

1 **Including filter-feeding gelatinous macrozooplankton in**  
2 **a global marine biogeochemical model: model-data**  
3 **comparison and impact on the ocean carbon cycle**

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10 **Key Points:**

- 11 • FFGM large carcasses and fecal pellets affect the balance between particulate ex-  
12 port and remineralization of total POC in the upper ocean  
13 • FFGM contribution to deep total POC export increases with depth to reach 70%  
14 at 5000 m while they contribute to 6% at 100 m.  
15 • FFGM-driven POC fluxes have a particular spatial structure as FFGM better ex-  
16 ploit low productivity environments than other macrozooplankton

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**Abstract**

Filter-feeding gelatinous macrozooplankton (FFGM), namely salps, pyrosomes and doliolids are increasingly recognized as an essential component of the marine ecosystem. Unlike crustacean zooplankton (eg., copepods) which feed on prey that is an order of magnitude smaller, filter-feeding allows FFGM access to a wider range of organisms, with predator over prey ratios as high as  $10^5:1$ . In addition, most FFGM produce carcasses and/or fecal pellets that sink 10 times faster than those of copepods. This implies a rapid and efficient export of organic matter to depth. Even if these organisms represent  $<5\%$  of the overall planktonic biomass, the induced organic matter flux could be substantial. Here we present a first estimate of the influence of FFGM organisms on the export of particulate organic matter to the deep ocean based on a marine biogeochemical earth system model: NEMO-PISCES. In this new version of PISCES, two processes characterize FFGM: the preference for small organisms due to filter feeding, and the rapid sinking of carcasses and fecal pellets. To evaluate our modeled FFGM distribution, we compiled FFGM abundances observations into a monthly biomass climatology using a taxon-specific conversion. FFGM contribute strongly to carbon export at depth ( $0.4 \text{ Pg C yr}^{-1}$  at 1000m), particularly in low-productivity region (up to 40% of POC export at 1000m) where they dominate macrozooplankton by a factor of 2. This export increases in importance with depth, with a simulated transfer efficiency close to one.

**Index terms and keywords**

Gelatinous zooplankton, Large pelagic tunicates, Filter-feeders, particulate carbon export, biogeochemical model

**1 Introduction**

Pelagic tunicates, i.e., salps, doliolids, pyrosomes and appendicularians, are free-swimming open ocean gelatinous zooplankton that are increasingly recognized as key-components of marine ecosystems and biogeochemical cycles (Henschke et al., 2016; Luo et al., 2020). All pelagic tunicates, with the exception of appendicularians, are part of the macrozooplankton (2-20 mm), and are filter-feeding organisms. They will be referred to hereafter as filter-feeding gelatinous macrozooplankton (FFGM). Although they are not part of the same phyla, FFGM which are urochordates share functional and morphological similarities with ctenophores and cnidarians (jellyfish). They have therefore been placed in the functional group of gelatinous zooplankton (GZ): FFGM are indeed water-rich free-swimming transparent animals.

The fragility of all GZ bodies partly explains the rarity of observations (Henschke et al., 2016). Nevertheless, it has been hypothesized that increasing anthropogenic pressures on the global ocean favor gelatinous zooplankton in most regions due to eutrophication, overfishing, or climate change (A. J. Richardson et al., 2009; Purcell, 2012). Research effort focusing on GZ have increased dramatically during the last two decades, particularly on cnidarians ("true-jellyfish") that contribute significantly to biological carbon cycling through "jelly-falls" events (ie. the accumulation of gelatinous zooplankton carcasses in the water column following a swarming event; Lebrato et al., 2012; A. K. Sweetman et al., 2014; A. Sweetman & Chapman, 2015; Luo et al., 2020). Similarly, many recent studies have focused on pelagic tunicates (namely salps (e.g. Phillips et al., 2009; Henschke et al., 2020; Henschke, Cherel, et al., 2021; Henschke, Blain, et al., 2021; Luskow et al., 2020; Ishak et al., 2020; Stone & Steinberg, 2016) , appendicularians (e.g. Berline et al., 2011) and doliolids (e.g. Stenvers et al., 2021)), revealing their importance in carbon cycling and for ecosystem structure, at least on a regional scale. Yet, despite this growing interest, Their importance on global scale remains uncertain.

65 Pelagic tunicates are capable of swarming, which means that their population can  
66 reach a high abundance in a very short time and can therefore represent a significant part,  
67 or even dominate, the zooplankton community during massive proliferation events (Everett  
68 et al., 2011; Henschke et al., 2016). Three mechanisms have been hypothesized to trig-  
69 ger these swarms: i) FFGM use a mucus structure to filter feed, which gives them ac-  
70 cess to a wide range of preys, from bacteria to mesozooplankton (Acuña, 2001; Suther-  
71 land et al., 2010; Bernard et al., 2012; Ambler et al., 2013; Sutherland & Thompson, 2022)  
72 This feeding strategy might allow them to proliferate in response to the bloom of a wide  
73 variety of organisms, in contrast to typical zooplankton with prey-to-predator size ra-  
74 tios ranging from 1:10 to 1:100 (B. Hansen et al., 1994). ii) FFGM generally have high  
75 clearance and growth rates (Alldredge & Madin, 1982; Henschke et al., 2016) that pro-  
76 mote rapid proliferation. The densest FFGM swarms can sweep over 200% of their res-  
77 ident water volume per day (Ishak et al., 2020). iii) Some FFGM, such as salps, have  
78 life cycles characterized by the alternation between a sexual phase (the blastozoid) and  
79 an asexual phase (the oozoid). During the asexual phase, oozoids produce long chains  
80 of blastozooids clones that can number several hundreds individuals and give rise to swarm-  
81 ing processes (Loeb & Santora, 2012; Kelly et al., 2020; Groeneveld et al., 2020). Based  
82 on their potential to form large swarms, FFGM can significantly affect ecological pro-  
83 cesses, at least locally.

84 FFGM could also have an impact on the ocean carbon cycle. Indeed, many FFGM  
85 produce fast sinking carcasses and/or fecal pellets that induce a very efficient carbon ex-  
86 port during swarming events (Henschke et al., 2016). Large fecal pellets and carcasses  
87 of salps are carbon-rich (more than 30% of dry weight (DW)) and sink at speeds up to  
88 2700m d<sup>-1</sup> for fecal pellets and 1700m d<sup>-1</sup> for carcasses (Henschke et al., 2016; Lebrato  
89 et al., 2013). In areas where salps proliferate, they can induce a carbon transfer to the  
90 seafloor 10 times faster than in their absence (Henschke et al., 2016). For pyrosomes, knowl-  
91 edge on their impact and the nature of their carcasses and fecal pellets remains very lim-  
92 ited (Décima et al., 2019). Intense carcass fall events have been described as responsi-  
93 ble for large carbon exports due to their high carbon content (35% DW, one of the high-  
94 est among GZ) (Lebrato & Jones, 2009). Although their fecal pellets sink 30 times slower  
95 than those of large salps (70m d<sup>-1</sup> Drits et al. (1992) vs 1700m d<sup>-1</sup> (Henschke et al.,  
96 2016)), they are able to export a significant amount of carbon in combination with ac-  
97 tive transport through diurnal vertical migrations (Stenvers et al., 2021; Henschke et al.,  
98 2019). Because of their rapidly sinking fecal pellets (over 400m/d) and high clearance  
99 rates, doliolids also affect carbon fluxes (Takahashi et al., 2013, 2015; Ishak et al., 2020)  
100 but their impact remains poorly documented.

101 Overall, most studies to date have focused on the regional scale. But Luo et al. (2020)  
102 have estimated the contribution to the global carbon cycle of three categories of gelati-  
103 nous zooplankton: ctenophores, cnidarians and pelagic tunicates. Using a data-driven  
104 carbon cycle model, they found that pelagic tunicates contribute three quarters of the  
105 particulate organic carbon (POC) flux induced by gelatinous zooplankton or one quar-  
106 ter of the total POC exported at 100m. A more recent study by the same team (Luo et  
107 al., 2022) revised this estimate to 0.57 Pg C yr<sup>-1</sup>, representing 9% of total export past  
108 100 m, by explicitly representing FFGM in the Cobalt-v2 biogeochemical model (FFGM  
109 refer to Large pelagic tunicates in their study).

110 Marine biogeochemical models have repeatedly shown their usefulness in under-  
111 standing marine processes on a global scale: in particular on the role of plankton in ecosys-  
112 tem processes (e.g. Saillely et al., 2013; Le Quéré et al., 2016; Kearney et al., 2021) and  
113 biogeochemical fluxes (e.g. E. Buitenhuis et al., 2006; Kwiatkowski et al., 2018; Aumont  
114 et al., 2018). Their complexity has been greatly increased by the addition of multiple  
115 limiting nutrients and multiple functional groups or size classes of phytoplankton and  
116 zooplankton (e.g. Le Quéré et al., 2005; Follows et al., 2007; Ward et al., 2012; Aumont  
117 et al., 2015). In particular, Plankton Functional Type (PFT) models have been intro-

118 duced as a way of grouping organisms that keeps overall biological complexity at a man-  
 119 ageable level (Moore et al., 2001; Gregg et al., 2003; Le Quéré et al., 2005). Wright et  
 120 al. (2021) showed that the introduction of a jellyfish PFT (cnidarians only) into the PLANK-  
 121 TOM model has a large direct influence on the biomass distribution of the crustacean  
 122 macrozooplankton PFT and indirectly influences the biomass distributions of protozoo-  
 123 plankton and mesozooplankton through a trophic cascade. This influence could be ex-  
 124 plained by the specific diet of jellyfish that differs from other zooplankton PFTs. Sim-  
 125 ilarly, due to their specific filter feeding mode, their likely significant role in carbon cy-  
 126 cling via carcasses and fecal pellet falls, and their potentially large biomass via swarm-  
 127 ing processes, the inclusion of FFGM as a new PFT in a PFT-based model is relevant  
 128 and has been recently achieved by Luo et al. (2022).

129 Here, we use the PISCES-v2 model (Aumont et al., 2015) which is the standard  
 130 marine biogeochemistry component of NEMO (Nucleus for European Modelling of the  
 131 Ocean).

132 In this study, a new version of PISCES was developed (PISCES-FFGM) in which  
 133 two new PFTs were added: a generic macrozooplankton (GM) based on an allometric  
 134 scaling of the existing mesozooplankton and a filter-feeding gelatinous macrozooplank-  
 135 ton (FFGM). Two processes characterize the FFGM in this version of the model: access  
 136 to a wide range of prey through filter feeding and rapid sinking of carcasses and fecal  
 137 pellets. We first examine how the model succeeds in reproducing the surface distribu-  
 138 tion of FFGM by providing a new compilation of abundance observations converted to  
 139 carbon biomass via taxonomy-specific conversion functions to make this assessment. Sec-  
 140 ond, because the modeling study by Luo et al. (2022) focused on the impact of FFGM  
 141 on surface processes, we investigated these same impacts to investigate whether our mod-  
 142 eling framework and formulations produce results consistent with theirs. Our study pro-  
 143 vides also some new insights: 1) we explore the FFGM-specific spatial patterns of or-  
 144 ganic matter production, export and particles composition in the top 100 m; 2) we in-  
 145 vestigate the impacts of FFGM on the export of particulate organic carbon to the deep  
 146 ocean via an explicit representation of fast-sinking fecal pellets and carcasses.

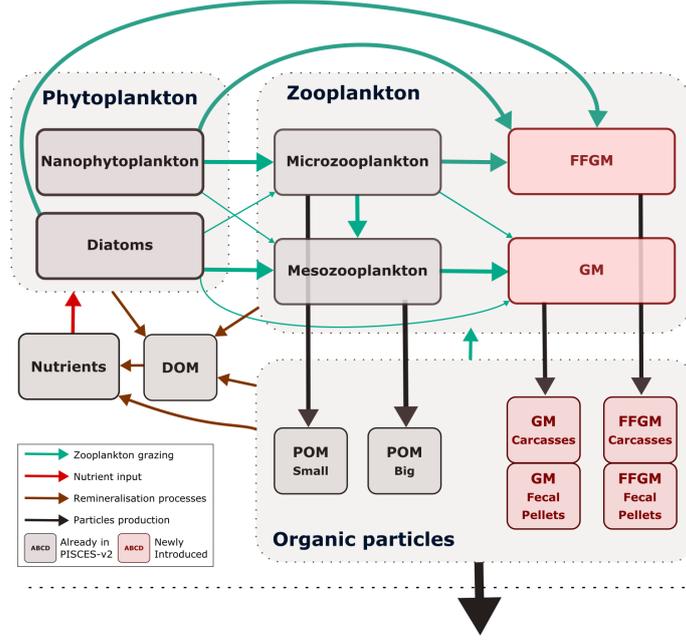
## 147 2 Materials and method

### 148 2.1 Model description

#### 149 2.1.1 Model structure:

150 The marine biogeochemical model used in the present study is a revised version of  
 151 PISCES-v2 (gray boxes in fig. 1). It includes five nutrient pools ( $Fe$ ,  $NH_4^+$ ,  $Si$ ,  $PO_4^{3-}$   
 152 and  $NO_3^-$ ), two phytoplankton groups (Diatoms and Nanophytoplankton, denoted  $D$   
 153 and  $N$ ), two zooplankton size classes (Micro- and Mesozooplankton, denoted  $Z$  and  $M$ )  
 154 and an explicit representation of particulate and dissolved organic matter, reaching a to-  
 155 tal of 24 prognostic variables (tracers). A full description of the model is provided in (Aumont  
 156 et al., 2015).

157 In the version used here, two groups of macrozooplankton were added, one corre-  
 158 sponding to generic macrozooplankton organisms (hereafter referred to as GM, see fig.  
 159 1) and the other to salp-like filter-feeding gelatinous macrozooplankton organisms (here-  
 160 after referred to as FFGM, see fig. 1). As with micro- and mesozooplankton in the stan-  
 161 dard version of PISCES, the C:N:P stoichiometric composition of the two macrozooplank-  
 162 ton groups is assumed to be constant. In addition to their carbon biomass, two additional  
 163 tracers were introduced into the model for each macrozooplankton group correspond-  
 164 ing to fecal pellets and carcasses in carbon units, respectively (GM Carcasses, GM Fe-  
 165 cal Pellets, FFGM Carcasses and FFGM Fecal Pellets, see fig. 1). Because both macro-  
 166 zooplankton groups have a constant Fe:C stoichiometry and feed on phytoplankton that  
 167 have a flexible Fe:C stoichiometry (Eq. 16 to 20 in (Aumont et al., 2015)), two compart-



**Figure 1. Architecture of PISCES-FFGM.** This figure only shows the organic components of the model omitting thus oxygen and the carbonate system. This diagram emphasizes trophic interactions (turquoise arrows) as well as particulate organic matter production (black arrows), two processes strongly impacted by the introduction of two new zooplankton groups in PISCES-FFGM (pink boxes). FFGM is for Filter-Feeding Gelatinous Macrozooplankton, GM is for Generic Macrozooplankton, POM is for Particulate Organic Matter, DOM is for Dissolved Organic Matter.

168 ments representing the iron content of the fecal pellets of the two macrozooplankton groups  
 169 were added. Figure 1 summarizes the tracers and interactions newly introduced into PISCES  
 170 for this study (referred to as PISCES-FFGM hereafter).

171 The tracers considered for particulate and dissolved organic matter are (organic  
 172 particles in fig. 1): *sPOC* which refers to small organic carbon particles, *bPOC* which  
 173 refers to large organic carbon particles, *DOC* which refers to dissolved organic carbon,  
 174 *DIC* which refers to dissolved inorganic carbon,  $Ca_{FFGM}$  which refers to the carbon con-  
 175 tent of FFGM carcasses,  $Fp_{FFGM}$  which refers to the carbon content of FFGM fecal pel-  
 176 lets,  $Ca_{GM}$  which refers to the carbon content of GM carcasses and  $Fp_{GM}$  which refers  
 177 to the carbon content of GM fecal pellets.

### 178 **2.1.2 Macrozooplankton (FFGM and GM) dynamics**

179 We first present the generic equation describing the dynamics of the two groups  
 180 of macrozooplankton, and then focus on the modeling choices we made to differentiate  
 181 the two groups of organisms. All symbols and definitions are summarized in Table 1.

182 The temporal evolution of the two compartments of macrozooplankton is governed  
 183 by the following equation:

$$\frac{\partial X}{\partial t} = e^X G_X (1 - \Delta(O_2)) f_X(T) X - (m^X + m_c^X) f_X(T) (1 - \Delta(O_2)) X^2$$

$$-r^X f_X(T) \left( \frac{X}{K_m + X} + 3\Delta(O_2) \right) \quad (1)$$

Symbol	Description
<b>I. STATE VARIABLES</b>	
$P$	Nanophytoplankton
$D$	Diatoms
$Z$	Microzooplankton
$M$	Mesozooplankton
$GM$	GM
$FFGM$	FFGM
$Ca_{FFGM}$	FFGM Carcasses
$Fp_{FFGM}$	FFGM Fecal Pellets
$Ca_{GM}$	GM Carcasses
$Fp_{GM}$	GM Fecal Pellets
<b>II. PHYSICAL VARIABLES</b>	
$T$	Temperature
<b>III. GROWTH</b>	
$e^X$	growth efficiency of $X$
$a^X$	unassimilation rate of $X$
$g_m^X$	maximal $X$ grazing rate
$K_G^X$	half saturation constant for $X$ grazing
$p_Y^X$	$X$ preference for group $Y$
$Y_{\text{thresh}}^X$	group $Y$ threshold for $X$
$F_{\text{thresh}}^X$	feeding threshold for $X$
$w_X$	sinking velocity of $X$ particles
$ff_m^X$	$X$ flux feeding rate
$m^X$	$X$ quadratic mortality
$m_c^X$	$X$ non predatory quadratic mortality
$r^X$	$X$ linear mortality
$K_m$	half saturation constant for mortality
$\alpha$	remineralsation rate
<b>CLOGGING</b>	
$C_{th}$	clogging threshold
$C_{sh}$	clogging sharpness

**Table 1.** Variables and parameters used in the set of equations governing the temporal evolution of the state variables

184 This equation is similar to the one used for micro- and mesozooplankton in PISCES-  
185 v2 (Aumont et al., 2015). In this equation,  $X$  is the considered macrozooplankton biomass  
186 ( $GM$  or  $FFGM$ ), and the three terms on the right-hand side represent growth, quadratic  
187 and linear mortalities.  $e^X$  is the growth efficiency. It includes a dependence on food qual-  
188 ity as presented in PISCES-v2 (Eq. 27a and 27b in Aumont et al. (2015)). Quadratic  
189 mortality is divided between mortality due to predation by unresolved higher trophic lev-  
190 els (with a rate  $m^X$ ) and mortality due to disease (with a rate  $m_c^X$ ). All terms in this  
191 equation were given the same temperature sensitivity  $f_X(T)$  using a Q10 of 2.14 (Eq.  
192 25a and 25b in Aumont et al. (2015)), as for mesozooplankton in PISCES-v2 and accord-  
193 ing to E. Buitenhuis et al. (2006). Linear mortality is enhanced and growth rate is re-  
194 duced at very low oxygen levels, as we assume that macrozooplankton are not able to  
195 cope with anoxic waters ( $\Delta(O_2)$  varies between 0 in fully oxic conditions and 1 in fully  
196 anoxic conditions, see Eq. 57 in Aumont et al. (2015)).

197 The difference between the two macrozooplankton groups lies in the description  
 198 of the term  $G_X$ , i.e. the ingested matter. A full description of the equations describing  
 199  $G_X$  is provided in the supporting information section TextS2 (Eq. S1 to Eq. S12). Be-  
 200 low we present the two different choices of feeding representation that differentiate the  
 201 dynamics of the two macrozooplankton groups, GM and FFGM, in the model.

202 GM, namely generic macrozooplankton, is intended to represent crustacean macro-  
 203 zooplankton, such as euphausiids or large copepods. Their parameterization is similar  
 204 to that of mesozooplankton (Eq. 28 to 31 in Aumont et al. (2015)). Therefore, in ad-  
 205 dition to conventional suspension feeding based on a Michaelis-Menten parameterization  
 206 with no switching and a threshold (Eq. S1, S2 and S3), flux-feeding is also represented  
 207 (Eq. S4) as has been frequently observed for both meso- and macrozooplankton (Jackson,  
 208 1993; Stukel et al., 2019). GM can flux-feed on small and large particles as well as on  
 209 carcasses and fecal pellets produced by both GM and FFGM (Eq. S6). We assume that  
 210 the proportion of flux-feeders is proportional to the ratio of the potential food available  
 211 for flux feeding to the total available potential food (Eq. S7 and S8). Suspension feed-  
 212 ing is supposed to be controlled solely by prey size, which is assumed to be about 1 to  
 213 2 orders of magnitude smaller than that of their predators (Fenchel, 1988; B. Hansen et  
 214 al., 1994). Thus, GM preferentially feed on mesozooplankton, but also, to a lesser ex-  
 215 tent on microzooplankton, large phytoplankton and small particles (Eq. S5 and S10, Fig.  
 216 1).

217 FFGM represent the large pelagic tunicates (i.e. salps, pyrosomes and doliolids but  
 218 not appendicularians). Pelagic tunicates are all highly efficient filter feeders and thus have  
 219 access to a wide range of prey sizes, from bacteria to mesozooplankton (Acuña, 2001;  
 220 Sutherland et al., 2010; Bernard et al., 2012; Ambler et al., 2013). There is no strong  
 221 evidence that FFGM feed on mesozooplankton in the literature. Therefore, we assume  
 222 in our model that FFGM are solely suspension feeders (*i.e.* with concentration-dependent  
 223 grazing based on a Michaelis-Menten parameterization with no switching and a thresh-  
 224 old, see Eq. S1, S2 and S3) feeding with identical preferences on both phytoplankton groups  
 225 ( $D$  and  $N$ ) as well as on microzooplankton ( $Z$ ) (Eq. S11 and S12, Fig. 1). They can also  
 226 feed on small particles ( $sPOC$ , Sutherland et al. (2010)) (Eq. S11, Fig. 1).

### 227 **2.1.3 Carcasses and fecal pellet dynamics:**

228 Carcasses  $Ca_{FFGM}$  and  $Ca_{GM}$  are produced as a result of non predatory quadratic  
 229 and linear mortalities of GM and FFGM, respectively. The  $Fp_{FFGM}$  and  $Fp_{GM}$  are pro-  
 230 duced as a fixed fraction of the total food ingested by the two macrozooplankton groups.  
 231 Remineralization of fecal pellets and carcasses by bacteria is modeled using the same tempera-  
 232 ture-dependent specific degradation rate with a  $Q_{10}$  of 1.9, identical to that used for small  
 233 and large particles. In addition to remineralization, carcasses and fecal pellets undergo  
 234 flux feeding by GM as explained in the previous subsection. The sinking speeds of these  
 235 particle pools are assumed to be constant. A complete description of the equations gov-  
 236 erning the temporal evolution of fecal pellets and carcasses is provided in the support-  
 237 ing information section TextS2 (Eq. S14 and S15).

## 238 **2.2 Model experiments**

239 The biogeochemical model is run in an offline mode with dynamical fields identi-  
 240 cal to those used in Aumont et al. (2015). These climatological dynamic fields (as well  
 241 as the input files) can be obtained from the NEMO website ([www.nemo-ocean.eu](http://www.nemo-ocean.eu)) and  
 242 were produced using an ORCA2-LIM configuration (Madec, 2008). The spatial resolu-  
 243 tion is about  $2^\circ$  by  $2^\circ \cos(\phi)$  (where  $\phi$  is the latitude) with a meridional resolution en-  
 244 hanced at  $0.5^\circ$  in the equator region. The model has 30 vertical layers with increased  
 245 vertical thickness from 10 m at the surface to 500 m at 5000 m. PISCES-FFGM was ini-  
 246 tialized from the quasi-steady-state simulation presented in Aumont et al. (2015). The

247 two macrozooplankton groups, their fecal pellets and carcasses were set to a small uni-  
 248 form value of  $10^{-9}$  mol CL<sup>-1</sup>. The model was then integrated for the equivalent of 600  
 249 years, forced with 5-day averaged ocean dynamic fields and with a three-hour integra-  
 250 tion time step.

251 To investigate the spatial pattern and depth gradient of particulate organic car-  
 252 bon fluxes and the modeled distribution of GM and FFGM, three additional simulations  
 253 were performed: PISCES-GM ("Generic Macrozooplankton"), PISCES-LOWV ("LOW  
 254 Velocity") and PISCES-CLG ("Clogging").

255 The first experiment (PISCES-GM) was designed to investigate the impact of an  
 256 explicit FFGM representation (with a different grazing parameterization than GM) on  
 257 the spatial and vertical distribution of POC fluxes: In PISCES-GM, the FFGM inges-  
 258 tion rate ( $g_m^{FFGM}$  defined in table 1 and used in Eq. S3) was set to 0 which is equiva-  
 259 lent to running the model with a single generic macrozooplankton group.

260 The second experiment (PISCES-LOWV) was designed to evaluate the impact of  
 261 the high sinking speeds of particles from GM and FFGM. In PISCES-LOWV, the sink-  
 262 ing speeds of all fecal pellets and carcasses produced by GM and FFGM ( $w_{FPX}$  and  $w_{CaX}$ ,  
 263 defined in table 1 and used in Eq. S14 and S15) were assigned the same values as for large  
 264 particles in PISCES-v2, i.e. 30 m d<sup>-1</sup>.

265 The third experiment (PISCES-CLG) was designed to explore the impacts of clog-  
 266 ging. Clogging, defined as the saturation of an organism's filtering apparatus with high  
 267 levels of particulate matter, is a poorly documented mechanism for FFGM but has been  
 268 observed (Harbison et al., 1986; Perissinotto & Pakhomov, 1997) or suggested (Perissinotto  
 269 & Pakhomov, 1998; Pakhomov, 2004; Kawaguchi et al., 2004) for some salps species. Un-  
 270 like other macrozooplankton groups, it has been shown that salps biomass remain rel-  
 271 atively low at high chlorophyll concentrations (Heneghan et al., 2020). In PISCES-CLG,  
 272 the achieved ingestion rate of FFGM ( $G_{FFGM}$ , see Eq. S13) is modulated by a clogging  
 273 function  $F_C(Chl)$  inspired by the parameterization proposed by Zeldis et al. (1995):

$$F_C(Chl) = 1 - \frac{1}{2} (1 + \text{ERF}(C_{sh}(NCHL + DCHL - C_{th}))) \quad (2)$$

274 In this equation,  $C_{th}$  is the clogging threshold,  $C_{sh}$  is the clogging sharpness and ERF  
 275 is the Gauss error function.

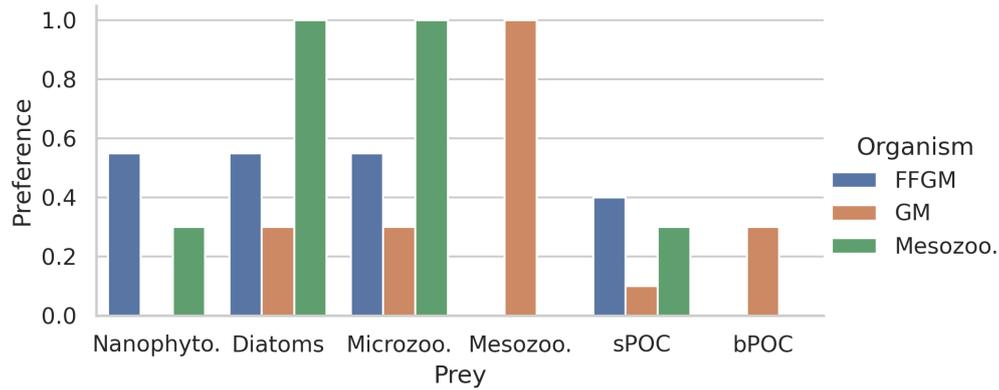
276 All three sensitivity experiments were initialized with the year 500 output fields  
 277 from the baseline PISCES-FFGM experiment. They were then run for 100 years. All re-  
 278 sults presented in this study are average values over the last 20 years of each simulation.

### 279 2.3 Model parameters

280 Each zooplankton group is characterized by a size range, assuming that sizes within  
 281 the group are distributed along a spectrum of constant slope -3 in log-log space, accord-  
 282 ing to the hypothesis of Sheldon et al. (1972). The ranges are: 10-200  $\mu\text{m}$  for microzoo-  
 283 plankton, 200-2000  $\mu\text{m}$  for mesozooplankton and 2000-20000  $\mu\text{m}$  for macrozooplankton  
 284 (GM and FFGM).

285 All parameters in PISCES-FFGM have identical values to those in Aumont et al.  
 286 (2015). The only exception is the mesozooplankton quadratic mortality rate, whose value  
 287 has been greatly reduced from its standard value of  $0.03 (\mu\text{mol CL}^{-1})^{-1} \text{day}^{-1}$  to  $0.004$   
 288  $(\mu\text{mol CL}^{-1})^{-1} \text{day}^{-1}$  since predation by higher trophic levels is now explicitly repre-  
 289 sented.

290 The values of the parameters that were introduced in PISCES-FFGM to represent  
 291 the evolution of GM and FFGM are shown in Table 2. Metabolic rates are assumed to  
 292 vary with size according to the allometric relationship proposed by P. J. Hansen et al.



**Figure 2. Histogram of the preferences of secondary consumers for their respective prey.** Secondary consumers are mesozooplankton, FFGM and GM, and preys are nanophytoplankton, diatoms, microzooplankton, mesozooplankton, small organic particles and large organisms particles. A preference of 1 indicates that any prey reached is consumed, a preference of 0 indicates that the prey is never consumed.

(1997). Therefore, maximum grazing, respiration and flux-feeding rates were calculated from their values for mesozooplankton using a size ratio of 10. The preferences of GM and FFGM for their different prey are detailed in section 2.1.2. Their values are shown in Figure 2. The sinking speed of FFGM carcasses (resp. fecal pellets) is set to  $800 \text{ m day}^{-1}$  (resp.  $1000 \text{ m day}^{-1}$ ) (Henschke et al., 2016). The sinking speeds of GM fecal pellets and carcasses are set rather arbitrarily to  $100 \text{ m day}^{-1}$  and  $300 \text{ m day}^{-1}$  respectively, within the wide range of values found in the literature (Small et al., 1979; Fowler & Knauer, 1986; Lebrato et al., 2013; Turner, 2015). A low clogging threshold  $C_{th}$  of  $0.5 \mu\text{g Chl L}^{-1}$  is chosen to limit FFGM growth in all moderate and high productivity regions. Clogging sharpness  $C_{sh}$  is set to  $5 \mu\text{g Chl L}^{-1}$ , the value proposed by Zeldis et al. (1995). The quadratic mortality rates have been adjusted by successive simulations evaluated against the observations presented in the next section.

## 2.4 Observations

### 2.4.1 Observations data for validating the modeled FFGM biomass estimates

We compiled an exhaustive dataset of in situ pelagic tunicates (i.e., Thaliaceans) concentrations from large scale plankton monitoring programs and previous plankton data compilations to derive monthly field of pelagic tunicates biomass (in  $\text{mg C m}^{-3}$ ) that can be used as a standard data set to evaluate the FFGM biomass estimated by PISCES-FFGM. First, five main data sources were retrieved: NOAA’s Coastal and Oceanic Plankton Ecology, Production, and Observation Database (COPEPOD; O’Brien (2014)), the Jellyfish Database Initiative (JeDI; Lucas et al. (2014)), KRILLBASE Atkinson et al. (2017), the Australian Continuous Plankton Recorder (CPR) survey (AusCPR; IMOS (2021)) and the Southern Ocean CPR survey (SO-CPR; (Hosie, 2021)). This compilation gathered planetary scale plankton concentration measurements collected through a broad variety of sampling devices over the last 150 years, with taxonomic identification of varying precision and scientific names, some of which changed through time. Therefore, we curated the scientific names and the taxonomic classification of each observation to harmonize names across all data sets and to correct deprecated names and syn-

Symbol	Source	GM ( $X = GM$ )	FFGM ( $X = FFGM$ )	Unit
$e_{max}^X$	★	0.35	0.35	-
$a^X$	★	0.3	0.3	-
$g_m^X$	●	0.28	0.28	$d^{-1}$
$K_G^X$	★	2e-5	2e-5	$mol\ L^{-1}$
$p_X^P$	‡	0	0.55	-
$p_X^D$	‡	0.3	0.55	-
$p_X^Z$	‡	0.3	0.55	-
$p_X^M$	‡	1	0	-
$p_X^{POC}$	‡	0.1	0.4	-
$p_X^{GOC}$	‡	0.3	0	-
$P_X^{thresh}$	★	1e-8	1e-8	$mol\ L^{-1}$
$D_X^{thresh}$	★	1e-8	1e-8	$mol\ L^{-1}$
$Z_X^{thresh}$	★	1e-8	1e-8	$mol\ L^{-1}$
$M_X^{thresh}$	★	1e-8	1e-8	$mol\ L^{-1}$
$POC_X^{thresh}$	★	1e-8	1e-8	$mol\ L^{-1}$
$F_X^{thresh}$	★	3e-7	3e-7	$mol\ L^{-1}$
$w_{Cax}$	‡	300	800	$m\ d^{-1}$
$w_{Fpx}$	‡	100	1000	$m\ d^{-1}$
$ff_m^H$	●	5e5	-	$m^2\ mol^{-1}$
$m^X$	†	1.2e4	1.2e4	$L\ mol^{-1}\ d^{-1}$
$m^X$	†	4e3	4e3	$L\ mol^{-1}\ d^{-1}$
$r^X$	●	0.003	0.005	$d^{-1}$
$K_m$	★	2e-7	2e-7	$mol\ L^{-1}$
$\alpha$	★	0.025	0.025	$d^{-1}$

**Table 2. Parameter values used in PISCES-FFGM.** The symbols in the "Source" column indicate how the parameter value was determined: (★) parameters for which we assumed that both GM and FFGM share the same characteristics as mesozooplankton, (●) metabolic rates assumed to vary with size, thus scaled using an allometric scaling conversion of mesozooplankton value based on (P. J. Hansen et al., 1997), (†) parameters tuned to fit PISCES-v2 general biology dynamics, and (‡) indicates parameters whose values have been arbitrarily set based on information available in the literature and/or of the authors expertise.

322 onyms based on the backbone classification of the World Register of Marine Species (WoRMS;  
323 Horton et al. (2022)) using the 'worms' R package version 0.2.2 (Holstein, 2018). Then,  
324 only those observations corresponding to an organism belonging to the Class Thaliacea  
325 were kept. Observations without a precise sampling date and at least one sampling  
326 depth indicator (usually maximum sampling depth, in meters) were discarded. All data  
327 sets provided concentrations in  $ind\ m^{-3}$  except KRILLBASE which provided salp (mostly  
328 *Salpa thompsoni*) densities in  $ind\ m^{-2}$  which we converted to  $ind\ m^{-3}$  based on the max-  
329 imum sampling depth of the corresponding net tows. In KRILLBASE, 5'186 observa-  
330 tions of Thaliaceans with missing density values were discarded (35.6% of the original  
331 14'543 observations). In COPEPOD, concentrations are standardized as if they were all  
332 taken from a plankton net equipped with a 330  $\mu\ m$  mesh (Moriarty & O'Brien, 2013).  
333 862 point observations with missing concentration values were discarded (3.5% of the  
334 original 24'316 observations). We examined the composition of the original data sources  
335 compiled within JeDI and COPEPOD by assessing the recorded institution codes as well

336 as their corresponding spatio-temporal distributions to evaluate the observations over-  
 337 lapping between these two previous data syntheses. We logically observed a very high  
 338 overlap between COPEPOD and JeDI as the former data set was the main data con-  
 339 tributor to the latter. Therefore, overlapping records were identified based on their sam-  
 340 pling metadata, scientific names, concentration values, the recorded institution codes and  
 341 recorded data sources, and they were removed from JeDI. This removed 14'198 (74.1%)  
 342 of the JeDI's original Thaliaceans observations.

343 This synthesis of Thaliaceans concentrations gathered globally distributed 491,529  
 344 point observations (Figure S1), collected at a mean ( $\pm$  std) maximum sampling depth  
 345 of 23.1 ( $\pm$  70.5) m over the 1926-2021 time period (mean  $\pm$  std of the sampling year =  
 346 2006.6  $\pm$  11.5). The shallow average sampling depth was driven by the dominant con-  
 347 tribution of the two CPR surveys, which represented 93% of all point observations. Re-  
 348 moving the CPR surveys deepened the mean maximum sampling depth of the observa-  
 349 tions to 189.3 ( $\pm$  196.1) m. The range of observed Thaliacean concentration ranged from  
 350 0.0 ind  $m^{-3}$  to 10,900 ind  $m^{-3}$  with an average of 1.3 ( $\pm$  45.4) ind  $m^{-3}$ .

351 Most of the records showed a fairly precise taxonomic resolution as 39% of the data  
 352 was species- resolved (mostly *S. thompsoni*, *Soestia zonaria*, *S. fusiformis* and *Thalia*  
 353 *democratica*), 0.19% genus-resolved (mostly *Thalia*, *Doliolum* and *Salpa*) and 38% family-  
 354 resolved (mostly Salpidae and Doliolidae). Therefore, we were able to perform taxon-  
 355 specific conversions from individual concentrations to biomass concentrations (in mg C  
 356  $m^{-3}$ ) for each point observation (see Table S1). We used the taxon-specific carbon weights  
 357 (mg C ind $^{-1}$ ) summarized by Lucas et al. (2014) which were based on the group-specific  
 358 length-mass or mass-mass linear and logistic regression equations of Lucas et al. (2011).  
 359 Not all the observations had a precise counter part in the carbon weights compilation  
 360 of Lucas et al. (2014) because they were not identified at the species or the genus level  
 361 (e.g., Class-level, Order-level or Family-level observations). In these cases, we computed  
 362 the median carbon weight of those taxa reported in Lucas et al. (2014) and which com-  
 363 posed the higher level taxonomic group (i.e., the carbon weight of Salpidae corresponded  
 364 to the average carbon weight of all Salpidae species), and used this average carbon weight  
 365 to convert the individual concentrations to carbon concentrations. The resulting point  
 366 biomass measurements ranged between 0.0 mg C  $m^{-3}$  and 19'451 mg C  $m^{-3}$ , with an  
 367 average of 0.63  $\pm$  48 mg C  $m^{-3}$ . However, this range is largely zero-inflated (94.6% of  
 368 the observations corresponded to a biomass of 0.0 mg C  $m^{-3}$ ) due to the high relative  
 369 contribution of both CPR surveys whose data only comprised 1.1% of non null values.  
 370 Such strong zero inflation can be attributed to sampling artifacts due to the specificities  
 371 of the CPR and thus very likely do not reflect real absences (A. Richardson et al.,  
 372 2006). Indeed, the CPR continuously collects plankton at standard depth of 7 m and at  
 373 a speed of nearly 0.2 m  $s^{-1}$ , as seawater flows in through a square aperture of 1.61  $cm^2$ ,  
 374 which is too narrow to adequately sample large gelatinous macrozooplankton such as salps  
 375 and doliolids, especially in the Southern Ocean (Pinkerton et al., 2020). Consequently,  
 376 we decided to remove the observations from the AusCPR and the SO-CPR from our final  
 377 validation data set. Biomass observations larger than two times the standard deviation  
 378 were considered as outliers and were excluded as well. Then, we only retained this  
 379 observations taking on the upper 300 m depth to exclude really deep water samples and  
 380 focus on zooplankton communities that inhabit the euphotic layer. The biomass levels  
 381 of this subset ranged between 0.0 and 488 mg C  $m^{-3}$  (4.9  $\pm$  25.7 mg C  $m^{-3}$ ). Thali-  
 382 acean concentrations issued from single net sample were summed when necessary (e.g.,  
 383 when species and/or genera counts were sorted within one plankton sample) to be rep-  
 384 resentative of a Thaliacea-level point measurement. At this point, the dataset contains  
 385 18'875 single observation of Thaliacean biomass. Hereafter, we will refer to this dataset  
 386 as "AtlantECO dataset".

387 Ultimately, monthly Thaliacean biomass fields were computed for validating the  
 388 monthly FFGM biomass fields of PISCES-FFGM. Thaliacea biomass concentrations were

389 averaged per months on a 36x72 grid to obtain the 12 monthly climatological fields of  
 390 Thaliacea biomass needed for evaluating our model. A low resolution grid (5x5) has been  
 391 used to counterbalance patchiness of data, as suggested by (Lilley et al., 2011). After  
 392 this final step, the monthly climatological values of Thaliacea biomass concentrations ranged  
 393 between 0.0 and 454 mg C m<sup>-3</sup> ( $6.53 \pm 26.21$  mg C m<sup>-3</sup>). Hereafter, we will refer to  
 394 this climatology as "AtlantECO climatology".

### 395 **2.4.2 Additional datasets**

396 We also used the monthly fields derived from the observations as a standard data  
 397 set to evaluate some of the other PISCES-FFGM compartments: total macrozooplankton,  
 398 mesozooplankton, total chlorophyll, nutrients and oxygen.

399 *2.4.2.1 Total macrozooplankton* As with FFGM, for total macrozooplankton ob-  
 400 servations, a low resolution grid has been used. We use a monthly macrozooplankton abun-  
 401 dances binned on a 72x36 grid (ind m<sup>-3</sup>, vertically integrated between 0 and 100m) from  
 402 MARine Ecosystem DATA (MAREDAT) (Moriarty et al., 2013), and then convert abun-  
 403 dances to carbon-based concentration to evaluate our modeled distribution of total macro-  
 404 zooplankton biomass (*i.e.* FFGM and GM). Conversion of abundance to carbon concen-  
 405 tration requires an average individual weight. An average individual weight of 588  $\mu$ g  
 406 was chosen by considering an individual with a mean size of 6.3 mm (the geometric mean  
 407 of the macrozooplankton size class) and applying the relationship proposed for copepods  
 408 by Watkins et al. (2011).

409 *2.4.2.2 Mesozooplankton* We use the monthly mesozooplankton database binned  
 410 on a 360x180 grid (mmol m<sup>-3</sup>, vertically integrated between 0 and 300m) from MARine  
 411 Ecosystem DATA (MAREDAT) (Moriarty & O'Brien, 2013) to evaluate our modeled to-  
 412 tal mesozooplankton biomass distribution.

413 *2.4.2.3 Nutrients and Oxygen* We use the World Ocean Atlas (Garcia et al., 2019a,  
 414 2019b) 360x180 monthly climatological distribution for  $PO_4^{2-}$  (Surface),  $NO_3^-$  (Surface)  
 415 and  $O_2$  (vertically integrated between 100 and 600m and between 2000 and 4000m) to  
 416 evaluate our modeled nutrient and oxygen distributions .

417 *2.4.2.4 Chlorophyll* We use a 360x180 gridded monthly average of the long-term  
 418 multi-sensor time-series OC-CCI (Ocean Colour project of the ESA Climate Change Ini-  
 419 tiative, Sathyendranath et al. (2019)) of satellite phytoplankton chlorophyll-*a* sea sur-  
 420 face concentration converted into mmol m<sup>-3</sup> to evaluate our modeled total chlorophyll  
 421 distribution. The same product regridded on a 36x72 grid is used to compare observed  
 422 and modeled relationships between chlorophyll and FFGM abundance (Fig. 5).

### 423 **2.4.3 Model evaluation**

424 The model evaluation is based on monthly fields averaged over the last 20 years  
 425 of the PISCES-FFGM reference.

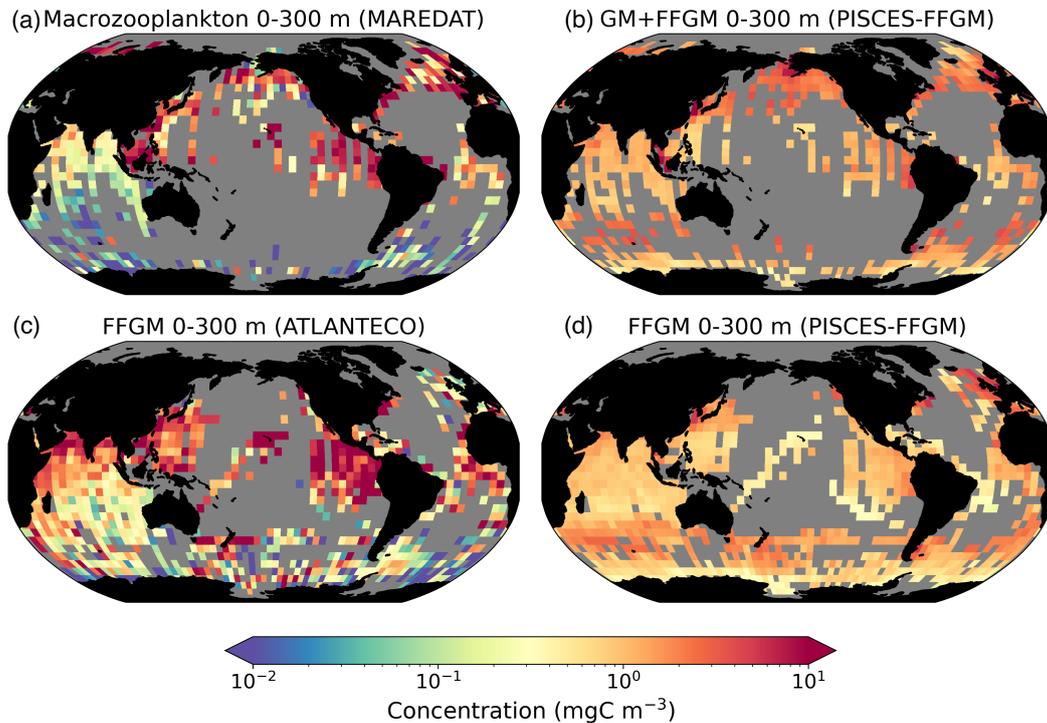
426 FFGM: For each unique observation in the AtlantECO dataset, we sampled the mod-  
 427 eled FFGM biomass from the PISCES-FFGM climatology at the corresponding coordi-  
 428 nates (latitude,longitude), month, and depth range (minimal depth and maximal depth),  
 429 so that each observed biomass can be compared to a "model-sampled" biomass. When  
 430 compared to AtlantECO climatology, the annual mean FFGM biomass fields and the statis-  
 431 tics (Table 3) are calculated from these "model-sampled" biomasses to avoid bias due  
 432 to different sampling.

433 Other variables : The other model outputs used in this evaluation ( $NO_3^-$ ,  $PO_4^{2-}$ ,  
 434 Chl, Mesozooplankton, GM+FFGM) were regridded horizontally and vertically on the  
 435 same grid as the corresponding observations (see previous section). The macrozooplank-

436 ton and mesozooplankton fields were integrated vertically on the appropriate vertical range.  
 437 When compared to observations, model outputs are sampled at exactly the same loca-  
 438 tion and month as the observations. Annually averaged fields as well as statistics (Ta-  
 439 ble 3) are computed from these sampled fields to avoid bias due to different sampling.

### 440 3 Results

#### 441 3.1 Evaluation of simulated biomasses



**Figure 3. Comparison between observed and modeled macrozooplankton biomasses.** Annual means of carbon concentrations ( $\text{mg C m}^{-3}$ , log-scale), averaged over the top 300 meters on a  $5^\circ$  resolution grid. (a) macrozooplankton from MAREDAT (b) "model-sampled" total macrozooplankton (GM+FFGM) (c) FFGM from AtlantECO climatology (d) "model-sampled" FFGM. As described in section 2.4.3, modeled biomasses were sampled where observations were available.

442 We focus here on the evaluation of the new components added in this version of  
 443 PISCES, i.e. GM and FFGM. In the supporting information, we present an evaluation  
 444 of nitrate, chlorophyll and mesozooplankton (See Text S1 and Fig. S2). For these trac-  
 445 ers, note that the performance of PISCES-FFGM is similar to that of PISCES-v2 (Aumont  
 446 et al., 2015). The total integrated biomass of all living compartments simulated by PISCES-  
 447 FFGM is 1.4 Pg C for the upper 300 meters of the global ocean. Primary producers ac-  
 448 count for 51% of this biomass. Total macrozooplankton accounts for 12% of the total  
 449 biomass. Our model predicts that FFGM and GM contribute roughly equally to macro-  
 450 zooplankton biomass, each having a biomass of about 0.08 Pg C.

451 The annual mean distributions of total macrozooplankton (FFGM and GM) and  
 452 FFGM only, averaged over the top 300 m of the ocean, are compared to available ob-  
 453 servations (Figure 3). A quantitative statistical evaluation of the model performance for

	Experiment	Total Macrozooplankton PISCES-FFGM	FFGM PISCES-FFGM	FFGM PISCES-CLG
Model	Mean (mg C m <sup>-3</sup> )	1.65	1.18	0.69
	Median (mg C m <sup>-3</sup> )	1.56	0.80	0.30
	Std (mg C m <sup>-3</sup> )	1.29	0.96	0.69
Observation	Mean (mg C m <sup>-3</sup> )	11.01	8.22	7.79
	Median (mg C m <sup>-3</sup> )	0.52	1.11	0.99
	Std (mg C m <sup>-3</sup> )	128	26.9	26.3
comparison	Bias (mg C m <sup>-3</sup> )	-9.36	-7.04	-7.53
	Bias (log10)	0.57	0.04	-0.18
	R Spearman	0.26 ( $p < 10^{-5}$ )	0.17 ( $p < 10^{-5}$ )	0.34 ( $p < 10^{-5}$ )
	High biomasses match	94 %	91 %	84 %
	Low biomasses match	2 %	14 %	41 %

**Table 3. Macrozooplankton model vs. observation statistics.** "Mean", "median" and "standard" deviation are computed on all the non-zero biomass values of the annual climatologies (as defined in section 2.4.3 of the methods) weighted by their respective cell areas. "Bias" is computed as the difference between modeled and observed means. "Bias (log10)" is computed on log10 converted observed and modeled climatologies. "R Spearman" is the Spearman correlation coefficient computed on non zero values of the climatologies. "High biomasses match" is the percentage of observed area where biomasses are greater than 0.5 mg C m<sup>-3</sup> that correspond to area where model biomasses are greater than 0.5 mg C m<sup>-3</sup>. "Low biomasses match" is the percentage of observed area where biomasses are lower than 0.5 mg C m<sup>-3</sup> that correspond to area where model biomasses are lower than 0.5 mg C m<sup>-3</sup>.

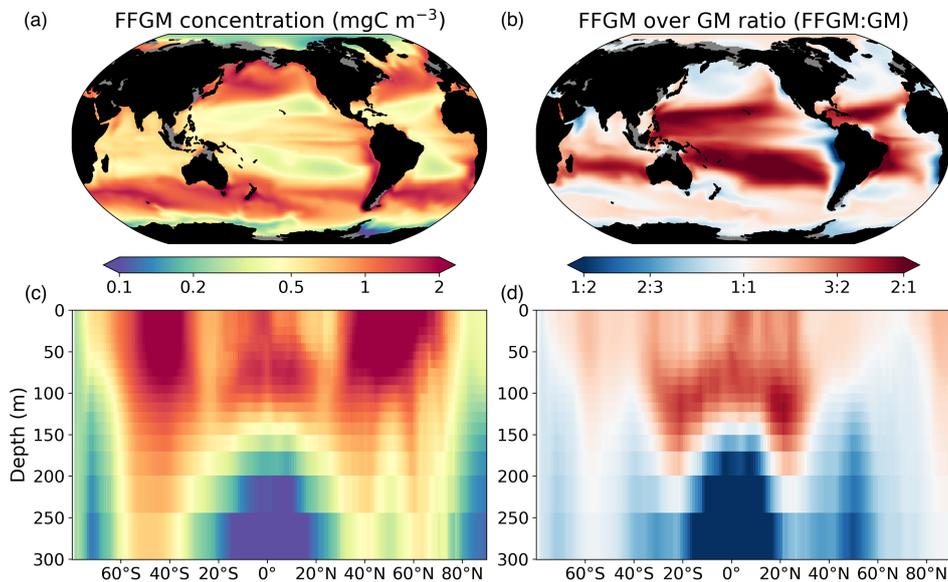
454 these two fields is presented in Table 3. The Spearman correlation coefficient between  
455 observed and modeled total macrozooplankton biomasses is 0.26 (p-value < 0.001). Ar-  
456 eas of high macrozooplankton biomass are correctly simulated in the northern hemisphere  
457 by our model: 94% of the area in which observed concentrations are greater than 0.5 mg  
458 C m<sup>-3</sup> correspond to areas in which the concentration is greater than 0.5 mg C m<sup>-3</sup> in  
459 the model. On the other hand, observations suggest moderate biomass in the Indian Ocean  
460 (between 0.05 and 0.5 mg C m<sup>-3</sup>) and low biomass in the Southern Ocean (lower than  
461 0.05 mg C m<sup>-3</sup>). These low and moderate biomasses are not captured by our model which  
462 simulates values greater than 0.5 mg C m<sup>-3</sup> in both areas: 98% of the area in which ob-  
463 served concentrations are lower than 0.5 mg C m<sup>-3</sup> correspond to areas in which mod-  
464 eled concentrations are greater than 0.5 mg C m<sup>-3</sup>. Overall, the simulated distribution  
465 of macrozooplankton is too homogeneous with respect to what the observations suggest.  
466 This is confirmed by the much smaller standard deviation in our model simulation than  
467 in the observations, 1.3 and 128 mg C m<sup>-3</sup> respectively.

468 Our model simulates a distribution of FFGM in the upper ocean that correlates  
469 with observation with a Spearman correlation coefficient of 0.17 (p-value < 0.001). The  
470 simulated FFGM biomass is high (>0.5 mg C m<sup>-3</sup>) in the equatorial domain of the Pa-  
471 cific and Atlantic oceans and in the mid latitudes of both hemispheres. Conversely, FFGM  
472 biomass is moderate (between 0.05 and 0.5 mg C m<sup>-3</sup>) in the oligotrophic subtropical  
473 gyres and in the high latitudes (>60°). Compared to observations, the spatial patterns  
474 of high biomasses are better reproduced than for total macrozooplankton: 91% of the  
475 area in which observed concentrations are greater than 0.5 mg C m<sup>-3</sup> correspond to ar-  
476 eas in which modeled concentrations are greater than 0.5 mg C m<sup>-3</sup>. However, the max-  
477 imum observed values are strongly underestimated: the 95th percentile of the modeled  
478 values is 2.6 mg C m<sup>-3</sup> while it is 32 mg C m<sup>-3</sup> in the observations. In the Southern  
479 Ocean, the simulated distribution is much more zonally homogeneous than suggested by

480 observations (Fig. 3). Overall, the predicted median biomass of FFGM is similar to that  
 481 of observations, 0.80 vs. 1.11 mg C m<sup>-3</sup>. As with macrozooplankton, but to a lesser ex-  
 482 tent, the simulated standard deviation is significantly lower than in the observations, 0.96  
 483 and 26.9 mg C m<sup>-3</sup> respectively. The standard and log10 biases are closer to 0 than those  
 484 calculated for macrozooplankton (Table 3).

485 The addition of clogging in PISCES-CLG doubled the model-data spatial correlation  
 486 (Spearman's correlation coefficient is 0.34 compared to 0.17 previously, see Table 3).  
 487 This improvement is explained by a better representation of areas with moderate and  
 488 low biomass in PISCES-CLG (concentrations <0.5 mg m<sup>-3</sup>), especially in the southern  
 489 part of the Southern Ocean (see fig. S3). Indeed, 41% of the areas where observations  
 490 give values below 0.5 mg C m<sup>-3</sup> correspond to areas where the model predicts values be-  
 491 low 0.5 mg C m<sup>-3</sup> (vs only 14% in PISCES-FFGM). However, the simulated spatial vari-  
 492 ability remains strongly underestimated (std = 0.69 mg C m<sup>-3</sup> in PISCES-CLG and 26.9  
 493 mg C m<sup>-3</sup> in the AtlantECO climatology). Furthermore, biases are increased when clog-  
 494 ging is added (see Table 3).

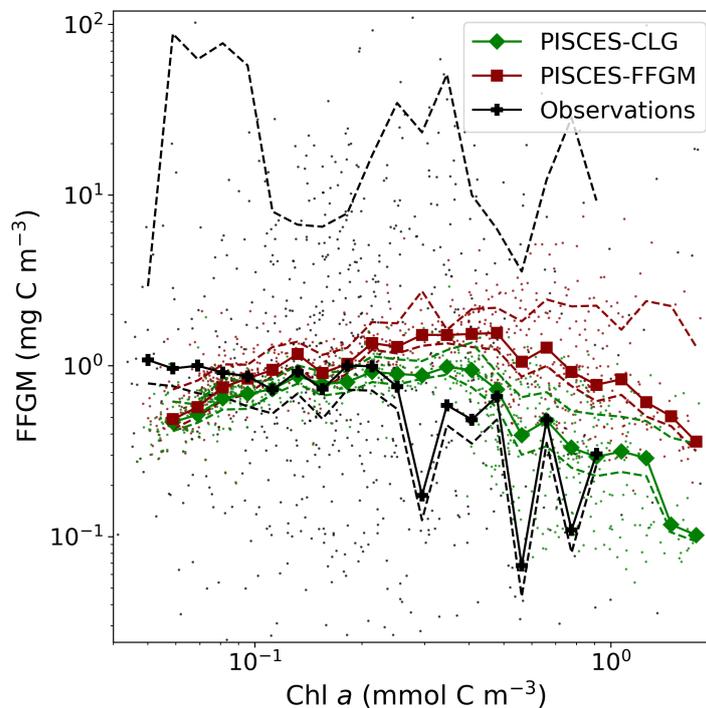
### 495 3.2 Simulated FFGM distribution



**Figure 4. FFGM and FFGM:GM ratio.** Annual mean of FFGM carbon concentrations (mg C m<sup>-3</sup>, log-scale), averaged over the top 300 meters (a), and zonally averaged (c). Annual of mean FFGM:GM ratio, averaged over the top 300 meters (b), and zonally averaged (d). Red tones indicate FFGM dominance, blue tones indicate GM dominance.

496 In this section, we first compare the simulated spatial distributions of FFGM and  
 497 GM. Figure 4 displays the annual mean FFGM to GM ratio averaged over the top 300  
 498 m of the ocean. It also shows the zonally averaged distribution of this ratio. The most  
 499 striking feature is the reverse distribution of the ratio as compared to the simulated ab-  
 500 solute biomass of both GM and FFGM. The ratio exceeds 2 in oligotrophic subtropical  
 501 gyres while it is minimal in the most productive regions. In the eastern boundary up-  
 502 welling systems, FFGM biomass can be more than two times lower than GM biomass.  
 503 In terms of the vertical distribution, the ratio is on average larger than 1 in the euphotic  
 504 zone. Below the euphotic zone, it sharply decreases as GM become dominant. In the mesopelagic  
 505 domain, flux-feeding has been shown to be a very efficient mode of predation (Jackson,

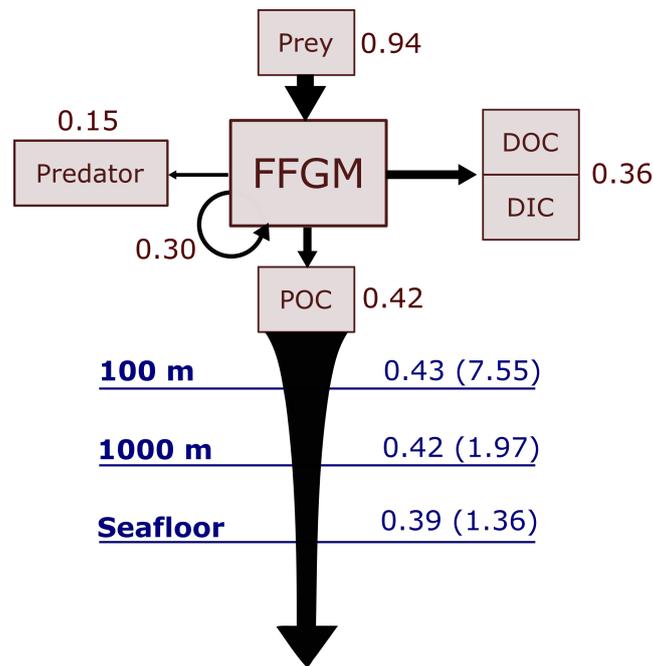
1993 ; Stukel, Ohman, et al., 2019). Since FFGM are not able to practice this feeding mode, they are outcompeted by GM. FFGM:GM ratio is maximum in the lower part of the euphotic zone in the subtropical domain where deep chlorophyll maxima are located.



**Figure 5. Chlorophyll-FFGM relationship.** Log-log scatter plot showing FFGM concentration versus total chlorophyll concentration for PISCES-FFGM, PISCES-CLG clogging run, and for the AtlantECO vs OC-CCI chlorophyll datasets. The datasets were gridded into an annual climatology with a spatial resolution of  $5^\circ$ . Each small dot corresponds to one grid cell of these climatologies. Large dots connected by a line represent the median per 0.07-wide log-bins of chlorophyll, dashed lines represent standard deviations below and above the median for each bin.

We then analyse the distribution of FFGM biomasses as a function of chlorophyll levels. Black dotted line and points on figure 5 show the FFGM biomass from the AtlantECO database plotted against the corresponding chlorophyll concentrations from OC-CCI (see section 2.4.2). Despite considerable scatter, this data-based analysis suggests a modest decrease of FFGM biomass for chlorophyll concentrations above about  $0.3 \text{ mg Chl m}^{-3}$ . Yet, this decrease is far from systematic, since even at high chlorophyll concentrations, FFGM biomass can be very high ( $>10 \text{ mg Chl m}^{-3}$ ). In our reference PISCES-FFGM simulation (red dotted-line and points on figure 5), the median values of FFGM biomass appear to be consistent with observations at intermediate chlorophyll concentrations between  $0.08$  and  $0.3 \text{ mg Chl m}^{-3}$ . However, as already mentioned in the previous section, our model predicts a much weaker variability of FFGM biomass. For higher chlorophyll concentrations, median FFGM levels become significantly larger than in the observations (up to one order of magnitude larger, see fig. 5). Here again, the addition of clogging in PISCES-CLG (green dotted line and points in fig. 5) reduces the bias and thus better reproduces the observed relationship between FFGM biomass and chlorophyll  $a$  concentration.

### 3.3 Carbon cycle



**Figure 6. Schematic representation of carbon fluxes induced by processes related to FFGM.** Values are in  $\text{Pg C yr}^{-1}$ . The upper part of the diagram represents the sources and sinks of FFGM integrated globally over the first 100 meters. The source is the grazing on the different prey. The arrow going from FFGM to FFGM corresponds to the flux related to growth due to assimilated food. The sinks are : i) the remineralization, non-assimilation and linear mortality that go into the dissolved organic carbon (DOC) and dissolved inorganic carbon (DIC) ii) the quadratic predatory mortality term (directly remineralized in PISCES-FFGM because of the lack of explicit representation of upper level predators) and iii) the production of particular organic carbon (POC) via carcasses and fecal pellets. The lower part of the diagram corresponds to the export of POC linked to the fall of carcasses and fecal pellets of FFGM. The values in blue correspond to the global annual FFGM-driven POC flux through the corresponding depth, the values in parenthesis representing the total POC flux (i.e. related to FFGM, GM, bPOC and sPOC).

#### Carbon export from the surface ocean :

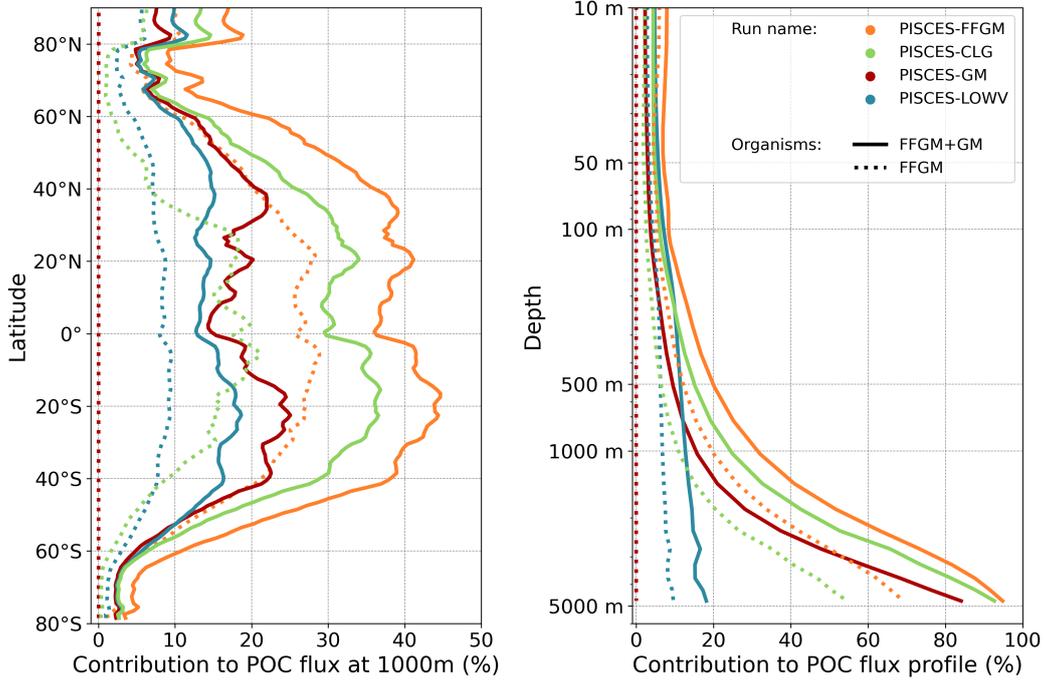
We first discuss the role of macrozooplankton in shaping the carbon cycle in the upper ocean, focusing on differences between GM and FFGM-related surface processes. Table 4 shows the globally integrated sinking flux of organic carbon particles at 100 m and 1000 m, while Figure 6 focuses on the FFGM-driven carbon fluxes. The total export flux from the upper ocean (at 100 m) is  $7.55 \text{ Pg C yr}^{-1}$  (Table 4). This value is relatively similar to previous estimates using different versions of PISCES (Aumont et al., 2015, 2017, 2018). It is also within the range of published estimates, *i.e.*  $4\text{--}12 \text{ Pg C yr}^{-1}$  (e.g. Laws et al., 2000; Dunne et al., 2007; Henson et al., 2011; DeVries & Weber, 2017). Small and large particles produced by phytoplankton, microzooplankton and mesozooplankton account for 91% of this carbon flux. The remaining 9% ( $0.69 \text{ Pg C yr}^{-1}$ , Table 4) is due to macrozooplankton, with one third of this amount coming from carcasses

Experiment	Depth (m)	bPOC (Pg C yr <sup>-1</sup> )	sPOC (Pg C yr <sup>-1</sup> )	$F_{pGM}$ (Pg C yr <sup>-1</sup> )	$Ca_{GM}$ (Pg C yr <sup>-1</sup> )	$F_{pFFGM}$ (Pg C yr <sup>-1</sup> )	$Ca_{FFGM}$ (Pg C yr <sup>-1</sup> )	Total (Pg C yr <sup>-1</sup> )	GM+FFGM contribution	FFGM contribution
PISCES-FFGM	100	4.49	2.37	0.09	0.17	0.29	0.14	<b>7.55</b>	34%	21%
PISCES-CLG	100	4.70	2.42	0.10	0.19	0.14	0.07	<b>7.62</b>	27%	12%
PISCES-GM	100	4.92	2.49	0.11	0.20	0.00	0.00	<b>7.73</b>	17%	0%
PISCES-LOWV	100	4.72	2.41	0.08	0.15	0.24	0.12	<b>7.71</b>	13%	7%
PISCES-FFGM	1000	1.18	0.12	0.11	0.14	0.27	0.15	<b>1.97</b>	9%	6%
PISCES-CLG	1000	1.22	0.12	0.12	0.15	0.14	0.08	<b>1.83</b>	7%	3%
PISCES-GM	1000	1.27	0.13	0.12	0.16	0.00	0.00	<b>1.68</b>	4%	0%
PISCES-LOWV	1000	1.23	0.13	0.04	0.06	0.07	0.04	<b>1.56</b>	8%	5%

**Table 4. Particulate carbon flux composition at 100 and 1000 m.** Units are in Pg C yr<sup>-1</sup>. sPOC (resp. bPOC) is for small (resp. large) particulate organic carbon.  $Ca_{GM}$  (resp.  $Ca_{FFGM}$ ) is for GM (resp. FFGM) carcasses.  $F_{pGM}$  (resp.  $F_{pFFGM}$ ) is for GM (resp. FFGM) fecal pellets.

538 and the remaining from fecal pellets. FFGM are responsible for an export of 0.46 Pg C yr<sup>-1</sup>  
 539 (Table 4), which represents 62% of the total macrozooplankton contribution.

540 The particularly large contribution from FFGM compared to GM comes from higher  
 541 production (grazing of 0.94 Pg C yr<sup>-1</sup> compared to 0.63 Pg C yr<sup>-1</sup> for GM, figures 6 and  
 542 S4) while both groups shows similar export efficiency: 45% of the grazed matter is ex-  
 543 ported at 100m, with the remaining 55% being split between implicit predation by up-  
 544 per trophic levels and loss to dissolved inorganic and organic carbon.



**Figure 7. Macrozooplankton relative contribution to particulate organic carbon fluxes.** The color indicates the PISCES configuration considered (see sensitivity section). The figure on the left shows the relative contribution of FFGM (dash) and macrozooplankton (FFGM+GM, solid) to the POC export at 1000m averaged zonally. The figure on the right shows the globally averaged vertical profile of these relative contributions.

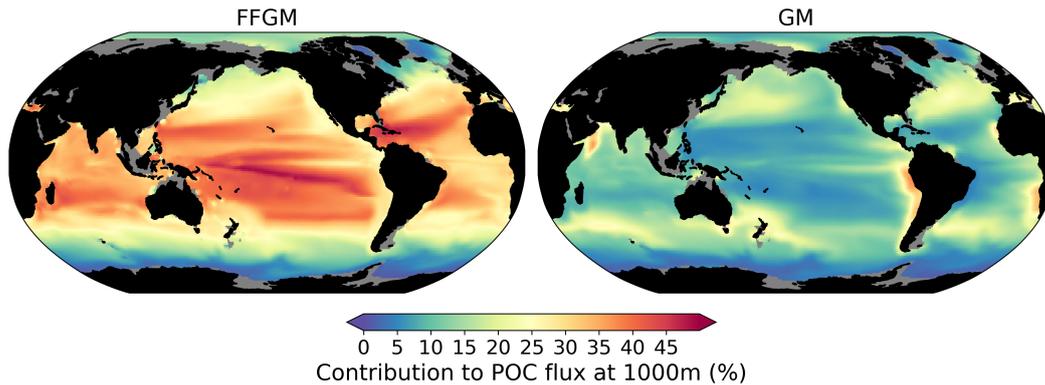
545 **Carbon transfer efficiency in the deep ocean :** We then analyze how the rep-  
 546 resentation of the two new macrozooplankton groups influences the fate of particulate  
 547 organic carbon in the deep ocean. At 1000 m, the total simulated POC flux is 1.97 Pg  
 548  $\text{C yr}^{-1}$  (Table 4). This flux is about 26% of the flux at 100 m. Most of this strong de-  
 549 crease is due to the loss of small and large organic particles. Macrozooplankton-driven  
 550 export is very effective because it remains almost unchanged from 100 m to 1000 m, 0.69  
 551 and 0.67 Pg  $\text{C yr}^{-1}$ , respectively (Table 4). Therefore, the contribution of macrozooplank-  
 552 ton increases strongly with depth to 34% of the total carbon export at 1000 m (Fig. 7).  
 553 The respective contribution of particles produced by GM and FFGM (carcasses and fecal  
 554 pellets) to this flux is almost identical at both depth horizons. At 5000 m, more than  
 555 90% of the carbon flux is due to macrozooplankton.(Fig. 7).

556 The PISCES-LOWV experiment, in which carcasses and fecal pellets sinking speeds  
 557 of both macrozooplankton groups are reduced to  $30 \text{ m d}^{-1}$ , shows a much greater at-  
 558 tenuation of POC fluxes with depth: while the total export of organic carbon at 100 m  
 559 increases slightly to  $7.71 \text{ Pg C yr}^{-1}$ , it is reduced by 20% at 1000m compared to the stan-  
 560 dard PISCES-FFGM run ( $1.56 \text{ Pg C yr}^{-1}$ , see table 4). The macrozooplankton contri-  
 561 bution is similar to that found in the standard model at 100m (8%) but the contribu-  
 562 tion is reduced to 13% at 1000m and to 20% at 5000m (Fig. 7). This confirms that the  
 563 strong contribution of macrozooplankton to POC fluxes at depth in the standard run  
 564 is explained by the very high sinking speeds of carcasses and fecal pellets. These high  
 565 sinking speeds prevent any significant remineralization of these particles as they sink to  
 566 the seafloor.

567 The PISCES-GM experiment, in which FFGM are not allowed to grow, shows a  
 568 similar depth gradient of the macrozooplankton contribution (Fig. 7, red curve) com-  
 569 pared to the standard run, but a lower contribution at each depth (by 10%). Indeed, the  
 570 transfer efficiency from 100 to 1000 m differs by only 2% between the two groups in the  
 571 standard model (97% for FFGM, 95% for GM) so that particles produced at the surface  
 572 by both groups have a similar fate towards the deep ocean. However, the estimated trans-  
 573 fer efficiency is biased as both groups of organisms produce particles below 100m. Be-  
 574 cause they can adopt a flux feeding strategy of predation, GM occupy the whole water  
 575 column whereas FFGM remain confined to the upper ocean (see section 3.2 and Figure  
 576 4). As a result, GM also produce particles below 100 m which contribute to the flux at  
 577 1000 m and explains the computed higher transfer efficiency. This is confirmed by the  
 578 PISCES-LOWV experiment: the efficiency of FFGM is reduced to 30% in this simula-  
 579 tion while that of GM is only reduced to 40%, even though the carcasses and fecal pel-  
 580 lets sinking velocities of both groups are identical. As the remineralization processes are  
 581 identical in the two runs, we can reasonably assume that the difference comes from the  
 582 relatively higher productivity below 100m of GM compared to FFGM.

583 **POC flux spatial patterns :** Although the processes underlying the efficient  
 584 sequestration of the particulate carbon issued from the two groups of macrozooplank-  
 585 ton are similar, we investigate how the spatial and temporal patterns of the induced deep  
 586 POC export differ between GM and FFGM.

587 The relative contribution of FFGM and GM to the POC flux at 1000 m presented  
 588 in Figure 8 is very contrasted between the two macrozooplankton groups. The POC flux  
 589 due to FFGM is maximal at about 40% of the total flux in the oligotrophic subtropical  
 590 gyres. In the productive areas of the low and mid-latitudes, it has intermediate values  
 591 close to 25%. It is minimal (<15%) at high latitudes, especially along the Antarctic. In  
 592 contrast, POC fluxes due to GM are maximal in the productive regions of the low and  
 593 mid-latitudes, especially in boundary upwelling systems where they can exceed 35% of  
 594 the total flux. These patterns are consistent with the respective spatial distribution of  
 595 FFGM and GM (ratio shown in figure 4).



**Figure 8. Relative contribution of macrozooplankton to particulate organic carbon flux at 1000m.** On the left (resp. right): relative importance at 1000m of FFGM (resp. GM) carcasses and fecal pellets driven POC flux to total POC flux (incl. GM and FFGM carcasses and fecal pellets as well as small and large particles).

596 We further investigate the importance of GM and FFGM for the spatial patterns  
 597 of the export of carbon to the deep ocean by contrasting PISCES-FFGM and PISCES-  
 598 GM experiments (see Section 2.2). Figure 7 shows the relative contribution of macro-  
 599 zooplankton to POC flux as a function of latitude. By comparing the standard model  
 600 (orange curve) with the experiment without FFGM (PISCES-GM, red curve), we de-  
 601 deduce that the explicit representation of FFGM alters strongly the latitudinal distribu-  
 602 tion of this relative contribution. It is significantly increased at all latitudes. This increase  
 603 is particularly important in the low latitudes where the contribution goes from less than  
 604 20% when FFGM are not allowed to grow (PISCES-GM) to more than 45% in the refer-  
 605 ence simulation PISCES-FFGM. Furthermore, export due to GM is maximal at about  
 606 40°N and S. When FFGM are included, the contribution of total macrozooplankton is  
 607 relatively constant between these latitudes. This result highlights the strong efficiency  
 608 of FFGM at exporting organic matter to the deep ocean, in particular in oligotrophic  
 609 regions with low productivity. The addition of FFGM reduces the contribution of GM  
 610 at all latitudes, especially at mid and low latitudes in which the contribution losses 15  
 611 to 20% (7). This reduction results from the competition between FFGM and GM.

612 Clogging reduces the contribution of FFGM to total export of carbon (from 21 to  
 613 12% at 1000m, table 4, fig. 7). It was also shown to improve the agreement of the sim-  
 614 ulated FFGM distribution with observations (Figure 5). In contrast, the latitudinal and  
 615 vertical distributions of total macrozooplankton contribution to particulate carbon ex-  
 616 port are not strongly affected by this process (green curves Figure 5, spatially homoge-  
 617 neous reduction of the contribution by  $\approx 5\%$ ).

## 618 4 Discussion

619 We added explicit representation of two macrozooplankton groups in PISCES-FFGM:  
 620 a generic macrozooplankton group, for which the parameterization is based on an allo-  
 621 metric scaling of the mesozooplankton group already existing in PISCES-v2 ((Aumont  
 622 et al., 2015), see section 2.3) and which feed mainly on the latter, and an FFGM group  
 623 that can feed on phytoplankton as microzooplankton. The introduction of FFGM into  
 624 PISCES, based solely on the representation of their specific diet due to the filter-feeding  
 625 mode, provided some insights into the potential impacts of FFGM on planktonic com-

munities and carbon cycling at the global scale through trophic effects (e.g. competition with generic macrozooplankton) and efficient carbon export.

#### 4.1 FFGM distribution and biomass

To evaluate the modeled FFGM biomasses, we compiled data from different sources (section 2.4) to produce a gridded climatology of large pelagic tunicates. Our AtlantECO dataset is based on similar observations as the previously compiled dataset (Luo et al., 2020, 2022), but we used a different approach to convert abundances to biomasses by taking into account the taxonomic information available on the samples, even when the species is not given.

Our model predicts a median biomass of FFGM similar to our dataset (0.80 vs. 1.11 mg C m<sup>-3</sup>), and reproduces 91% of the areas where biomass is high ( $\geq 0.5$ ) (Table 3). The introduction of a clogging mechanism, which would represent a saturation of the salp filtering apparatus for high prey concentrations, improves the representation of low biomass areas (section 2). In PISCES-CLG, a sensitivity experiment in which the clearance rate is decreased for chlorophyll concentrations above 0.5  $\mu\text{mol L}^{-1}$ , the Spearman correlation coefficient is doubled when comparing simulated and observed FFGM concentrations. Note however that this clogging mechanism and its impact on pelagic tunicates growth is largely under-documented, and rely on 30-yr old publications (Harbison et al., 1986; Fortier et al., 1994).

However, our modeled variability of the spatial distribution of FFGM was 25 times lower than the observed variability (Table 3). This large variability in observations has already been described in previous compilations of pelagic tunicates observations (Luo et al., 2020, 2022). Numerous aspects may contribute to the high variability of observations compared to models: scarcity of the observations, design of the sampling strategy (Hjøllo et al., 2021), biases in the sampling and enumeration methods (Frank, 1988; Mack et al., 2012), use of species- and location-dependent conversion factors (Arhonditsis & Brett, 2004), differing definitions of the compared groups or communities and the scale of investigation (local measurements are compared to average 5x5° estimates). Indeed, zooplankton patchiness increases with organism size (E. T. Buitenhuis et al., 2013). Physical (mesoscale and submesoscale processes) and biological (diel vertical migrations, predator avoidance, food patches, mate search) drivers combine to drive zooplankton patchiness (Folt & Burns, 1999). Although the introduction of a macrozooplankton compartment (namely cnidarian jellyfish) has been shown to increase patchiness in a recent modeling study (Wright et al., 2021), the spatial resolution ( $\tilde{2}$  degrees) of our model setup, and the lack of key biological processes (e.g., complex life cycle and high clearance rates) in our model likely preclude representation of such patchiness.

After the addition of FFGM in PISCES, our simulation results consistently show that FFGM dominate macrozooplankton in low-productivity regions, but that absolute abundances of FFGM are nonetheless higher in productive areas of the world ocean (Fig. 4). In a recent study using the COBALTv2 biogeochemical model, Luo et al. (2022) explored the role of pelagic tunicates in the marine ecosystem, with the addition of two new plankton functional groups, *i.e.* a large salp/doliolid group similar to our FFGM, and a small appendicularian group (Luo et al., 2022). They showed that the FFGM:GM ratio in their model follows a decreasing relationship with chlorophyll, consistently with our modeled FFGM:GM ratio patterns. To better reproduce the relationship between AtlantECO FFGM biomass and chlorophyll from the OC-CCI product, the addition of clogging was needed in our model (Fig. 5 and section 3.2). Given the paucity of data, it is currently difficult to evaluate these model insights from macrozooplankton databases alone. Heneghan et al. (2020) showed that salps dominate other macrozooplankton groups in low-productivity regions, but, contrary to our model results, these authors also showed that these organisms are more abundant in absolute terms in these low-productivity re-

Source		PISCES-FFGM (Luo et al., 2022)		(Luo et al., 2020)	(Luo et al., 2020)	PISCES-LOV	(Lebrato et al., 2019)	(Henschke et al., 2016)
		model	model	data-driven	data-driven	model	data-driven	data-driven
	Type of study			1000	800	30	800-1200	0-1700
	Ca_FFGM sinking speed	m d <sup>-1</sup>	800	100	-	-	-	-
	Fp_FFGM sinking speed	m d <sup>-1</sup>	1000	100	650	100	30	490-4000
Biomasses	Vertically integrated biomass	TgC	133	102*	-	-	-	-
	Upper 100m biomass	TgC	48.5	81.5	-	-	-	-
Surface Ocean POC export	Total grazing by FFGM	Pg C yr <sup>-1</sup>	0.94	-	6.6	-	-	-
	Predation on FFGM by UTL	Pg C yr <sup>-1</sup>	0.15	0.1	0.94	-	-	-
	FFGM POC Prod. top 100 m	Pg C yr <sup>-1</sup>	0.42	0.79	3.91	3.91	0.44	<0.04*
	Ca_FFGM contrib. to POC	%	35%	20%	20%	-	-	-
	Fp_FFGM contrib. to POC	%	65%	80%	80%	-	-	-
	FFGM driven POC exp. 100m	Pg C yr <sup>-1</sup>	0.43	0.57	2.7	1.3	0.36	-
	FFGM export efficiency	%	100%	72%	69%	33%	82%	-
	FFGM contrib. to POC100	%	6%	9%	20%	10%	5%	-
	Dif. in POC100 (with vs without FFGM †)	%	-2%	+2%	-	-	-	-
	Dif. in tot MAC contrib. to POC100 (with vs without FFGM †)	%	+55%	+41%	-	-	-	-
	Dif. in GM contrib. to POC100 (with vs without FFGM †)	%	-19%	-11%	-	-	-	-
	FFGM driven POC exp. 1000m	Pg C yr <sup>-1</sup>	0.42	-	1.4	0.33	0.11	<0.02-0.03*
	FFGM driven POC exp. Seafloor	Pg C yr <sup>-1</sup>	0.39	-	0.86	0.17	0.002	<0.01*
	FFGM POC Teff 100m to 1000m	%	97%	-	52%	25%	30%	46-54%
Yearly max. FFGM POC exp. ‡	mg C m <sup>-2</sup>		141 (min : 0.34 , max : 1580)	-	-	38 (min : 0.30 , max : 323)	-	128 - 6725 (min : 0.6 - 1171 , max : 656 - 77 143)

**Table 5. Comparison of parameters related to the impact of FFGM on the carbon cycle between different global scale studies based on data and/or models.** Ca\_FFGM is for FFGM carcasses. Fp\_FFGM is for FFGM fecal pellets. UTL is for Upper Trophic Levels. POC is for Particulate Organic carbon. Prod. is for Production. Contrib. is for contribution. Dif. is for Difference. Export efficiency is the ratio between the POC export below 100 m and the POC production in the upper 100 m. POC100 is for total POC export below 100m. exp. is for export to. Teff is for transfer efficiency. Tot MAC is for total macrozooplankton (GM + FFGM). \* Lebrato et al. (2019) consider also cnidarians and ctenophores. † Luo et al. (2022) integrate FFGM biomass includes appendicularians. ‡ We assume that our comparison between PISCES-FFGM and PISCES-GM is consistent with Luo et al. (2022)’s comparison between GZ-COBALT and COBALTv2. ‡ (Henschke et al., 2016) provides an estimate of POC export at 1000 m during a localized 1-month duration swarm event, the range is based on the spread of the results considering different species. We compare those values to the yearly maximum FFGM-driven POC export at 1000 m in our model, the range is based on the spread of the results considering all different grid cells.

677 regions than elsewhere in the ocean. Yet, they don’t explore the processes that could drive  
 678 this distribution. As evidenced by our PISCES-CLG experiment, clogging may be a po-  
 679 tential explanatory mechanism but the evidence for this process is weak. Future stud-  
 680 ies are needed to determine the processes involved in limiting FFGM biomass at high  
 681 chlorophyll concentrations.

## 682 4.2 FFGM contribution to the biological pump

683 Our modeled FFGM have a weak impact on phytoplankton and microzooplank-  
 684 ton biomasses, due to the low predation pressure they exert on these low-trophic levels  
 685 (grazing flux of 1 Pg C yr<sup>-1</sup>, which represents less than 3% of primary productivity).  
 686 Nevertheless, due to the high sinking speed of FFGM-derived fecal pellets and carcasses,  
 687 FFGM substantially increase the carbon export ratio and transfer efficiency. We com-  
 688 piled results from distinct studies on global biogeochemical impacts of FFGM in table  
 689 5 to support our results.

690 *4.2.0.1 Surface ocean particulate organic carbon production and export:* The  
 691 overall PISCES-FFGM modeled production of POC by FFGM in the upper 100 m is 0.42  
 692 Pg C yr<sup>-1</sup> (Table 5). This value falls within the range of data-driven estimates (Table  
 693 5). It is an order of magnitude above the value of 0.03 Pg C yr<sup>-1</sup> from Lebrato et al.  
 694 (2019)’s study, presented as a lower bound estimate due to their conservative assump-  
 695 tion of equivalence between GZ annual production and total GZ biomass. On the other  
 696 hand, our simulated FFGM POC production within the top 100 m is 10 times lower than  
 697 the estimate of 3.9 Pg C by Luo et al. (2020). In this study, FFGM production was forced

698 offline by modeled phytoplankton and zooplankton climatologies, so that FFGM preda-  
 699 tion had no feedback on their prey biomass. Luo et al. (2020)'s production estimate can  
 700 be seen as an upper estimate as GZ-induced predation pressure would affect the biomass  
 701 of other trophic levels in a fully-coupled model, thus affecting the gelatinous biomass it-  
 702 self and the induced carbon fluxes. Indeed the higher FFGM POC production is mostly  
 703 due to a higher FFGM grazing in their study ( $6.6 \text{ Pg C yr}^{-1}$  compared to our modeled  
 704 value of  $1 \text{ Pg C yr}^{-1}$ , Table 5). Finally, our modeled FFGM impacts on upper ocean POC  
 705 are similar to those by (Luo et al., 2022) based on COBALT-GZ: the simulated produc-  
 706 tion of detritus by FFGM in the first 100 m in our model is twice lower than in Luo et  
 707 al. (2022)'s model and the effective export of this detritus at 100 m is 30% lower (Ta-  
 708 ble 5). The smaller difference in export than in production lies in the use of a 10 times  
 709 lower particle sinking speed and a 20 times higher remineralization rate in COBALT-  
 710 GZ (Stock et al., 2014) compared to PISCES-FFGM, resulting in a lower production ex-  
 711 port efficiency in COBALT-GZ than in PISCES-FFGM (Table 5). Note that appendic-  
 712 ularians in GZ-COBALT produced 4 times less detritus in the upper 100m than large  
 713 tunicates, which supports our choice to represent only FFGM (i.e. macrozooplankton)  
 714 and not filter-feeding mesozooplankton in our biogeochemical model.

715 The impact of an explicit representation of FFGM on POC export is negligible in  
 716 both models when compared to a version without FFGM (+/- 2%, Table 5). But the  
 717 contribution of macrozooplankton to POC fluxes increases significantly in both models  
 718 (GZ-COBALT: +41%, PISCES-FFGM: +55%, Table 5) and this despite the simulated  
 719 decrease in export by GM (-11% in GZ-COBALT, -19% in PISCES-FFGM, Table 5),  
 720 so that the contribution of FFGM only to POC export at 100 m in both models is more  
 721 than 5% (Table 5). Thus, we can reasonably state that the representation of FFGM in  
 722 a biogeochemical model redistributes the carbon particles between the different compart-  
 723 ments over the top 100 m (more of very large particles from macrozooplankton, less of  
 724 small particles from smaller organisms) without significantly altering the total amount.  
 725 This change in particles composition is key to the major role that FFGM play in the ex-  
 726 port of carbon to the deep ocean.

727 *4.2.0.2 Deep ocean particulate organic carbon export:* FFGM have a modest im-  
 728 pact on subsurface export (less than 10 % of the global POC export at 100 m depth),  
 729 but this impact is highly increasing with depth, reaching much higher values at the seafloor  
 730 (>40%) and suggesting that FFGM play a key role in carbon storage in the deep ocean.  
 731 We also demonstrated that surface FFGM productivity and the transfer efficiency of FFGM-  
 732 driven POC are key processes that strongly affect the magnitude and distribution of deep  
 733 POC export.

734 The FFGM-driven export of POC at 1000 m (resp. seafloor) of 0.42 (resp. 0.39)  
 735  $\text{Pg C yr}^{-1}$  falls between the low value of 0.02 (resp. 0.01)  $\text{Pg C yr}^{-1}$  proposed by (Lebrato  
 736 et al., 2019) and the much larger estimate of 1.4 (resp. 0.86)  $\text{Pg C yr}^{-1}$  given by (Luo  
 737 et al., 2020) (Table 5). The quite large differences between these estimates are mainly  
 738 explained by the evaluation of surface FFGM productivity: FFGM productivity is 10  
 739 times higher in (Luo et al., 2020)'s study than in ours. In contrast, Lebrato et al. (2019)  
 740 used for gelatinous zooplankton a biomass estimate of 38  $\text{TgC}$  provided by Lucas et al.  
 741 (2014), which resulted in low export values ( $<0.04 \text{ Pg C yr}^{-1}$ ) at all levels of the wa-  
 742 ter column.

743 In addition to surface productivity, the efficiency of POC transfer is critical to the  
 744 absolute value of POC export at depth. The sinking velocity of particles is a key factor  
 745 that strongly controls this efficiency. In the studies of (Lebrato et al., 2019) and (Luo  
 746 et al., 2020), where the sinking velocities are greater than  $650 \text{ m d}^{-1}$ , the transfer effi-  
 747 ciency is about 50% (Table 5). It is reduced to 25% when the FFGM fecal pellets (which  
 748 account for 80% of FFGM detritus in their study) velocity is reduced to  $100 \text{ m d}^{-1}$ . The  
 749 same finding was observed in reducing the velocity from  $800\text{-}1000 \text{ m d}^{-1}$  to  $30 \text{ m d}^{-1}$   
 750 in our experiment PISCES-LOWV, where the transfer efficiency from 100 to 1000 m de-

751 creases from 97% to 30%. However, due to the use of a low remineralization rate, our  
 752 simulated transfer efficiency from 100 to 1000 m is very high compared to (Luo et al.,  
 753 2020) for similar carcasses and fecal pellets sinking speeds (Table 5). Still, our transfer  
 754 efficiency in PISCES-FFGM fits the vertical profiles of depth attenuation of jelly-driven  
 755 organic matter export proposed by Lebrato et al. (2011) for high sinking velocities and  
 756 low remineralization rates.

757 Last but not least, PISCES-FFGM seems to capture the intensity and part of the  
 758 variability of the intense carbon export events described by Henschke et al. (2016) linked  
 759 to short time proliferation events of FFGM: they estimated the export potential at 1000  
 760 m of different salps species during a 1 month swarm. Mean values ranged from 128 to  
 761 6725 mg C m<sup>-2</sup> depending on the species, the minimum from 0.6 to 1171 mg C m<sup>-2</sup> and  
 762 the maximum from 656 to 77 143 mg C m<sup>-2</sup>. We compare these results to the annual  
 763 maxima of the FFGM carbon export simulated at each grid point by our model (Table  
 764 5). The values obtained range from 0.34 to 1580 mg C m<sup>-2</sup> with a spatial mean of 141  
 765 mg C m<sup>-2</sup>, which is consistent with the species-range of mean, min and max in their study  
 766 (Table 5). This also supports our choice of a very low remineralization rate and high fall  
 767 rates. The latter is confirmed with the PISCES-LOV experiments in which modeled ex-  
 768 port maxima fall below the min, mean and max ranges of Henschke et al. (2016)'s study.

### 769 4.3 Model limitations in representing GM and FFGM

770 *4.3.0.1 Representation of patchiness.* Patchiness is particularly strong for gelati-  
 771 nous zooplankton. Indeed, they present very high growth and clearance rates and can  
 772 therefore efficiently and rapidly exploit their environment under favorable conditions,  
 773 with localized swarming and thus patchiness (Graham et al., 2001; Purcell, 2009; Lilley  
 774 et al., 2011; Lucas et al., 2014). However, in the current model, increasing clearance rates  
 775 or growth rates of FFGM without adequate modifications of FFGM mortality rates would  
 776 inevitably cause the generic macrozooplankton population to collapse because it would  
 777 be outcompeted by FFGM everywhere except in the mesopelagic and deep ocean. To  
 778 further investigate the effect of high growth rates and clearance rates of FFGM, a bet-  
 779 ter understanding of the physiological and environmental drivers of the FFGM mortal-  
 780 ity processes triggering the end of their swarms seems essential, as their causes are mul-  
 781 tiple and too poorly documented to be currently modeled (Pitt et al., 2014).

782 *4.3.0.2 Representation of seasonal variability* Our standard PISCES-FFGM sim-  
 783 ulation shows an approximate one-month lead in the seasonal biomass peak of FFGM  
 784 compared to GM, this lag being consistent at the global scale to that of the food of the  
 785 two groups (Figure S6). This suggests that the filter-feeding mode of FFGM may have  
 786 an impact on the temporal dynamics of the FFGM-driven POC flux. However, it is dif-  
 787 ficult to give a high confidence level to this statement because the spatial distributions  
 788 between the lags of the organisms and their food are very patchy and the temporal vari-  
 789 ability of the prey does not correspond to that of the corresponding groups when focus-  
 790 ing on specific regions (Figure S6).

791 Furthermore, the data temporal resolution is insufficient to validate these seasonal  
 792 patterns: only 7% of the grid points in the AtlantECO climatology are derived from data  
 793 covering at least 6 distinct months.

794 Also, life-cycle are currently not represented in the model despite that it can sig-  
 795 nificantly affect the temporal dynamics of a BGC-model (Clerc et al., 2021): most FFGM  
 796 have a complex life-cycle, with an alternation between a sexual and asexual phase that  
 797 could be a major driver of their population dynamics (Henschke et al., 2016). A single-  
 798 species observation based study on *Thalia democratica* in South-East Australia suggested  
 799 that life history characteristics such as asexual reproduction and growth are associated  
 800 with inter-annual variations in abundance and thus may be major factors determining  
 801 population dynamics, in particular swarm magnitude (Henschke et al., 2014). Inclusion

802 of such life cycle traits in a single-species model of *Salpa thompsoni* in the Southern Ocean  
 803 helped understand the seasonal and interannual variability of salp abundance (Henschke  
 804 et al., 2018). These studies are focused on one species and one region, and the inclusion  
 805 of their life cycle in a global model where FFGM constitute a single compartment would  
 806 require a multispecies large scale evaluation of the FFGM life cycle role in the tempo-  
 807 ral dynamics of the swarming process.

808 *4.3.0.3 Representation of deep carbon export* One of the greatest sources of un-  
 809 certainty about the export of carbon from FFGM to the deep ocean is the transfer ef-  
 810 ficiency (see Table 5), which depends primarily on remineralization rates and sinking speeds.  
 811 This raises questions about the processes that could affect the fate of carcasses and fe-  
 812 cal pellets (CAFP) as they sink. At a given temperature, our simple FFGM represen-  
 813 tation includes constant remineralization of CAFP and consumption through filter feed-  
 814 ing by GM (Eq. S14 and S15). The induced losses are very low compared to FFGM’s  
 815 CAFP production rates (<5%). However, predation by scavengers could significantly af-  
 816 fect CAFP during their fall (Dunlop et al., 2018; Scheer et al., 2022). Benthic consump-  
 817 tion by scavengers is well documented for jellyfish carcasses (A. K. Sweetman et al., 2014;  
 818 Henschke et al., 2013), but their fate in the vertical column is largely unknown. Also,  
 819 most measured sinking speed values are based on small (a few meters) sinking column  
 820 experimental setup and thus do not account for any degradation process (Lebrato et al.,  
 821 2013). A clear understanding of FFGM carcasses and fecal pellets fate is needed to prop-  
 822 erly estimate their deep ocean impacts.

823 *4.3.0.4 Deep nutrient fields* Our model results suggest that export values of the  
 824 order of what we found here and of those reported in (Luo et al., 2020, 2022) could con-  
 825 siderably affect nutrient fields in the deep ocean. This effect would be apparent only in  
 826 long spinup simulations of a global biogeochemical model. Indeed, in our PISCES-FFGM’s  
 827 500-year-long simulation, deep nutrient fields keep drifting away from the initial state  
 828 after hundreds of years, ending up in degraded bottom nutrients fields as compared to  
 829 observations (Figure S5).

830 *4.3.0.5 Conclusion* We explicitly represented large pelagic tunicates in the global  
 831 marine biogeochemistry model PISCES and evaluated the simulated distribution of FFGM  
 832 by compiling available observations of FFGM abundance into a FFGM biomass clima-  
 833 tology using a taxonomy-resolving biomass-abundance conversion. Representation of FFGM  
 834 in a marine biogeochemical model has a small impact on total detritus production in the  
 835 first 100 m. Yet, 6% of this production is due to FFGM, a small yet significant number.  
 836 Due to their high sinking speeds, almost all of the organic matter produced by FFGM  
 837 is transported to the deep ocean. Therefore, FFGM carcasses and fecal pellets dominates  
 838 the export of organic matter in the deep ocean (e.g. 70% at 5000m). The spatial dis-  
 839 tribution of FFGM-driven export differs from that of the other macrozooplankton group,  
 840 GM, which also contributes significantly to export at depth (25% at 5000m). Indeed, due  
 841 to their filter-feeding mode of predation, access to preys of variable size allows FFGM  
 842 to better exploit low productivity environments than GM, especially in subtropical olig-  
 843 otrophic gyres, where FFGM are twice as abundant as GM and thus contribute 5 times  
 844 more to POC export at 1000m.

845 A more detailed inclusion of the processes involved in the bloom-and-burst dynam-  
 846 ics of FFGM (e.g. life cycle, clogging, high clearance rates) will be necessary to better  
 847 understand the spatial and temporal variability of their impacts on carbon export and  
 848 ecosystem structure. Still, a promising perspective would be to run our PISCES-FFGM  
 849 model forced by climate projections. As climate change could favor small phytoplank-  
 850 ton (Peter & Sommer, 2013), we could expect an amplification of the spatial pattern we  
 851 currently described: FFGM could be even more favored in low productive regions.

## 5 Open Research

### Availability statement

This section needs to be completed. All raw and gridded data sets will be made publicly available in open access within the framework of the European H2020 project AtlantECO (grant agreement no 862923). Preliminary DOIs can be made available to the reviewers upon request. All model outputs necessary to reproduce the results in this manuscript will be made publicly available.

### Acknowledgments

We are very grateful to Lars Stemmann, Olivier Maury, Jean-Christophe Poggiale and Fabien Lombard for insightful comments during the development of this manuscript and to Christian Ethé and Olivier Torres for setting up the model configuration.

This project used the HPC resources of TGCC under the allocation NUMBER (project gen0040) provided by GENCI (Grand Equipement National de Calcul Intensif). This study benefited from the ESPRI (Ensemble de Services Pour la Recherche l'IPSL) computing and data center (<https://mesocentre.ipsl.fr>) which is supported by CNRS, Sorbonne Université, Ecole Polytechnique, and CNES and through national and international grants.

This study has received funding from the Agence Nationale de la Recherche grant agreement ANR-17-CE32-0008 (CIGOEF).

MV and FB acknowledge funding from the European Union's Horizon 2020 research and innovation programme under grant agreement no. 862923. This output reflects only the author's view, and the European Union cannot be held responsible for any use that may be made of the information contained therein.

LB acknowledges support from the European Union's Horizon 2020 research and innovation COMFORT (grant agreement No 820989), ESM2025 (grant agreement No 101003536) and from the Chaire ENS-Chanel.

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