

# The Energy Decay of Warm-core Eddies in the Gulf of Mexico

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## Key Points:

- Time evolution of the energy of Warm-core rings in the Gulf of Mexico is assessed using empirical methods and satellite altimetry.
- The vast majority of mechanical energy (kinetic plus available potential) is lost early in the eddies life cycles, far from the western boundary.
- Wind-current feed back effects such as Ekman buoyancy flux and wind stress work play an important role in energy conversion and decay.

## Abstract

The Gulf of Mexico (GoM) is home to some of the most energetic eddies in the ocean. They detach from the Loop-Current and drift through the basin, transporting large amounts of heat and salt. These eddies, known as Loop Current rings (LCRs) have a crucial role in the GoM's dynamics and in the weather of the eastern US, and this role is largely conditioned by their longevity and decay properties. Here, we use an empirical method to estimate the energy evolution of all LCRs detached since 1993. We found that, contrary to the commonly accepted idea that LCRs conserve their energy as they drift through the GoM and decay suddenly against the western platform, LCRs' energy decays faster in the eastern basin, and they typically lose three-quarter of their energy before encountering the continental shelf. We also show that wind-current feedback largely contributes to the energy decay and conversion.

## Plain Language Summary

Ocean eddies can be long-lived and carry large amounts of heat and salt across ocean basins and marginal seas. This is the case of Loop Current rings (LCRs), which are large warm-core eddies drifting through the Gulf of Mexico (GoM). Understanding how these eddies lose their energy is key to understand their longevity and transport properties. Here, we use a previously validated empirical method based on *in situ* observations to estimate the time evolution of LCRs energy using satellite observations. We show that LCRs decay continuously during their life cycle, contrary to the previously accepted idea that they decay when collapsing against the western GoM's continental shelf. LCRs have already lost three-quarter of their energy before they reach any topographic obstacle. Using wind observations, we also show that wind-current interactions are key to the energy loss of these eddies.

## 1 Introduction

Coherent eddies can carry tracers across oceanic basins and participate in the transport of water-masses that impact large scale circulation and climate as well as ecosystems. For instance, Loop Current rings (LCRs), detaching from the Loop Current (LC), carry warm subtropical underwater (SUW) originating from the Caribbean through the Gulf of Mexico (Elliott, 1982; Leben, 2005) and have a direct impact on the basin's water mass properties (Vidal et al., 1994; Hamilton et al., 2018; T. Meunier et al., 2020), hurricane intensification (Shay et al., 2000; Jaimes et al., 2016), and thunderstorm occurrence east of the Rocky mountains (Molina et al., 2016). Similarly, Agulhas rings carry anomalously warm and salty Indian Ocean water through the South Atlantic and participate in the upper limb of the Atlantic Meridional Overturning Circulation (Beal et al., 2011; Biastoch et al., 2008; Marsh et al., 2007). The ability of such mesoscale eddies to be efficient in tracer transport, as well as the geographical characteristics of the heat, salt, or biogeochemical properties redistribution they induce, is directly controlled by their longevity and energy decay properties.

Although the processes responsible for the formation of LCRs have been (and remains) the focus of intense research (e.g. Candela et al. (2002); Oey et al. (2003); Lugo-Fernández et al. (2016); Donohue et al. (2016); Le Hénaff et al. (2023)), the processes responsible for their decay have received less attention. As of now, the decay of LCRs has been largely attributed to two processes : vortex-splitting (Forristall et al., 1992; Biggs et al., 1996; Lipphardt et al., 2008) and topographic interactions along the western GoM's continental slope, which was consequently nicknamed the *eddy graveyard* (Biggs, 1992; Vidal et al., 1994; Hamilton et al., 1999). However, these studies were all mostly descriptive and limited to a small number of eddies, and did not provide large-number statistics on the decay of LCRs based on a solid metric such as energy.

Using 25 years of satellite altimetry and a convenient empirical relationship between sea surface height (SSH) and heat content, T. Meunier et al. (2020) showed that, on average, LCRs have already lost over two thirds of their heat content before they reach the so-called *eddy graveyard*, and that heat decays at an inverse exponential rate right from the start of LCRs' life cycle, so that some other processes have to be invoked for the decay of LCR's heat content.

Beyond vortex-splitting and topographic effects, many processes can participate in the decay of a mesoscale eddy. These include Frontal instability (Brannigan et al., 2017; Pérez et al., 2022), Interaction with submesoscale eddies (de Marez et al., 2020; Jouanno et al., 2016; Tedesco et al., 2019), Mesoscale straining (Mariotti et al., 1994), Layering and double diffusion (Schmitt et al., 1986; Armi et al., 1989; Meunier et al., 2015; Middleton et al., 2021), interaction with internal gravity waves (Kunze, 1986; Polzin, 2008; Joyce et al., 2013), surface heat fluxes (Dewar, 1987), or wind-current interactions (Dewar & Flierl, 1987; Duhaut & Straub, 2006; Renault, Molemaker, Gula, et al., 2016).

Of all these processes, wind-current interactions in the form of current feedback on the wind stress are of special interest, since the latter was shown numerically to yield a reduction of 20 to 35 % of eddy kinetic energy in eddying currents (Duhaut & Straub, 2006; Renault, Molemaker, Gula, et al., 2016; Renault et al., 2017). This process, based on the simple idea that wind stress depends on relative wind speed in a frame of reference moving with the current, rather than absolute wind speed, has consequences in the wind-stress distribution over a mesoscale eddy. The wind stress is increased where the current opposes the wind, and decreased where the current flows in the direction of the wind. This results in a systematic negative wind stress work integrated over an eddy's surface, which extracts kinetic energy (KE) (Dewar & Flierl, 1987), as well as Ekman pumping which induces a negative buoyancy flux that converts available potential energy (APE) into KE (Gaube et al., 2015; Wilder et al., 2022). While these processes were studied in depth in eddying current systems using regional numerical models (Renault, Molemaker, Gula, et al., 2016; Renault et al., 2017), their impact on individual mesoscale eddies remains largely unknown apart from the extremely idealized studies of Dewar and Flierl (1987) and Wilder et al. (2022), which suggest a significant impact of current-feedback on numerical, isolated, idealized eddies. However, observational evidence of the impact of current-feedback on mesoscale eddies are still lacking.

Recently, T. Meunier et al. (2022) applied an empirical method known as the gravest empirical mode method (GEM) (Watts et al., 2001; Sun & Watts, 2001) to reconstruct the daily three-dimensional temperature and salinity structure of all LCRs detached in the GoM between 1993 and 2022. Their method was validated against *in situ* glider observations, and showed a striking accuracy (coefficient of determination  $R^2$  greater than 0.93 between the GEM-reconstructed and the directly observed fields).

In this paper, we take advantage of this validated reconstruction method to estimate the energy decay properties, in time and space, of 40 LCRs detached between 1993 and 2023. Using scatterometer and reanalysis wind products, we also estimate the effect of current feedback on the decay and conversion of the eddies' energy. Beyond a regional study of some particular class of eddy, to the best of our knowledge, this work is the first systematic observation-based statistical study of the energy decay of mesoscale eddies, as well as the first observation-based estimate of the relative impact of wind-current interactions on the decay and conversion of eddies' energy. The results of this study could therefore provide some insight on the processes controlling mesoscale eddy decay in the ocean.

## 2 Data

This work is largely based on altimeter-derived sea surface height (SSH) observations. We use daily AVISO absolute dynamic topography (ADT) gridded fields. The grid has a  $1/4^\circ$  degree resolution and the data is available from January 1st 1993 until now.

We also use nearly 7000 Argo profiles distributed over the entire deep GoM (Fig. 1a). Most of these floats were launched as part of the National Academies of Sciences Understanding Gulf Ocean Systems (UGOS) program.

The main wind product used is IFREMER CERSAT Global Blended Mean Wind Fields, which combines scatterometer observations with ECMWF operational wind analyses. The product is available on a  $1/4^\circ$  grid every 6 hours. A 24 hours averaging was performed to coincide with the ADT observations. Detailed information as well as a retrospective study of the product's performance can be found in Desbiolles et al. (2017). Because Scatterometers estimate wind from the sea-state, and the latter depends on relative wind speed rather than absolute wind speed, their wind products are known to retain some signal of the current feedback (Plagge et al., 2012). Although this effect is attenuated by the gridding process, we compared the scatterometer product's results with data from ECMWF's ERA5 reanalysis, which provides estimates of the absolute wind. ERA5 is based on a 4DVAR ensemble data assimilating atmospheric model and is distributed as hourly outputs on a  $1/4^\circ$  grid. Here, a 24 hours averaging was also performed. Full details about the methods can be found in Hersbach et al. (2020).

For both wind products, wind stress was computed using the COARE3.5 parameterization (Edson et al., 2013).

### 3 Methods

#### 3.1 Theoretical background

In this study, we seek to describe the time evolution of the total mechanical energy of individual eddies, which is the three-dimensional integral of energy density over the eddy's volume.

The evolution equation for kinetic energy density ( $E_k = \frac{1}{2}\rho|\mathbf{u}|^2$ ) reads (Gill, 1982):

$$\frac{DE_k}{Dt} + \rho'gw = -\nabla \cdot P'\mathbf{u} + \mathbf{u} \cdot \frac{\partial \boldsymbol{\tau}}{\partial z} + \nabla \cdot \mathbf{K}\nabla E_k + \rho\epsilon, \quad (1)$$

where  $E_k$  is kinetic energy,  $\frac{D}{Dt}$  is the material derivative,  $\rho'$  is density anomaly referenced to a minimum potential energy profile,  $w$  is vertical velocity,  $P'$  is pressure anomaly,  $\mathbf{u}$  is the velocity vector,  $\boldsymbol{\tau}$  is the wind stress,  $z$  is the vertical coordinate,  $\mathbf{K}$  is a diagonal diffusivity coefficient tensor and  $\epsilon$  is the energy dissipation rate.

The second term on the left-hand side is the buoyancy flux, and is equal to the material rate of change of available potential energy density (Holliday & McIntyre, 1981):

$$\frac{DE_p}{Dt} = \rho'gw, \quad (2)$$

so that the left-hand side of equation (1) represents the material rate of change of total mechanical energy density ( $E_m = E_k + E_p$ ). Still following Holliday and McIntyre (1981), available potential energy density is defined as:

$$E_p = -\int_0^\eta g\tilde{\eta} \frac{\partial \rho_r}{\partial z}(z - \tilde{\eta})d\tilde{\eta}, \quad (3)$$

where  $\rho_r$  is the reference minimum potential energy profile and  $\eta$  is the isopycnal displacement, relative to the reference profile. In this work, we study the effects of the (relative) wind on energy decay and energy conversion, and we compute the buoyancy flux of equation (2) using the non-linear Ekman pumping vertical velocity (Stern, 1965) :

$$w_e = \frac{1}{\rho_0} \nabla \times \left( \frac{\boldsymbol{\tau}}{f + \zeta} \right) \quad (4)$$

Note that computing the non-forced vertical velocity using the Omega equation was also considered, but, since the estimated fields are not necessarily solutions of any equation

of motion, and given the relatively low spatial resolution, the omega-derived vertical velocity is essentially noise, with zero-mean vertical velocity.

As mentioned above, the wind stress to be considered here is the relative wind stress, which takes into account the feedback of the current and reads :

$$\boldsymbol{\tau} = \rho_a C_d |\mathbf{u}_a - \alpha \mathbf{u}_g| (\mathbf{u}_a - \alpha \mathbf{u}_g), \quad (5)$$

where  $\rho_a$  is air density,  $C_d$  is the drag coefficient computed using the COARE3.5 parameterization (Edson et al., 2013),  $\mathbf{u}_a$  is the absolute wind velocity 10 m above the sea surface and  $\mathbf{u}_g$  is the velocity of the surface current, computed from the altimetry-derived SSH using the geostrophic balance assumption.  $\alpha$  is a coefficient applied to the current velocity to account for the feedback of the stress increase/reduction by the current-feedback through frictional effects, that results in a reduction/increase of the absolute wind speed (Renault, Molemaker, McWilliams, et al., 2016). We used values of 1 and 0.7, following empirical results of Renault et al. (2019).

We are interested in quantifying the evolution of the total energy ( $\mathbb{E}_m$ ) transported by an eddy whose boundary is defined by the closed line  $\mathcal{C}$  which encircles the surface  $\mathcal{S}$  :

$$\mathbb{E}_m = \iint_{\mathcal{S}} \int_{-H}^0 E_m dz ds, \quad (6)$$

where  $ds$  is a surface element and  $H$  is the eddy's thickness. We do not make the hypothesis that the boundary  $\mathcal{C}$  is a material line. Integrating the left hand side of equation (1) in its flux form, using the Ostrogradski theorem, the Leibniz theorem, and the Reynolds transport theorem, we get an exact equation for the evolution of total mechanical energy, under the assumption that the thickness does not vary :

$$\begin{aligned} \frac{d}{dt} \mathbb{E}_m = \int_{-H}^0 \left\{ \underbrace{\oint_{\mathcal{C}} (\mathbf{u}_c - \mathbf{u}) \cdot \mathbf{n} E_m dl}_{(a)} - \underbrace{\oint_{\mathcal{C}} \mathbf{u} \cdot \mathbf{n} P' dl}_{(b)} + \underbrace{\iint_{\mathcal{S}} \mathbf{u} \cdot \frac{\partial \boldsymbol{\tau}}{\partial z} ds}_{(c)} \right. \\ \left. + \underbrace{\oint_{\mathcal{C}} \mathbf{K} \nabla E_k \cdot \mathbf{n} dl}_{(d)} + \underbrace{\iint_{\mathcal{S}} \rho \epsilon ds}_{(e)} \right\} dz \quad (7) \end{aligned}$$

The first term on the right hand side represents the energy flux through the eddy's boundary that is caused by the relative flow in the eddy's moving referential. In the case of a Lagrangian coherent vortex, the boundary  $\mathcal{C}$  is a material line, so that it is exactly advected by the flow ( $\mathbf{u}_c = \mathbf{u}$ ) and the term vanishes. the second term is the work of the pressure force. Under the geostrophic approximation, it is null whatever the boundary. The third term is the wind stress work. Since the wind stress is dependent on the relative wind speed, even an homogeneous wind will exert some work, whose integral over the eddy's surface will be negative, whatever the vorticity sign (Dewar & Flierl, 1987). The viscous terms (d) represent all the unresolved (sub-grid scale) advective turbulent processes that might cause an energy flux through the eddy's boundary (which is only defined using coarse resolution observations) and the dissipation term (e) is essentially related to small scale turbulence yielding conversion of mechanical energy to internal energy.

Because we are interested in examining and comparing the energy evolution of a set of eddies of different sizes, we need to use a normalized metric. The surface-averaged energy is a convenient variable:

$$\overline{E}_m = \frac{\mathbb{E}_m}{S}. \quad (8)$$

Integrating equation (7) with respect to time and dividing by the instantaneous area of the eddy, we get an equation for the surface-averaged energy density.

$$\overline{E}_m(t) = \frac{1}{S} \int_0^t \int_{-H}^0 \{(a) + (b) + (c) + (d) + (e)\} dz d\tau \quad (9)$$

Since our purpose is to compute statistical properties of energy decay by ensemble-averaging over all the detached eddies, using the surface-averaged energy density is convenient. Ensemble averaging using the integrated total energy would introduce a bias by increasing the weight of large eddies in the average. Normalizing total energy by each eddy's area avoids this bias. More interestingly, normalizing energy by the eddy's area also removes the effects of energy loss due to direct loss or gain of area (hence mass), as can happen during filamentation or splitting/merging events. For the sake of concision, in the rest of the paper, the terms KE, APE and TE will be used to designate the surface averaged kinetic, potential and total mechanical energy density. When studying the impact of the current-modified wind stress on the eddies' energy, the two variables we will focus on are the surface-averaged, time-integrated wind stress work, normalized by the TE's initial value of each eddy that we will casually refer to as the wind stress work (WSW), and the surface-averaged, time-integrated Ekman buoyancy flux, normalized by the initial value of APE of each eddy, that we will casually refer to as the Ekman buoyancy flux (EBF) :

$$\text{WSW}(t) = \frac{1}{\overline{E}_m(0)} \frac{\int_0^t \iint_{\mathcal{S}} \boldsymbol{\tau}(\tilde{t}) \cdot \mathbf{u}_g(\tilde{t}) ds d\tilde{t}}{S(t)}, \quad (10)$$

$$\text{EBF}(t) = \frac{1}{\overline{E}_p(0)} \frac{\int_0^t \int_{-H}^0 \iint_{\mathcal{S}} \rho' g w_e ds dz d\tilde{t}}{S(t)}. \quad (11)$$

### 3.2 The gravest empirical mode method (GEM)

To estimate the daily 3D structure of temperature and salinity (hence geostrophic velocity, KE and APE density) from satellite altimetry, we use an empirical method known as the gravest empirical mode projection (GEM) (Watts et al., 2001; Sun & Watts, 2001; T. Meunier et al., 2022). It consists of establishing an empirical relationship between the vertical thermohaline structure of the ocean and the dynamic height to build transfer functions that associate one single value of temperature and salinity for each couple {pressure, SSH}. The procedure used here was used and validated in the Gulf of Mexico and is described in details in T. Meunier et al. (2022). The mean yearly transfer functions for salinity and temperature are shown in Figure 1c and d, respectively : the downward sloping of the isotherms and of the subsurface salinity maximum with increasing SSH, associated with the subtropical underwater (SUW) carried by LCRs is evident. Note that, to account for the seasonality of surface conditions, which affects the accuracy of the three-dimensional reconstruction in the top 200 dbar, the GEM fields were constructed on a monthly basis. An example of APE and KE density cross-section reconstructed using the GEM along with an average LCR's SSH anomaly is shown in Figure 1e and f, respectively. APE is concentrated in the core of the eddy, while KE is stronger near the edges, where density gradients are sharper. The APE maximum is 2.5 times larger than the KE maximum and, when integrated over the whole LCR, APE largely dominates over KE (T. Meunier et al., 2022).

### 3.3 Eddies' edge definition

In this work, we aim to follow the whole LCR, and not only its coherent part, which is usually confined to a small portion of the core (Beron-Vera et al., 2018; Andrade-Canto et al., 2020). Hence, rather than tracking so called Lagrangian coherent vortices (Haller & Beron-Vera, 2013; Beron-Vera et al., 2013), the eddy as we wish to track it should be defined as a compact rotating body of water detached from the LC, regardless of its *a priori* conservation properties. The detection and tracking method is based on ADT. First we define the LC's edge as the smallest-value ADT contour linking the Yucatan channel and the Florida strait. This contour usually encloses closed ADT contours, which are the signature of unborn LCRs. We define LCR detachment as the instant at which none of the closed contours remain enclosed within the LC anymore. Once detached, to define the LCR's boundary, we use a mixed Eulerian-Lagrangian procedure : at  $t=0$  (detachment date), we initialize a cluster of virtual Lagrangian particles within the outermost closed SSH contour, which coincides with the boundary of the water mass detached from the LC. These particles are distributed on a  $1\text{ km} \times 1\text{ km}$  regular grid. They are then integrated in the geostrophic surface velocity field inferred from the altimetry product, using a fourth-order Runge-Kutta scheme (Butcher, 1996). The particle concentration is then computed and interpolated on a regular  $1/48^\circ$  grid. The LCR's edge is defined as the concentration contour with the original value at  $t=0$ , that is, the closed isopleth that encloses a compact surface of constant concentration equal to the original concentration. A sequence of particle concentration maps during the drift of LCR Poseidon in April and November 2016 is shown on figure 1g and h. The edge contour based on the concentration criterion is compared to two other Eulerian criteria : the Maximum velocity contour and the last closed SSH contour. While particle loss through filamentation of the eddy is evident, concentration within the eddy's boundary remain largely homogeneous. It is important to note that, while the method is based on Lagrangian particles integration, it does not belong to the class of *Lagrangian methods*, that seek coherent eddy boundaries, such as the null geodesics rings of Haller and Beron-Vera (2013) and Beron-Vera et al. (2018). Here, the boundary does not ensure the coherence of the eddy, in the sense that tracer conservation is not guaranteed (and our detected eddies actually do loose tracer). However, contrary to Lagrangian Coherent Vortices detected using proper *Lagrangian methods*, this method allows us to follow the entire body of water detached from the LC, and not only a small fragment of it. The initial boundaries of all eddies are shown in Figure 1b, where the SSH at the eddies' centers is color-coded.

## 4 Results

Individual trajectories of the 40 LCRs are shown on Figure 2a. TE, normalized by the initial values at the time of detachment, is color-coded. While a few LCRs maintain high levels of energy ( $>80\%$ ) past the Campeche bank, west of  $93^\circ\text{W}$ , the vast majority start losing energy soon in their life cycle in the eastern GoM. The ensemble average of normalized TE for all LCRs is shown in Figures 2b. The circles' size represents the average diameter of LCRs during their drift. After 200 days, the average TE left in the eddies is only 26 % of the initial value at detachment. This value is of 21 and 28 % for KE and APE, respectively (not shown).

The evolution of the mean KE, APE and TE (non-normalized by initial values) is shown against time and longitude on figure 2c. The 95% confidence interval, computed using the bootstrap method, is shown as the light shaded areas. APE largely dominates over KE during the entire eddies' life cycle. On average, both KE and APE start decaying right from the first days after detachment and keep on decaying continuously with time until 200 days, resulting in a similar decay of TE. The average halving time for KE, APE and TE is of 101 days, 120 days, and 117 days, respectively. Looking at the longitude-dependence of energy, a similar pattern appears : APE and KE start decaying in the eastern basin near  $88.5^\circ\text{W}$  and decay continuously across the entire GoM. The halving lon-



gitude is of  $91^\circ\text{W}$  for KE and  $91.6^\circ\text{W}$  for APE and TE. The energy decay rates (Figure 2d) not only confirm that energy loss occurs all along the eddies' drift through the GoM, but also that LCRs lose APE (and TE) at a faster rate in the eastern GoM, during the first month of their life cycle. KE does not exhibit this increased decay rate in the eastern basin, and decays regularly all along the eddies life cycle.

Since previous works have pointed out some critical longitude ( $\approx 92\text{--}93^\circ\text{W}$ ) where eddies seem to decay at a faster rate, mostly because of instability and eddy-splitting (direct loss of mass; see Lipphardt et al. (2008) and references therein), we also investigated the evolution of the total LCRs energy (the energy density integrated over the full eddy's surface and not the energy per unit area discussed above) as well as the evolution of the eddies area, which would clearly indicate regions of more frequent splitting. The surface integrated KE, APE and TE, normalized by their initial values and ensemble-averaged are shown against time and longitude in Figures 2e and f, respectively. As for the surface-averaged KE, TE and APE, the surface integrated KE, TE and APE smoothly and regularly decay with time during the whole LCRs life cycle. Similarly, the decay against longitude shows no obvious discontinuity, except for a slightly faster decay between  $90$  and  $91^\circ\text{W}$ . The evolution of the ensemble-averaged LCRs areas against time and longitude shows the same regular decay, without any evident discontinuity.

Equation (7) shows that for a mesoscale geostrophic eddy, where the work of the pressure force is negligible, TE can only decay through the action of wind stress work, an energy flux through the boundary, and diffusivity. As mentioned above, the energy flux through the boundary is impossible to estimate, since the boundary's velocity is undefinable (because it is not defined by a series of material points, or material line), and the diffusive term can not be estimated with satellite altimetry and Argo data only. On the other hand, the effects of wind stress work can be assessed using wind observations. Similarly, the decay of APE is directly linked to buoyancy fluxes (Equation 3), and while the vertical velocity can not be fully estimated, we still can estimate the contribution of Ekman pumping to it (Equation 4). In other words, while this work does not allow for a full budget of the energy equation, the effects of relative wind stress can be assessed effectively. Figure 2g compares the average evolution of TE (normalized by initial values) to the effects of wind stress work for 4 different wind stress parametrization : absolute wind stress obtained from scatterometer products (SCAT abs), relative wind stress obtained from scatterometer products and computed by removing the full current velocity from the absolute wind value (SCAT rel), relative wind stress obtained from the ERA5 reanalysis and computed by removing the full current velocity from the absolute wind value (ERA rel), and relative wind stress obtained from scatterometer products and computed by removing 70 % of the current velocity to the absolute wind value (SCAT rel07). Relative wind stress work appears to contribute to the decay of the LCRs' TE, with an average of 18% and 15% of the original TE extracted when using Scatterometer and ERA5 wind products, respectively. This corresponds to respectively a quarter and a fifth of the energy lost by LCRs in 200 days. When using Renault et al. (2019)'s parameterization which uses a current velocity reduced by 30% to compute relative wind stress (SCAT rel07), we find that, on average, wind stress work then only accounts for about 10 % of the TE loss in LCRs. In all three cases, energy decay through wind stress work is faster in the beginning of the LCR's life cycle, and after about 120 days, wind stress work does not extract energy any more. For comparison, the work of the absolute wind was also computed. Surprisingly, it represents an energy source for LCRs with an average increase of the original TE by about 20% after 200 days. This shows that the mean spatial distribution of absolute wind in the GoM tends to increase LCRs energy, showing that the energy sink is really related to wind-current interactions.

Time evolution of APE is compared to the effects of Ekman buoyancy fluxes in figure 2h. The wind parameterizations shown are the same as for Figure 2g. The close coincidence of the decay of APE and that expected from Ekman buoyancy fluxes computed from scatterometer-derived relative wind stress is striking. It suggests that Ekman buoyancy fluxes might account for the whole observed APE decay. Using ERA5 relative wind



stress, we find that Ekman buoyancy flux accounts for a slightly reduced part of the APE decay (70 % of the APE loss after 200 days), and using the reduced current parameterization for relative wind stress computation (SCAT abs07), Ekman buoyancy flux accounts for about 65% of the total APE loss in 200 days. Using absolute wind stress, we find that Ekman buoyancy flux corresponds to an APE increase of nearly 20% during the first 170 days, followed by a small decrease in the last 30 days.

## 5 Discussion and conclusion

Using an observation-based method, we assessed the statistical properties of the energy decay of all LCRs detached since 1993. To the best of our knowledge, this is the first time such statistics on the energy decay of mesoscale eddies are obtained anywhere in the World ocean based on an observations.

We showed that LCRs' energy decay is a fast process occurring continuously during the whole eddies' life cycles. This observation is in contradiction with the commonly accepted idea that LCRs steadily drift through the GoM, retaining their hydrographic properties until they collapse against the western platform, as explicitly shown in the numerical simulations of Romanou et al. (2004), and suggested in a number of other numerical studies listed in Lipphardt et al. (2008) (Sturges et al., 1993; Dietrich et al., 1997; Kantha et al., 2005). However, it should be pointed out that our results do not contradict the idea of an *eddy graveyard*, but clearly show that the decay of LCRs does not happen only there, but starts right from the first month after detachment from the LC and continues all along their life cycle. As they reach the so-called graveyard, LCRs are already old and weak eddies that have typically lost over 3/4 of their energy. Topographic effects in the western basin can therefore not be considered a major cause of LCRs decay, but rather the ultimate process dispersing their remnants. Moreover, the energy decay rate was shown to be faster in the eastern basin soon after detachment than in the western basin, and the halving longitude is found as far East as 91.5°W.

Our results also contradict another long-standing claim that there exists some critical longitude in the GoM, between 92 and 93°W, beyond which LCRs suddenly decay at a fast rate, mostly through low wavenumber instability (the observed elliptization of the eddies seems consistent with the development of an azimuthal mode 2 vortex Rossby wave) yielding vortex-splitting (Hamilton et al., 1999; Vukovich, 2007; Lipphardt et al., 2008). Although we acknowledge that LCRs might become unstable, loose coherence or split as directly observed by Biggs et al. (1996), so that increased decay rate might occasionally occur in the central basin, our results show no sign of a critical longitude between 92 and 93 °W (nor anywhere) where LCRs loose mass or energy at a faster rate, and the decay of the eddies' area and surface-integrated KE, APE and TE is a gradual process occurring all through their life cycles.

Beyond the description of the decay of energy in LCRs, we also investigated the possible role of wind-current feedback on that decay. We showed that the effect of relative wind stress work is non-negligible, but likely not a leading order mechanism in the decay of kinetic energy (and total mechanical energy) as it accounts for 25 and 20 % of the total energy loss in 200 days when using scatterometer and ERA5 wind fields, respectively. When using Renault et al. (2019)'s parameterization to account for the diminution of absolute wind when the current opposes the wind (and *vice versa*) (by reducing the current velocity by 30 % when computing relative wind stress), we found that wind stress work then only accounts for 10 % of the total energy loss in 200 days.

However, the impact of current-feedback was shown to be important for the decay of APE, which would be entirely driven by Ekman buoyancy flux when using relative wind stress computed from scatterometer data. Beyond this striking result, the control of APE loss by Ekman buoyancy flux has implications on the decay of KE. Buoyancy flux does not extract TE: it converts APE to KE (it is often referred to as the baroclinic conversion term), which means that all along the LCRs life cycle, KE is fueled by this conversion. Given the modest role of wind stress work, this means that some other pro-

cesses, such as turbulent diffusivity at the edges of the eddies (subgrid-scale advective processes that are unresolved by the altimetry grid) might be important in the decay of LCRs' TE through the decay of KE.

Of course, our study does not allow estimation of the impact of direct energy flux through the eddies' boundary which is related to the fact that the edges we consider are not material lines (term (a) in equation 7). This is clearly a caveat of not using Lagrangian coherent boundaries, which would ensure the flux term to be zero, and would allow us to estimate the diffusive term as it would become the only unknown of Equation (7). The choice not to use such Lagrangian metrics was motivated by the need to estimate the evolution of the full detached patch of LC water, and not only a small piece of its core. In particular, relative wind stress work becomes important near the edge of the eddies, where the velocity is maximum, while long-horizon coherent LCRs usually exclude this peripheral part. However, working with three-dimensional Lagrangian objective eddy framing (Beron-Vera et al., 2013; Haller & Beron-Vera, 2013) would be complementary to this study. Although it would only allow to estimate the properties and fate of the inner core of LCRs, it could allow to study more accurately the impact of Ekman buoyancy fluxes in baroclinic conversion. Such study is currently in progress.

Note that, while this study is focused on the GoM, similar processes are expected to occur in other warm-core rings (e.g. Agulhas rings, Gulf stream rings, Mozambic channel rings, North Brazil Current rings, Kuroshio rings etc.) and a similar study, dedicated to the decay of Agulhas rings is also in progress.

## 6 Open Research

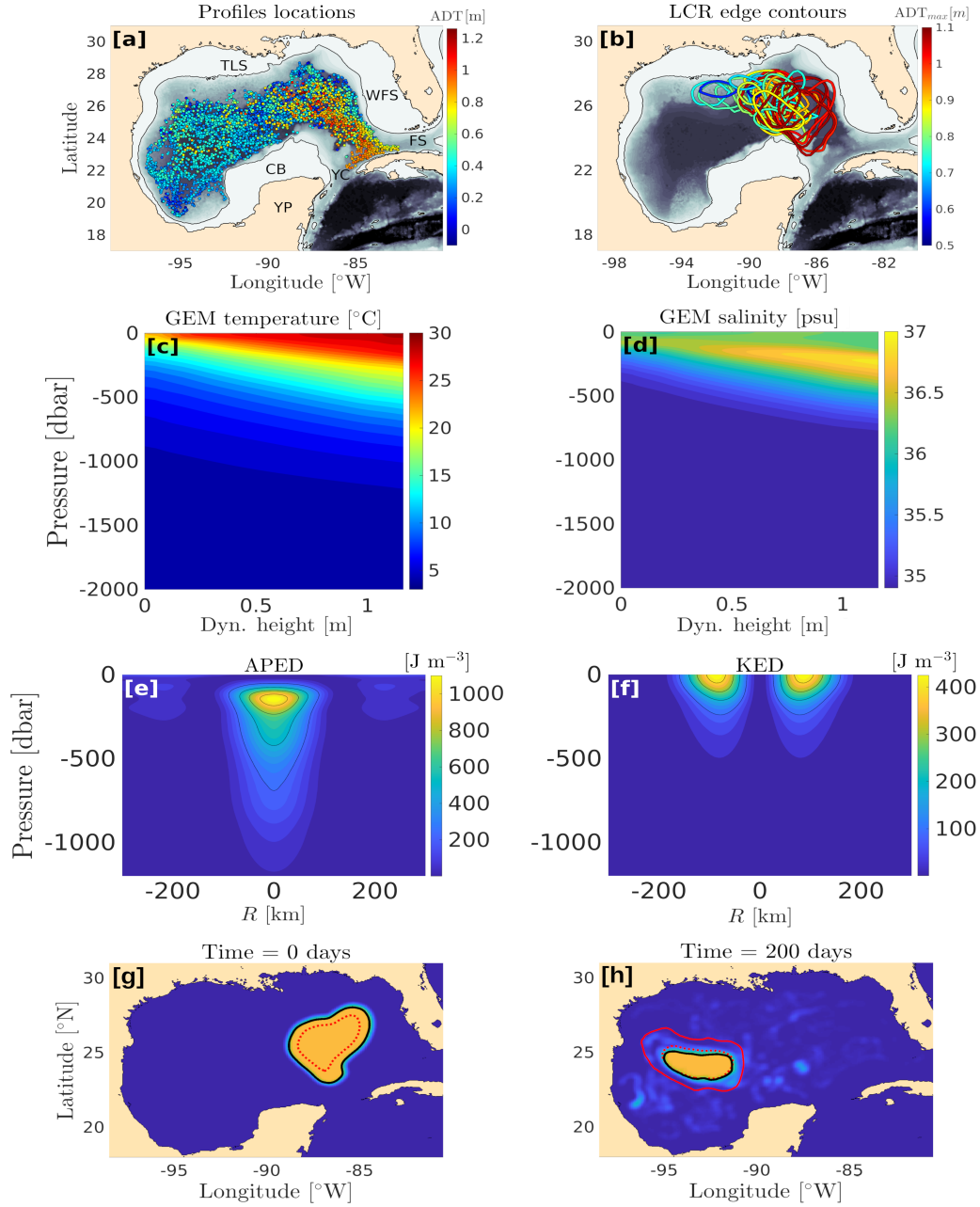
All data used in this study are publicly available. The gridded scatterometer wind product and the absolute dynamic topography SSH product are available from Copernicus (<https://data.marine.copernicus.eu>). Direct link to the wind product dataset is also available at <https://www.pigma.org/geonetwork/srv/api/records/85c907d3-98fc-4ce7-b7e4-7332aa3fe660>. Direct link to the ADT product is also available at [https://data.marine.copernicus.eu/product/SEALEVEL\\_GLO\\_PHY\\_CLIMATE\\_L4\\_MY\\_008\\_057/download](https://data.marine.copernicus.eu/product/SEALEVEL_GLO_PHY_CLIMATE_L4_MY_008_057/download). Data is available after users create a Copernicus account. Argo data are available from any data assembly center (e.g. <https://dataselection.euro-argo.eu/>), specifying the geographical limits of  $[-99 -80]^{\circ}W$  and  $[17 31]^{\circ}N$ .

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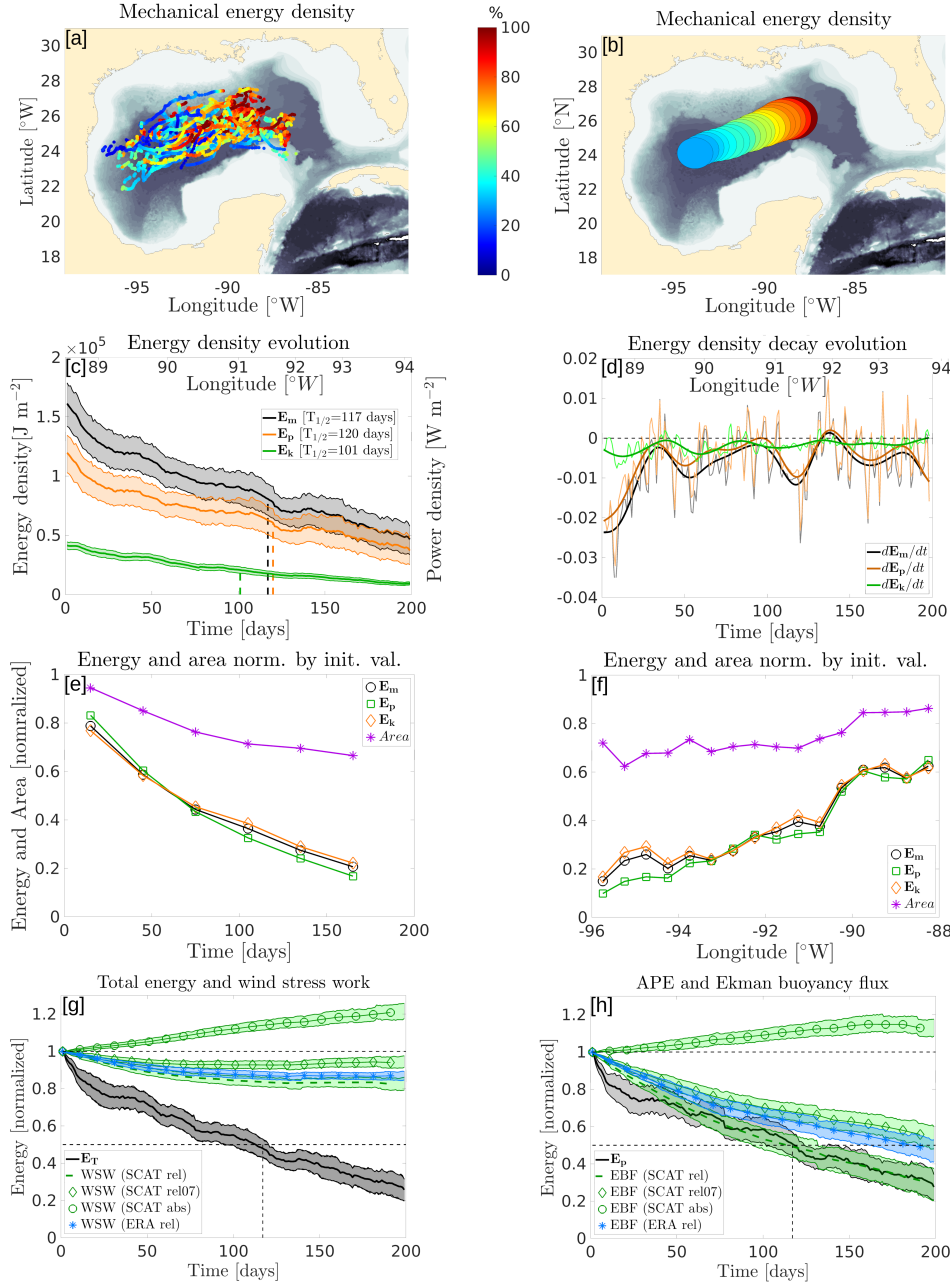
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**Figure 1.** a: Location of the ARGO profiles used to build the Gravest Empirical Modes (GEM) for the Gulf of Mexico (GoM). The color code represents the local Absolute Dynamic Topography (ADT). The names of the GoM's main topographic features are indicated. YC stands for Yucatan channel, FS stands for Florida strait, WFS stands for West Florida shelf, CB stands for Campeche bank, YP stands for Yucatan peninsula, and TLS stands for Texas-Louisiana shelf. b: Edge contours on the first day after detachment for all the Loop Current Rings (LCR) detached between 1993 and 2023. The color code represents ADT at the center of the eddy. c: GEM transfer function for temperature. The x-axis is dynamic height, the y-axis is pressure, and the color map is temperature. d: Same as (c) for salinity. e: Available potential energy density section across an average LCR reconstructed using the GEM method. f: same as (e) for kinetic energy density. g: Initial particle concentration in Loop Current Ring Poseidon on detachment date. The plain and dashed red lines represent the last closed ADT contour and the maximum velocity contour, respectively. Normalized passive tracer concentration is color-coded (yellow is 1 and dark blue is zero). The concentration-based edge contour of Poseidon is materialized as a thick black line. h: Same as (g) after 200 days—11—



**Figure 2.** a: Trajectories of the center of the 40 Loop Current Rings (LCR). Total mechanical energy (TE) normalized by its initial value is color-coded. b: Ensemble-averaged TE normalized by its initial value. The position and diameter of the circles represent the average position and size of the LCRs. c: Time evolution of the average energy density. The black, orange and green lines represent total mechanical energy (TE), available potential energy (APE), and kinetic energy (KE), respectively. The light shaded areas around the lines represent the 95% confidence interval. The vertical dashed lines represent the energy halving time. The average longitude corresponding to the time is indicated on the top x-axis. d: Time evolution of the ensemble averaged rate of change of energy density (time derivative of energy density). The smooth dark thick lines represent the 30-days low pass filtered values, while the thin light lines represent the raw results. The color code is the same as in panel (c). e: Ensemble average of the non-surface-averaged KE (orange diamonds), TE (black circles) and APE (green squares) and eddy's area (purple stars) normalized by their initial values against time. f: Same as (e) against longitude. g: Energy extraction by the effects of wind stress work. The dashed line represents wind stress work computed from the scatterometer data. The green diamonds represent wind stress work computed by reducing the current by a factor of a 0.7 when computing relative wind. The red circles represent the effect of wind stress work computed using the absolute wind stress. The blue stars represent the effect of wind stress work computed using the ERA5 reanalysis. Time evolution of the TE is indicated as a black line. h: Energy conversion by the effect of the Ekman buoyancy flux. The Color code is the same as in panel (g). Time evolution of APE is indicated as the black line.

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