

# The recent emergence of Arctic Amplification

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## Key Points:

- Arctic Amplification only started to occur in the second half of the 20th century. Prior to this the Arctic cooled under global warming.
- Without the increases in both aerosols and greenhouse gas emissions, Arctic Amplification would have occurred throughout the past century.
- Internal variability also plays an important role in setting the observed trends.

## Abstract

Arctic Amplification is robustly seen in climate model simulations of future warming and in the paleoclimate record. Here, we show that in the instrumental record Arctic Amplification is only a recent phenomenon, and that for much of the 20th century the Arctic cooled while the global-mean temperature rose. To investigate why this occurred, we analyze large ensembles of comprehensive climate model simulations under different forcing scenarios. Our results suggest that the global warming from greenhouse gases was largely offset in the Arctic by regional cooling due to aerosols, with internal climate variability also contributing to Arctic cooling and global warming trends during this period. This suggests that the disruption of Arctic Amplification was due to a combination of factors unique to the 20th century, and that enhanced Arctic warming should be expected to be a consistent feature of climate change over the coming century.

## Plain Language Summary

Arctic Amplification is the phenomenon by which the Arctic warms at a faster rate than the global average. Evidence for the occurrence of Arctic Amplification is widely found in climate model simulations as well as in paleoproxy reconstructions of past climate changes. In this study, we investigate the extent to which Arctic Amplification has occurred in observations from the past century. We show that Arctic Amplification is only a recent phenomenon, and that for much of the 20th century, the Arctic cooled while the global-mean temperature rose. We investigate why this happened using a range of climate model simulations, and we find that there were two main causes for these opposing trends. The first is that regional cooling from aerosols counteracted the warming from greenhouse gases in the Arctic. However, this cannot fully explain the observed trends. The second is that natural fluctuations of the climate system manifested in a pattern of Arctic cooling under global warming. This suggests that the disruption of Arctic Amplification was due to a combination of factors unique to the 20th century, implying that enhanced Arctic warming should be expected to be a consistent feature of climate change over the coming century.

## 1 Introduction

Arctic Amplification, the phenomenon whereby the surface air temperature in the Arctic warms at an enhanced rate relative to the rest of the globe, is believed to be one of the most robust features of the climate system's response to external forcings. First identified in Manabe and Stouffer (1980), it has long been known that global climate models simulate amplified warming in the Arctic in response to climate forcing. For example, in response to a quadrupling of CO<sub>2</sub>, CMIP5 models simulate a mean surface warming of over 10K in the Arctic, more than double the global mean surface warming (Pithan & Mauritsen, 2014). Amplified warming in the Arctic is a consistent feature of climate model projections of the coming century (Barnes & Polvani, 2015; Laine et al., 2016; Davy & Outten, 2020), and is found in both idealized and comprehensive climate models (Holland & Bitz, 2003; Langen & Alexeev, 2007; Franzke et al., 2017; Merlis & Henry, 2018; Beer et al., 2020). Furthermore, multiple lines of paleoclimate evidence indicate Arctic Amplification of warming and cooling in past warm and cold climates (CAPE, 2006; Masson-Delmotte et al., 2006; Miller et al., 2010; Park et al., 2019).

Many previous studies have investigated the potential causes of Arctic Amplification (Winton, 2006; Screen & Simmonds, 2010; Hwang et al., 2011; Serreze & Barry, 2011; Pithan & Mauritsen, 2014; Stuecker et al., 2018; Henry & Merlis, 2019; Previdi et al., 2020; Henry et al., 2021; Beer et al., 2020). A large array of mechanisms have been identified, including the sea ice albedo feedback, the lapse rate feedback, and changes in energy transports, however the precise contributions of these varied mechanisms remain

62 unclear. For example, numerous studies have pointed to the surface albedo feedback as  
63 a fundamental process in driving Arctic Amplification (Screen & Simmonds, 2010; Dai  
64 et al., 2019; Chung et al., 2021), yet enhanced Arctic warming is found in climate model  
65 simulations even when the surface albedo feedback is disabled (Graversen & Wang, 2009).  
66 Untangling the drivers of Arctic Amplification is further complicated because different  
67 forcing agents, including CO<sub>2</sub>, ozone depleting substances, black carbon, and industrial  
68 aerosols, can also have different imprints on the spatial pattern of surface warming (Navarro  
69 et al., 2016; Stuecker et al., 2018; Stjern et al., 2019; Polvani et al., 2020).

70 While many studies have focused on the causes and effects of Arctic Amplification  
71 in climate model simulations, relatively little attention has been paid to this phenomenon  
72 in the observational record. In this study, we focus on the extent to which Arctic Am-  
73 plification has occurred in the instrumental record over the past century. We begin by  
74 analyzing several observational products to quantify the nature of Arctic Amplification  
75 during the past century, and then we use large ensembles of comprehensive climate model  
76 simulations to quantify how industrial aerosols, greenhouse gases, and internal climate  
77 variability have contributed to the observed trends.

## 78 **2 Data and methods**

### 79 **2.1 Observations**

80 We focus on the past century (1921-2020), eschewing earlier times due to the relative  
81 lack of temperature measurements. We analyze five different observational prod-  
82 ucts for our analysis of near surface air temperature (SAT) trends over the instrumen-  
83 tal record: GISTEMPv4 (Lenssen et al., 2019; GISTEMP, 2021), HadCRUT5 (Morice  
84 et al., 2020), the Cowtan and Way (2014) update to HadCRUT4 referred to as HadCRUT4-  
85 hybrid, the ERA-20C reanalysis (Poli et al., 2016), and the ERA5 reanalysis (Hersbach  
86 et al., 2020). For ease of presentation, we focus on results from GISTEMPv4, which is  
87 based on NOAA-GHCN-v4 station data over land and ERSSTv5 over ocean (Huang et  
88 al., 2017). Note that only GISTEMPv4 and HadCRUT5 cover the entire period of inter-  
89 est: the HadCRUT4-hybrid dataset ends in 2018, ERA-20C ends in 2010, and ERA5  
90 only starts in 1950. Lastly, we analyze the 200 members of the HadCRUT5 ensemble to  
91 assess the impacts of observational uncertainty, which arises from statistical infilling of  
92 sparsely observed areas, measurement uncertainty, and changes in SST measurement prac-  
93 tices, among other factors. More information about these datasets is given in Table S1.

### 94 **2.2 Climate model simulations**

95 We utilize the 40 members of the Community Earth System Model v1 Large En-  
96 semble (CESM1-LE), introduced by Kay et al. (2015). CESM1 is a CMIP5-class climate  
97 model, using the CAM5 atmospheric model and the POP2 ocean model. Each of the 40  
98 members uses identical historical forcing (Lamarque et al., 2010) for the period 1920-  
99 2005 and future emissions under the RCP8.5 scenario from 2006 onwards (Meinshausen  
100 et al., 2011). The only difference between the members arises from chaotic fluctuations  
101 generated by round-off level ( $10^{-14}$  K) perturbations to the atmospheric initial temper-  
102 ature in 1920. As such, each ensemble member is a realization of the climate system over  
103 the past century, with the ensemble mean isolating the forced response to external forc-  
104 ing and the spread among the ensemble members being solely due to internal variabil-  
105 ity of the climate system. The CESM1-LE has been widely used to investigate the roles  
106 of anthropogenic forcing and internal variability in driving observed trends in Arctic SAT  
107 and sea ice (Ding et al., 2017; England et al., 2019; Ding et al., 2019; Landrum & Hol-  
108 land, 2020; Krishnan et al., 2020; Polvani et al., 2020).

109 To investigate the contributions of anthropogenic aerosols and greenhouse gases to  
110 observed SAT trends over the past century, we analyze two CESM1 single-forcing en-

111 sembles, each containing 20 members, introduced by Deser, Phillips, et al. (2020). The  
 112 first, CESM1 x-aer, is identical to the CESM1-LE except that industrial aerosol concen-  
 113 trations are fixed at 1920 values. All other forcings evolve as in the CESM1-LE. In the  
 114 same fashion, the second, CESM1 x-ghg, is identical to the CESM1-LE except that green-  
 115 house gas concentrations are held fixed at 1920 values. Taking the difference between  
 116 the ensemble mean of the CESM1-LE (which features all forcings) and either CESM1  
 117 x-aer or CESM1 x-ghg (which feature all but one forcing) isolates the roles of aerosols  
 118 and greenhouse gases in driving historical temperature trends in CESM1. More details  
 119 on the large ensembles analyzed here are given in Table S2.

120 We compare the results from the CESM1-LE with three other large ensembles which  
 121 also participated in the recent CLIVAR large ensemble collection (Deser, Lehner, et al.,  
 122 2020). Specifically, we analyse the 30 members of CSIRO-Mk3-6-0 (Jeffrey et al., 2013),  
 123 the 20 members of GFDL-CM3 (Sun et al., 2018) and the 100 members of MPI-ESM (Maher  
 124 et al., 2019). We focus on the results from the CESM1-LE because of the availability of  
 125 single forcing ensembles with this model.

### 126 **3 Prolonged periods of observed Arctic cooling during global warm-** 127 **ing**

128 We begin by examining the timeseries of observed Arctic (60°N - 90°N) and global  
 129 average SAT anomalies relative to a baseline period of 1951-1980, as shown in Figure 1.  
 130 These show the much larger interannual variability and multi-decadal changes in the Arc-  
 131 tic temperature (blue line) as compared to the global temperature (red line). As doc-  
 132 umented by previous studies (Johannessen et al., 2004; Gillett et al., 2008; Serreze et al.,  
 133 2009; Semenov & Latif, 2012), the timeseries of Arctic surface temperatures can be char-  
 134 acterized by a period of anomalous warmth in the 1930s and 1940s, and then a substan-  
 135 tial period of cooling lasting until the 1980s, which was followed by four decades of rapid  
 136 Arctic warming up to the present day. These results remain qualitatively unchanged if  
 137 the domain is limited to either land or ocean regions only in the Arctic (Fig. S2). In con-  
 138 trast to the Arctic, global temperatures have generally risen throughout the past cen-  
 139 tury, with muted warming for much of the mid-20th century and an increased rate of warm-  
 140 ing after approximately 1980.

141 The two timeseries in Figure 1 demonstrate that Arctic Amplification did not oc-  
 142 cur for large parts of the 20th century. To illustrate this, we focus on the 50-year period  
 143 1935-1984, indicated by the trendlines in Figure 1. During this period, the global-mean  
 144 temperature showed a small warming trend of  $0.03\text{ }^{\circ}\text{C}/\text{decade} \pm 0.02\text{ }^{\circ}\text{C}/\text{decade}$ , yet the  
 145 Arctic cooled at a rate of  $-0.15\text{ }^{\circ}\text{C}/\text{decade} \pm 0.08\text{ }^{\circ}\text{C}/\text{decade}$ , where the uncertainties rep-  
 146 resent the 95% linear regression confidence interval. These results from GISTEMPv4 are  
 147 consistent with the other observational and reanalysis products analyzed here (Fig. S1).

148 Next, we examine all 50-year trends in observed Arctic (blue) and global (red) SAT  
 149 over the past century (Fig. 2a) to investigate exactly when Arctic Amplification began.  
 150 The red shading in Figure 2a shows the 50-year SAT trends that feature Arctic Ampli-  
 151 fication (i.e. where the rate of Arctic warming is greater than the rate of global warm-  
 152 ing), while the blue shading denotes 50-year SAT trends of simultaneous Arctic cooling  
 153 and global warming. We highlight three important points: (i) the first occurrence of Arc-  
 154 tic Amplification in the observed record was the years 1951-2000, (ii) during the past cen-  
 155 tury, periods of Arctic cooling under global warming occurred approximately as frequently  
 156 as periods of Arctic Amplification, and (iii) the switch from Arctic cooling to Arctic am-  
 157 plified warming (white shading) was rapid, transitioning in less than five years. The same  
 158 broad features are found in the other observational and reanalysis datasets that we an-  
 159 alyzed (Fig. S3a). In addition, the conclusions are robust to the impacts of observational  
 160 uncertainty as represented by the 200 ensemble members of HadCRUT5 (Fig. S3b) and

161 remain approximately unchanged if we examine the non-infilled version of HadCRUT5  
162 (compare panels b and c in Fig. S3).

163 We turn to the 40 members of the CESM1-LE to examine the extent to which this  
164 observed behaviour is replicated by climate model simulations (Figs. 2b,c). There are  
165 individual ensemble members that largely capture the evolution of observed trends in  
166 Arctic and global SAT. As an example we show ensemble member #33 (Fig. 2b, crosses),  
167 which exhibits Arctic cooling for much of the early and mid 20th century followed by  
168 Arctic amplified warming thereafter. There are several differences between this ensemble  
169 member and the observations, namely that the Arctic cooling trends are approximately  
170 30% smaller and the transition to Arctic Amplification occurs a few years earlier than  
171 observed. Overall, however, ensemble member #33 replicates the main features of the  
172 observed trends. But, the majority of ensemble members fail to simulate the observed  
173 trends in both Arctic and global-mean temperatures. As an example, ensemble member  
174 #31 exhibits Arctic-amplified surface warming trends throughout the past century  
175 (Fig. 2b, circles), which is inconsistent with the observations. The substantial differences  
176 in the running 50-year Arctic trends of the two members indicate the important role of  
177 internal variability in driving observed trends.

178 To separate the roles of the forced response and internal variability in contributing  
179 to the observed trends, we show the 50-year SAT trends of all 40 members of the CESM1-  
180 LE (Fig. 2c, dots) in addition to the ensemble mean (Fig. 2c, thick lines). The observed  
181 global-mean trends are roughly consistent with the forced trends in the CESM1-LE, with  
182 weakly positive trends over the early and middle-20th century, followed by more rapid  
183 warming in the second half of the century. In contrast, the observed 50-year Arctic cooling  
184 trends over the mid-20th century are at the edge of the CESM1-LE distribution (compare  
185 panels a and c of Fig. 2), suggesting that the observed trends are either a low probability  
186 trajectory of the climate system or that the CESM1-LE systematically underestimates  
187 the Arctic cooling during this period. More recently, the strong Arctic warming  
188 trends over the past 50 years are reproduced by the ensemble-mean of the CESM1-  
189 LE (compare panels a and c of Fig. 2).

190 Although most CESM1-LE members do not reproduce the Arctic cooling trends  
191 seen over much of the 20th century, the ensemble mean does show a slight Arctic cooling  
192 for trends centred in the 1940s and 50s, during a period in which the ensemble-mean  
193 global trend is positive. This indicates a potential role for the ensemble mean response  
194 alone to drive periods of Arctic cooling concurrent with global warming. Moreover, the  
195 majority of the members (31 out of 40) include at least some periods of Arctic cooling  
196 concurrent with global warming over the past century, suggesting that this observed phenomenon  
197 can be reproduced in part by climate models. The 50-year Arctic and global  
198 SAT trends simulated by the CESM1-LE are broadly consistent with the three other large  
199 ensembles we analyzed (Fig. S4) although we note that MPI-ESM does not simulate the  
200 periods of observed Arctic cooling and CSIRO-Mk3-6-0 and GFDL-CM3 do not replicate  
201 the observed global warming in the mid-20th century. In addition, GFDL-CM3 simulates  
202 Arctic warming trends over the second half of the 20th century which are much  
203 stronger than observed. Overall, we find that of the CESM1-LE performs the best at capturing  
204 the major characteristics of the observed trends. Next, we explore why there was  
205 no Arctic Amplification for much of the 20th century.

#### 206 **4 What caused the lack of Arctic Amplification?**

207 In order to investigate what caused the opposing trends in observed Arctic and global  
208 SAT, we focus on three central potential drivers: industrial aerosols, greenhouse gases,  
209 and internal climate variability. We investigate the forced response to aerosols and greenhouse  
210 gases by analyzing the 20 member CESM1 single forcing ensembles with fixed aerosols  
211 (x-aer) and fixed greenhouse gases (x-ghg). Previous studies have shown the importance

212 of industrial aerosols (Fyfe et al., 2013; Navarro et al., 2016; Gagne et al., 2017; Deser,  
 213 Phillips, et al., 2020) and greenhouse gases (Gillett et al., 2008; Nafaji et al., 2015; Deser,  
 214 Phillips, et al., 2020; Polvani et al., 2020) in contributing to the observed Arctic SAT  
 215 trends over the past century. However, the contributions of greenhouse gases, aerosols  
 216 and internal variability to the observed Arctic SAT trends have yet to be quantified.

217 Figure 3a shows that in the absence of changes in the concentrations of industrial  
 218 aerosols since 1920, CESM1 indicates that both Arctic and global SAT would have risen  
 219 monotonically throughout the past century. Hence, in the absence of extra industrial aerosol  
 220 emissions, the Arctic would not have experienced any 50-year cooling trends over the past  
 221 century. In contrast, without the increase in greenhouse gases since 1920, CESM1 indi-  
 222 cates that both the Arctic and the global mean surface would have cooled throughout  
 223 the century (Fig. 3b). This suggests that greenhouse gases are necessary to explain the  
 224 global surface warming throughout the century as well as the rapid Arctic warming since  
 225 the second half of the 20th century, while aerosols are required to explain the Arctic cool-  
 226 ing over much of the 20th century. It is important to note that in both ensembles, x-aer  
 227 and x-ghg, Arctic Amplification is present in nearly every member throughout the en-  
 228 tire century. Therefore it is the specific combination of greenhouse gas emissions and in-  
 229 dustrial aerosol emissions, and the extent to which their effects offset each other at a global  
 230 and regional scale, that created the conditions for Arctic Amplification not to occur for  
 231 much of the 20th century.

232 Next we focus again on the period 1935-1984 and quantify the contributions to the  
 233 observed trends from aerosols, greenhouse gases, and internal variability (Fig 4). This  
 234 50-year period was chosen to broadly represent the Arctic and global mid-20th century  
 235 SAT trends. Similar results are found for a decade earlier (1925-1974, Fig. S5). To cal-  
 236 culate the contribution of aerosols (light blue bar) we take the difference between the  
 237 ensemble mean of the CESM1-LE and the ensemble mean of the CESM1 x-aer ensem-  
 238 ble and then calculate the trends over the period of interest. To calculate the contribu-  
 239 tion of greenhouse gases (light red bar), we repeat this process but with the CESM1 x-  
 240 ghg ensemble rather than x-aer. The role of internal variability is estimated as the resid-  
 241 ual (orange bar) after the effects of aerosols and greenhouse gases have been subtracted  
 242 from the observed trend. This is compared with the range of trends attributable to in-  
 243 ternal variability in the CESM1-LE (black bars), which is computed by subtracting the  
 244 ensemble mean to remove the forced response and then calculating the central 95% range  
 245 of the 50-year SAT trends across the 40-members (i.e. the difference between the 39th  
 246 and 2nd member after ranking).

247 These results allow us to estimate what caused the observed Arctic cooling of  $-0.15^{\circ}\text{C}/\text{decade}$   
 248 during 1935-1984. As shown in Figure 4a, CESM1 implies that industrial aerosols caused  
 249 a cooling trend of  $-0.27^{\circ}\text{C}/\text{decade}$ , and greenhouse gases caused a warming trend of  $+0.20^{\circ}\text{C}/\text{decade}$ ,  
 250 which resulted in a net anthropogenic impact of  $-0.07^{\circ}\text{C}/\text{decade}$ . Thus the residual needed  
 251 to account for the observed trend is  $-0.08^{\circ}\text{C}/\text{decade}$ . This is well within the range of 50-  
 252 year SAT trends due to internal variability simulated by CESM1-LE [ $-0.14^{\circ}\text{C}/\text{decade}$ ,  
 253  $+0.11^{\circ}\text{C}/\text{decade}$ ], implying that the residual can be plausibly attributed to internal vari-  
 254 ability. The results therefore suggest that in the absence of internal variability, the ob-  
 255 served Arctic cooling trend would have been half as large. In addition, some realizations  
 256 of internal variability would have overcome the net anthropogenic cooling effect and re-  
 257 sulted greater Arctic warming over this period than the observed global warming. That  
 258 is to say, the lack of Arctic Amplification was not an inevitable response to the anthro-  
 259 pogenic forcing.

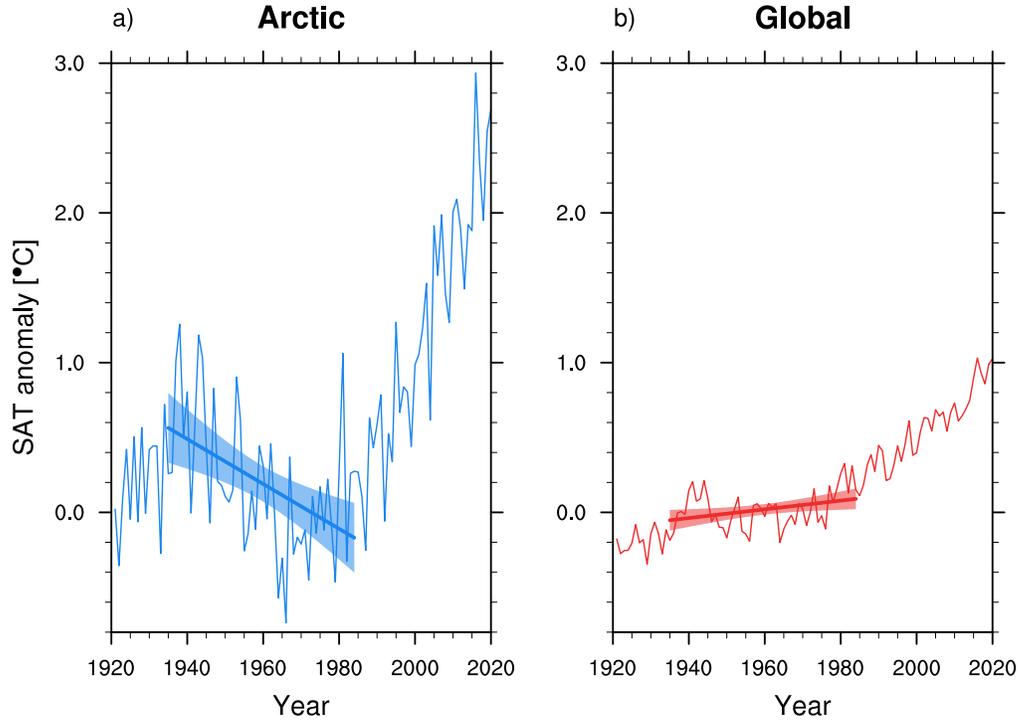
260 We can similarly estimate what drove the observed global mean warming trend of  
 261  $+0.03^{\circ}\text{C}/\text{decade}$  (Fig. 4b) during 1935-1984. We find that there is a near perfect can-  
 262 cellation between the aerosol cooling trend of  $-0.09^{\circ}\text{C}/\text{decade}$  and the greenhouse gas in-  
 263 duced warming trend of  $+0.09^{\circ}\text{C}/\text{decade}$ . Thus the residual needed to account for the  
 264 observed trend is  $+0.03^{\circ}\text{C}/\text{decade}$ . This is at the edge of the range of internal variabil-

265 ity simulated by CESM1-LE [ $\pm 0.03^\circ\text{C}/\text{decade}$ ]. This suggests that we could have experi-  
266 enced global mean cooling during this period under a different realization of internal  
267 variability, and as such we could have experienced a period of Arctic amplified cooling.  
268 We note that the Arctic Amplification factor, commonly defined as the Arctic SAT trend  
269 divided by the global SAT trend, is only 2.2 for greenhouse gases (a driver of Arctic and  
270 global warming) over this period, but 3.0 for industrial aerosols (a driver of Arctic and  
271 global cooling). This difference likely arose because aerosol emissions primarily occurred  
272 over North America and Northern Europe (Navarro et al., 2016; Deser, Phillips, et al.,  
273 2020; Krishnan et al., 2020). This helps to explain why the simulated (i.e. the CESM1-  
274 LE ensemble mean) features a small cooling trend in the Arctic SAT and a weak warm-  
275 ing trend in the global SAT (compare ensemble mean lines in Fig. 2c).

## 276 5 Conclusions

277 In this study we have investigated the extent to which Arctic Amplification occurred  
278 over the past century in the observed record. We found that Arctic Amplification is a  
279 relatively recent phenomenon during this period, first occurring in 50-year trends centered  
280 in the second half of the 20th century. We showed that 50-year periods with Arctic  
281 cooling concurrent with global warming occurred as frequently as periods with Arctic  
282 amplified warming during the past century. We then used CESM1 to investigate why  
283 Arctic Amplification was not ubiquitous throughout the past century. We showed that  
284 CESM1 single forcing experiments imply that without historical changes in greenhouse  
285 gases or aerosols, Arctic Amplification would have consistently occurred. We found that  
286 it is the cancellation of these two forcings, with aerosols having an outsized effect on the  
287 Arctic compared to the global mean, that created the conditions for Arctic cooling during  
288 global warming. Lastly we used CESM1 results to estimate the contributions of aerosols,  
289 greenhouse gases, and internal variability to the observed SAT trends during 1935-1984.  
290 These results imply that the lack of Arctic Amplification during this period, which is re-  
291 produced by many members of the CESM1-LE, was made more likely due to anthropogenic  
292 forcing, and that internal variability also played a key role. Different realizations of in-  
293 ternal variability could have caused the global mean or Arctic SAT trend to switch sign,  
294 and thus the lack of Arctic Amplification was not inevitable under the anthropogenic  
295 forcing.

296 Arctic Amplification is thought to be one of the most robust features of global warm-  
297 ing. Yet, we have shown that the emergence of this phenomenon in the observed record  
298 has only occurred relatively recently. We reconciled this by demonstrating that the lack  
299 of Arctic Amplification over much of the early and mid 20th century arose from a par-  
300 ticular combination of factors: a cancellation between the forced response to greenhouse  
301 gases and aerosols, the stronger Arctic amplified response to aerosols than to greenhouse  
302 gases over this time period, and the specific trajectory of internal variability. Moving for-  
303 ward, it is unlikely that this set of factors will manifest at any point in the 21st century  
304 for three main reasons: (i) the overall aerosol burden is expected to decrease over this  
305 century (Szopa et al., 2013; Fiedler et al., 2019) and so the forced response to increas-  
306 ing greenhouse gas concentrations will dominate, (ii) if increases in aerosol emissions do  
307 occur then it is expected they will originate from low latitudes (Fiedler et al., 2019) and  
308 thereby have limited cooling effects on the Arctic, and (iii) the levels of internal variabil-  
309 ity simulated by the four large ensembles studied here are not large enough to overcome  
310 the forced Arctic amplified warming response to projected increases in greenhouse gas  
311 concentrations. Hence, Arctic Amplification is likely to be a robust and persistent fea-  
312 ture of climate change over the coming century, despite not occurring for much of the  
313 past century.



**Figure 1.** Timeseries of (a) Arctic and (b) global annual mean surface air temperature anomalies, relative to the mean temperature during the period 1951-1980, from GISTEMPv4. The trends during 1935-1994 are indicated with shading to represent the 95% linear regression confidence interval.

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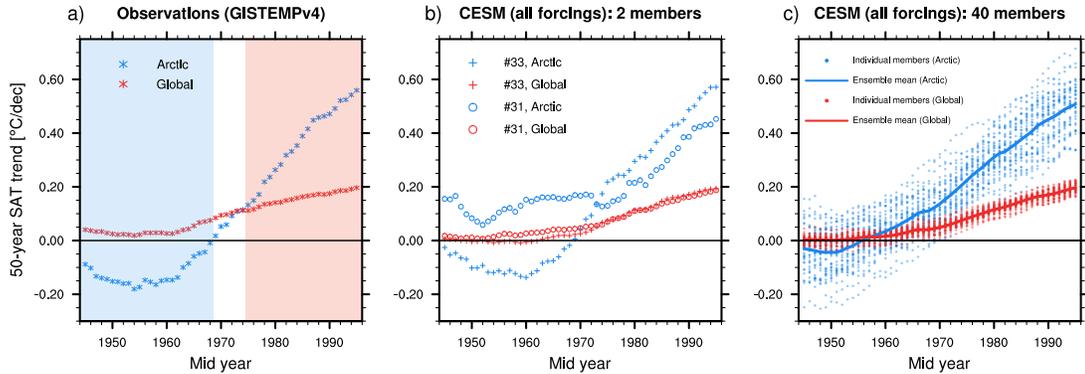
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The authors declare that they have no competing interests. Support for this work came from the National Science Foundation. The GISTEMPv4 observational dataset can be downloaded from <https://data.giss.nasa.gov/gistemp/>. The HadCRUT5 observational dataset can be download form <https://www.metoffice.gov.uk/hadobs/hadcrut5/>. The HadCRUT4-hybrid observational dataset can be downloaded from <https://www-users.york.ac.uk/kdc3/papers/coverage2013/series.html>. ERA20C and ERA5 reanalysis data can be downloaded from the Copernicus Climate Change Service Data Store at <https://cds.climate.copernicus.eu/>. The multiple large ensemble archive can be found at <http://www.cesm.ucar.edu/projects/community-projects/MMLEA/>. The CESM1-CAM5 single forcing runs are accessible via the NCAR Climate Data Gateway.

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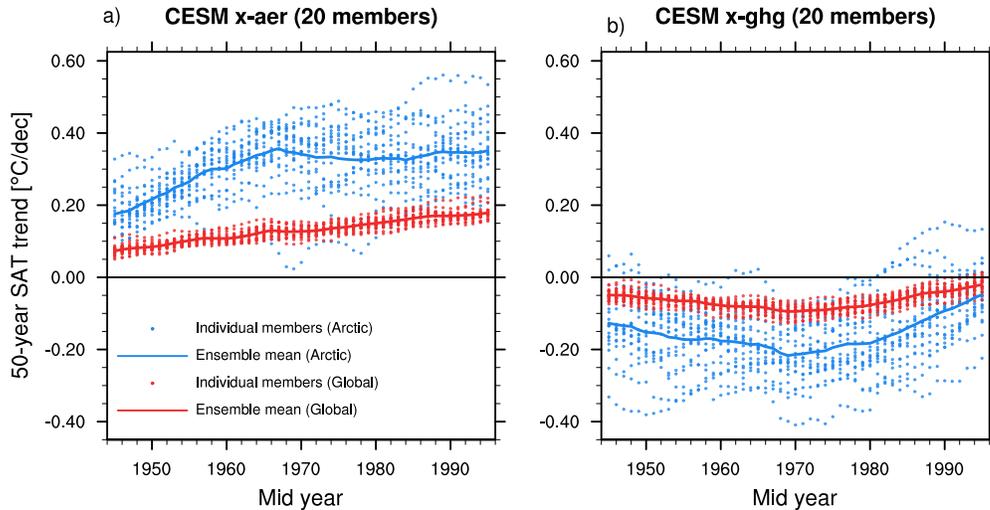
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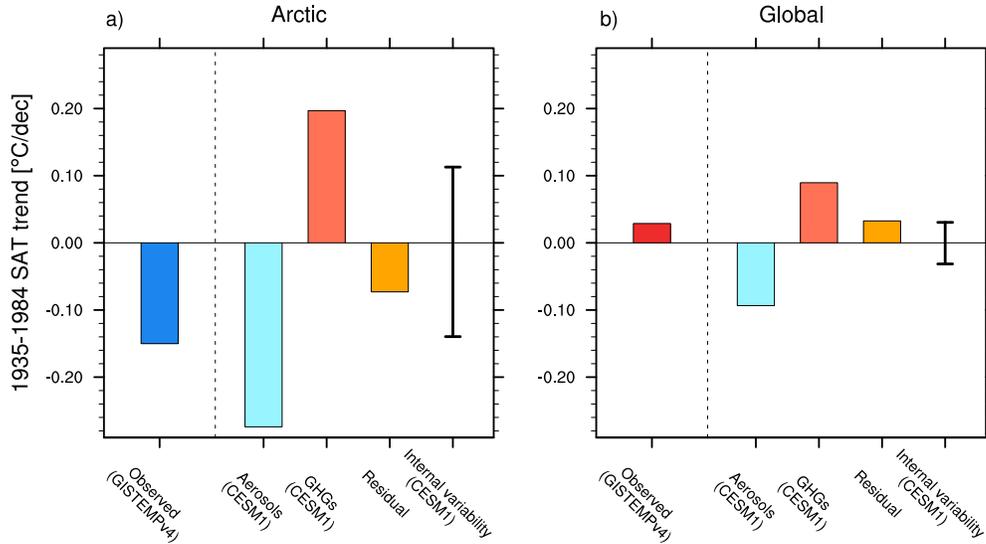
**Figure 2.** 50-year trends in Arctic (blue) and global (red) SAT over the years 1921-2020, with the mid-year shown on the horizontal axis. (a) The observed trends using GISTEMPv4, with blue shading indicating periods of concurrent Arctic cooling and global warming and red shading indicating periods of Arctic amplified warming. (b) The trends of two individual ensemble members from the CESM1-LE: ensemble member #31 (circles) and ensemble member #33 (crosses). (c) The dots show the trends for all 40 members of the CESM1-LE and the solid lines shows the ensemble mean trends.

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**Figure 3.** As in Fig. 2c, but for (a) the 20 members of the CESM1 x-aer ensemble and (b) the 20 members of the CESM1 x-ghg ensemble.

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**Figure 4.** Decomposition of the observed (a) Arctic and (b) global SAT trends during 1935-1984 into contributions from aerosols, GHGs, and internal variability. The forced responses to aerosols and GHGs are calculated using the CESM1 ensembles. The residuals after the aerosol and GHG forced responses are subtracted from the observed trends are also indicated. The black bars indicate the 95% confidence intervals for the trends which could be explained by internal variability according to the CESM1-LE.

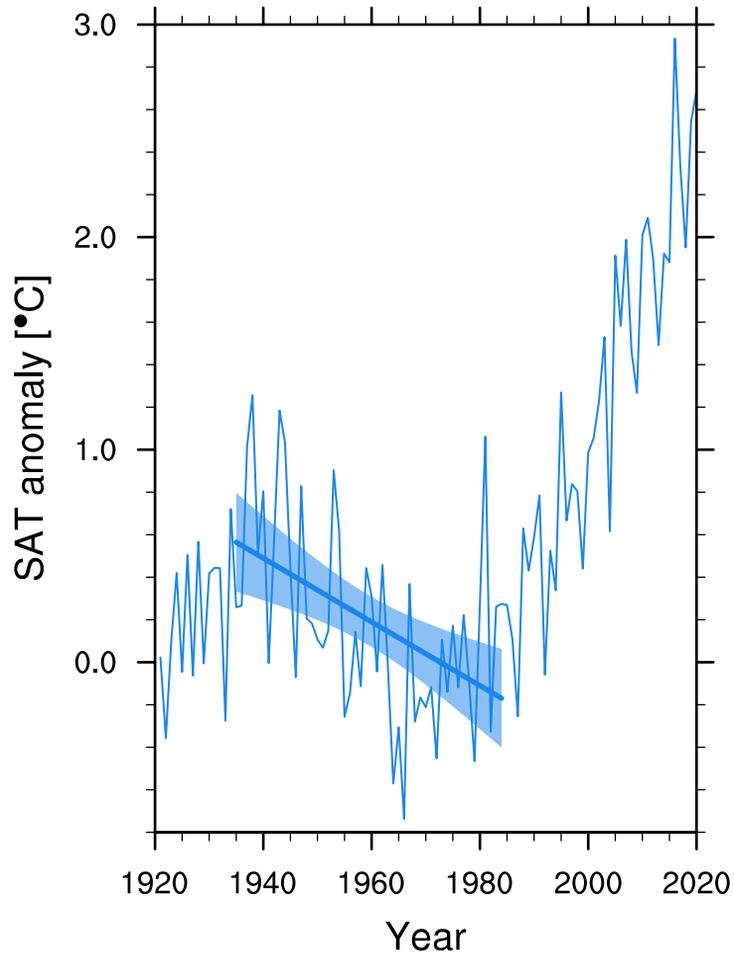
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Figure 1.

a) **Arctic**



b) **Global**

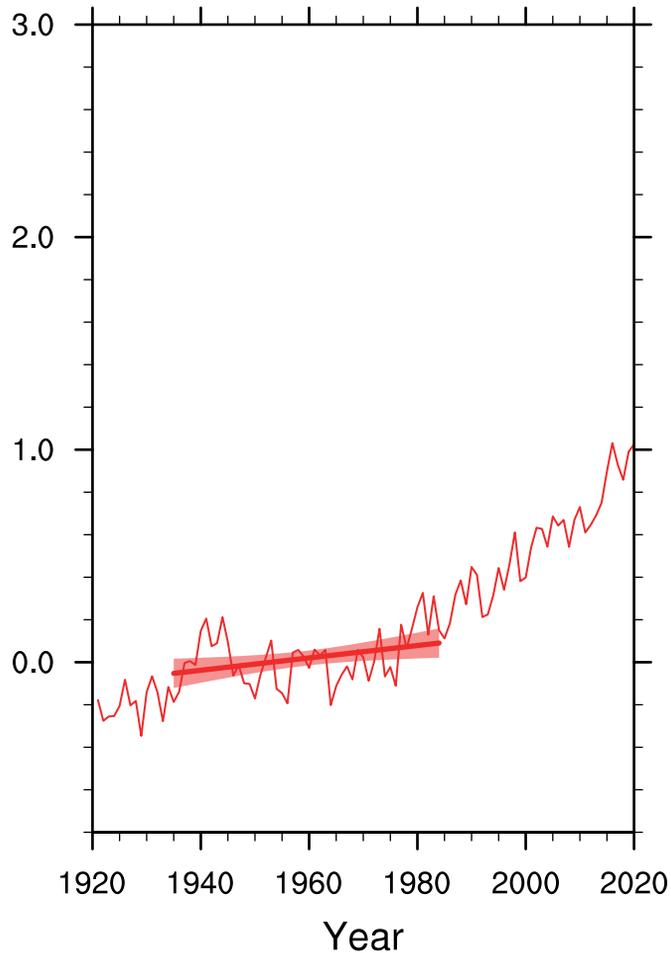
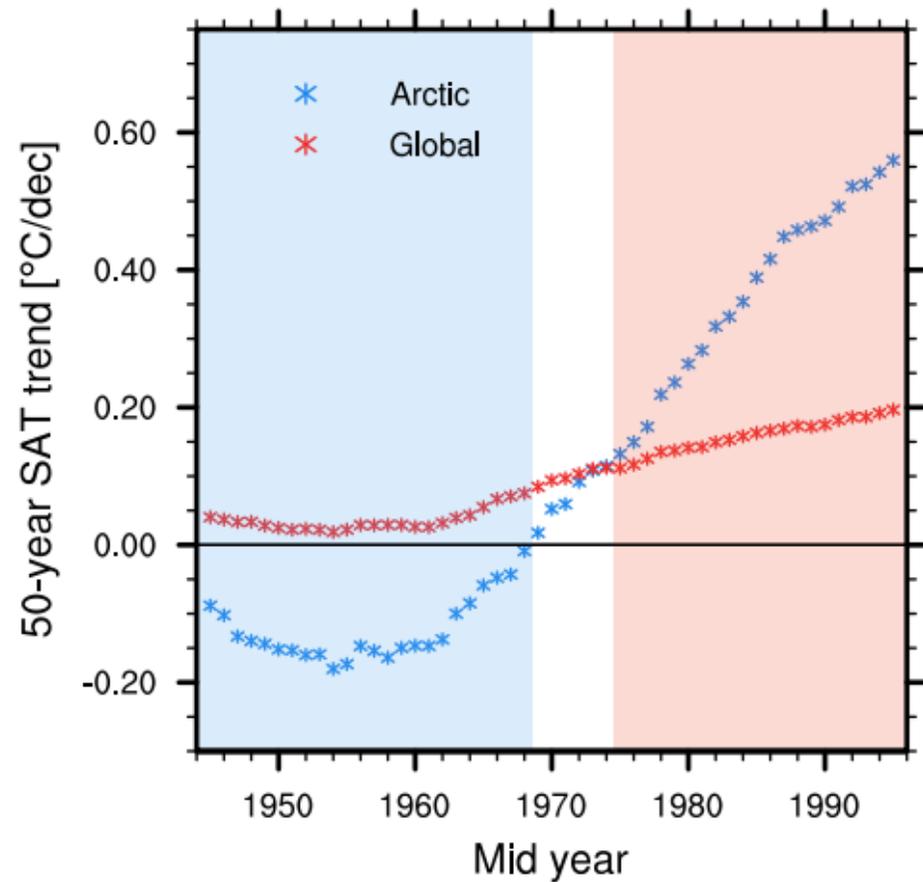
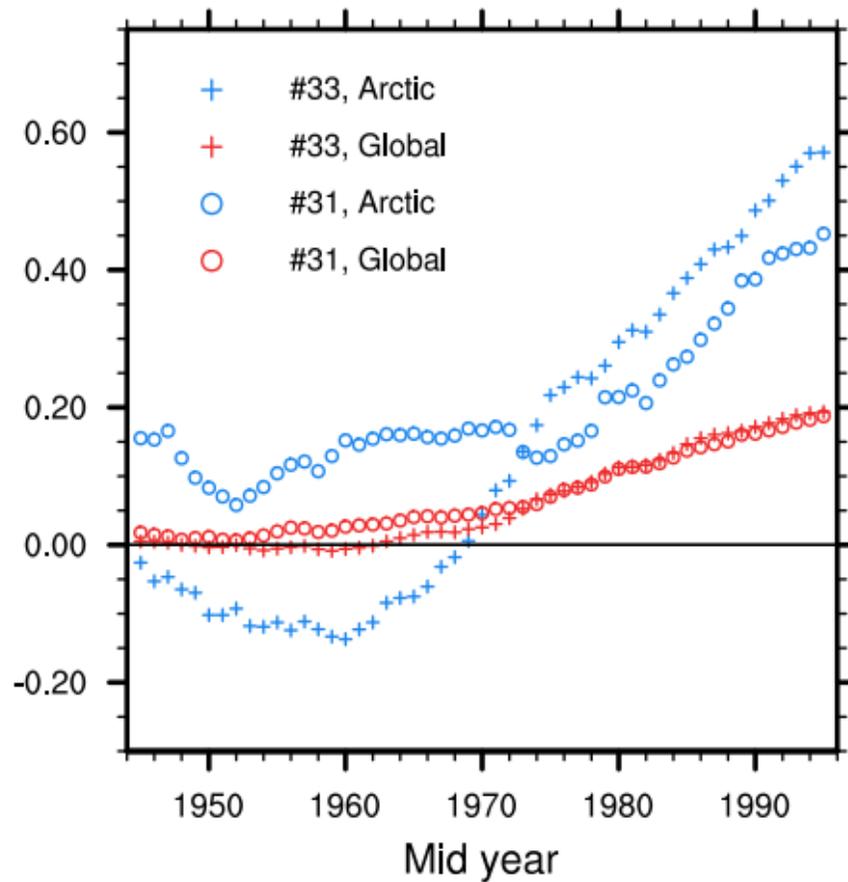


Figure 2.

a) Observations (GISTEMPv4)



b) CESM (all forcings): 2 members



c) CESM (all forcings): 40 members

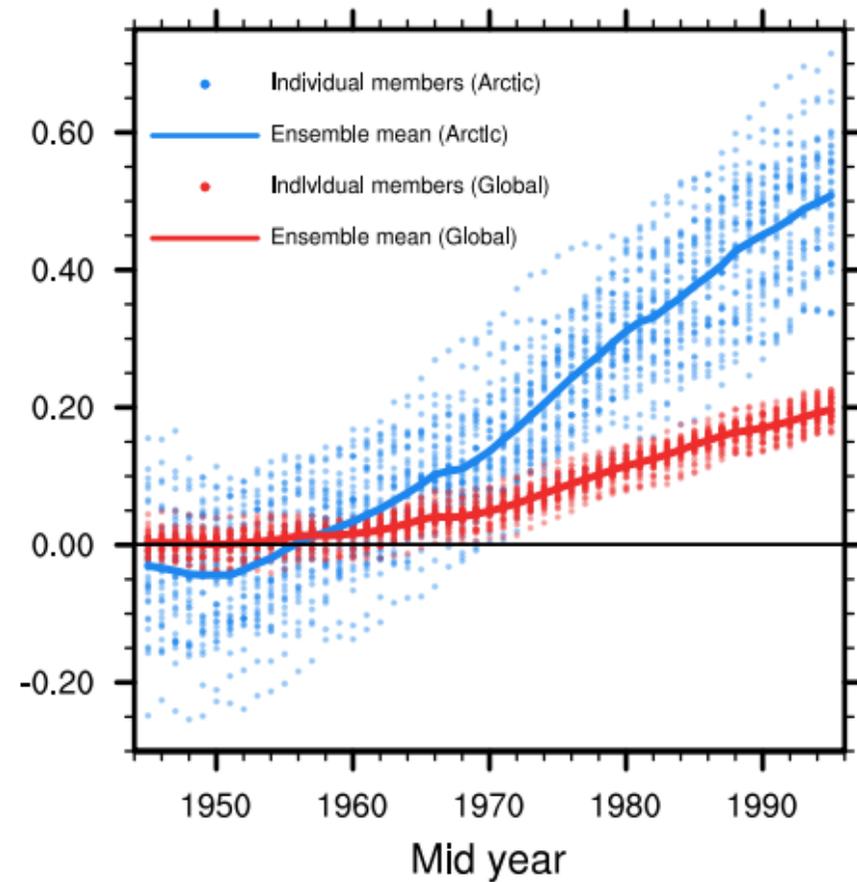
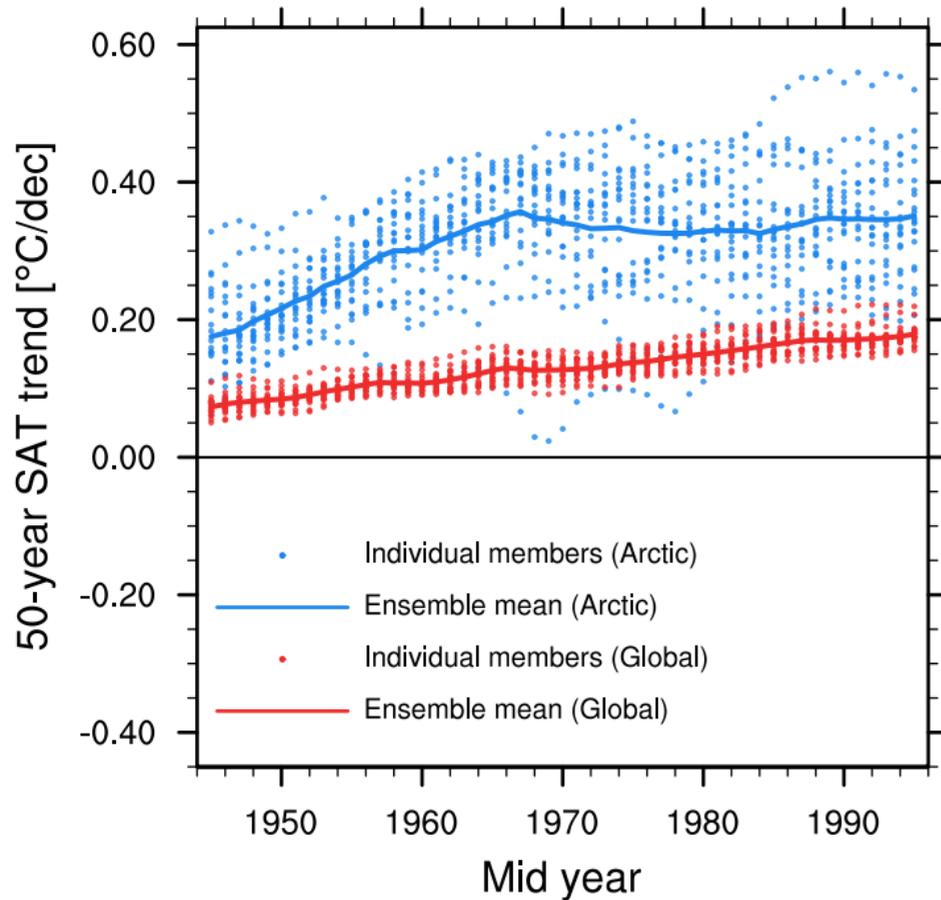


Figure 3.

a) **CESM x-aer (20 members)**



b) **CESM x-ghg (20 members)**

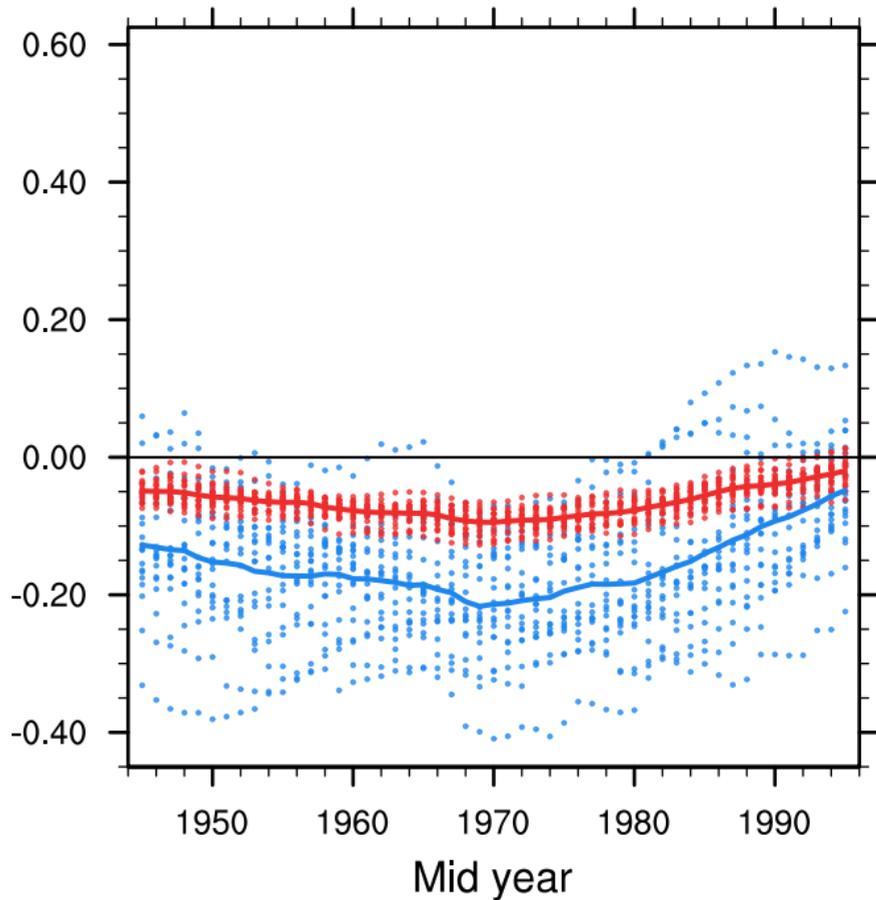


Figure 4.

