

Title: Variation of Jupiter's Aurora Observed by Hisaki/EXCEED: 4. Quasi-Periodic Variation

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Key points (<=140 characters):

1. Quasi-periodic variations of a few to several days seen in Jupiter's polar-integrated northern aurora observed by Hisaki

2. Auroral bursts <10 h sometimes seen at peak of periodic variation, whose occurrence increases with Io's volcanic activity

3. This periodic variation additionally seen in aurora intensity enhancements associated with solar wind variations

Abstract

[1] Quasi-periodic variations of a few to several days are observed in the energetic plasma and magnetic dipolarization in Jupiter's magnetosphere. Variation in the plasma mass flux related to Io's volcanic activity is proposed as a candidate of the variety of the period. Using a long-term monitoring of Jupiter's northern aurora by the Earth-orbiting planetary space telescope Hisaki, we analyzed the quasi-periodic variation seen in the auroral power integrated over the northern pole for 2014–2016, which included monitoring Io's volcanically active period in 2015 and the solar wind near Jupiter

40 during Juno's approach phase in 2016. Quasi-periodic variation with periods of 0.8–8 days was
41 detected. The difference between the periodicities during volcanically active and quiet periods is not
42 significant. Our dataset suggests that a difference of period between this volcanically active and quiet
43 conditions is below 1.25 days. This is consistent with the expected difference estimated from a
44 proposed relationship based on a theoretical model applied to the plasma variation of this volcanic
45 event. The periodicity does not show a clear correlation with the auroral power, central meridional
46 longitude, or Io phase angle. The periodic variation is continuously observed in addition to the auroral
47 modulation due to solar wind variation. Furthermore, Hisaki auroral data sometimes shows particularly
48 intense auroral bursts of emissions lasting <10 h. We find that these bursts coincide with peaks of the
49 periodic variations. Moreover, the occurrence of these bursts increases during the volcanically active
50 period. This auroral observation links parts of previous observations to give a global view of Jupiter's
51 magnetospheric dynamics.

52 **1. Introduction**

53 [2] Jupiter's huge magnetosphere shows quasi-periodic variations with periods of a few to several-day.
54 Long-term observation by the Galileo spacecraft shows the periodic variation of the energetic ion flux
55 and spectral slope in the vast magnetospheric region beyond $20 R_J$ (R_J is Jupiter's radius) and various
56 local time [e.g., Woch et al., 1998; Kronberg et al., 2009]. Periodic reconfiguration of the
57 magnetosphere between a loading phase involving thinning of the magnetospheric current sheet for ~2
58 days and an unloading phase associated with dipolarization of the magnetic field for ~1 days was
59 proposed by Woch et al. [1998]. Statistical analysis of the magnetic field observed by Galileo showed
60 a reconnection-like variation with 1–4 days intervals for some orbits [Vogt et al., 2010]. Polar-
61 integrated aurora observed by International Ultraviolet Explorer (IUE) showed variations by a factor of
62 2–4 in time scales of 5–10 days [Prangé et al., 2001]. They found that this periodic variation
63 corresponds to the variation of the magnetic field between quiet and disturbed days observed in
64 Jupiter's magnetotail by the Galileo magnetometer (MAG). Louarn et al. [2014] reported the

65 enhancement of auroral radio flux (hectometric emission, HOM) and the initiation of a radio source in
66 the Io plasma torus (IPT) (narrow-band kilometric emission, nKOM) almost simultaneously with the
67 periodic events of energetic ions and plasma injection features. The appearance and disappearance of
68 the auroral spot observed by Hubble Space Telescope (HST) in the poleward region of the dawnside
69 main emission also show variation with a period of 2–3 days [Radioti et al., 2010]. A similar spot has
70 been suggested to be a precursor of the auroral intensification [Gray et al., 2016].

71 [3] There are various periodicities with a similar time scale, but it is unknown what controls the
72 variation. Io's volcanic activity is one of the candidates via changing the plasma mass flux in the
73 magnetosphere. Io provides a massive plasma outflow consisting of sulfur and oxygen ions as main
74 contributors to the plasma pressure, which balances with the magnetic pressure in Jupiter's
75 magnetosphere. Kronberg et al. [2007] proposed a theoretical model to quantitatively explain the
76 variation of periodicity that was based on magnetic field and plasma observations. They found
77 theoretically that the time constant of the Jovian magnetosphere needed for mass loading until
78 reconnection onset decreases with increasing plasma mass flux, although this has not been confirmed
79 by observations. The contribution of solar wind variation is also under debate. Kronberg et al. [2008]
80 and Yao et al. [2019] suggested that the periodic variation seen in the energetic plasma flux and
81 magnetic field is independent of the solar wind variation. Vogt et al. [2019] analyzed plasma and
82 magnetic field observations by Galileo in Jupiter's magnetotail and suggested two types of variations,
83 (i) magnetospheric compression events due to variation of solar wind dynamic pressure and (ii) tail
84 reconnection and plasmoid release, most likely internally driven by the Vasyliunas cycle [Vasyliunas,
85 1983].

86 [4] The Hisaki Earth-orbiting space telescope monitors both the IPT and Jupiter's northern aurora
87 simultaneously [Yoshioka et al., 2013; Yoshikawa et al., 2014; Yamazaki et al., 2014; Kimura et al.,
88 2019]. Although the EXCEED (Extreme Ultraviolet Spectroscope for Exospheric Dynamics)
89 spectrometer cannot resolve the auroral structure due to its moderate spatial resolution (about 1 R_J at

90 Jupiter's opposition), it provides auroral spectra continuously for ~40 min during each 106-min orbit.
91 Since its launch in 2013, Hisaki has observed Jupiter for several months around its oppositions. In this
92 study, we analyze the periodic variation of the aurora observed by EXCEED using 2014–2016 data,
93 which includes monitoring a volcanically active event in 2015 [e.g., Yoshikawa et al., 2017; Tsuchiya
94 et al., 2018; Tao et al., 2018; Kimura et al., 2018] and solar wind during Juno's approaching phase.

95 **2. Observations and Data Procedure**

96 [5] The Hisaki observations and auroral analysis are outlined briefly here; for details of the
97 observations and data reduction, see Kimura et al. [2019], and for the analysis of the auroral spectra,
98 see Tao et al. [2016a; 2016b]. The northern auroral region is covered by the central thin part of a dawn-
99 dusk directed dumbbell-shaped 140 arcsec slit with an effective spatial resolution of 17 arcsec.
100 EXCEED detects part of the H₂ Lyman and Werner band emissions over the 80–148 nm wavelength
101 range with full width at half maximum (FWHM) resolution of 0.4 nm. The auroral signals within the
102 20 arcsec aperture of the slit width are integrated over specific wavelength ranges. The waveband
103 138.5–144.8 nm is used to estimate the total auroral emission and input power. The far-ultraviolet color
104 ratio (CR) is defined as the ratio of the intensity of the waveband absorbed least by atmospheric
105 hydrocarbons (138.5–144.8 nm) to that absorbed most (123–130 nm), which for EXCEED is defined
106 as CR_{EXCEED}. As the CR reflects the depth of the auroral electron precipitation into the hydrocarbon
107 layer, the auroral electron energy can be estimated assuming the atmosphere model. The total number
108 flux is derived from the electron energy and energy flux. The source current density can be estimated
109 with reference to the auroral electron acceleration theory [Tao et al., 2016b]. We analyze the
110 observation when the Jupiter northern aurora was facing to Earth, i.e., when the central meridional
111 longitude (CML) was 45–345° system III longitude. Since the northern auroral oval is non-
112 axisymmetric surrounding the magnetic pole, which is shifted from Jupiter's rotational pole, the
113 auroral power detectable from Earth varies with Jupiter's rotation. This power variation due to the
114 appearance is scaled by multiplying by the factor (auroral area integrated over all longitude)/(visible

115 auroral area at instantaneous CML), assuming a typical auroral location. The appearance-corrected
116 power is obtained as shown in Figure 1a. Auroral observation is integrated over 10 minutes to increase
117 the signal to noise ratio.

118 [6] Quasi-periodic variation of the aurora is detected automatically as follows. First, we obtain a
119 temporal sequence of the median of the power in the waveband 138.5–144.8 nm within a 0.5-day
120 window shifted by 0.25 day (green line in Figures 1a and 1b). Then we take a 3-point running average,
121 i.e., over 0.75 day (thick grey line in Figure 1b) and obtain its time deviation, $d(\text{Power})/dt$ (black line
122 in Figure 1c). We select events with positive $d(\text{Power})/dt$ with a duration of 0.5 day or more (orange
123 points in Figures 1b and 1c) and negative $d(\text{Power})/dt$ with a duration of 0.5 day or more (blue points
124 in Figures 1b and 1c). In order to exclude small perturbations, such as those around day of year (DOY)
125 25 in 2015 (Figure 1), whose amplitudes are insufficient to discuss the periodicity, only cases satisfying
126 $\Sigma(|d(\text{Power})/dt|) > 28$ GW/day summed over the positive and negative deviations are picked up. The
127 detected events are shown by vertical purple lines at the peak of each event in Figures 1a and 1b. After
128 excluding events with lacking data of ≥ 0.5 day in the interval, we obtain the temporal interval between
129 the brightness peaks of the quasi-periodic events (“ dt ” hereinafter).

130 [7] We also investigate the amplitude of the periodic variation and the existence of bursty auroral
131 brightening with short durations of < 10 h. The amplitude of each variation is estimated from the
132 difference between the maximum and minimum of the running averaged power, as shown by thick
133 black lines in Figure 2b. If the maximum value during each periodic brightness peak (diamonds in
134 Figure 2b) is above the maximum of the running average (green line) by 1.5σ or more, where σ is the
135 error estimated from the photon statistics, we label it as a periodic event with a significant auroral burst.
136 For example, enhancements on DOY ~4, 11, 15, and 17 in 2014 are detected as significant auroral
137 bursts as shown by red vertical lines in Figure 2b, while the others on DOY 1, 5, 21, and 23 are periodic
138 variations without significant bursts as shown by blue vertical lines.

139 [8] We compare the periodic variation with the external solar wind observed by Juno during its
140 approaching phase toward Jupiter. Solar wind dynamic pressure is considered to be an important
141 parameter that affects Jupiter's magnetosphere, as investigated in many studies [e.g., Vogt et al., 2019;
142 Nichols et al., 2017, Kita et al., 2019]. During Juno's solar wind plasma observation from May to July
143 2016, the continuity of Hisaki observation was not adequate for automatic analysis. Compressed
144 magnetic field structures of the interplanetary magnetic field (IMF) are often associated with
145 enhancements of solar wind dynamic pressure. We refer to the IMF observation by the magnetometer
146 (MAG) [Connerney et al., 2017] on board Juno for the solar wind information to cover January and
147 February 2016. We use MAG data with a time resolution of 60 s taken from the NASA Planetary Data
148 System (PDS) website. The IMF variation observed at the Juno spacecraft is shifted to the location of
149 Jupiter assuming a solar wind velocity of 400 km/s [e.g., Wilson et al., 2018] and a constant structure
150 during the solar rotation. This simple estimation is applicable since Juno was close to Jupiter, within
151 0.12 AU and 6.2° separation in heliospheric radius and longitude, respectively, for DOY 20–63 in
152 2016.

153 **3. Results**

154 [9] Figure 3 shows an overview of the dataset analyzed in this study. Hisaki is continuing its
155 observation of Jupiter's aurora (even now in August 2020), while we use highly continuous data until
156 the middle (DOY 241) of 2016 to detect the periodic variation automatically. Top plots show the
157 auroral power in the 138.5–144.8 nm band, which reflects the total input power. Detected periodic
158 variations are indicated by red or blue lines at their peaks in the top plots and their separation interval
159 dt is shown in middle plots. This dataset covers quiet (from DOY 1 in 2014 to DOY20 in 2015) and
160 large active volcanic event (DOY 20–100 in 2015) as seen in the variation of S^+ emission from the IPT
161 (bottom plots). Since some sporadic volcanic activities occurred in 2016, i.e., DOY ~140 [Kimura et
162 al., 2017; Tsuchiya et al., 2019], we exclude the 2016 dataset from the comparison of behaviors
163 between volcanically quiet and active time.

164 [10] Figure 4 shows a histogram of the separation interval of the auroral periodic variation dt . The
165 interval over the whole analyzed period varies in the range of 0.8–11.5 days with a peak at 2 days. The
166 analyses applied to Io's volcanically quiet (from DOY 1 in 2014 to DOY 40 in 2015) and active (DOY
167 40–140 in 2015) periods are shown by dotted and dot-dashed lines, respectively, which are
168 concentrated at a similar separation time. The mean and standard deviation during the quiet (active)
169 period are 3.0 (2.6) and 1.3 (1.0) days, respectively. For the quiet period, we excluded the extreme
170 event at $dt=11.5$ days. We use the Mann–Whitney U-test to investigate whether two independent
171 samples taken from non-normal populations have the same distribution. Excluding the extreme event,
172 the null hypothesis, i.e., (dt during the active period) = (dt during the quiet period), is not rejected by
173 the Mann–Whitney U-test (test statistics: $U=249.5$, $z=1.05$, $p=0.290$, sample size $n=40$). See Section
174 4.3 for the power analysis.

175 [11] An interesting finding from this analysis is that the auroral bursts sometimes occurred at the peaks
176 of the periodic variation, several examples of which are shown in Figure 2. The events on DOY ~4,
177 11, and 15 in 2014 are auroral bursts reported by Kimura et al. [2015]. The first two events were almost
178 simultaneously observed with HST. The auroral images taken by HST show low-latitude expansion
179 and blobs along the main aurora [Kimura et al., 2015; Badman et al., 2016]. These events were seen at
180 the peak of the periodic variation. There are also periodic variations that are not associated with
181 significant auroral bursts: e.g., DOY 1, 5, 21, and 23 in 2014 in Figure 2.

182 [12] The existence (red) and absence (blue) of the auroral bursts over the whole period shows
183 concentrations of the occurrence of these events, e.g., DOY ~10 and ~355 in 2014, 40–120 in 2015,
184 and 20–50 in 2016. On the other hand, the quasi-periodic variation is seen almost all the time. The
185 longest period in which the periodic variation coincided with the auroral burst, DOY 40–120 in 2015,
186 corresponds to Io's volcanically active event. The number of events associated with significant power
187 enhancements is 16 (17) within 39 (26) periodic variations, i.e., an occurrence ratio of 41% (65%), for
188 the volcanically quiet (active) period from DOY 1 in 2014 to DOY 40 in 2015 (DOY 40–140 in 2015).

189 [13] Figure 5 shows the relationship between the separation interval and geometric parameters and
190 auroral powers. There is not clear correlation between dt , CML (Figure 5a), I_0 phase angle (Figure 5b),
191 and the amplitude of the periodic variation (Figure 5c) which would reflect the size of magnetospheric
192 reconfiguration (see Section 4.2). The same analysis using different I_0 volcanic activity levels also
193 shows no clear correlation if the extreme event $dt > 8$ is excluded. On the other hand, we found a
194 significant positive correlation between the amplitude power and auroral burst power. The amplitude
195 power corresponds to the maximum difference of power within a periodic variation (e.g., the size of
196 thick black lines in Figure 2b), while the auroral burst power is the excess of auroral burst (e.g.,
197 diamonds in Figure 2b) from the peak power of the periodic variation. The correlation coefficient is
198 0.49 for the dataset using the whole period, and 0.64 and 0.69 for I_0 volcanically active and quiet times,
199 respectively.

200 [14] Superposed-epoch analysis is carried out for the observed power and the estimated parameters
201 from the spectral analysis. The timing of the power peaks of the quasi-periodic variation is set to
202 time=0 as enhancement is seen in the auroral power (Figure 6a). The mean value of all events within
203 each time bin is shown in red. If we exclude the periodic events associated with the auroral bursts, the
204 mean value (blue) at time=0 decreases, while this purely reflects the periodic variation. CR_{EXCEED}
205 shows a slight decrease from ~ 1.4 to ~ 1.3 ; this decrease is smaller than their variance ~ 0.4 (Figure 6b).
206 This decrease around time=0 is less clearly seen if the auroral burst events are excluded (blue, Figure
207 6b). In contrast to the variation in CR_{EXCEED} , the source current is enhanced from ~ 3 to ~ 7 nA/m² with
208 increasing auroral power (Figure 6c). Since the absolute values of these parameters vary among events,
209 we conducted similar analysis using the variation ratio of each parameter normalized by the initial
210 value of each periodic variation, as shown in Figures 6d–6f. Increasing and decreasing trends are more
211 clearly seen in the power and current density. This periodic variation of auroral power is mainly related
212 with the change in source current. The source current varies with the periodic variation by a factor of
213 ~ 1.6 (Figure 6f).

214 [15] We statistically investigate the durations of increasing and decreasing power over in the quasi-
215 periodic variation. The duration of increasing (decreasing) vary from 0.5 to 2.25 days (0.25 to 2.75
216 days) with mean and standard deviation values of 0.96 ± 0.39 (0.78 ± 0.52) days for the whole period as
217 shown in Figure 7a (Figure 7b). The difference between the durations of increasing and decreasing
218 auroral power is significant according to the Mann–Whitney U-test ($U=4401.5$, $z=4.03$, $p=5.4\times 10^{-5}$,
219 $n=225$). The histogram of the duration differences shows a slightly longer increasing period by -
220 0.17 ± 0.66 day on average (Figure 7c).

221 [16] Finally we show a comparison between the auroral power and the variation of the interplanetary
222 magnetic field from the Juno observation in Figure 8. The auroral power over wavelengths of 80–170
223 nm without absorption, estimated from the observation at 138.5–144.8 nm [Tao et al., 2016b], is shown
224 on the right y-axis. There are significant solar wind variations on DOY 22–27, DOY 39–43, and DOY
225 50–59 in 2016. The lower envelope of the auroral power, i.e., the background of the periodic peaks, is
226 correlated with the IMF variations. For example, the auroral power increases from 1.5 TW on DOY
227 22–23 to 3.5 TW on DOY 24–25 and then decreases to ~ 1 TW on DOY 30. The power variation trend
228 is similar to that of the IMF. Periodic variations are seen in addition to these variations, e.g., DOY 22,
229 23, and 25 in the first enhancement. These periodic variations are continuously observed in periods of
230 both quiet and enhanced IMF. From this observation, the auroral power amplitude associated with the
231 solar wind is estimated to be 1–3 TW. This is comparable with the typical amplitude of the periodic
232 variation of ~ 0.8 TW and that of the auroral burst of ~ 1 TW up to 6 TW, which are estimated from the
233 whole dataset.

234 **4. Discussion**

235 **4.1. Comparison with Other Studies for a Global View**

236 [17] Quasi-periodic variations have been reported for various parameters along with their
237 characteristics. We focus on the periodicity, asymmetric increasing and decreasing time durations, and

238 the time scale of auroral power variation. Here we compare our results with those of previous studies
239 and construct a global view based on the magnetospheric reconfiguration model proposed by Woch et
240 al. [1998] (Figure 9).

241 [18] The separation time of the periodic variation seen in the aurora is scattered over 0.8–8 days with
242 a peak at 2 days. This is comparable with previous reports, i.e., 5–10 days seen in the aurora by IUE
243 observation [Prangé et al., 2001], 1.5–7 days in plasma spectra [Kronberg et al., 2009] and in the
244 signatures of magnetic field stretching and depolarization [Kronberg et al., 2008], and 1–4 days in
245 magnetic-reconnection-like features [Vogt et al., 2010] and in wave power spectra [Vogt et al., 2019],
246 and ~ 3 days in both magnetic field and plasma taken by Juno [Vogt et al., 2020].

247 [19] Our observation shows increases for a duration of 0.96 ± 0.39 days and decreases for a duration of
248 0.78 ± 0.52 days. Asymmetric durations of increases and decreases were found by in-situ plasma
249 observations [e.g., Woch et al., 1998]. The decrease in the energetic ion flux and the increase in the
250 spectral slope take ~ 2 days, while the flux increases and the slope decreases within ~ 1 day with
251 disturbed features.

252 [20] To compare the intrinsic durations in detail, we also statistically investigated the duration of the
253 periodic variation in the energetic ions observed by Galileo using the dataset of Kronberg et al. [2009].
254 Referring to the time variation of the spectral index γ of energetic ion distributions observed by
255 Energetic Particle Detector (EPD) on board Galileo, intervals of increasing and decreasing spectral
256 index are detected for the 71 events from 1996 to 2002. As a result, we found that the duration of
257 increasing spectral index is 1.84 ± 0.97 days and the duration of decreasing spectral index is 1.24 ± 0.87
258 days (Figure 10). The difference between durations of spectral hardening and softening is significant
259 according to the Mann–Whitney U-test ($U=3695$, $z=4.79$, $p=1.65 \times 10^{-6}$, $n=141$). The difference, i.e.,
260 the duration of decreasing subtracted by the duration of increasing, is -0.60 ± 0.92 days. Therefore, the
261 significant asymmetry in intervals of increasing and decreasing is confirmed in both the auroral power

262 (Section 3) and energetic ion spectral index γ that is related to the thinning of the plasma sheet
263 [Kronberg et al., 2007], while the difference between the durations seen in aurora, -0.17 ± 0.66 day, is
264 still much smaller than the difference related to variation of energetic ion spectra.

265 [21] Magnetic field dipolarization and plasma sheet thinning have been observed with the periodic
266 variation of energetic particles [e.g., Kronberg et al., 2007, Vogt et al., 2020]. They also found that the
267 magnetic field ratio of the southward component to the radial component reaches the threshold for the
268 ion tearing instability at the end of the stretching phase. Energetic ion bursts were sometimes but not
269 always observed during this disturbed time. Yao et al. [2019] found the magnetic reconnection-like
270 features, probably linked to small-scale drizzle reconnection, occur during both loading and unloading
271 variation seen in magnetic field and plasma observed by Juno. According to their Figure 2, the
272 occurrence of the reconnection-like feature seems to be concentrated around the end of the stretching
273 phase and beginning of the dipolarization phase. Prangé et al. [2001] found magnetic field disturbance
274 in the magnetospheric tail around the peak of the auroral power. Interestingly, our Hisaki observation
275 sometimes detected auroral bursts, and we found in this study that they occur at the peak of the periodic
276 variation. These aurora bursts are associated with auroral blobs and low-latitude expansion of the main
277 auroral oval on the basis of auroral imaging by HST [Kimura et al., 2015; Badman et al., 2016], an
278 example of which is shown in Figure 9. Bonfond et al. [2012] reported a months-long expansion of the
279 main emissions at the same time as the occurrence rate of intense equatorward emissions strongly
280 increased in 2007. Yao et al. (accepted) reported that signatures of larger scale reconnection have been
281 related to large auroral brightening seen in the dawnside which is called dawn storms. These auroral
282 structures are considered to represent the Jupiter's reconfiguration events. The stretching of the
283 magnetosphere and energy exploration process in the tail region (e.g., reconnections) initiate auroral
284 bursts. Inversely, auroral bursts provide an opportunity for monitoring reconfiguration events.

285 [22] Note that the magnetospheric reconnection-like feature and in-situ ion bursts are observed several
286 times within one periodic variation [e.g., Kronberg et al., 2007, Yao et al., 2019]. This multiple feature

287 would be related with the several auroral spots which appearing and disappearing with a period of 2–
 288 3 days [Radioti et al., 2010]. On the other hand, the auroral burst observed by Hisaki’s polar-integrated
 289 view would be sum and/or their developed feature of them.

290 [23] In the following sections, we will quantitatively discuss the auroral variation and Io’s volcanic
 291 activity and solar wind effects within this global view.

292 4.2. Quantitative Analysis of Auroral Variation

293 [24] The results of superposed-epoch analysis shown in Figure 6 suggests that the periodic variation
 294 of auroral emission is associated with the increase in auroral source current. Tao et al. [2016b]
 295 quantitatively evaluated the variation of auroral emission due to (i) a magnetospheric compression and
 296 (ii) a change in the relative contribution of different components in the auroral structures as possible
 297 explanations of the auroral variation during solar wind compressions and/or plasma injections. On the
 298 other hand, the periodic variation in the global feature (Section 4.1) is considered to correspond to the
 299 plasma sheet thinning phase rather than the radial compression for (i). The change in the auroral
 300 components, (ii), is also unlikely to be the cause of this variation. Here we consider a quantitative
 301 estimation for this case of plasma sheet thinning.

302 [25] The source current density $j_{\parallel 0} (2.5/k_B T_0 [\text{keV}]) \propto N_0 T_0^{-1/2}$ (see Tao et al. [2016b] for details) is
 303 the current density conveyed without acceleration by magnetospheric electrons with density N_0 and
 304 temperature T_0 . Here we also assume adiabaticity, i.e., $PV^\gamma = \text{constant}$ with $\gamma = 5/3$, where $P =$
 305 $N_0 k_B T_0$ is the plasma pressure, V is the flux tube volume (i.e., the volume per unit magnetic flux), and
 306 k_B is the Boltzmann constant. From the mass conservation, $VN_0 = \text{constant}$, we obtain $j_{\parallel 0} \propto$
 307 $N_0 (N_0^{\gamma-1})^{-1/2} = N_0^{2/3}$. Referring to the observed ~ 1.6 -fold increase in the source current (Figure
 308 6f), the plasma density is estimated to increase by a factor of $1.6^{3/2} = 2.0$ and the pressure variation by
 309 a factor of $1.6^{3/2+1} = 3.2$. From the mass conservation, $VN_0 = \text{constant}$, the volume will be decreased by
 310 50%. This can be achieved by, for example, a change in the dimensions of the initial region from $\Delta 15$

311 R_J in the radial direction with width $\Delta 4 R_J$ in the north-south direction to those of $\Delta 30 R_J$ and $\Delta 1 R_J$
312 width, respectively, in the thinning phase at the similar radial distance.

313 [26] Kronberg et al. [2007] evaluated the magnetic field variation during the periodic variation from
314 the results of in-situ observation. They obtained a radial component of $B_r=3.5$ nT and a meridional
315 component of $B_\theta=1.1$ nT in the mass-unloaded phase and values of $B_r=4.5$ nT and $B_\theta=0.1$ nT in the
316 reconnection phase. This suggests an increase in magnetic pressure by a factor of 1.5. Some events
317 showed a variation from $B_r=3$ nT to 6 nT (Figure 1 of Kronberg et al. [2007]) resulting in the magnetic
318 pressure increasing by a factor of ~ 3.5 . The plasma thermal pressure is almost balanced with the
319 magnetic pressure in the Jupiter magnetotail, as also shown by Kronberg et al. [2007]. Note that the
320 magnetic field variation was observed at magnetotail $\sim 120 R_J$, while the auroral source current mainly
321 reflects the middle magnetosphere $\sim 30\text{--}50 R_J$. Referring to the periodic variation in the plasma
322 pressure investigated by Kronberg et al. [2008], the pressure varies by a factor of 2.5–5.5 at 30–60 R_J .
323 The 3.2-fold pressure enhancement estimated from this study is comparable with the observed
324 variation. Therefore the auroral periodic variation is quantitatively linked with the source current
325 variation due to magnetospheric plasma thinning and dipolarization.

326 **4.3. Modulations by Io Volcanic Activity: Periodicity**

327 [27] Our analysis does not show a significant difference in the periodicity of the volcanic activity of
328 Io. On the other hand, decreasing time constant of the Jovian magnetosphere needed for mass loading
329 with increasing plasma mass flux has been proposed by Kronberg et al. [2007] on the basis of a
330 quantitative relationship. Here we estimate the expected variation of the periodicity from the
331 relationship and its detectability using our dataset.

332 [28] Assuming a pressure balance with appropriate simplifications for the Jupiter magnetotail
333 environment, Kronberg et al. [2007] defined a parameter representing the plasma sheet topology. They

334 obtained the periodic time constant τ from the time variation of the parameter. One of their proposed
 335 relationships relating the τ with the plasma mass flux is as follows:

$$336 \quad \tau \simeq \frac{\rho_{rec} - \rho_0}{\dot{\rho}} \propto \frac{\delta n}{\dot{\rho}}, \quad (1)$$

337 where the ρ_{rec} and ρ_0 are the plasma mass density just before the reconnection and that at the start of
 338 the mass-loading phase, respectively; $\dot{\rho} = \dot{m}/V_{ps}$, where \dot{m} is the mass-loading rate and V_{ps} is the
 339 mass-loaded plasma sheet volume; and $\delta n = \frac{(\rho_{rec} - \rho_0)}{16m_p}$ is the number density, where m_p is the proton
 340 mass. For $\delta n = 0.05$, referring to the plasma observation by Frank et al. [2002], the time constant is
 341 estimated to be 6.5–1 days for the probable mass-loading rate of 100–600 kg/s and ~2.5 days for the
 342 most likely value of the mass-loading rate of 250 kg/s.

343 [29] Io's volcanic activity in 2015 was distinct from the past events seen in the sodium nebula reaching
 344 60 kR at 50 R_J compared with 20–25 kR before this event [Yoneda et al., 2015]. From IPT spectral
 345 analysis combined with a chemical model, it was found that the net production of S and O increases
 346 from 700 ± 130 kg/s to 3000 ± 300 kg/s (~4.3 times) and the electron density increases from 2350 ± 340
 347 cm^{-3} to 2860 ± 260 cm^{-3} (~1.2 times) at ~6 R_J around the peak of the volcanic event compared with a
 348 quiet time [Yoshioka et al., 2018]. Their analysis also suggested that plasma outflow velocity increases
 349 by ~3.4 times during the volcanically active time. Hikida et al. [2020] applied the plasma diagnosis
 350 method to the Hisaki data with the 140 arcsec slit and obtained a similar electron variation from
 351 1790 ± 80 /cc to 2400 ± 100 /cc (~1.3 times) during the volcanic event. An analytic method considering
 352 conservations of the magnetic flux and energy in the interchange motion at the IPT associated with the
 353 IPT emission observed by Hisaki suggests an increase in the plasma mass-loading rate from 300 to
 354 500 kg/s (1.66 times) during this volcanic event [Kimura et al., 2018]. Auroral spectral analysis
 355 combined with the auroral particle acceleration theory suggests that the source plasma density around
 356 the middle magnetosphere also increases from 0.0019 to 0.0027 / cm^3 (1.4 times) [Tao et al., 2018].

357 Increases in the plasma density and mass-loading rate by factors of 1.2–1.7 are estimated from these
358 various methods.

359 [30] For the variation of the mass-loading rate from 300 to 500 kg/s [Kimura et al., 2018], relationship
360 (1) with $\delta n = 0.05$ corresponds to a decrease in the time constant from 2 to 1.2 days. The difference
361 between the maximum and minimum values is 0.8 day. For the increase in the plasma density and
362 mass-loading rate by a factor of 1.2–1.7, the decrease in the time constant is ~83–60%. If the time
363 constant at the volcanically quiet condition is 3 days, that at the active condition is expected to be in
364 the range of 2.5–1.8 days. The difference between quiet and active conditions is 0.50–1.2 days.

365 [31] Here, we analyzed the power, i.e., probability to detect the significant difference correctly, of our
366 test using the `wmwpow` package (ver. 0.1.2, R). This package evaluates the exact power of the Mann–
367 Whitney U-test using a Monte Carlo approach [Mollan et al., 2019]. The obtained detection power was
368 0.83 (0.81) with a potential difference of 1.25 (1.2) days for the event number of our dataset, which is
369 comparable to a generally acceptable value of 0.8. From our analysis, the difference between the
370 volcanic quiet ($dt=3.0$ day) and active conditions ($dt=2.6$ day) was 0.4 days (Section 3), which is less
371 than 1.25 day. Therefore, a difference of greater than 1.25 days is unlikely to exist between the active
372 and quiet conditions. This also indicates that our dataset is not adequate for detecting a difference of
373 less than 1.25 days. The expected difference of 0.50–1.2 days for this volcanic event is beyond this
374 detection ability.

375 [32] The power for a smaller difference improves with increasing number of samples. If the observed
376 separation times on DOY 10–200 in 2016 are added as the quiet period, the number of samples for the
377 quiet time increases to 43. With the 19 samples during the volcanically active time, the dataset has a
378 large power (0.889) for detecting a difference of 1.2 days but insufficient power to detect a difference
379 of 0.5 days (power of 0.261) according to the `wmwpow` analysis. In addition, the obtained mean values
380 of the two groups become closer, 2.65 and 2.62, for the quiet and active conditions, respectively.

381 [33] Therefore, a significant difference in periodicity between volcanically quiet and active conditions
382 is not derived from our dataset. From the detection analysis, we cannot conclude whether no difference
383 exists or whether a difference of less than 1.25 days exists. Further observations are expected to answer
384 this remaining question.

385 **4.4. Modulations by Io Volcanic Activity: Auroral Burst**

386 [34] The occurrence of aurora bursts increased significantly during enhanced volcanic activity as also
387 previously reported [Yoshikawa et al., 2017; Tsuchiya et al., 2018; Kimura et al., 2018, Tao et al.,
388 2018]. In addition, a new finding in this study is the correlation between the auroral burst power and
389 the power of the periodic amplitude. This correlation indicates that the explosion of the magnetospheric
390 power is related to the activity of the background periodic variation. These bursts are considered to be
391 the main contributor to the plasma mass release via magnetospheric reconnection.

392 **4.5 Modulations by Solar Wind**

393 [35] As seen in the comparison of the periodic variation obtained with Juno's IMF observation, the
394 periodic variation continues under solar wind compression events. This supports the independent
395 periodic variation of the energetic particle flux and spectral slope proposed by Kronberg et al. [2008]
396 and Vogt et al. [2019]. From the statistical analysis using the plasma and magnetic field datasets
397 measured by Galileo, Vogt et al. [2019] found that increases in the solar wind dynamic pressure are
398 statistically associated with magnetospheric compression events while tail reconnection and plasmoid
399 release are most likely internally driven by the Vasyliunas cycle. Our results of auroral observation
400 also reflect these two characteristic dynamics. As shown in Figure 8, the increasing trend of the auroral
401 base over several days closely reflects the variation of the IMF strength. This power modulation is
402 probably due to magnetospheric compression. Similar auroral variation was reported in Hisaki
403 observation by e.g., Kita et al. [2016] and Tao et al. [2016b], referring to the solar wind variation
404 estimated by model [Tao et al., 2005]. Using HST image taken in May-June 2016, Nichols et al. [2017]

405 reported that main emissions and duskside poleward region are brightened during the solar wind
406 compressions observed by Juno. On the other hand, the quasi-periodic variation and auroral bursts at
407 these peaks sometimes correspond to the auroral reconnection and plasmoid release as discussed in
408 Section 4.1. Our dataset of polar total auroral power is unique in its reflection of both types of dynamics.
409 The relative contribution of both dynamics to the total power is derived from this study, i.e., the
410 intrinsic periodic variation provides ~ 0.8 TW amplitude with an auroral burst of 1–6 TW and is
411 comparable to the 1–3 TW contribution from solar wind variation. This auroral power modulated by
412 solar wind is comparable with those observed in May–June 2016 [e.g., Gladstone et al., 2017, Nichols
413 et al., 2017, Kita et al., 2019].

414 **5. Summary**

415 [36] We have investigated the quasi-periodic variation of polar-*integrated* auroral power with a period
416 of a few to several days using observation by the Hisaki space telescope from the end of 2013 to the
417 middle of 2016. From our analysis, we obtained the following results.

418 [37] (1) The detected periodicity of the auroral power is 0.8–8 days with a peak at 2 days. The
419 increasing duration of the periodic auroral variation is slightly but significantly longer than the
420 decreasing duration on average, as seen with the in-situ plasma observation by Galileo.

421 [38] (2) Significant difference in the periodicity depending on the volcanic activity for the active period
422 in early 2015 was not detected in our dataset, partly because of the insufficient amount of data to detect
423 the expected difference from the theoretical estimation applied for this volcanic event. On the other
424 hand, our dataset suggests that a difference greater than 1.25 days is unlikely to exist between the
425 volcanically active and quiet conditions, which is consistent with the expected difference estimated
426 from a proposed relationship applied to the plasma variation of this volcanic event.

427 [39] (3) The periodic variation is mainly caused by the total auroral electron flux variation rather than
428 the averaged auroral energy variation. This variation is associated with magnetospheric thinning by
429 quantitative comparison with the in-situ observation.

430 [40] (4) Auroral bursts within short durations <10 h and a large amplitude were sometimes found at
431 the peaks of the periodic variation. A positive correlation was found between the auroral burst power
432 and the periodic amplitude. The occurrence of the auroral bursts was 41% of periodic peaks during the
433 volcanically quiet time, which increased to 65% during the volcanically active time.

434 [41] (5) The periodic variation associated with the auroral bursts was continuously seen when solar
435 wind structures hit the magnetosphere. The variation associated with solar wind is 1–3 TW, the
436 periodic variation is ~0.8 TW, and the auroral burst varies from ~1 TW to 6 TW.

437 [42] The time variation of the aurora suggests a link to other previous observations and theoretical
438 models associated with the magnetospheric reconfiguration. Remaining and newly proposed questions
439 for future works are as follows. Which spatial component(s) of the aurora is responsible for the periodic
440 variation? Does the periodicity depend on the variation in the plasma density and/or the mass-loading
441 rate? What determines the occurrence and absence of the bursts? For the third question, one possibility
442 is the amount of accumulated plasma [e.g., Kimura et al., 2018], and another is the geometry of the
443 plasma sheet and its condition towards reconnection-associated instabilities. Why is the asymmetry of
444 the increasing and decreasing durations in auroral power less than that of the periodic variation of in-
445 situ energetic ions? The reflection of global regions in auroral observations compared with the locality
446 for in-situ observations and/or the time variation between Hisaki and Galileo observations might be
447 related to this difference.

448 [43] These Hisaki observations provide a total power variation without resolving auroral spatial
449 distribution as achieved by Juno and HST. In spite of limited spatial resolution, this study revealed that

450 this Hisaki dataset can monitor the global internal dynamics of periodic variations and associated
451 auroral bursts.

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456 NASA PDS website (<https://pds.nasa.gov/>). Auroral images was taken from observations made with
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