Western Africa, the United States, Western and Central Europe, South America,  
 154 India+, Southeast Asia, Indonesia+, Oceania and Southern Africa, plus China which only  
 155 includes the national boundaries.

156

## 157 2.1 Calculation of trends and percent differences

158

159 We evaluate recent trends in emissions from the five SSPs based on the years 2010-2020. The  
 160 trends are derived from ordinary linear regression and expressed in units of percent change per  
 161 year relative to the mean over the data period. The SSP trends are compared to trends which are  
 162 based on the emissions from the inventories and inversion estimates. These trend estimates are  
 163 calculated by averaging the trends from all available inventories for a particular region, including  
 164 all inverse modelling estimates and bottom-up inventories which are available from 2010 until at  
 165 least 2015 which are statistically significant ( $p > 0.05$ ). A list of the individual trends is provided  
 166 in supplementary material. Note that MACCity and CAMS-GLOB-ANT are excluded because  
 167 their recent years are based on projections of past trends from inventories and are therefore not  
 168 considered independent. Unless otherwise specified, percentage differences in the range of  
 169 emission estimates are calculated based on the average between the highest and lowest values.

170

171

## 172 3 Data

### 173 3.1 Top-down Emissions

174 **Global**

175  
176 ***Chen et al. (2019)***: Global daily carbonaceous aerosol (OC and BC) emissions for the period  
177 2006-2011 are derived using the GEOS-Chem adjoint model (Henze et al., 2007) at a spatial  
178 resolution of  $2.5^\circ \times 2^\circ$  (Chen et al., 2019). OC and BC emission sources are constrained using  
179 retrievals of aerosol optical depth (AOD) and aerosol absorption optical depth (AAOD) (Chen et  
180 al. 2018, 2019) from the multi-angular and polarimetric POLDER/PARASOL (Polarization &  
181 Anisotropy of Reflectances for Atmospheric Sciences coupled with Observations from a Lidar)  
182 measurements retrieved via the GRASP (General Retrieval of Atmosphere and Surface  
183 Properties) algorithm ([www.grasp-open.com](http://www.grasp-open.com)) (Dubovik et al., 2011, 2014). Note that the  
184 anthropogenic contribution of the total derived black and organic carbon emissions is estimated  
185 based on a ratio which is applied in the GEOS-Chem (v10-1) model.

186  
187 ***Jiang et al (2017)***: Global CO emissions are constrained over the period 2001-2015 by  
188 assimilating multi-spectral CO measurements from MOPITT (version 6) using the 4D-Var data  
189 assimilation system (adjoint) in the GEOS-Chem model at a spatial resolution of  $5^\circ \times 4^\circ$ . The  
190 initial conditions and land boundary conditions (CO concentrations) are optimized using a  
191 sequential sub-optimal Kalman filter. Here we include two of the inversions described in Jiang et  
192 al. (2017), one in which emissions are constrained using MOPITT CO profiles and another using  
193 total CO columns. Anthropogenic CO emissions are separated from other sources using a single  
194 scaling factor, based on the a priori emissions, to adjust all emissions in a grid.

195  
196 ***Tropospheric Chemistry Reanalysis (TCR-2)***: Global  $\text{NO}_x$  and  $\text{SO}_2$  emissions are constrained  
197 over the period of 2005-2018 by assimilating multiple satellite data sets for multiple species  
198 ( $\text{NO}_2$ , CO,  $\text{O}_3$ ,  $\text{SO}_2$ ) using the global CTM (Chemistry Transport Model) MIROC-Chem  
199 (Watanabe, et al., 2011) based on the ensemble Kalman filter (EnKF) technique performed at  
200  $1.125^\circ \times 1.125^\circ$  resolution in the TCR-2 framework (Miyazaki et al. 2017, 2019, 2020a, 2020b).  
201 The assimilated measurements were obtained from OMI (the Ozone Monitoring Instrument),  
202 GOME-2 (Global Ozone Monitoring Experiment-2) and SCIAMACHY (SCanning Imaging  
203 Absorption SpectroMeter for Atmospheric CHartography) for  $\text{NO}_2$ , TES for  $\text{O}_3$ , MOPITT for  
204 CO, and MLS for  $\text{O}_3$  and  $\text{HNO}_3$ , and OMI for  $\text{SO}_2$ .

205  
206 **Müller et al. (2018):** Global CO emissions for 2013 are derived at  $2^\circ \times 2.5^\circ$  resolution based on  
207 the adjoint of the IMAGESv2 model constrained by satellite IASI CO column data. The model  
208 uses prescribed OH fields constrained by methylchloroform measurements. An ensemble of top-  
209 down simulations is carried out and the top-down emissions are compared with various  
210 independent CO observations and evaluated. The inversion adopting the lowest average OH level  
211 in the Northern Hemisphere provides the best agreement with all tested independent observation  
212 data sets, and the corresponding top-down emissions are used in our study. Emissions are derived  
213 for three categories (anthropogenic, biogenic and biomass burning) using a technique which  
214 relies on the spatio-temporal patterns of the a priori emissions, through assumed correlations  
215 between a priori emission errors. In essence, the inversion tries to preserve the patterns of the a  
216 priori. The strength of that constraint is determined by the assumed correlations.

217  
218  
219 **Zheng et al. (2019):** Global CO emissions for the period 2000-2017 are derived using the global  
220 3-D transport model of the Laboratoire de Météorologie Dynamique (LMDz) coupled with a  
221 simplified chemistry module, Simplified Atmospheric Chemistry assimilation System (SACS)  
222 based on a multi-species atmospheric Bayesian inversion approach (Zheng et al., 2018a, 2018b)  
223 at a spatial resolution of  $3.75^\circ \times 1.9^\circ$ . Zheng et al. (2019) perform the following three global  
224 inversions which are used in this study: (i) an inversion constrained only by CO total column  
225 data from the MOPITT version 7 over 2000-2017 (inversion 1); (ii) an inversion also constrained  
226 by HCHO column data from Ozone Monitoring Instrument (OMI) version 3 on the basis of  
227 inversion 1 for the period 2005-2017 (inversion 2); (iii) an inversion further constrained by  
228 column-averaged dry air mole fractions of  $\text{CH}_4$  ( $\text{XCH}_4$ ) from Greenhouse gases Observing  
229 SATellite (GOSAT) on the basis of inversion 2 for the period 2010-2017 (inversion 3).  
230 Emissions are derived for four categories (anthropogenic, biomass burning, biogenic and  
231 oceanic) by multiplying the optimized 8-daily surface CO total fluxes by the proportion of each  
232 sector in each model grid cell as given by the prior.

233

234 **Regional**

235

236 **Cao et al. (2018):** NMVOC emissions in China for the year 2007 are derived using the GEOS-  
237 Chem CTM (version 8.2.1) and its adjoint at a spatial resolution of  $5^\circ \times 4^\circ$ . The GEOS-Chem  
238 CTM was updated to include improved NMVOC chemical schemes. Emissions are constrained  
239 by using HCHO and glyoxal columns observed by the GOME-2A and OMI satellite instruments.  
240 Four inversion experiments using different combinations of these satellite observations were  
241 conducted in order to explore their impacts on the top-down emission estimates. The  
242 anthropogenic NMVOC emission estimates for 2007 range from 16.4-23.6 Tg yr<sup>-1</sup>. In this study,  
243 we show the average of the four estimates which is 20.2 Tg yr<sup>-1</sup>.

244

245 **Qu et al. (2019a):** NO<sub>x</sub> and SO<sub>2</sub> emissions in East Asia for the period 2005-2012 are derived  
246 simultaneously using the GEOS-Chem adjoint model (Henze et al. 2007, 2009) and a hybrid 4D-  
247 Var / mass balance approach at  $0.5^\circ \times 0.667^\circ$  resolution (Qu et al., 2017). The emissions are  
248 constrained by the OMI NO<sub>2</sub> NASA standard product (Krotkov et al., 2017) and the OMI Royal  
249 Belgian Institute for Space Aeronomy (BIRA) SO<sub>2</sub> product (Theys et al., 2015). We refer to this  
250 inventory as Qu-joint. An inversion is also performed constraining only NO<sub>2</sub> in order to assess  
251 the benefits of constraining multiple species. This inversion is referred to as Qu-single. The  
252 inversions were performed on a limited domain which does not include all of the countries of  
253 the India+ region shown in Figure 1 of the Supplement, therefore the averages for Qu-single and  
254 Qu-joint are for India only.

255

256 **Qu et al. (2019b):** Global SO<sub>2</sub> emissions for the period 2005-2017 are derived using the GEOS-  
257 Chem adjoint model (Henze et al. 2007, 2009) and a hybrid 4D-Var / mass balance approach at a  
258  $2^\circ \times 2.5^\circ$  horizontal resolution. In order to assess the uncertainty related to satellite retrievals,  
259 two inversions are performed using different OMI satellite retrievals to constrain the emissions;  
260 the NASA standard product OMSO2 (Li et al., 2013) and the BIRA product (Theys et al., 2015).  
261 Note that these SO<sub>2</sub> emissions are shown only for China and India because the estimates are less  
262 accurate in comparatively cleaner areas with lower SO<sub>2</sub> columns due to negative column  
263 densities in the OMI SO<sub>2</sub> retrievals (Qu et al., 2019). These inversions are referred to as Qu-  
264 BIRA and Qu-NASA.

265

266 **Stavrakou et al. (2017)**: Global NMVOC emissions for the period 2005-2014 are derived using  
 267 the adjoint model technique in the IMAGESv2 global CTM at a spatial resolution of  $2^\circ \times 2.5^\circ$ .  
 268 Emissions from open fire vegetation and human activities are constrained using vertical columns  
 269 of formaldehyde (HCHO) retrieved from the Ozone Monitoring Instrument (OMI-BIRA). The  
 270 anthropogenic VOC sources are found to be weakly constrained by the inversions on a global  
 271 scale due to their small contribution to the total HCHO columns (Stavrakou et al. 2009), except  
 272 over strongly polluted regions, like China.

273

274

275 **Table 1:** Description of inverse modelling estimates considered in this study.

	<i>Species</i>	<i>Region</i>	<i>Res. (lon x lat)</i>	<i>Period</i>	<i>CTM</i>	<i>Satellite constraint</i>	<i>Inversion method</i>
<b>Global</b>							
<i>Chen et al. (2019)</i>	OC/BC	Global	$2.5^\circ \times 2^\circ$	2006-2011	GEOS-Chem	PARASOL/ GRASP* AOD/AAOD	Adjoint
<i>TCR-2</i>	NO <sub>x</sub>	Global	$2.8^\circ \times 2.8^\circ$	2005-2018	MIROC-Chem	OMI/GOME-2/SCIAMACHY NO <sub>2</sub> , TES O <sub>3</sub> , MOPITT CO, MLS O <sub>3</sub> /HNO <sub>3</sub>	Ensemble Kalman filter
<i>Müller et al. (2018)</i>	CO	Global	$2.5^\circ \times 2^\circ$	2013	IMAGESv2	IASI CO and OH levels based on methylchloroform (MCF) obs.	Adjoint
<i>Zheng et al. (2019)</i>	CO	Global	$3.75^\circ \times 1.9^\circ$	2000-2017	LMDz-SACS	MOPITT CO, OMI HCHO, GOSAT XCH4	Bayesian
<i>Jiang et al. (2017)</i>	CO	Global	$5^\circ \times 4^\circ$	2001-2015	GEOS-Chem	MOPITT CO and OH levels based on MCF obs.	Adjoint
<b>Regional</b>							
<i>Cao et al. (2018)</i>	NMVOC	China	$5^\circ \times 4^\circ$	2007	GEOS-Chem	GOME-2/OMI HCHO/ CHOCHO	Adjoint
<i>Qu et al. (2019a)</i>	NO <sub>x</sub> , SO <sub>2</sub>	East Asia	$0.5^\circ \times 0.667^\circ$	2005-2012	GEOS-Chem	OMI SO <sub>2</sub> , NO <sub>2</sub>	Adjoint
<i>Qu et al. (2019b)</i>	NO <sub>x</sub>	East Asia	$2.5^\circ \times 2^\circ$	2005-2017	GEOS-Chem	OMI NO <sub>2</sub>	Adjoint
<i>Stavrakou et al. (2017)</i>	NMVOC	China	$2.5^\circ \times 2^\circ$	2005-2014	IMAGESv2	OMI HCHO	Adjoint

276

277

278

### 3.2 Bottom-up Emissions

279 In this study we compare 13 of the most recent anthropogenic global and regional bottom-up  
 280 inventories, respectively. Included in the global inventories are EDGARv4.3.2 (JRC, Crippa et  
 281 al., 2018) and CEDS (Hoesly et al., 2018) which are both traditional bottom-up inventories. We  
 282 also compare CAMS-GLOB-ANTv4.1 and MACCity which are based on bottom-up inventories  
 283 and projections, thereby provide emissions to the current year. The CAMS global emissions are  
 284 based on both the EDGARv4.3.2 and CEDS emissions. The standard version of CAMS-GLOB-  
 285 ANT (version 4.1) used in this study applies the monthly profiles provided by CAMS-GLOB-  
 286 TEMPO (Granier et al., 2019) to the annual emissions from EDGARv4.3.2 for the years 2000-  
 287 2012. After 2012, the data are linearly extrapolated to 2020 using trends derived from the CEDS  
 288 emissions for the years 2011-2014. A detailed description of all the global and regional bottom-  
 289 up inventories are presented in the Supplement, along with a brief discussion regarding the  
 290 uncertainties among global inventories.

291

292 **Table 2:** Description of global and regional bottom-up inventories considered in this study.

	<i>Species</i>	<i>Region</i>	<i>Res.</i> (lon x lat)	<i>Period</i>	<i>Reference</i>
<b><i>Global Inventories</i></b>					
<i>CAMS-GLOB-ANT</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Global	0.1° x 0.1°	2000-2019	Granier et al. (2019)
<i>EDGARv4.3.)</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Global	0.1° x 0.1°	1970-2012	Crippa et al. (2018)
<i>CEDS</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Global	0.5° x 0.5°	1950-2014	Hosley et al. (2018)
<i>HTAPv2</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Global	0.1° x 0.1°	2008, 2010	<a href="http://www.htap.org">http://www.htap.org</a>
<i>MACCity</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Global	0.5° x 0.5°	2007	<a href="https://eccad3.sedoo.fr">https://eccad3.sedoo.fr</a>
<b><i>Regional Inventories</i></b>					
<i>CAMS-REG-AP</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Europe	0.1° x 0.05°	2000-2017	Kuenen et al. (2014) Granier et al. (2019)
<i>DACCIWA</i>	NO <sub>x</sub> , CO, NMVOC,	Africa	0.1° x 0.1°	2000-2015	Keita et al. (2020)

	SO <sub>2</sub> , OC, BC				
<i>DICE-Africa</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Africa	0.1° x 0.1°	2000-2015	Marais et al. (2016)
<i>US NEI</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub>	United States		2000-2016	<a href="https://www.epa.gov/air-emissions-inventories">https://www.epa.gov/air-emissions-inventories</a>
<i>MEICv1.3</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	China	0.25° x 0.25°	2008-2017	<a href="http://www.meicmodel.org">http://www.meicmodel.org</a> Zheng et al. (2018)
<i>REASv3.1</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub> , OC, BC	Asia	0.25° x 0.25°	1950-2015	Kurokawa et al. (2019)
<i>Sharma et al. (2019)</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub>	India		2011	Sharma et al. (2019)
<i>Sun et al. (2018)</i>	NO <sub>x</sub> , CO, NMVOC, SO <sub>2</sub>	China		1949-2015	Sun et al. (2018)

293

294

### 295 **3.3 Shared Socioeconomic Pathways (SSPs)**

296

297 The projected emissions from five SSPs were developed for use in the current Coupled Model

298 Intercomparison Project phase 6 (Eyring et al., 2016). The SSPs are global scenarios which

299 describe how the future emissions might evolve according to socioeconomic development,

300 demographics, technological advances within the context of climate change mitigation and

301 adaptation during the period 2015-2100 (van Vuuren et al., 2014; O'Neill et al., 2014; Kreigler et

302 al., 2012). The air pollutant emission trajectories associated with the SSP scenarios have been

303 harmonized with the CEDS historical global inventory for the year 2015 and are described in

304 Rao et al., (2017) and Gidden et al. (2019). The pathways are based on five narratives describing

305 alternative socioeconomic developments. SSP1 and SSP5 assume strong pollution control

306 scenarios, and therefore emissions which are substantially lower than current levels, whereas

307 according to SSP3 and SSP4 future emissions are higher than current levels. SSP2 is based on a

308 medium pollution control scenario where emissions remain close to current levels (Rao et al.

309 2017). In addition, mitigation policies are added to each scenario in order to achieve a

310 prescribed radiative forcing by the end of the 21<sup>st</sup> century (i.e. 2.6, 4.5, 6.0 and 8.5 W m<sup>-2</sup>).

311

312 In this study, we focus on the four Tier 1 SSPs (SSP126, SSP245, SSP370 and SSP585), which  
313 have the same radiative forcing as the RCPs used in CMIP5, but combine socioeconomic and  
314 technological developments, and have been given priority in CMIP6 (O'Neill et al., 2016). In  
315 addition, we also analyze SSP460, which has been designated as a Tier 2 scenario in order to  
316 complement and extend the Tier 1 scenarios and the RCPs (O'Neill et al., 2016). The emissions  
317 associated with these scenarios are available at a  $0.5^\circ \times 0.5^\circ$  spatial resolution ([https://esgf-  
318 node.llnl.gov/projects/input4mips](https://esgf-node.llnl.gov/projects/input4mips)).

#### 319 **4 Comparison of emission inventories and inverse modelling estimates**

320 In this section, we compare inverse modelling anthropogenic emission estimates of CO, NO<sub>x</sub>  
321 NMVOC, SO<sub>2</sub>, BC and OC to global and regional inventories which were developed using  
322 bottom-up estimation methods over the six regions of interest described in Section 2 and shown  
323 in Figure 1.

##### 324 **4.1 CO**

325 The annual average CO emissions for each region are displayed in Figure 2. The inversion  
326 estimates are shown for Zheng et al. (2019), Müller et al. (2018) and Jiang et al. (2017). We  
327 compare two estimates from Zheng; Inv1 which uses satellite constraints of only columns of CO  
328 for the time period 2000-2017, and Inv2 which, in addition, is constrained by both CO and  
329 HCHO columns for 2005-2017. Two estimates from Jiang et al. (2017) are also compared; Jiang-  
330 prof which assimilates profile data and Jiang-colum which assimilates total column data. More  
331 specific details can be found in the Supplement.

332

333 In regions with high emissions such as China, the U.S. and India, the inversion estimates are  
334 generally within the range of the bottom-up inventories in terms of magnitude and have a similar  
335 range of uncertainty. In all regions, Müller's estimates are significantly lower compared to those  
336 of Zheng or Jiang. The lower estimates of Müller et al. (2018) are due to the use of prescribed  
337 modelled OH levels which are based on observations of methylchloroform (MCF) and further  
338 modulated based on comparisons with aircraft profiles and ground-based data. The OH field  
339 prescribed in their inversion setup is at the lower end of the range of what has been reported in  
340 the literature and calculated in most CTMs, implying less CO oxidation. The resulting higher CO  
341 lifetime is in turn compensated by lower optimized emissions. Another factor which could  
342 partially explain the discrepancy among the top-down estimates is the satellite data used in the

343 inversion system. Müller assimilates IASI satellite CO data, while Zheng and Jiang, both of  
344 which have higher values, use MOPITT.

345  
346 In Western and Central Europe, Müller's estimate for 2013 is within the narrow range of CO  
347 emission reported by the bottom-up inventories, and is quite close to the regional CAMS-REG  
348 inventory (Figure 2). However, regarding the U.S., Müller's estimate of CO emission is  
349 considerably below (approx. 35%) the regional U.S. EPA inventory. Müller et al. (2018) attribute  
350 this large discrepancy to an overestimation in the U.S. bottom-up inventory, which has also been  
351 suggested in other studies (Hudman et al., 2008; Anderson et al., 2014; Jiang et al., 2015). In  
352 other regions such as China, India and Western Africa, it is difficult to comment on the accuracy  
353 of the magnitude of the inversion (or bottom-up) emission estimates due to the lack of  
354 measurements and regional information.

355  
356 The two inversion estimates of CO emissions from Zheng are quite similar in magnitude and  
357 trend in most regions, but differ in terms of inter-annual variability (Figure 2). The two estimates  
358 from Jiang show similar trends, but significant differences in magnitude over most regions,  
359 indicating that the type of data assimilated (e.g. profile vs. total column) has a large impact on  
360 the emission estimates. In most regions, the Jiang-prof estimates are higher than Jiang-colum. In  
361 China, there is a sharp decline in emissions after 2010 as a result of the stringent clean air  
362 policies that have been implemented in recent years due to the severe air quality issues which  
363 have been documented in other studies (Zheng et al., 2018). With the exception of the regional  
364 inventories MEIC v1.3 and REASv3.1, which show a decrease of 4.4 and 2.7 % yr<sup>-1</sup> for the  
365 period 2011-2015, most of the inventories do not show the declining CO emissions after 2010.  
366 However, this trend is captured quite well by all of the inversion estimates which show a  
367 decrease in emissions ranging from 1 to 3 % yr<sup>-1</sup> over 2010-2015. While both of Zheng's  
368 estimates capture the magnitude and decreasing trend in China after 2011, Inv1 stabilizes after  
369 2014, whereas Inv2 continues to show a declining trend similar to the MEICv1.3 emissions.  
370 Between 2011 and 2017, the Inv2 estimates decrease at an annual rate of 5.6%, while the Inv1  
371 estimates decrease at a slower rate of 3.3%. The only difference between these two inversions is  
372 that, in addition to constraining columns of CO, formaldehyde (HCHO) is also constrained in  
373 Inv2. Constraining HCHO has a significant influence on the chemical production of CO and the

374 trend in Inv2. Tropospheric columns of OMI HCHO have been reported to keep increasing over  
375 China, likely due to significant increases in NMVOC emissions (Shen et al., 2019; Li et al.,  
376 2019). This could explain why CO emission estimates continue to decrease in Inv2, while Inv1  
377 flattens out since the optimized emissions are overestimated to compensate for the  
378 underestimation of CO photochemical production.

379  
380 In the U.S., the CO inversion estimates from Zheng and Jiang are similar to the U.S. NEI  
381 regional inventory in terms of trends, and slightly higher in magnitude (Figure 2). Both estimates  
382 follow the continuing decreasing trend of the regional inventory up to 2010, after which Jiang's  
383 estimates diverge and indicate a stabilization in CO emissions in the United States. Jiang et al.  
384 (2017) attribute this slowdown of U.S. pollutant reduction to factors such as diminished returns  
385 on improved catalytic converters which they suggest are unaccounted for in the U.S. inventory  
386 (Jiang et al., 2018). Jiang et al. (2018) demonstrate that satellite retrievals and surface  
387 measurements also indicate a significant reduction in the decreasing trends of CO and NO<sub>x</sub>  
388 concentrations after 2010 as compared to the previous years, corroborating the trend shown in  
389 their CO inverse estimates. However, there is not always a direct linear relationship between  
390 anthropogenic NO<sub>x</sub> emissions and measurements of tropospheric NO<sub>2</sub>. This point is further  
391 discussed in the following section on NO<sub>x</sub> emissions.

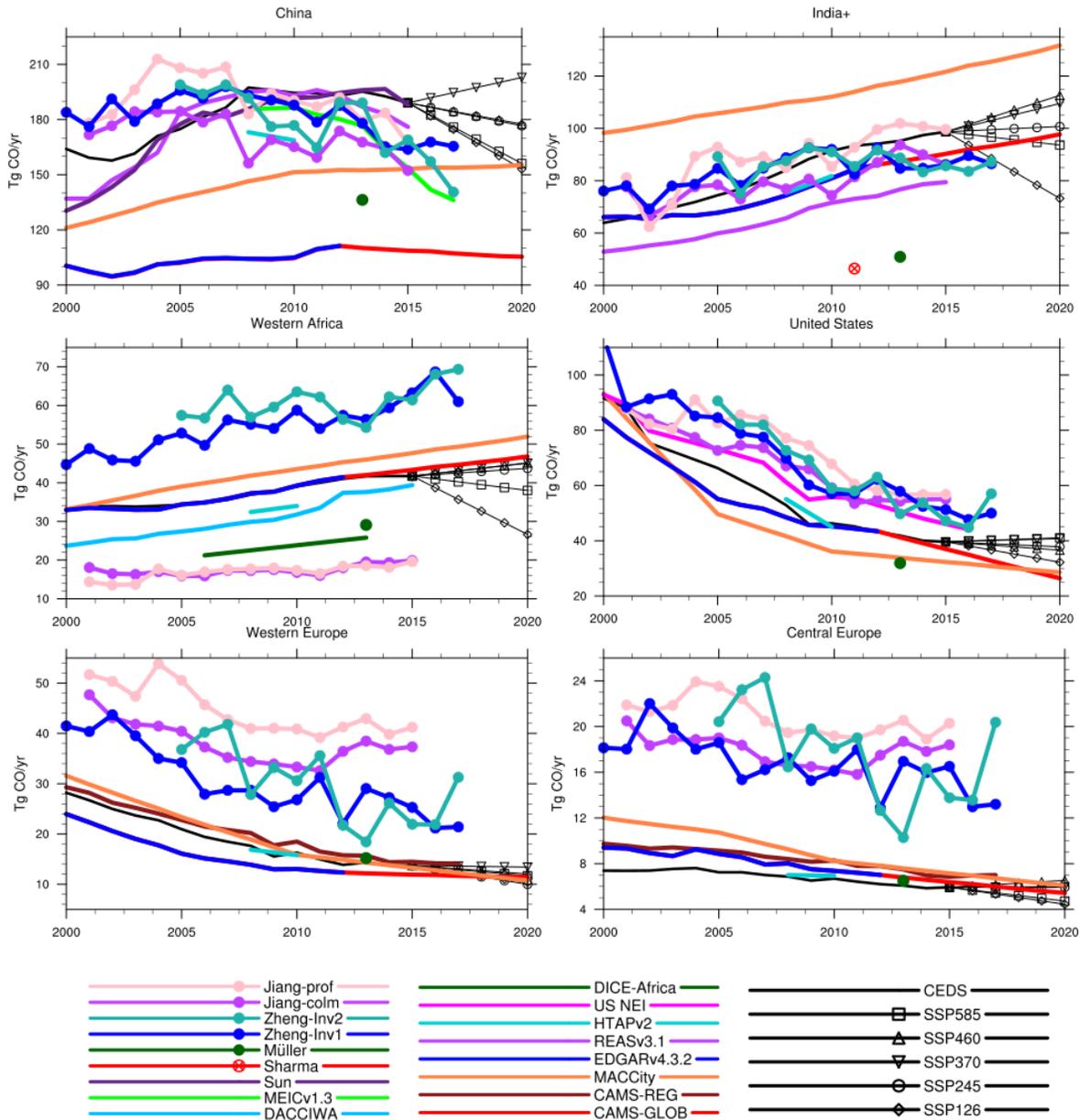
392  
393 In Europe, Zheng's CO inversion estimates match fairly well the general trends from the bottom-  
394 up inventories, which show a steady declining trend from the year 2000 that slows down after  
395 2009. Zheng's estimates show a lot of inter-annual variability, while Jiang's estimates show  
396 higher emissions in 2015 than in 2009, indicating a slightly positive recent trend that is contrary  
397 to the inventories. In terms of magnitude of emissions, all of the inversion estimates are  
398 considerably higher, except for that of Müller which is very close to the CAMS-REG-AP regional  
399 inventory as discussed above (Figure 2). Zheng et al. (2019) suggest that the bottom-up  
400 inventories underestimate emissions in these regions, and show that their model biases are  
401 reduced compared to independent ground-based CO measurement when the posterior emissions  
402 are used. However, given that Europe is a region where there is a relatively a high level of  
403 detailed emission information, the higher inverse estimates of Zheng and Jiang may be due to  
404 higher OH levels in their model, since the estimates by Müller et al. (2018), using

405 methylchloroform-constrained OH abundances, provide an excellent agreement with the bottom-  
406 up inventories.

407  
408 The largest uncertainty among the inversion estimates is found in Western Africa, where the CO  
409 emissions range from 17 Tg (Jiang-colm) to 64 Tg (Zheng-Inv2) in 2010 (Figure 2). However,  
410 Jiang's estimates are likely biased towards the low-end because they are influenced by the low a  
411 priori emissions used in their inversion method. The scheme used to partition emissions into  
412 separate categories (e.g. anthropogenic, biomass burning, etc.) is different in each inversion  
413 system (see the Supplement and references therein) and based on information from the a priori  
414 emissions. Therefore, this can account for small differences between inversion estimates,  
415 particularly in regions such as Western Africa where there is a large contribution from non-  
416 anthropogenic sources, such as biomass burning. In addition, the data used to construct bottom-  
417 up inventories (e.g. emission factors, activity data, etc) in Africa also have quite large  
418 uncertainties. The DACCIWA and DICE-Africa regional inventories both show similar  
419 increasing trends in CO emissions. The DACCIWA estimate is about 33% higher than DICE-  
420 Africa in 2010, however, there are some sectors not included in DICE-Africa (e.g. on-grid  
421 energy and formal industry, see the Supplement for details) which accounts, at least in part, for  
422 the discrepancy.

423  
424 In India, the inversion estimates are largely within the range of the global inventories and the  
425 regional inventory REASv3.1, however, they are almost twice as high as Sharma's estimate. The  
426 trends are similar to the inventories until about 2012, after which the inversion estimates begin to  
427 stabilize (Figure 2). Because India is one of the regions in which updated regional emission data  
428 is not easily available, it is difficult to evaluate the trend in CO emissions seen in the estimates.  
429 Although not shown here, Sharma projects a value of 53.9 Tg for the year 2021, indicating an  
430 increasing trend in CO emissions in India.

431



432

433 Figure 2. CO emissions from different inventories, inversion estimates and the SSPs. Note that  
 434 the Sharma estimate is for the national boundaries of India only.

435

436

437 **4.2 NO<sub>x</sub>**

438 Figure 3 displays annually averaged NO<sub>x</sub> emissions for each region. Inversion estimates are  
 439 shown for TCR-2 in all regions and Qu in China and India. Two estimates for Qu are both  
 440 derived using the same hybrid 4D-Var / mass balance inversion method. The ‘single’ estimates  
 441 are constrained by only NO<sub>2</sub> observations and the ‘joint’ estimates are constrained by NO<sub>2</sub> and  
 442 SO<sub>2</sub>, as described in section 2. We recall from Section 2 that Qu’s joint and single estimates are

443 for India only, while the TCR-2 and inventory emission averages are for India+ (Figure 1 in the  
444 Supplement) which is a larger region encompassing neighboring countries such as Pakistan,  
445 Nepal and Bangladesh. Qu's estimates would likely be 15-25% higher if they included India+.  
446 Although there is a slight difference in the magnitude between the 'single' and 'joint' species  
447 estimates, the results in China and India indicate that this difference is minor compared to the  
448 spread in the top-down emission estimates, at least at the regional scale. Qu et al. (2019a) report  
449 large differences at grid-scale, but not in the national budget between the 'single' and 'joint'  
450 inversions.

451  
452 In terms of trends in NO<sub>x</sub> emissions, the degree of agreement between the inversion estimates  
453 and the inventories varies regionally (Figure 3). In China, the trends are consistent with the  
454 regional inventories (MEICv1.3, REASv3.1 and Sun) and the CEDS global inventory, all of  
455 which show a broad peak around 2011-12, followed by a sharp decline. This declining trend is  
456 not found in the other global among the bottom-up inventories than the top-down estimates. For  
457 example, the Qu-single estimate for 2012 is the lowest of the inversion estimates (~17 Tg NO),  
458 while the TCR-2 estimate is about 25% higher (~21 Tg NO). Among the bottom-up inventories,  
459 CEDS is the highest (~23 Tg NO), about 45% higher than the lowest value given by Sun et al.  
460 (2018) (~16 Tg NO<sub>x</sub>). In China, the range of uncertainty is about 20% greater in the bottom-up  
461 emissions than the top-down emissions.

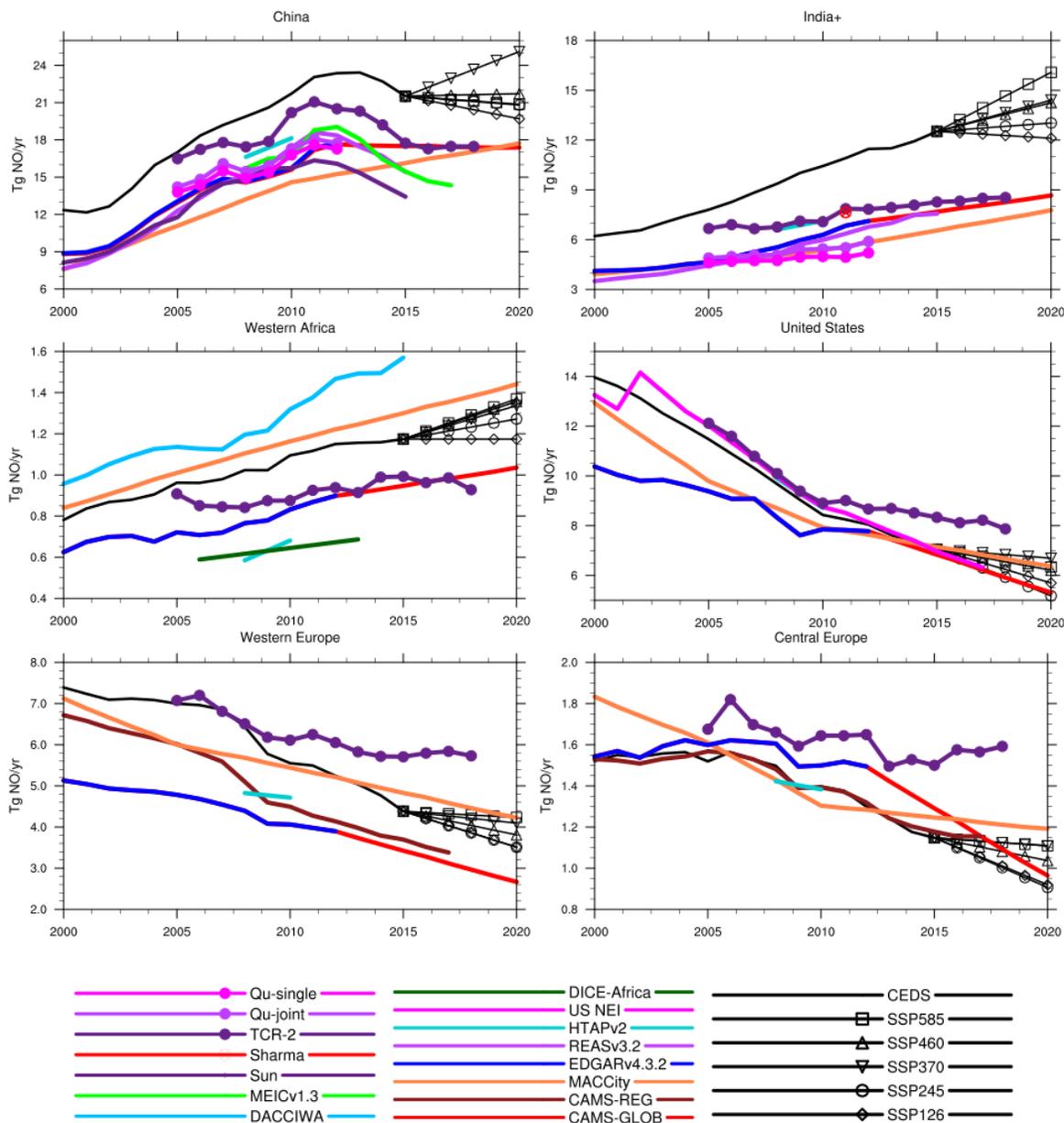
462  
463 In the United States and Europe, the TCR-2 inversion estimates of NO<sub>x</sub> emissions generally  
464 follow the decreasing trends shown in the bottom-up inventories, except in recent years where  
465 TCR-2 shows a tendency towards a stabilization of emissions, while the inventories indicate a  
466 continuing significant decline (Figure 3). In the U.S. between 2010 and 2016, the TCR-2 NO<sub>x</sub>  
467 emissions decreased by about 3.8% per year on average, while the U.S. regional inventory shows  
468 a continuously strong reduction of almost 5.5% per year. Jiang et al. (2018) also report a  
469 slowdown of the decreasing trend in CO and NO<sub>x</sub> top-down emissions in the U.S. inferred by  
470 satellite NO<sub>2</sub> data and suggest that the method used to calculate the NEI U.S. emissions leads to  
471 an overestimation of the reduction of emissions after 2010, whereas an alternative fuel-based  
472 method could produce trends more consistent with the top-down estimates. Furthermore, the  
473 top-down emissions were found to be more consistent with observed ozone concentrations than

474 the U.S. inventory (He et al., 2019). In contrast, Silvern et al. (2019) assert that the continued  
475 decreasing trend in the U.S. national inventory is consistent with the trends in NO<sub>2</sub> surface  
476 observations, and that the stabilization in the OMI NO<sub>2</sub> trend, and the TCR-2 NO<sub>x</sub> estimates,  
477 could be attributed to an underestimate of free tropospheric NO<sub>2</sub> background in the models. They  
478 suggest that increases in the relative contribution of non-anthropogenic background sources of  
479 NO<sub>x</sub> (e.g. lightning and soils) explains the discrepancy between trends in OMI NO<sub>2</sub> and the NEI  
480 inventory. Furthermore, Li and Wang (2019) report that in rural areas, strong nonlinear  
481 relationships exist between anthropogenic NO<sub>x</sub> emissions and satellite NO<sub>2</sub> columns, implying  
482 that constraining NO<sub>x</sub> emissions by only NO<sub>2</sub> satellite retrievals might lead to inaccurate  
483 estimates. Laughner and Cohen (2019) reported changes in NO<sub>x</sub> lifetime for many U.S. cities  
484 and suggested that accounting for this change in lifetime more accurately (for example, by using  
485 models with higher resolution) should help to improve emission trend estimates. Clearly, a better  
486 understanding of the relationship between anthropogenic NO<sub>x</sub> emissions and NO<sub>2</sub> satellite  
487 retrievals, as well as uncertainties in bottom-up methods used to estimate NO<sub>x</sub> emissions, is  
488 urgently needed.

489

490 In India and West Africa, the inversion estimates fall within the range of the inventories, but they  
491 show a weaker increasing trend than the inventories (Figure 3). For example, in Western Africa  
492 for the period 2005-2015, the TCR-2 estimates increase by slightly less than 1.5% yr<sup>-1</sup>, while the  
493 regional inventory DACCIWA increases by more than 3.5% yr<sup>-1</sup> and DICE-Africa increases by  
494 more than 2% yr<sup>-1</sup> (2006-2013). Similarly, in India during the period 2010-2015 NO<sub>x</sub> emissions  
495 given by REASv3.1 increase by almost 5% yr<sup>-1</sup>, while the TCR-2 estimates increase at a slower  
496 rate of about 2.5% yr<sup>-1</sup>. In 2011, the TCR-2 estimates 7.9 Tg NO<sub>x</sub> which is in very good  
497 agreement with Sharma's value of 7.6 Tg NO<sub>x</sub>.

498



499

500 Figure 3. NO<sub>x</sub> emissions from different inventories, inversion estimates and the SSPs. Note that  
 501 the Sharma estimate is for the national boundaries of India only.

502

### 503 4.3 NMVOCs, SO<sub>2</sub> and Carbonaceous Aerosols

504 Fewer inventories and inverse estimates are available for these species compared to CO and  
 505 NO<sub>x</sub>. Figure 4a,b shows the annually averaged NMVOC inversion emission estimates in China  
 506 from Stavrakou and Cao, along with the global and regional inventories. Note that Cao's  
 507 estimate in Figure 3 represents an average of four inversion experiments, ranging from 16.4 to  
 508 23.6 Tg NMVOC, using different combinations of satellite observations as described in Cao et

509 al. (2018). The significant spread reflects uncertainty in the satellite data and its impact on top-  
510 down NMVOC estimates. Both the inversion estimates (Stavrakou and Cao) and the inventories  
511 lie within a narrow range (Figures 2 and 3), although this does not necessarily indicate less  
512 uncertainty. Cao's estimate for 2007 (20.2 Tg NMVOC) is slightly below the range of the  
513 inventories (21.0-27.5 Tg NMVOC) and is quite close to the REASv3.1 estimate for 2013 (21.0  
514 Tg NMVOC). Stavrakou's estimates are similar in magnitude to the inventories and exhibit a  
515 stronger interannual variability. The inventories show an increasing trend in NMVOC emissions  
516 in China until ca. 2012, after which they start to level out. Over 2005-2014, Stavrakou's  
517 estimates indicate a weak positive trend of ca.  $1.4\% \text{ yr}^{-1}$ , which is lower than those of REASv3.1  
518 ( $5.2\% \text{ yr}^{-1}$  over 2005-2014), Li et al. ( $3.2\% \text{ yr}^{-1}$ ) and MEICv1.3 ( $3.4\% \text{ yr}^{-1}$  over 2008-2014). The  
519 top-down emission trend is primarily driven by the tropospheric HCHO column trend,  
520 amounting to  $1.3\% \text{ yr}^{-1}$  over Northern China between 2005 and 2014 (Stavrakou et al., 2017).  
521 Note that Shen et al. (2019) recently reported comparable summertime HCHO trends over  
522 Northern China over 2005-2016, and they found a good consistency with corresponding HCHO  
523 trends calculated by the GEOS-Chem CTM driven by MEICv1.3 emissions. The reasons for the  
524 apparent discrepancy with Stavrakou's results are unclear but might be related to intermodel  
525 differences.

526  
527 Annually averaged  $\text{SO}_2$  emissions for China and India are illustrated in Figure 4c,d with  
528 inversion estimates from Qu and TCR-2 for both regions. The two estimates from Qu are derived  
529 using identical methods, except the emissions are constrained using different OMI satellite  
530 retrievals (Qu-BIRA and Qu-NASA, Qu et al., 2019). The impact of the two different retrievals  
531 on the magnitude of emissions is more significant in China where the average difference in  
532 annual  $\text{SO}_2$  emission estimates is about 20%. In India, this difference is around 6%, a  
533 consequence of the fact that differences in the satellite retrievals are not the same in all regions.  
534 Both estimates indicate relatively similar trends in China and India between the two OMI  $\text{SO}_2$   
535 products. A third estimate of  $\text{SO}_2$  emission from Qu (Qu-joint) was derived using satellite  $\text{SO}_2$   
536 and  $\text{NO}_2$  constraints simultaneously, as described in section 2. We recall that the Qu-joint  
537 inversion estimates is for India only, unlike the other inversion and bottom-up estimates which  
538 are for India+. The Qu-joint estimate would be roughly 25% higher if it included India+.

539

540 In China, the range of SO<sub>2</sub> emissions is large in both the inversion estimates and the bottom-up  
541 inventories (Figure 4c,d). For example, in 2010 the inversion estimates range from  
542 approximately 17 (TCR-2) to 24 (Qu-BIRA) Tg SO<sub>2</sub> and the inventories range from about 16  
543 (Sun) to 28 (MEICv1.3) Tg SO<sub>2</sub>. While the inversion estimates show a significant decrease in  
544 Chinese emissions after 2010, this decline is not as steep as in the regional REASv3.1 and  
545 MEICv1.3 inventories in which SO<sub>2</sub> emissions decreased by about 35 and 40% between 2011  
546 and 2015, respectively. In general, the inversion estimates are significantly lower than in the  
547 regional inventories until after 2014.

548

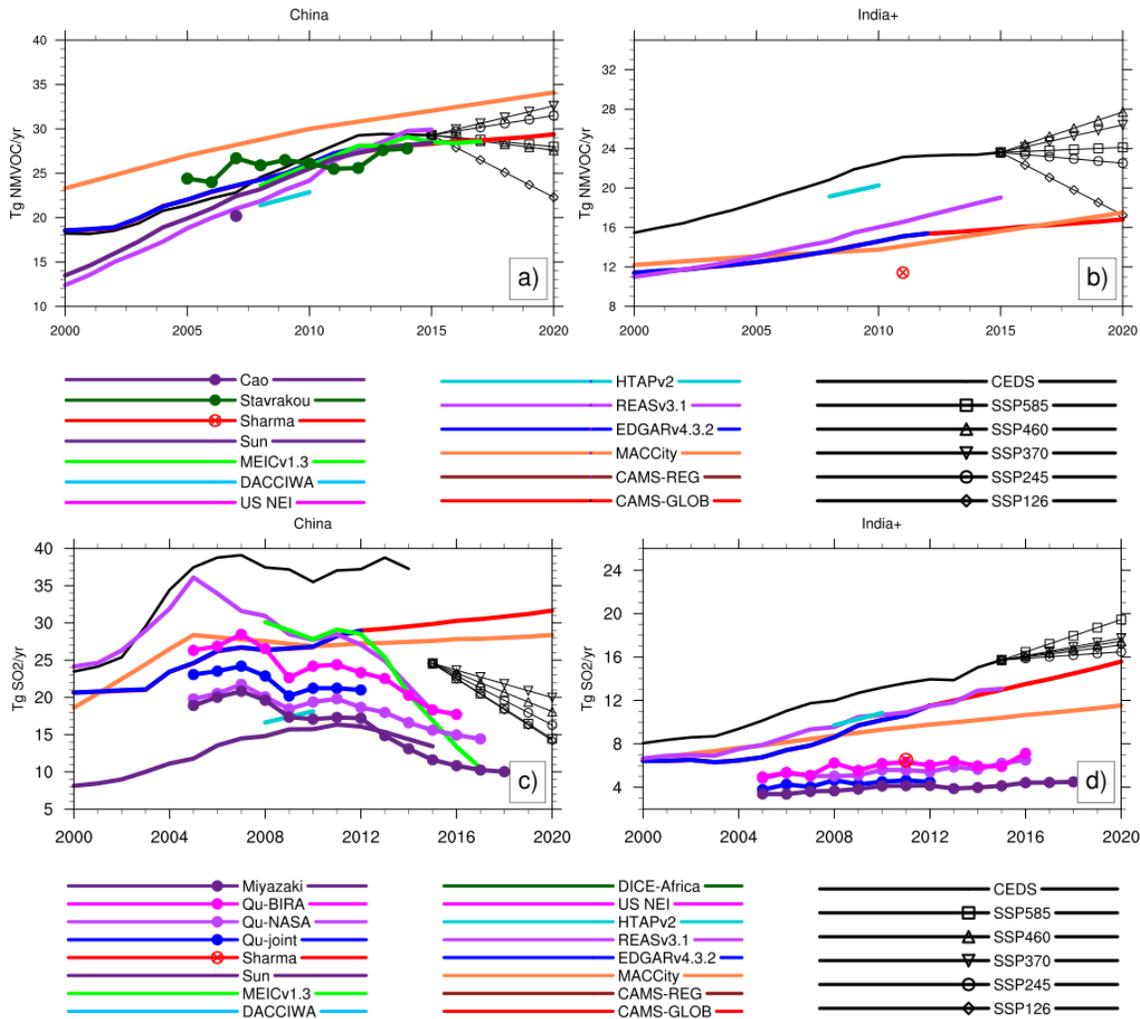
549 In India, SO<sub>2</sub> emissions derived from the inversion systems are all considerably lower than what  
550 is reported in most inventories (Figure 4c,d). The exception is Sharma's bottom-up estimate for  
551 2011 of 6.5 Tg SO<sub>2</sub> which is considerably lower than the other inventories and closer to the  
552 inversion estimates. Even if we add 25% to Sharma's estimates for India to account for the other  
553 countries included in the India+ region, it is still the less than the other bottom-up estimates. The  
554 inventories all indicate that emissions have been steadily increasing in India since about 2005,  
555 however, this trend is weaker in the inversion estimates. The weaker trends in the inversion  
556 estimates could in part be due to satellite sensors which are not sensitive enough to pick some of  
557 the diffuse sources. Posterior SO<sub>2</sub> emissions in less polluted areas are harder to estimate due to  
558 the large amount of negative retrievals (Qu et al., 2019).

559

560

561

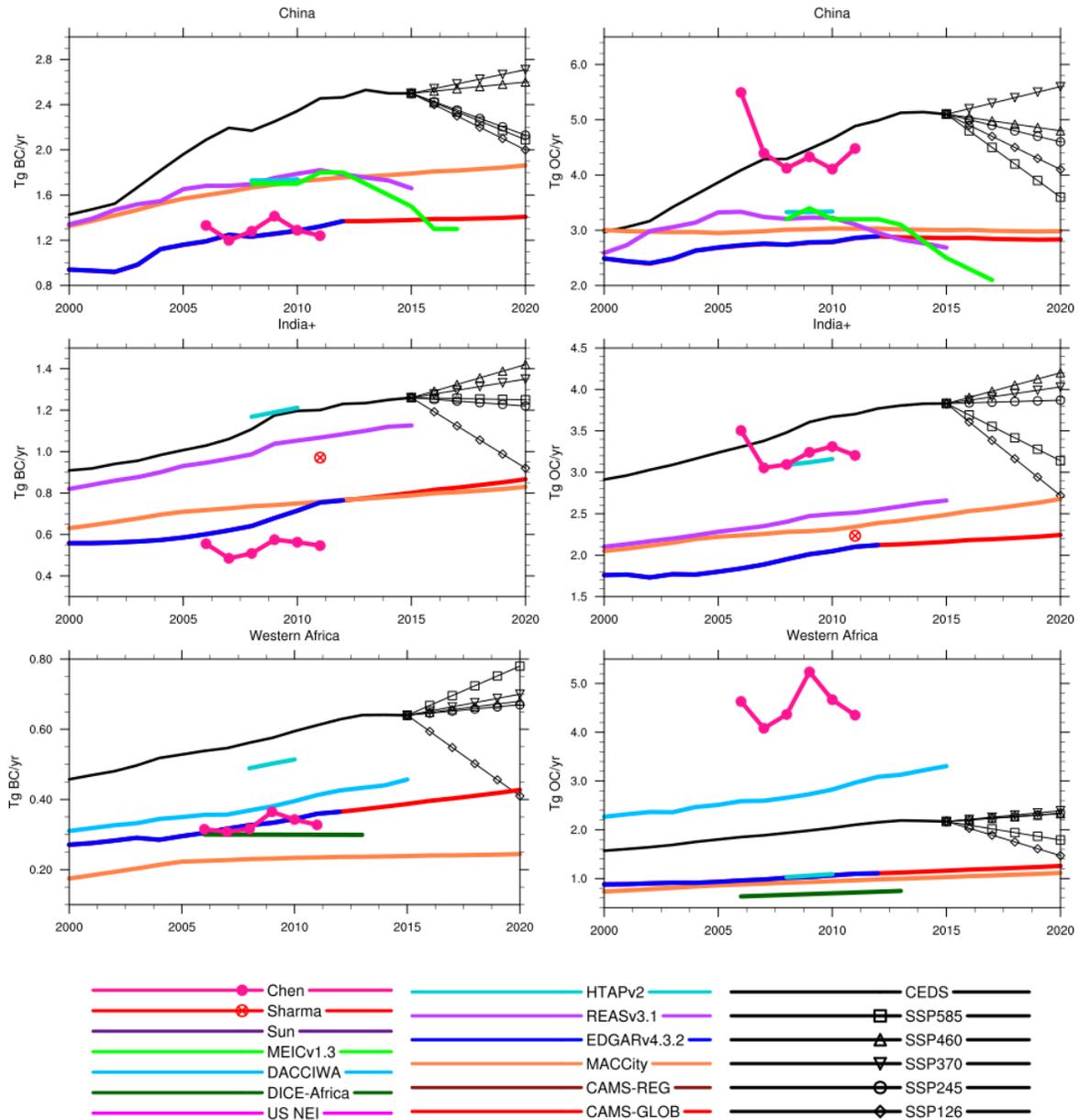
562



563  
 564 Figure 4. NMVOC (a, b) and SO<sub>2</sub> (c,d) emissions from different inventories, inversion estimates  
 565 and the SSPs. Note that the Qu-joint and Sharma estimates are for the national boundaries of  
 566 India only.

567  
 568 Figure 5 shows annually averaged black and organic carbon emissions in China, India and  
 569 Western Africa. Chen's inverse emission estimates provide a unique global time-series for the  
 570 period 2006-2011, as a first attempt in using an inversion approach to indirectly derive black and  
 571 organic carbon emissions by fitting satellite measurements of spectral aerosol extinction (AOD)  
 572 and absorption (AAOD) (Chen et al., 2019). The inversion estimates of BC emissions are within  
 573 the range of the bottom-up inventories in China and Western Africa, and slightly below in India  
 574 (Figure 5). In contrast, the inversion estimates of anthropogenic OC emissions are considerably  
 575 higher than most inventories in China and Western Africa, and at the upper end of the range in

576 India. It is interesting to note that Chen's estimates and the DACCIWA regional inventory show  
577 considerably higher OC emissions than the global inventories and the regional inventory DICE-  
578 Africa, while BC emissions are quite similar. This could suggest an underestimation in OC  
579 emissions in the global inventories in these regions. Other studies have identified inaccuracies in  
580 global inventories in West Africa where there is much uncertainty related to the emission factors  
581 and activity data from region-specific emissions sources which are unaccounted for (Lioussé et  
582 al., 2014; Marais et al., 2016). For example, wood burning for cooking and heating is a large  
583 source of OC and BC emissions in Africa but is likely misrepresented in inventories due to lack  
584 of representative data. Inversion estimates could potentially be used to improve inventory  
585 emissions from these and other misrepresented sources. However, uncertainties in the inversion  
586 estimation of black and organic carbon by fitting indirectly measurements of AOD/AAOD need  
587 to also be better quantified. Both the BC and OC inversion estimates show a stronger interannual  
588 variability than the bottom-up inventories, but in most cases the general trends agree reasonably  
589 well.



590

591 Figure 5. BC and OC emissions from different inventories, inversion estimates and the SSPs.  
 592 Note that the Sharma estimate is for the national boundaries of India only.

593

594

595 **5 Comparison of recent emission trends with SSPs for different regions**

596 In this section, we compare recent trends from available inventories and inversion estimates to  
 597 the five SSP scenarios under consideration. Linear trends of NO<sub>x</sub> and CO emissions from the  
 598 SSPs for the years 2010-2020, and the average of the HTAP trends from the bottom-up inventories and  
 599 inverse modelling estimates (INV) which have data until at least 2015, are summarized in Table  
 600 3. All trends are expressed in % yr<sup>-1</sup> relative to the mean over the data period. Individual trends

601 and their statistical significance is given in supplementary material. Note that for CO INV trends,  
602 we include only the Zheng et al. (2019) estimates from the inversion using the most recent  
603 satellites constraints for CO, HCHO and CH<sub>4</sub> (inversion 3, cf. Section 2), and the Jiang-prof  
604 estimates which assimilates profile data. For each region, the SSP trends which differ by less  
605 than 1% compared to the INV trend are in bold, indicating the scenarios with recent trends that  
606 are closest to those of the inventories and inverse modelling emissions. This type of analysis is to  
607 a certain extent speculative, therefore the results only indicate suggestions of the scenarios which  
608 match most closely the estimations of actual recent trends. Furthermore, recent trends in the last  
609 10 years are not necessarily indicative of how emissions will evolve in the future. Recall that  
610 SSP1 and SSP5 are both strong pollution control scenarios; SSP5 assumes an energy-intensive,  
611 fossil fuel based economy, while SSP1 represents a future shifted towards more sustainable  
612 practices. SSP3 and SSP4 represent futures with more pessimistic development trends, where  
613 there is little investment in health and education and fast growing populations. In SSP3, national  
614 and regional security is prioritized, whereas in SSP4 large inequalities dominate both within and  
615 across countries. SSP2 is the moderate pollution control scenario, in which trends more or less  
616 follow their historical patterns.

617  
618 In China, SSP126 is the scenario which matches most closely to the actual recent trends in CO  
619 emissions. This is a strong pollution control scenario that reflects the stringent air quality  
620 standards China has enforced in recent years. All of the SSP NO<sub>x</sub> trends agree with the declining  
621 emissions reported by the inventories (REASv3.1, Sun, MEICv1.3) and inverse modelling  
622 estimates (Qu-joint, TCR-2 and DECISOv5.1qa), but for the period considered, 2010-2020, the  
623 decrease is not as strong. However, considering only the last few years since 2015, we can see in  
624 Figure 2 that the MEICv1.3 regional inventory and the inverse modelling estimates from  
625 DECISOv5.1qa both show a decline in the decreasing trend of NO<sub>x</sub> emissions, which is  
626 qualitatively similar to that of SSP126.

627  
628 The results indicate that recent emission trends in the United States are also similar to SSP126  
629 (Table 3). The INV trends, which are based on the NEI regional inventory and inverse modelling  
630 estimates, show recent declines in NO<sub>x</sub> and CO emissions on the order of 3% yr<sup>-1</sup>, consistent  
631 with the strong pollution control scenario. However, it should also be noted that the CO

632 inversion estimates from Jiang et al. (2017) which show a slowdown in declining CO emissions  
633 in the US are not included in the INV average because there is not a statistically significant trend.  
634 If the Jiang et al. (2017) inversion estimates are included, the INV trend would show less of a  
635 negative trend in CO emissions more similar to that of SSP585. In Western and Central Europe,  
636 as in the U.S., there is a smaller spread in emissions among the SSPs, especially during the early  
637 century, therefore the trends are more similar and indicate the same sign. Trends from SSP126  
638 and SSP585 match most closely to the INV trends for Europe, both of which are strong pollution  
639 control scenarios. However, it should be noted that the inversion estimates of CO (Jiang et al.  
640 (2017) and Zheng et al.(2019)) and NO<sub>x</sub> (TCR-2) emissions do not show a statistically  
641 significant trend in recent years in Europe, and therefore are not included in the INV trend  
642 average which is based largely on the CAMS-REG-AP regional inventory (Figures 1 and 2). In  
643 Southeast Asia, the strong pollution control scenario SSP126 also matches the INV trend most  
644 closely, which indicates a positive trend in NO<sub>x</sub> and declining trend in CO emissions.

645  
646 Our results, based on the trends given in Table 3 as well as a qualitative visual interpretation of  
647 Figures 1 and 2, indicate that the most likely scenario representing recent emission trends in  
648 India is SSP370 which is a low pollution control scenario. In Western Africa and Indonesia+,  
649 emission trends from SSP460 match most closely to the recent INV trends in NO<sub>x</sub> and CO  
650 emissions. SSP4, also a low pollution control scenario, represents a scenario with a fast-growing  
651 population with increasing inequalities, leading to societies that are highly vulnerable to climate  
652 change. Indeed, air quality is deteriorating in many West African countries due to rapid increases  
653 in population, economic growth and coinciding lack of regulations, especially in megacities  
654 (Lioussé et al. 2014). For some of the regions the results are ambiguous, either because none of  
655 the SSP emission trends are close to the actual trends (i.e. Southern Africa), or because there is  
656 not a statistically significant INV trend (i.e. Middle East and South America). In general, and  
657 especially in regions where there is a lack of reliable emission data, the results should be  
658 interpreted with caution in terms of robustness.

659

660

661

662

663  
664  
665  
666  
667  
668  
669

**Table 3:** Linear trends of NO<sub>x</sub> and CO emissions from SSPs for the years 2010-2020, and the averaged trend from all regional inventories and inversion estimates (INV) which have data until at least 2015 and are statistically significant ( $p > 0.05$ ). Trends are expressed in units of % yr<sup>-1</sup> relative to the mean over the data period. The number of available datasets with statistically significant trends used to calculate the average is indicated in parenthesis. SSP trends shown in bold correspond most closely to recent observed trends for NO<sub>x</sub> and CO.

		INV	SSP126	SSP245	SSP370	SSP460	SSP585	closest scenario
China	NO <sub>x</sub>	<b>-3.07 (4)</b>	-1.59	-1.06	0.86	-0.64	-1.03	SSP126
	CO	<b>-2.94 (3)</b>	<b>-2.36</b>	-1.03	0.36	-0.98	<b>-2.19</b>	
Middle East	NO <sub>x</sub>	*	0.02	0.44	1.29	-1.39	1.01	----
	CO	*	-2.12	-3.31	-0.12	-1.86	0.02	
Western Africa	NO <sub>x</sub>	<b>+2.03 (2)</b>	0.59	1.40	<b>1.92</b>	<b>2.08</b>	2.18	SSP460
	CO	<b>+3.35 (2)</b>	-3.52	0.97	1.27	1.30	-0.41	
India	NO <sub>x</sub>	<b>+3.33 (2)</b>	1.39	2.13	<b>3.17</b>	<b>3.05</b>	4.35	SSP370
	CO	<b>+1.32 (3)</b>	-2.03	<b>0.95</b>	<b>1.81</b>	2.10	0.23	
United States	NO <sub>x</sub>	<b>-3.16 (2)</b>	-3.98	-4.80	-2.47	<b>-3.19</b>	<b>-2.99</b>	SSP126
	CO	<b>-3.70 (3)</b>	<b>-3.54</b>	-2.02	-1.28	-2.34	-1.21	
Western Europe	NO <sub>x</sub>	<b>-2.47 (2)</b>	-4.86	-4.86	<b>-3.44</b>	-4.12	<b>-3.12</b>	SSP126
	CO	<b>-4.21 (1)</b>	<b>-3.87</b>	<b>-4.19</b>	-1.48	-2.50	-2.80	
Central Europe	NO <sub>x</sub>	<b>-3.44 (1)</b>	<b>-4.21</b>	<b>-4.34</b>	-2.46	-3.11	-2.46	SSP126
	CO	<b>-3.20 (1)</b>	-3.72	-1.01	-0.93	-0.03	<b>-3.18</b>	
South America	NO <sub>x</sub>	*	-0.07	-0.41	1.12	-0.52	0.94	---
	CO	*	-2.56	-2.85	-0.47	-1.31	-1.74	
Southeast Asia	NO <sub>x</sub>	<b>+2.07 (2)</b>	<b>1.84</b>	0.45	2.07	2.00	<b>1.35</b>	SSP126
	CO	<b>-2.72 (2)</b>	<b>-3.02</b>	-0.97	1.05	0.84	-0.08	
Indonesia	NO <sub>x</sub>	<b>+1.43 (2)</b>	0.61	0.08	2.01	<b>1.44</b>	<b>1.41</b>	SSP460
	CO	<b>+2.86 (2)</b>	-3.15	-1.83	0.50	0.65	-0.63	
Oceania	NO <sub>x</sub>	*	-0.75	-0.38	0.36	-0.32	0.17	---
	CO	<b>-5.90 (1)</b>	-3.17	-1.07	-1.21	-2.08	-0.91	
Southern Africa	NO <sub>x</sub>	<b>+2.91 (1)</b>	-0.19	0.21	0.76	0.32	0.32	---
	CO	<b>+3.04 (2)</b>	-2.53	0.80	1.11	1.28	-0.37	

670  
671  
672

673

674

## 675 **6 Conclusions**

676 In this study, we have presented a comprehensive overview and comparison of current state-of-  
677 the-art top-down and bottom-up emission estimates of CO, NO<sub>x</sub>, NMVOC, SO<sub>2</sub>, BC and OC for  
678 several world regions. The results show that the top-down estimates are generally within the  
679 range of bottom-up emission inventories and exhibit a similar level of uncertainty, or even less in  
680 certain regions such as China. In general, for all species the largest discrepancies are found  
681 outside of regions such as the U.S., Europe and Japan where the most accurate and detailed  
682 information on emissions (e.g. activity data, emission factors) is available. In terms of absolute  
683 magnitude, the largest spread in CO and NO<sub>x</sub> inventory emissions is found in China where in  
684 2010 differences are approximately 90 (~60%) and 6 (~33%) Tg, respectively. Significant  
685 differences are also seen for other compounds in China: ~20 Tg (~80%) for SO<sub>2</sub>, 2 Tg (~55%)  
686 for OC, ~1 Tg (+55%) for BC. Contrarily, NMVOC inventory emissions are in better agreement  
687 in China, with a difference of ~4 Tg (~15%) in 2010, as compared to most other regions (e.g. ~8  
688 Tg (~90%) in Western Africa and ~8 Tg (~45%) in India+). This agreement might be  
689 coincidental, however, and does not necessary imply that the emissions are less uncertain. In  
690 terms of percentage differences, the variation is not region- or species-dependent, and in general  
691 ranges from about 15% up to about 100%. In part, this likely reflects differences in the  
692 inclusion/exclusion of specific sectors, as well as in the methodologies used to construct the  
693 bottom-up inventories. A more in-depth detailed analysis at the sectoral level would give insight  
694 into how much of the uncertainty is due to these factors, but would be beyond the scope of this  
695 study.

696

697 Top-down emission estimates offer great potential and clear advantages, however, future work  
698 aimed at identifying, quantifying and reducing the uncertainties is needed.

699 Constraining multiple species in inversion modelling methods can lead to better consistency in  
700 simulated atmospheric chemical processes and thus more accurate optimized emissions (e.g.,  
701 Zheng et al., 2019a; Qu et al., 2019a). Correctly modelling OH fields are also important due to  
702 its significant impact on oxidation processes (Müller et al., 2018; Jiang et al., 2011; Miyazaki et  
703 al., 2020b). More generally, the representation of chemical and transport processes in model

704 should be improved. In addition, the type of satellite data assimilated (e.g. profile vs total  
705 column) has a large impact on inversions (e.g., Jiang et al., 2017) and should be further explored.  
706 Finally, the satellite retrievals have important uncertainties with significant impacts on the  
707 emission inversions. Much can be learned in terms of quantifying these and other sources of  
708 uncertainty in inverse modelling estimates through more collaborated inter-comparison projects  
709 such as the Global Carbon Project ([www.globalcarbonproject.org](http://www.globalcarbonproject.org)) which targets CO<sub>2</sub> and CH<sub>4</sub>.  
710 A first step has been made with the IGAC AMIGO project (Analysis of eMIssions usinG  
711 Observations, <https://amigo.aeronomie.be/>) which brings together the international scientific  
712 community with the common goal of better quantifying emissions for a variety of trace gases and  
713 at different spatio-temporal scales. Inverse modelling has been identified as an integral part of  
714 AMIGO.

715

716 Top-down emissions offer great potential to supplement or improve bottom-up inventories,  
717 particularly in regions where global inventories often lack the necessary up-to-date and accurate  
718 information regarding regional activity data and emission factors. For example, China has  
719 undergone rapid economic growth in addition to stringent pollution control policies in recent  
720 years, both of which have led to rapidly changing activity data and emission factors (Zheng et  
721 al., 2018). This evolution has had a large impact on emissions, which were estimated to have  
722 been decreasing substantially in China during the last several years, except for NMVOC. The  
723 downward trend in China's emissions is well captured by the inversion estimates, as well as  
724 (Rao, et al., 2016) by the detailed regional inventories, but is not represented in any of the global  
725 inventories. This is a clear example of where inversion estimates provide useful constraints to the  
726 global bottom-up inventories, particularly in countries that are undergoing rapid changes in  
727 economy, technology, and environmental policies, such as India and Africa. Finally, since  
728 inversion estimates become available more quickly than bottom-up inventories, they can be used  
729 to extrapolate bottom-up inventories to the most recent years, which would benefit air quality  
730 forecasting. This also adds information that can be used to tune baseline emissions in recent  
731 years in the development of future emission scenarios such as the RCPs and SSPs.

732

733 This study has also compared recent emission trends in regional inventories and inversion  
734 estimates to those of five SSP near-future projections for several world regions. For each region,

735 we identified the scenarios for which the recent CO and NO<sub>x</sub> trends for the recent years (2010-  
736 2020) match most closely the best estimates based on bottom-up and top-down estimates. This  
737 type of analysis can be helpful in updating inventories for the most recent years and can serve as  
738 a guide in selecting CMIP6 climate change projections to be used for regional downscaling in air  
739 quality forecasting, and near-future pollution control/mitigation and climate impact studies. In  
740 addition, highlighting inconsistencies between the SSPs and actual emissions can help improve  
741 in the development of future emission scenarios. Not surprisingly for China, which has  
742 experienced drastic reductions in emissions due to the enforcement of stringent air quality  
743 policies, the trends from the strong pollution control scenarios, SSP1 and SPP5, are most  
744 representative of the actual recent trend. In India and Western Africa, regions of rapid population  
745 growth and significant increases in unregulated emissions, SSP3 and SSP4 which represent  
746 futures with more pessimistic development trends (e.g. little investment in health and education  
747 and fast growing populations) match most closely the actual recent trends.

748

## 749 **Acknowledgments**

750 This study has received support from the CAMS (Copernicus Atmospheric Monitoring System)  
751 project and contributes to the AMIGO (Analysis of eMIssions usinG Observations,  
752 <https://amigo.aeronomie.be/>) project of the International Global Atmospheric Chemistry (IGAC)  
753 project. This research (MB) has also been supported by the projects PRODEX TROVA (2016-  
754 2018) and TROVA-E2 (2019) of the European Space Agency (ESA) funded by the Belgian  
755 Science Policy Office, as well as by the SOLFEO project funded by ESA. We acknowledge the  
756 use of data products from the NASA AURA and EOS Terra satellite missions. We also  
757 acknowledge the free use of tropospheric NO<sub>2</sub> column data from the SCIAMACHY, GOME-2,  
758 and OMI sensors from [www.temis.nl](http://www.temis.nl). Part of this work was conducted at the Jet Propulsion  
759 Laboratory, California Institute of Technology, under contract with the National Aeronautics and  
760 | Space Administration (NASA).

761

## 762 **References**

- 763  
764 Arellano, A. F., Kasibhatla, P. S., Giglio, L., Van der Werf, G. R., & Randerson, J. T. (2004).  
765 Top-down estimates of global CO sources using MOPITT measurements. *Geophys. Res.*  
766 *Lett.*, *31*, doi:10.1029/2002GL015609.  
767 Bey, I., Jacob, D., Yantosca, R., Logan, J. A., Field, B., Fiore, A., . . . Schultz, M. (2001). Global  
768 modeling of tropospheric chemistry with assimilated meteorology: Model description and  
769 evaluation. *Journal of Geophysical Research*, *106*, 23073-23095.  
770 <https://doi.org/10.1029/2001JD000807>.

- 771 Bouwman, A., Kram, T., & Klein Goldewijk (eds), K. (2006). *Integrated modelling of global*  
772 *environmental change. An overview of IMAGE 2.4*. Netherlands Environmental  
773 Assessment Agency (MNP) publications.
- 774 Cao, H., Fu, T.-M., Zhang, L., Henze, D., Miller, C., Lerot, C., . . . Zhao, Y. (2018). Adjoint  
775 inversion of Chinese non-methane volatile organic compound emissions using space-  
776 based observations of formaldehyde and glyoxal. *Atmos. Chem. Phys.*, *18*, 15017–15046,  
777 <https://doi.org/10.5194/acp-18-15017-2018>.
- 778 Chen, C., Dubovik, O., Henze, D. K., Chin, M., Lapyonok, T., Schuster, G. L., . . . Torres, B.  
779 (2019). Constraining global aerosol emissions using POLDAR/PARASOL satellite  
780 remote sensing observations. *Atmos. Chem. Phys.*, 14585-14606, doi:10.5194/acp-19-  
781 14585-2019.
- 782 Chen, C., Dubovik, O., Henze, D. K., Lapyonak, T., Chin, M., Ducos, F., . . . Li, L. (2018).  
783 Retrieval of desert dust and carbonaceous aerosol emissions over Africa from  
784 POLDER/PARASOL products generated by the GRASP algorithm. *Atmos. Chem. Phys.*,  
785 12551-12580, doi:10.5194/acp-18-12551-2018.
- 786 Crippa, M., Guizzardi, D., Muntean, M., Schaaf, E., Dentener, F., Aardenne, J. A., . . . Janssens-  
787 Maenhout, G. (2018). Gridded emissions of air pollutants for the period 1970-2012 with  
788 EDGARv4.3.2. *Earth Syst. Sci. Data*, *10*, 1987-2013, [https://doi.org/10.5194/essd-10-](https://doi.org/10.5194/essd-10-1987-2018)  
789 [1987-2018](https://doi.org/10.5194/essd-10-1987-2018).
- 790 Ding, J., A. R. v., Mijling, B., & Levelt, P. (2017). Space-based NO<sub>x</sub> emission estimates over  
791 remote regions improved in DECSO. *10*, 925-938.
- 792 Dubovik, O., Herman, M., Holdak, A., Lapyonok, T., Tanré, D., Deuzé, J. L., . . . Lopatin, A.  
793 (2011). Statically optimized inversion algorithm for enhanced retrieval of aerosol  
794 properties from spectral and multi-angle polarimetric satellite observations. *Atmos. Meas.*  
795 *Tech.*, *4*, <https://doi.org/10.5194/amt-4-975-2011>.
- 796 Dubovik, O., Lapyonok, T., Litvinov, P., Herman, M., Fuertes, D., Ducos, F., . . . Federspiel, C.  
797 (2014). GRASP: a versatile algorithm for characterizing the atmosphere. *SPIE*  
798 *Newsroom*, <https://doi.org/10.1117/2.1201408.005558>.
- 799 Elbern, H., Schmidt, H., & Ebel, A. (1997). Variational data assimilation for tropospheric  
800 chemistry modeling. *J. Geophys. Res. Atmos.*, *102*, doi:10.1029/97JD01213.
- 801 Eyring, V., Bony, S., Meehl, G., Senior, C., Stevens, B., Stouffer, R., & Taylor, K. (2016).  
802 Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental  
803 design and organization. *Geoscientific Model Development*, *9*,  
804 <https://doi.org/10.5194/gmd-9-1937-1026>.
- 805 Gidden, M. J., Riahi, K., Smith, S. J., Fujimori, S., Luderer, G., Kriegler, E., . . . Fricko, O.  
806 (2019). Global emissions pathways under different socioeconomic scenarios for use in  
807 CMIP6: a dataset of harmonized emissions trajectories through the end of the century.  
808 *Geosci. Model Dev.*, 1443–1475, <https://doi.org/10.5194/gmd-12-1443-2019>.
- 809 Granier, C., Bessagnet, B., Bond, T., & al., e. (2011). Evolution of anthropogenic and biomass  
810 burning emissions of air pollutants at global and regional scales during the 1980-2010  
811 period. *Climate Change*, *109*, <https://doi.org/10.1007/s10584-011-0154-1>.
- 812 Granier, C., Darras, S., Gon, H. D., Doubalova, J., Elguindi, N., Galle, B., . . . Sindelarova, K.  
813 (2019). *The Copernicus Atmosphere Monitoring Service global and regional emissions*.  
814 Copernicus Atmosphere Monitoring Service.

- 815 He, T.-L., Jones, D. B., Huang, B., Liu, Y., Miyazaki, K., Jiang, Z., . . . Worden, J. R. (2019).  
816 Recurrent U-net: Deep learning to predict daily summertime ozone in the United States.  
817 arXiv:1908.05841.
- 818 Henze, D. K. (2009). Inverse modeling and mapping US air quality influences of inorganic  
819 PM<sub>2.5</sub> precursor emissions using the adjoint of GEOS-Chem. *Atmos. Chem. Phys.*, *9*,  
820 5877-5903.
- 821 Henze, D. K., Hakami, A., & Seinfeld, J. H. (2007). Development of the adjoint of GEOS-Chem.  
822 *Atmos. Chem. Phys.*, *7*, doi:10.5194/acp-7-2413-2007.
- 823 Hosley, R., Smith, S., Feng, L., Klimont, Z., Janssens-Maenhout, G., Pitkanen, T., . . . al., e.  
824 (2018). Historical (1750-2014) anthropogenic emissions of reactive gases and aerosols  
825 from the Community Emissions Data System (CEDS). *Geosci. Model Dev.*, *11*,  
826 <http://doi.org/10.5194/gmd-11-369-2018>.
- 827 HTAPv2. (???) HTAPv2. [www.htap.org](http://www.htap.org).
- 828 Jiang, Z. D. (2011). Quantifying the impact of model errors on top-down estimates of carbon  
829 monoxide emissions using satellite observations. *J. Geophys. Res.*, *116*,  
830 D15306, doi:10.1029/2010JD015282.
- 831 Jiang, Z., Jones, D., Worden, H., Deeter, M., Henze, D., Worden, J., . . . Schuck, T. (2013).  
832 Impact of model errors in convective transport on CO source estimates inferred from  
833 MOPITT CO retrievals. *J. Geophys. Res. Atmos.*, *118*, doi:10.1002/jgrd.50216.
- 834 Jiang, Z., Jones, D., Worden, J., Worden, H., Henze, D., & Wang, Y. (2015). Regional data  
835 assimilation of multi-spectral MOPITT observations of CO over North America. *Atmos.*  
836 *Chem. Phys.*, 6801-6814, doi:10.5194/acp-15-6801-2015.
- 837 Jiang, Z., Worden, J., Worden, H., M. Deeter, D. J., Arellano, A., & Henze, D. (2017). A 15-year  
838 record of CO emissions constrained by MOPITT CO observations. *Atmos. Chem. Phys.*,  
839 *17*, 4565-4583.
- 840 Keita. (2020). DACCIWA. *Journal of ???*
- 841 Kriegler, E., Edmonds, J., Hallegatte, S., K.L., E., Kram, T., Riahi, K., . . . Vuuren, D. v. (2014).  
842 A new scenario framework for climate change research: the concept of shared climate  
843 policy assumptions. *122*, 401-414.
- 844 Krotkov, N. A., Lamsal, L. N., Celarier, E., Swartz, W., Marchenko, S., Bucsela, E. J., & al, e.  
845 (2017). The version 3 OMI NO<sub>2</sub> standard product. *Atmospheric Measurement*  
846 *Techniques*, *10*, <https://doi.org/10.5194/amt-10-3133-2017>.
- 847 Kuenen, J., Visschedijk, A. J., Jozwicka, M., & Denier van der Gon, H. (2014). TNO-MACC\_II  
848 emission inventory: a multi-year (2003–2009) consistent high-resolution European  
849 emission inventory for air quality modelling. *Atmos. Chem. Phys.*, 10963-10976,  
850 doi:10.5194/acp-14-10963-2014.
- 851 Kurokawa, J., & Ohara, T. (2019). Long-term historical trends in air pollutant emissions in Asia:  
852 Regional Emission inventory in ASia (REAS) version 3.1. *Atmos. Chem. Phys. Discuss.*,  
853 <https://doi.org/10.5194/acp-2019-1122>.
- 854 Laughner, J., & Cohen, R. (2019). Direct observation of changing NO<sub>x</sub> lifetime in North  
855 American cities. *Science*, *366*, 723-727, doi: 10.1126/science.aax6832.
- 856 Li, C., Joiner, J., Krotkov, N. A., & Bhartia, P. (2013). A fast and sensitive new satellite SO<sub>2</sub>  
857 retrieval algorithm based on principal component analysis: Application to the Ozone  
858 Monitoring Instrument. *Geophysical Research Letters*, *40*,  
859 <https://doi.org/10.1002/2013GL058134>.

- 860 Li, J., & Y. Wang. (2019). Inferring the anthropogenic NO<sub>x</sub> emission trend over the United States  
861 during 2003–2017 from satellite observations: Was there a flattening of the emission  
862 trend after the Great Recession? *Atmos. Chem. Phys. Disc.*, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-2019-472)  
863 2019-472.
- 864 Li, M. Z. (2019). Persistent growth of anthropogenic non-methane volatile organic compound  
865 (NMVOC) emissions in China during 1990–2017: drivers, speciation and ozone  
866 formation potential. *Atmos. Chem. Phys.*, *19*, 8897–8913, [https://doi.org/10.5194/acp-19-](https://doi.org/10.5194/acp-19-8897-2019)  
867 8897-2019.
- 868 Li, M., Zhang, Q., Streets, D. G., He, K. B., Cheng, Y. F., Emmons, L. K., . . . Zhang, Y. (2014).  
869 Mapping Asian anthropogenic emissions of non-methane volatile organic compounds to  
870 multiple chemical mechanisms. *Atmos. Chem. Phys.*, *14*, 5617–2014.
- 871 Liousse, C., Assamoi, E., Criqui, P., Granier, C., & Rosset, R. (2014). Explosive Growth in  
872 African combustion emissions from 2005 to 2030. *Environmental Research Letters*, *9*,  
873 doi:10.1088/1748-9326/9/3/035003.
- 874 Liu, F. Z., Li, M., Huo, H., & He, K. B. (n.d.).
- 875 Liu, F., Zhang, Q., Tong, D., Zheng, B., Li, M., Huo, H., & He, K. B. (2015). High-resolution  
876 inventory of technologies, activities, and emissions of coal-fired power plants in China  
877 from 1990 to 2010. *Atmos. Chem. Phys.*, *15*, doi:10.5194/acp-15-13299-2015.
- 878 Müller, J.-F. T.-F. (2018). Top-Down CO Emissions Based on IASI Observations and  
879 Hemispheric Constraints on OH Levels. *Geophysical Research Letters*, *45*,  
880 <http://doi.org/10.1002/2017GL076697>.
- 881 Müller, J.-F., & Stavrakou, T. (2005). Inversion of CO and NO<sub>x</sub> emissions using the adjoint of  
882 the IMAGES model. *Atmos. Chem. Phys.*, *5*, <https://doi.org/10.5194/acp-5-1157-2005>.
- 883 MACCity. (n.d.). MACCity. <https://eccad3.sedoo.fr>.
- 884 Marais, E., & Wiedinmyer, C. (2016). Air Quality Impact of Diffuse and Inefficient Combustion  
885 Emissions in Africa (DICE-Africa). *Environmental Science and Technology*, *50*,  
886 doi:10.1021/acs.est.6b02602.
- 887 Martin, R. V., Jacob, D. J., Chance, K., Kurosu, T. P., Palmer, P. I., & Evans, M. J. (2003).  
888 Global inventory of nitrogen oxide emissions constrained by space-based observations of  
889 NO<sub>2</sub> columns. *J. Geophys. Res. Atmos.*, *108*, doi:10.1029/2003JD003453.
- 890 MEICv1.3. (n.d.). MEICv1.3. <http://www.meicmodel.org>.
- 891 Menut, L., Bessagnet, B., Khovorostyanov, D., Beekman, M., Blond, N., Colette, A., . . .  
892 Vivanco, M. (2013). CHIMERE 2013: a model for regional atmospheric composition  
893 modelling. *Geosci. Model Dev.*, *6*, doi:10.5194/gmd-6981-2013.
- 894 Mijling, B., & A. v. (2012). Using daily satellite observations to estimate emissions of short-  
895 lived air pollutants on a mesoscopic scale. *J. Geophys. Res. Atmos.*, *117*,  
896 doi:10.1029/2012JD017817.
- 897 Miyazaki, K. B. (2020). Evaluation of a multi-model, multi-constituent assimilation framework  
898 for tropospheric chemical reanalysis. *Atmos. Chem. Phys. Discuss.*,  
899 <https://doi.org/10.5194/acp-2019-645>, in press.
- 900 Miyazaki, K., Bowman, K., Sekiya, T., Eskes, H., Boersma, F., Worden, H., . . . Ogochi, K.  
901 (2020). An updated tropospheric chemistry reanalysis and emission estimates, TCR-2, for  
902 2005–2018. *Earth Syst. Sci. Data Discuss.*, submitted.
- 903 Miyazaki, K., Eskes, H., Sudo, K., Boersma, K. F., Bowman, K., & Kanaya, Y. (2017). Decadal  
904 changes in global surface NO<sub>x</sub> emissions from multi-constituent satellite data  
905 assimilation. *Atmos. Chem. Phys.*, *17*, doi:10.5194/acp-17-807-2017.

- 906 Miyazaki, K., Eskes, H., Sudo, K., Boersma, K. F., Bowman, K., & Kanaya, Y. (2019). Balance  
907 of emission and dynamical controls on ozone during the Korea;United States Air Quality  
908 campaign from multiconstituent satellite data assimilation. *Journal of Geophysical*  
909 *Research: Atmospheres*, *124*, 387-413.
- 910 NEI. (2018). National Emission Inventory for the U.S. [https://www.epa.gov/air-emissions-](https://www.epa.gov/air-emissions-inventories/)  
911 [inventories/](https://www.epa.gov/air-emissions-inventories/).
- 912 O'Neill, B., Tebaldi, C., Vuuren, D. v., Eyring, V., Friedlingstein, P., Hurtt, G., . . . Sanderson, B.  
913 M. (2016). The Scenario Model Intercomparison Project (ScenarioMip) for CMIP6.  
914 *GeoSci. Model Dev.*, *9*, doi:10.5194/gmd-9-3461-2016.
- 915 Qu, Z., Henze, D. K., Capps, S. L., Wang, Y., Xu, X., Wang, J., & Keller, M. (2017). Monthly  
916 top-down NO<sub>x</sub> emissions for China (2005;2012): A hybrid inversion method and trend  
917 analysis. *J. Geophys. Res. Atmos.*, *122*, 4600-4625.
- 918 Qu, Z., Henze, D. K., Li, C., Theys, N., Wang, Y., Wang, J., . . . Ren, X. (2019). SO<sub>2</sub> Emission  
919 Estimates Using OMI SO<sub>2</sub> Retrievals for 2005-2017. *J. Geophys. Res. Atmos.*, *124*,  
920 <https://doi.org/10.1029/2019JD030243>.
- 921 Qu, Z., Henze, D. K., Theys, N., Wang, J., & Wang, W. (2019). Hybrid Mass Balance/4D-Var  
922 Joint Inversion of NO<sub>x</sub> and SO<sub>2</sub> Emissions in East Asia. *J. Geophys. Res. Atmos.*, *124*,  
923 doi:10.1029/2018JD030240.
- 924 Rao, S., Klimont, Z., J. Smith, S., Dingenen, R. V., Dentener, F., Bouwman, L., . . . Calvin, K.  
925 (2016). Future air pollution in the Shared Socio-economic Pathways. *Global*  
926 *Environmental Change*, <https://doi.org/10.1016/j.gloenvcha.2016.05.012>.
- 927 Riahi, K., Vuuren, D. v., & Kriegler, E. (2017). The Shared Socioeconomic Pathways and their  
928 energy, landuse, and greenhouse gas emissions implications : An overview. *Global*  
929 *Environmental Change*, *42*, 153-168.
- 930 Stavrou, T., Müller, J.-F., Bauwens, M., Smedt, I. D., Roozendaal, M. V., Mazière, M. D., . . .  
931 Guenther, A. (2015). How consistent are top-down hydrocarbon emissions based on  
932 formaldehyde observations from GOME2 and OMI? *Atmos. Chem. Phys.*, *15*, 11861-  
933 11884.
- 934 Stavrou, T., Müller, J.-F., Boersma, K. F., van der A, R. J., Kurokawa, J., Ohara, T., & Zhang,  
935 Q. (2013). Key chemical NO<sub>x</sub> sink uncertainties and how they influence top-down  
936 emissions of nitrogen oxides. *Atmos. Chem. Phys.*, *13*, 9057-9082,  
937 <https://doi.org/10.5194/acp-13-9057-2013>.
- 938 Stavrou, T., Müller, J.-F., De Smedt, I., Van Roozendaal, M., van der Werf, G. R., Giglio, L.,  
939 & Guenther, A. (2009). Evaluating the performance of pyrogenic and biogenic emission  
940 inventories against one decade of space-based formaldehyde columns. *Atmos. Chem.*  
941 *Phys.*, *9*, <https://doi.org/10.5194/acp-9-1037-2009>, 2009.
- 942 Stavrou, T., Müller, J.-F., Peeters, J., Razavi, A., Clarisse, L., Clerbaux, C., . . . Paton-Walsh,  
943 C. (2012). Satellite evidence for a large source of formic acid from boreal and tropical  
944 forests. *Nature Geosci.*, *5*, <https://doi.org/10.1038/ngeo1354>.
- 945 Sun, W., Shao, M., Granier, C., Liu, Y., Ye, C., & Zheng, J. (2018). Long-term trends of  
946 Anthropogenic SO<sub>2</sub>, NO<sub>x</sub>, CO, and NMVOCs Emissions in China. *Earth's Future*, *6*,  
947 1112-1133, <https://doi.org/10.1029/2018EF000822>.
- 948 Theys, N., Smedt, I. d., Gent, J. v., Danckaert, T., Wang, T., Hendrick, F., & al., e. (2015). Sulfur  
949 dioxide vertical column DOAS retrievals from the Ozone Monitoring Instrument: Global  
950 observations and comparison to ground-based and satellite data. *Journal of Geophysical*  
951 *Research: Atmospheres*, *120*, <https://doi.org/10.1002/2014JD022657>.

- 952 van Vuuren, D., Edmonds, J., Kainuma, M., & al., e. (2011). The representative concentration  
953 pathways: an overview. *Climate Change*, *109*, <https://doi.org/10.107/s10584-011-0148-Z>.
- 954 Watanabe, S., Hajima, T., Sudo, K., Nagashima, T., Takemura, T., Okajima, H., . . . Kawamiya,  
955 M. (2011). MIROC-ESM 2010: model description and basic results of CMIP5-20c3m  
956 experiments. *Geosci. Model Dev.*, 845–872, doi:10.5194/gmd-4-845-2011.
- 957 Zhang, Q., Streets, D. G., Carmichael, G. R., He, K. B., Huo, H., Kannari, A., . . . Yao, Z. L.  
958 (2009). Asian emissions in 2006 for the NASA INTEX-B mission. *Atmos. Chem. Phys.*,  
959 *9*, 5131-5153.
- 960 Zheng, B., Chevallier, F., Ciais, P., Yin, Y., & Wang, Y. (2018). On the role of the flaming to  
961 smoldering transition in the seasonal cycle of African fire emissions. *Geophys. Res. Lett.*,  
962 *45*, doi: 10.1029/2018GL079092.
- 963 Zheng, B., Chevallier, F., Ciais, P., Yin, Y., Deeter, M., Worden, H., . . . He, K. B. (2018). Rapid  
964 decline in carbon monoxide emissions and export from East Asia between years 2005 and  
965 2016. *Environ. Res. Lett.*, *13*, doi: 10.1088/1748-9326/aab2b3.
- 966 Zheng, B., Chevallier, F., Yin, Y., Ciais, P., Fortems-Cheiney, A., Deeter, M. N., . . . Zhao, Y.  
967 (2019). Global atmospheric carbon monoxide budget 2000;2017 inferred from multi-  
968 species atmospheric inversions. *Earth Syst. Sci. Data*, *11*, doi:10.5194/essd-2019-61.
- 969 Zheng, B., Huo, H., Zhang, Q., Yao, Z. L., Wang, X. T., Yang, X. F., . . . He, K. B. (2014).  
970 High-resolution mapping of vehicle emissions in China in 2008. *Atmos. Chem. Phys.*, *14*,  
971 doi:10.5194/acp-14-9787-2014.
- 972 Zheng, B., Tong, D., Li, M., Liu, F., Hong, C., Geng, G., . . . Zhang, Q. (2018). Trends in  
973 China's anthropogenic emissions since 2010 as the consequence of clean air actions.  
974 *Atmos. Chem. Phys.*, *18*, doi:10.5194/acp-18-14095-2018.
- 975 Zhu, L. M., Sheng, J., Hu, L., Gonzalez Abad, G., & Chance, K. (n.d.).
- 976 Zhu, L., Jacob, D. J., Kim, P. S., Fisher, J. A., Yu, K., Travis, K. R., . . . Wolfe, G. M. (2016).  
977 Observing atmospheric formaldehyde (HCHO) from space: validation and  
978 intercomparison of six retrievals from four satellites (OMI, GOME2A, GOME2B,  
979 OMPS) with SEAC4RS aircraft observations over the southeast US. *Atmos. Chem. Phys.*,  
980 *16*, <https://doi.org/10.5194/acp-16-13477-2016>.
- 981 Zhu, L., Mickley, L. J., Jacob, D. J., Marais, E. A., Sheng, J., Hu, L., . . . Chance, K. (2017).  
982 Long-term (2005-2014) trends in formaldehyde columns across North America as seen  
983 by the OMI satellite instrument; Evidence of changing emissions of volatile organic  
984 compounds. *J. Geophys. Res.*, *44*, 7079-7086.

985  
986

987

988  
989