

1 **Is Computational Oceanography Coming of Age?**

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ABSTRACT

13 Computational Oceanography is the study of ocean phenomena by numerical simulation, especially
14 dynamical and physical phenomena. Progress in information technology has driven exponential
15 growth in the number of global ocean observations and the fidelity of numerical simulations of the
16 ocean in the past few decades. The growth has been exponentially faster for ocean simulations,
17 however. We argue that this faster growth is shifting the importance of field measurements and
18 numerical simulations for oceanographic research. It is leading to the maturation of Computational
19 Oceanography as a branch of marine science on par with observational oceanography. One
20 implication is that ultra-resolved ocean simulations are only loosely constrained by observations.
21 Another implication is that barriers to analyzing the output of such simulations should be removed.
22 Although some specific limits and challenges exist, many opportunities are identified for the future
23 of Computational Oceanography. Most important is the prospect of hybrid computational and
24 observational approaches to advance understanding of the ocean.

25 *Capsule summary.* Fast growth in the fidelity of ocean general circulation models is driving the
26 maturation of Computational Oceanography as a branch of marine science on par with observational
27 oceanography.

28 **1. Introduction**

29 Computational Oceanography is the study of ocean phenomena by numerical simulation, es-
30 pecially dynamical and physical phenomena using ocean general circulation models (OGCMs).
31 One early pioneer of this field wrote of the 1960s, 1970s, and 1980s as the “birth”, “infancy”,
32 and “adolescence” of OGCMs, respectively (Bryan 2006, see also Holland and McWilliams 1987;
33 McWilliams 1996). Similarly, the authors of a comprehensive review of OGCMs wrote at the turn
34 of the century “this field...has entered an era of healthy adolescence” (Griffies et al. 2000). With
35 twenty more years of data, this essay explores the continued growth of OGCMs and speculates
36 on their prospects. We ask: Is Computational Oceanography entering a new era that signifies its
37 coming of age?

38 For motivation, Fig. 1 compares oceanographic measurements and results from a high-resolution
39 OGCM. The region of interest is a topographic constriction called the Denmark Strait, between
40 Greenland and Iceland. The Denmark Strait Overflow (DSO) flows south through this gap and is
41 an important current for the Atlantic Meridional Overturning Circulation and thus for the ocean’s
42 role in North Atlantic climate. The two timeseries in Fig. 1a show DSO volume flux (transport).
43 One timeseries is from *in situ* measurements, the other is from a high resolution regional OGCM
44 (and they have been processed similarly with similar smoothing). The question is this: Which is
45 which? Fig. 1b compares *in situ* hydrographic measurements along a section north of Denmark
46 Strait with a synthetic hydrographic section from the OGCM. And Fig. 1c shows the trajectories
47 of drifting oceanographic floats approaching Denmark Strait from the north and trajectories of

48 drifting particles in the OGCM released from the same locations. Again, the question is which is
49 the real data and which is the synthetic data? In each case, the field measurements and the OGCM
50 results are different, but identifying them is difficult.

51 These are examples of OGCM Turing tests. They are inspired by Alan Turing’s imitation game to
52 distinguish between, and correctly identify, a person and an intelligent machine. The game involves
53 asking questions through an interface that obscures whether the responses are from the person or
54 the machine (Turing 1950). The difficulty of the OGCM Turing tests in Fig. 1 reflects the small
55 systematic error in the OGCM and therefore its realism. Some OGCM solutions are reaching the
56 point that they are essentially indistinguishable from observations, so they pass Turing tests like
57 those in Fig. 1. In the words of Ed Lorenz, numerical experiments will eventually “duplicate the
58 circulation to any desired degree of accuracy” (Lorenz 1967).¹

59 With these themes in mind, this essay explores the growth of OGCMs and compares it to the
60 growth of ocean observations. The focus is on the computer science and information technology
61 improvements that contribute to the growth. We then speculate on limits, opportunities, and
62 prospects for OGCMs.

63 **2. Unequal Exponential Growth**

64 Two examples illustrate the growth of ocean observations. First, consider temperature observa-
65 tions in the global deep ocean over the last half century. Fig. 2a shows the cumulative number
66 of temperature observations deeper than 1000m. They have grown exponentially (notice the y
67 axis is logarithmic). Averaged over the last century, the exponential growth has a doubling period
68 of 10.4 years, giving an approximately 60-fold expansion in the deep temperature database since
69 1960. Technology transitions have maintained this exponential growth, specifically, advances

¹The prescient Lorenz was writing about atmospheric models in the late 1960s, but the message applies to OGCMs today.

70 in micro-electronics and information technology. In the 1990s conductivity-temperature-depth
71 (CTD) sensors on autonomous profiling floats took over from ship CTD sampling, for example,
72 leading in the 2000s to the transformative Argo global float network (Argo 2020).

73 Second, consider the history of sea level observations from satellite altimeters. Sea level data
74 have revolutionized physical oceanography by providing information on the surface circulation,
75 mesoscale eddies, tides, and sea level change. Fig. 2b shows the sequence of altimeter missions
76 (colored bars) and the cumulative number of observing days (black line). The number of observing
77 days reveals the growth in sea level observations (although there is great variety between missions).
78 The number of sea level observations has grown nearly exponentially since the mid 1980s with a
79 doubling time of about 8.1 years and a ≈ 20 -fold expansion in the sea level database since 1985.
80 Again, micro-electronic and information technology advances have maintained this growth.

81 Technology advances have also fueled growth in the fidelity of OGCMs. For example, Fig. 2c
82 shows the history of global OGCM resolution. The black dots show five pioneering (cutting-edge)
83 models over the last 40 years. The Bryan and Lewis (1979) model had a peak resolution of 2.4°
84 with 12 vertical levels and the Rocha et al. (2016) model had a peak resolution of 0.02° with 90
85 vertical levels. The growth in OGCM resolution (number of grid points) is exponential with a
86 doubling time of 2.2 years and a 10^5 -fold increase since 1980. We also show the global ocean
87 models from the Intergovernmental Panel on Climate Change (IPCC) reports. The peak resolution
88 of the ocean OGCMs in the first IPCC report was 2.7° with 9 vertical levels and the peak resolution
89 in the latest (sixth) IPCC report is 0.067° with 75 vertical levels.² This growth is also exponential
90 with a doubling time of 2.8 years. For the most highly-resolved models in each assessment, the
91 doubling time is close to the cutting-edge OGCM doubling time.

²The AR6 data points on Fig. 2 are from the HighResMIP experiments, which is a sub-project on high-resolution models that does not run the full suite of CMIP6 experiments.

92 Now compare the horizontal resolution of ocean measurements with OGCM resolution. The
93 Argo profiling float network operates about 4000 floats at any one time. Each float makes a vertical
94 profile from 2000m depth to the surface every ten days. The global average spacing of profiles
95 is therefore 300 km.³ The spacing between altimeter tracks for the TOPEX/Poseidon and Jason
96 satellite altimeters is also about 315 km (at the equator), with a repeat period of ten days. The
97 present day peak OGCM resolution of $0.02^\circ \approx 2$ km is therefore 140 times higher.⁴

98 **3. Prospects for Future Growth**

99 Looking ahead, the future is bright for the Argo network. The reason is that Argo is part of
100 the Global Ocean and Global Climate Observing System, which implements the Paris Agreement
101 on climate change and United Nations sustainable development goals. New capacities, like deep
102 profiling floats, and new technologies, like biogeochemical sensors, are planned over the next few
103 years (GCOS 2016). It is unclear how the network can double in size in the next decade and
104 maintain long term exponential growth, but it is plausible.

105 The future is also bright for sea level measurements. The Surface Water and Ocean Topography
106 (SWOT) mission, scheduled for launch in 2022, will start a new era of sea level observation. SWOT
107 will observe sea level over a swath, rather than over a single patch. It will have 15 km resolution,
108 or better, covering most of the global ocean every 21 days (Morrow et al. 2019). It will improve
109 the spatial resolution of sea level data by a factor of about ten. Therefore, the prospects for the
110 altimetry record to continue growing exponentially in the 2020s are good.

³The vertical resolution of Argo profile data is about 5m, which is about 7 times higher than the best AR6 OGCMs and about 3 times higher than the Poseidon Project run mentioned below.

⁴This comparison avoids the issue of time dependence in the circulation. It simply (and conservatively) imagines the Argo and altimetry data from one ten day period are used to constrain the time-mean OGCM state over that period.

111 For OGCMs, resolution improves as supercomputer technology advances. Historically, that
112 follows Moore’s “law,” which says that transistor density in microprocessors doubles every two
113 years (Moore 1975). For instance, machines first achieved petaflop speeds (10^{15} floating point
114 operations per second) in 2008 and exaflop speeds (10^{18}) in March 2020, a doubling every 1.1 years
115 (see Fig. 2d). Computers available to the oceanographic, atmospheric, and climate communities
116 are less powerful. Still, the machines at NCAR and ECMWF⁵ also show exponential growth over
117 recent decades with a doubling every 1.1 years, albeit lagging the cutting-edge machines by about
118 five years (Fig. 2d). On this basis, the OGCM resolution will probably continue to double every 2.2
119 years, at least for several more years (assuming funding remains at historic levels). It is reasonable
120 to expect cutting-edge exascale global OGCMs with horizontal resolution around one kilometer
121 by the mid 2020s. After that, with widespread anticipation that Moore’s law will end (Waldrop
122 2016), future growth is uncertain.

123 **4. Maturation of Computational Oceanography**

124 This evidence shows that information technology advances are driving exponential growth in
125 ocean observations and exponential growth in OGCM resolution. But the OGCM growth rate is
126 faster. Therefore, OGCM resolution is also growing exponentially faster than the growth in ocean
127 field data. In 1990, OGCMs were obviously biased compared to measurements, for example, of
128 deep temperatures or sea level. In 2020, OGCMs are achieving resolutions that are substantially
129 greater than the gaps between measurements, at least for some regimes, like deep and abyssal
130 ocean currents. We should expect this trend to continue for the foreseeable future (the next
131 several years). Therefore, the question arises: When, and in what ways, will OGCMs become as
132 important as observations for advancing knowledge in physical oceanography? Historically, most

⁵Meaning the U.S. National Center for Atmospheric Research and the European Centre for Medium Range Weather Forecasts.

133 knowledge came from observations of the real ocean.⁶ The growth of OGCMs suggests that the
134 field is approaching an era in which numerical circulation models are as important as observations
135 for advancing knowledge. For example, diagnosing and understanding the rectified effects of
136 mesoscale eddy variability on the large-scale, low-frequency circulation will probably rely heavily
137 on high-resolution OGCMs.

138 What are the criteria to claim that OGCM solutions should be treated, in some cases, as seriously
139 as real measurements? Realizing them would mark the maturation of Computational Oceanography.
140 These criteria are on our checklist:

141 **1. Confidence in the fidelity of the basic tools and methods.** Consider two types of tool:
142 First, consider the theoretical definition of the ocean circulation problem. Computational
143 Oceanography relies on software to compute approximations to the incompressible rotating-
144 stratified Navier Stokes equations, with equations for the conservation of dissolved salts and
145 heat (McWilliams 1996; Griffies 2004; Fox-Kemper et al. 2019). There is little doubt that
146 these are the right equations for ocean circulation. The software is mature, stable, and diverse.
147 The issue of unresolved processes, and parametrizing their effects remains an important area
148 of research. For example, it is still unclear how to represent the impacts of unresolved
149 submesoscale processes on the larger scale flow. Although much progress has been made on
150 this problem in the last 30 years (Gent 2011; Le Sommer et al. 2018; Fox-Kemper et al. 2019),
151 resolution improvements have surely played an essential part in refining OGCM accuracy
152 (Griffies et al. 2000). In other words, we believe that the problem of parametrizing unresolved

⁶For example, Stewart (2008) writes: “The theory describing a convecting, wind-forced, turbulent fluid in a rotating coordinate system has never been sufficiently well known that important features of the oceanic circulation could be predicted before they were observed. In almost all cases, oceanographers resort to observations to understand oceanic processes.”

153 scales is not so pathological that it contaminates all of the resolved scales.⁷ A corollary is
154 that OGCMs are less complicated than the real ocean, meaning that OGCM variability is a
155 lower bound on the variability in the real system. These are *de facto* working hypotheses of
156 all theoretical and numerical approaches to understanding the ocean circulation.

157 Second, we need tools to adjust OGCM solutions to agree with observations; that is, to solve
158 the data assimilation and state estimation problem (Bennett 1992; Wunsch 1996, 2006; Kalnay
159 2002). For example, state estimation is used to produce retrospective reanalyses (hindcasts) of
160 the time-evolving ocean state and data assimilation is used to initialize prospective forecasts
161 of the future. Although many questions remain open, these methods are also now mature,
162 stable, and diverse.

163 **2. The number of OGCM degrees of freedom exceeds the number of observational con-**
164 **straints.** This criterion concerns the state estimation and data assimilation problems. In
165 essence, it is about whether it is possible (in principle) to adjust an OGCM solution to fit
166 the observations exactly or not. If the OGCM can be adjusted to fit the data exactly, the
167 state estimation problem is under-determined. Otherwise, it is over-determined.⁸ The num-
168 ber of OGCM degrees of freedom scales as the number of grid points (for large numbers
169 of grid points). The number of observational constraints scales as the number of distinct
170 measurements. Fig. 2 shows evidence that the number of OGCM degrees of freedom per
171 observational constraint exceeds one because, loosely, the peak OGCM resolution is now 140

⁷It is likely that errors in parameterized physics *influence* all resolved scales, not least because of error growth due to deterministic chaos. But the issue is whether the errors in parameterized physics cause systematic errors in the resolved scales, such as biases in statistics of resolved quantities. It is reasonable to suppose that (i) resolution improvements and parameterization improvements reduce these systematic biases towards zero, and (ii) the systematic biases are not so bad as to preclude use of models to understand (and hindcast and predict) the natural system. Of course, these are quantitative (not qualitative) hypotheses that vary from case to case (models, parameterizations, resolved metrics, science questions).

⁸Ignoring the atypical case of the problem being exactly determined.

172 times higher than the Argo and Jason data spacing (see footnote 4). This gap is growing
173 exponentially because OGCM resolution is growing exponentially faster than data density.
174 Therefore, the state estimation problem is moving from (in principle) being over-determined
175 to being under-determined.⁹

176 Crossing this threshold has interesting implications: First, the systematic errors in OGCMs
177 disappear and they pass Turing or Feigenbaum tests (Turing 1950; Feigenbaum 2003; Harel
178 2005), like those in Fig. 1. That is, OGCM solutions become indistinguishable from obser-
179 vations of the real ocean and a subject-matter expert cannot tell them apart. Regional OGCM
180 simulations of the Denmark Strait Overflow (DSO) at resolutions of 0.5–2 km are approach-
181 ing this point (Magaldi and Haine 2015; Almansi et al. 2020; Saberi et al. 2020). Similarly,
182 regional high-resolution state estimates are nearly under-determined (Lea et al. 2006; Dwivedi
183 et al. 2011). A fair comparison (Turing test) requires that the space-time scales of the ob-
184 servations and the model results are the same, which means the power spectra should match.
185 This comparison is a necessary test to realize Lorenz’ vision quoted in the Introduction. It
186 is not a sufficient test, however, as the OGCM results can resemble the measurements for the
187 wrong reasons, but we take it as strong evidence of small OGCM bias.

188 Second, the OGCM solutions make accurate, testable predictions about the real ocean. His-
189 torically, advances from theoretical and numerical research in dynamical oceanography have
190 lagged advances from observational research (see footnote 6). Once OGCMs become under-
191 determined by data, it will be common for them to make predictions that can be tested by field

⁹It is possible to argue that any inverse problem with real observations is formally under-determined because the observational error can be considered as an unknown parameter to be solved for (Wunsch 1996; Stammer et al. 2002). Regardless, no global ocean circulation state estimate has characterized the null space associated with the indeterminacy (to our knowledge), or presented different solutions that fit the observations equally well. Instead, the practice has been to stop the state estimation once an acceptable fit has been achieved (Stammer et al. 2002; Nguyen et al. 2020).

192 programs. For example, DSO simulations show exchange of dense water out of the overflow
193 onto the east Greenland continental shelf, and vice versa (Magaldi et al. 2011). They also
194 show entrainment of near-surface waters south of Iceland into the DSO within a few months,
195 at least during hard winters (Saber et al. 2020). It remains to be seen if these predictions
196 occur in the real ocean.

197 **5. Limits to Computational Oceanography**

198 Although these opportunities are exciting, there are clear limits to Computational Oceanography.
199 First, direct numerical simulation (DNS) of the global ocean circulation is inconceivable today.
200 DNS in this context means running OGCMs that resolve all scales of motion; from the planetary
201 scale to the dissipation scale (around 1mm), and from centuries to seconds. DNS would avoid
202 the challenge of parametrizing the effects of the unresolved scales, but at vast computational cost.
203 Fig. 3 shows why. It shows the full range of space and time scales relevant to the ocean general
204 circulation, about ten orders of magnitude in both. It also shows the space time volumes accessible
205 to present-day supercomputers, including the best AR6 OGCMs shown in Fig. 2, the Poseidon
206 Project run,¹⁰ and turbulence simulations (DNS and large eddy simulations, LES). To span the
207 entire space time plane, supercomputers would need to resolve about 10^{25} grid points and 10^{10} time
208 steps. That is about 16 orders of magnitude more grid points than is possible today. Extrapolating
209 the doubling time of 2.2 years in Fig. 2c, it would take 120 years to achieve this increase, which is
210 impossible to envision. Clearly, the exponential growth must roll off at some point, and, clearly,
211 OGCM simulations cannot replace observations of the natural ocean.

¹⁰The Poseidon Project intends to run a global OGCM at (nominally) 1 km horizontal resolution (poseidon.idies.jhu.edu). The Poseidon Project is unrelated to the TOPEX/Poseidon altimeter.

212 Another potential limit concerns scalability of OGCM codes. Fig. 2 shows that the historic
213 doubling time for the number of OGCM grid points is about twice the doubling time for supercom-
214 puter speed. That value is close to the optimal ratio of $3/2$, which assumes that machine speedup is
215 spent on increasing horizontal resolution, that model timestep is inversely proportional to the grid
216 spacing (for numerical stability), and that all other factors are equal. In other words, the historic
217 OGCM growth has nearly maintained pace with the supercomputer acceleration. It is unclear how
218 this trend will continue, however, because of the overhead of communication from processor cores
219 to other cores, to memory, and to disk (Le Sommer et al. 2018). Moreover, exascale supercomput-
220 ers will not resemble petascale supercomputers: they will have different architectures and greater
221 diversity (Giles and Reguly 2014). These changes are driven by physical limits on clock speed and
222 power densities in silicon microprocessors, as well as economic forces. To harness exascale ma-
223 chines OGCM software must radically change (for discussion of this point for atmospheric general
224 circulation models, see Lawrence et al. 2018 and Gropp and Snir 2013). The developers of next
225 generation OGCMs should adopt collaborative, open community habits (Le Sommer et al. 2018).
226 Promising paths are to define and refine modular sub-components, and to develop domain-specific
227 languages, performance tools, and data models that separate different levels in the software stack
228 for optimization by experts (Lawrence et al. 2018). OGCM computational intensity (the fraction of
229 time spent performing floating point calculations versus memory operations) is low: Le Sommer
230 et al. (2018) estimate OGCMs run at 5% peak speed, for example. So there is potential to accelerate
231 OGCMs by reducing this bottleneck (for example by exploiting time parallelism, Schreiber et al.
232 2017; Hamon et al. 2020). Exploiting new application-specific hardware accelerators and new
233 OGCM solver paradigms, like lower precision (Palmer 2012; Palem 2014), will also be important.
234 These developments will mitigate the saturation of transistor density and the demise of Moore's
235 Law, and they offer hope to continue the refinement of OGCM meshes.

236 Finally, there are challenging issues to couple OGCMs to other parts of the Earth system at
237 horizontal resolutions around one kilometer. For example, air/sea interaction, sea ice dynamics,
238 and biogeochemistry are all poorly understood and hard to simulate at these scales.

239 **6. Opportunities for Computational Oceanography**

240 The opportunities for Computational Oceanography to advance marine science include:

- 241 • Migration from the study of specific instances of phenomena to the study of statistics of these
242 phenomena. The DSO is one of many currents that is affected by rotation, stratification,
243 and bathymetry. It is inconceivable to observe all of them, but they can all be simulated
244 in an exascale OGCM. Empirical characterization of these numerical overflows would be an
245 important step forward.

- 246 • Discovery and characterization of intermittent, time-dependent, three-dimensional phenom-
247 ena, which are hard to observe. Submesoscale currents are in this class, which occur at
248 horizontal scales shorter than several kilometers (Thomas et al. 2008). Diapycnal mixing is
249 another example, which occurs at scales shorter than meters (MacKinnon et al. 2017).

- 250 • Comprehensive and illuminating analyses of ocean mass, heat, salt, momentum, energy, and
251 vorticity budgets, in a way that is nearly impossible with direct observations.

- 252 • Discovery and characterization of ocean circulation regimes that cannot be observed. Ex-
253 amples include the circulation during the last glacial maximum (paleo-oceanography) or in
254 extra-terrestrial oceans (exo-oceanography). For these ocean circulation problems, the data-
255 sparseness challenge is much worse than for the modern ocean (LeGrand and Wunsch 1995;
256 Amrhein et al. 2018; Way et al. 2017). Criterion 2 was achieved with smaller computational

resources for these fields, and therefore they have already entered the era of Computational Oceanography by the rationale in section 4.

- Robust observing system design using OGCM solutions as synthetic data. These Observing System Simulation Experiments (Errico et al. 2012) should become the best-practice standard for fieldwork design. There are implications for making the OGCM output accessible and easy to work with (see below), but the payoff from engaging observational oceanographers is great.
- Insight from OGCM state estimation to support fieldwork, ideally in real time. The community should recognize the fact that the under-determined state estimates imply an infinite number of OGCM solutions that match the data exactly. This means that techniques are needed to characterize and handle the OGCM null space (indeterminacy). For example, observational oceanographers at sea could make decisions about where, when, and how to observe using OGCM information that captures the range of possible circulation states consistent with data. This practice is common in atmospheric science already.
- More efficient identification of interesting phenomena using automatic methods, like artificial intelligence and data mining (Kutz 2017; Lguensat et al. 2019). In fact, such automatic methods will become essential as the size of OGCM output grows exponentially and overwhelms manual feature identification (see below).
- Increasing transition of dynamical oceanography to an experimental (computational) science. It has long been recognized that idealized models isolate physical mechanisms relevant to the general circulation and thereby build dynamical understanding. We still require idealized models; in particular, we need a hierarchy of models that span the gap between geophysical fluid dynamics problems and realistic simulations of the circulation. This hierarchy will

280 ensure that the increasing OGCM realism does not outpace understanding of the basic physics
281 (Held 2005; Vallis 2016; Coveney et al. 2016; Emanuel 2020).

282 **7. Prospects for Computational Oceanography**

283 How can these priorities be achieved and what are the prospects for Computational Oceanogra-
284 phy? We should focus on these issues in the next several years:

- 285 • The indeterminacy of OGCM solutions by observations should be recognized—we should
286 “embrace the null space.” Imagine computing an ensemble of high resolution (high degrees of
287 freedom per observation) state estimates that fit the observations (exactly or equally well within
288 instrumental errors). These state estimates would differ, for example, in the characteristics
289 of their eddies, or in their deep circulations, or in their internal wave fields, or in their
290 diapycnal mixing. In such a situation, the different state estimates should all be treated
291 seriously. The ensemble would characterize the null space (indeterminacy) in the inverse
292 problem and therefore quantify the variety of ocean states consistent with observations and
293 ocean circulation physics. This vision for uncertainty quantification echoes the probabilistic
294 practice of ensemble atmospheric model runs to forecast the weather (see also McWilliams
295 2007; Le Sommer et al. 2018).
- 296 • Barriers to dissemination of OGCM simulation output should be lowered—we should “democ-
297 ratize the data.” The output should be freely available, including to non-professional users.
298 Traditionally, effort has focused on the challenges of calculating OGCM solutions with super-
299 computers. The OGCM output has become increasingly hard to use, because of the massive
300 data volume, and the technical complexities that attend the high-performance computation.
301 Access to high-resolution OGCM output is restricted to a few experts in practice.

302 The remedy is to build high-performance data science infrastructure to match the high-
303 performance compute infrastructure (Overpeck et al. 2011). These data portals should be
304 open and have low thresholds to getting started. We should be able to sample the simulations
305 the way that we sample the real ocean. For example, it should be easy for an observational
306 oceanographer to plot a synthetic hydrographic section or mooring timeseries. The data portals
307 should include open software and significant compute resources to process and analyze the
308 simulation data. We should avoid the inefficient practice in which users are forced to download
309 voluminous data to their local machines and then write their own code to analyze them.
310 Technologies and infrastructure to achieve these goals are under development, such as the
311 OceanSpy OGCM data analysis package (Almansi et al. 2019), the Pangeo community in
312 geoscience big data (pangeo.io), and the SciServer and JASMIN big data science platforms
313 (Medvedev et al. 2016, www.jasmin.ac.uk).

- 314 • “Benchmark” OGCM reference solutions should be computed using the best available com-
315 pute resources and served to the public. They are of intrinsic value to all oceanographers,
316 not just ocean modellers, for the reasons stated above. Benchmark solutions for regional
317 ocean circulation problems are valuable for the same reasons, as are idealized simulations
318 of specific ocean dynamical processes. The track record of other fields using this approach
319 is impressive. For instance, the Johns Hopkins Turbulence Database exposes cutting-edge
320 turbulence simulation data to researchers and provides easy-to-use interfaces to retrieve and
321 interact with the data using novel metaphors like immersing virtual sensors into the 4-D data
322 (turbulence.pha.jhu.edu; Perlman et al. 2007; Li et al. 2008).
- 323 • OGCMs will migrate to exascale compute resources in the next few years. This migration will
324 involve new paradigms to access the data. For example, with today’s petaflop supercomputers

325 only about 0.1% of the OGCM solution can be permanently stored for analysis. The problem
326 arises because of the prohibitive time needed to transfer the massive output volume to long-
327 term storage media, and the prohibitive expense of the media. This loss of OGCM data will
328 be much worse on exaflop machines.

329 To mitigate this problem consider the strategy adopted by the Large Hadron Collider (LHC),
330 the world’s most sophisticated experimental facility. The LHC provides a single source of
331 data on subatomic particle collisions. Several experiments tap the data stream in so-called
332 “beam-lines.” Within each experiment, customized hardware monitors the stream. Only
333 about one event in ten million is retained for storage and detailed analysis. In exascale
334 oceanography the analogous idea (see section 3.3.5 in Asch et al. 2018) is to enable automatic
335 identification of selected circulation events and trigger storage while the OGCM runs. For
336 example, we could target intermittent intense mixing events, plus their antecedents and fates.
337 An implication is that we should build a software interface for community-supplied software
338 plugins to implement the custom triggers. Also, we need to enable posterior re-computation
339 of small space-time chunks of the full solution, with customized diagnostics, and possibly at
340 higher resolutions.

341 It is instructive to compare Computational Oceanography with computational meteorology, which
342 is the analogous field in atmospheric sciences. Computational meteorology has somewhat different
343 science objectives. Numerical Weather Prediction (NWP) is an important task, for example. The
344 main advances in NWP attributable to growth in computer power are: (i) improved model resolution
345 (now also approaching global 1 km horizontal resolution; Fuhrer et al. 2018), and (ii) improved
346 forecast uncertainty quantification through larger ensembles of forecast runs. Computational
347 meteorology also concerns reanalysis products to hindcast the historical atmospheric state. The

348 reanalysis state estimation tools tolerate unphysical adjustments (increments), however, which give
349 more accurate fits to observations at lower computational cost. This practice is different to the
350 ocean state estimation tools discussed here, which firmly constrain the model solutions to satisfy
351 the model equations.

352 Nevertheless, there are several useful lessons from computational meteorology: First, NWP has
353 steadily improved since the 1980s (Bauer et al. 2015). The rate is an improvement in forecast
354 skill of about one day per decade (meaning a 2015 three-day forecast is about as skillful as a 2005
355 two-day forecast). The improvement derives mainly from better forecast initialization and better
356 atmospheric general circulation models (AGCMs; Magnusson and Källén 2013; see also Simmons
357 and Hollingsworth 2002). In this context, better AGCMs means models that have higher resolution,
358 have more accurate parameterizations and/or complexity, and have larger forecast ensembles that
359 better estimate forecast uncertainty. Computing advances have played an enormous role in these
360 improvements (Bauer et al. 2015). Second, as AGCM resolution increases, new phenomena begin
361 to be resolved. For example, with AGCM grid spacing of a few kilometers convective scales are
362 partly resolved (convective systems) but partly unresolved (convective cells). This partial resolution
363 of convection is called the “gray zone”, akin to eddy-permitting resolution in OGCMs. The best
364 approach to set up convection parameterization schemes in the AGCM gray zone is unclear and
365 forecast skill does not always improve at all lead times as resolution increases (Hong and Dudhia
366 2012). Moreover, at cloud-resolving resolution, data density is mismatched with AGCM resolution
367 (the number of degrees of freedom exceeds the number of observations) and the model solution is
368 not well constrained (Hong and Dudhia 2012).

8. Conclusion

Global OGCMs have a rich history that stretches back to the 1970s and regional OGCMs stretch back to the 1960s (models of the tides stretch back even further; see Cartwright 2012). OGCMs have been valuable to elucidate the ocean circulation since their inception. More broadly, numerical solution of rotating, stratified flow has roots in numerical weather prediction (NWP) from the early twentieth century (Abbe 1901; Bjerknes 1904; Richardson 1922 see also Lynch 2008 and Benjamin et al. 2019 for historical perspectives on NWP and climate models). Since 2000, global OGCMs have continued their exponential improvement in resolution. They are now becoming unconstrained by observations. Benchmark OGCM solutions have increasing value to a growing community and should be permanently archived and freely available. Clear limits, opportunities, and prospects for Computational Oceanography are in sight. For these reasons, our answer to the question posed in the title of this essay is yes: Computational Oceanography is entering a new era and is coming of age.

This field promises powerful new tools to address previously intractable problems. It does not aim to supplant observational oceanography. Indeed, observing the natural ocean must never cease. Instead, the greatest opportunity lies in merging these hitherto disparate branches of marine science. Lasting progress will require that we trust computational insights, verify them with real world observations, and understand them with fundamental theory.

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391 *Data availability statement.* Codes to make the figures are available at
392 github.com/hainegroup/Computational-Oceanography-Commentary. For Fig. 2,
393 the temperature data are from the National Centers for Environmental Information World
394 Ocean Database, the altimeter mission data are from www.altimetry.info, the IPCC data
395 are from the IPCC reports and pcmdi.llnl.gov/CMIP6, and the supercomputer data are from
396 en.wikipedia.org/wiki/List_of_fastest_computers. The data for AR6 are from the
397 HighResMIP project in July 2020, which was incomplete then. The ECMWF and NCAR machine
398 speed data are from www.top500.org.

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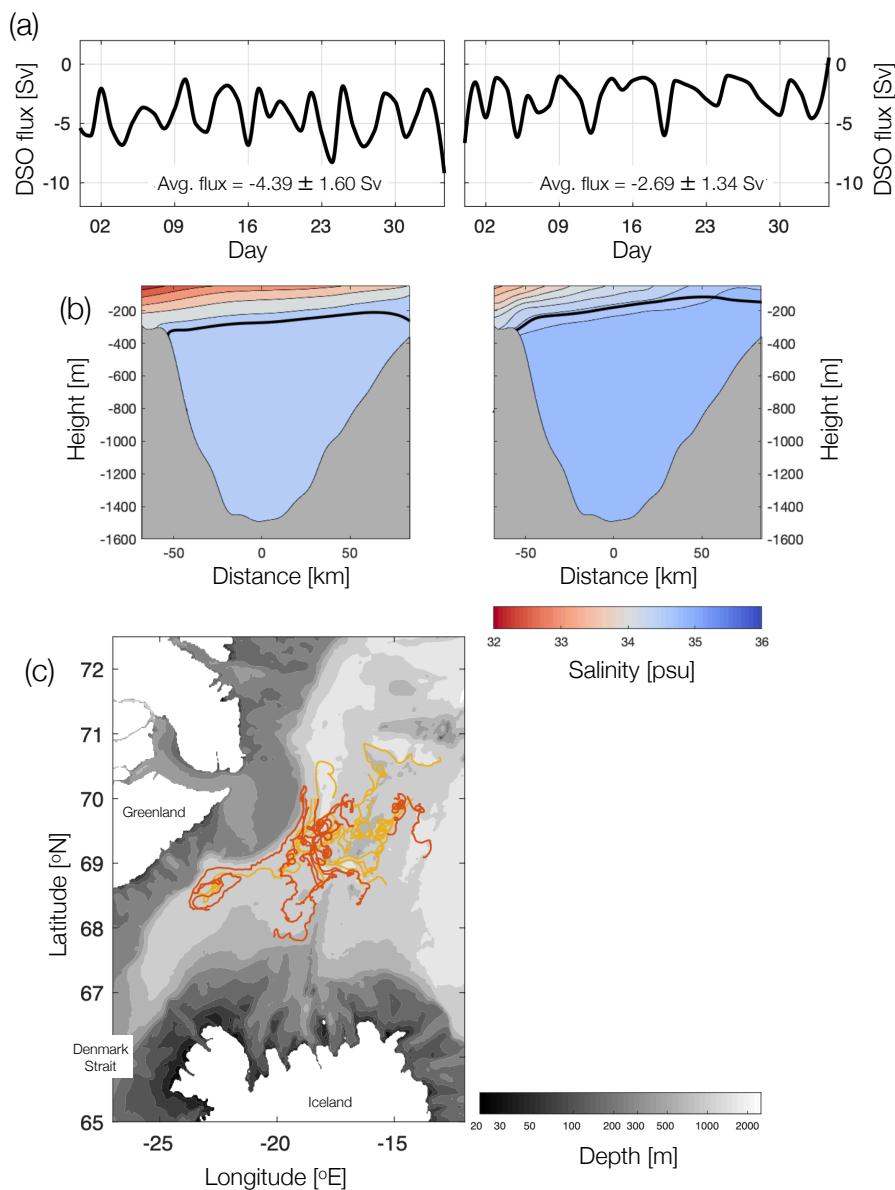
588 Wunsch, C., 2006: *Discrete Inverse and State Estimation Problems*. 1st ed., Cambridge University
589 Press, Cambridge, United Kingdom and New York, NY, USA, doi:10.1017/cbo9780511535949,
590 371 pp.

591 **LIST OF FIGURES**

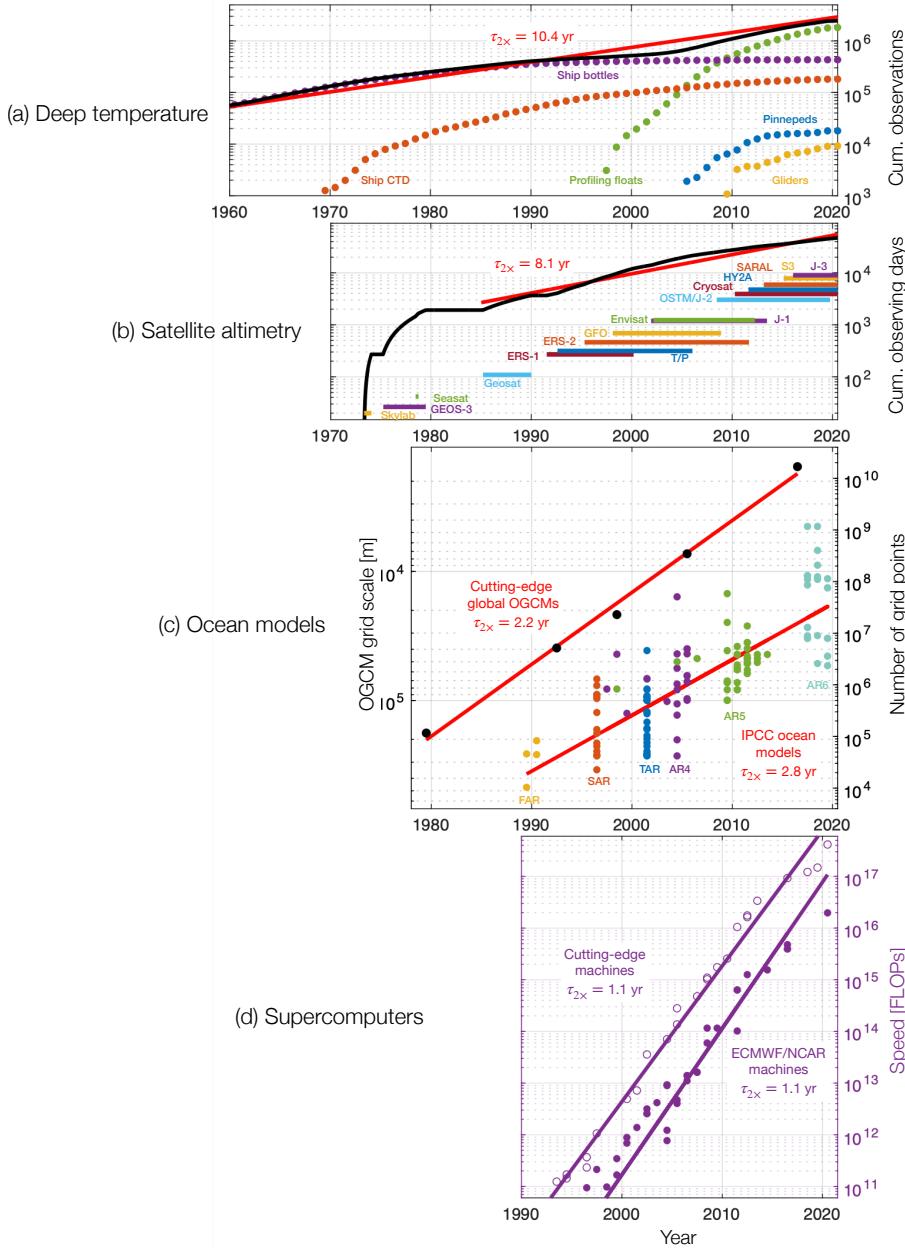
592 **Fig. 1.** OGCM Turing tests. In each of (a)–(c) oceanographic field measurements are compared
593 with OGCM results, but they are unlabeled (and processed similarly). The Turing test is
594 to identify which is which. (a) Denmark Strait Overflow (DSO) volume flux (S_v , $1S_v =$
595 $10^6\text{m}^3\text{s}^{-1}$, negative means equatorwards). Adapted from Haine (2010). (b) Salinity (colors)
596 on a section north of Denmark Strait (annual average; the heavy contour is the $27.80 \sigma_0$
597 density anomaly). (c) Lagrangian trajectories of RAFOS floats and synthetic RAFOS floats.
598 Adapted from Saberi et al. (2020). 31

599 **Fig. 2.** Unequal exponential growth. (a) History of deep (deeper than 1000 m) ocean temperature
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601 History of sea level measurements from satellite altimetry expressed by the cumulative
602 number of days of measurement. The satellite missions and their durations are indicated
603 with the colored bars. (c) History of cutting-edge global OGCM and IPCC ocean model
604 resolution expressed by the lengthscale of the horizontal grid and the number of model grid
605 points. Each dot represents one ocean model and the OGCMs are from Bryan and Lewis
606 (1979); Semtner and Chervin (1992); Maltrud et al. (1998); Maltrud and McClean (2005),
607 and Rocha et al. (2016). (d) History of top supercomputers using Rmax speed (FLOPS =
608 floating point operations per second) for fastest machines (open circles) and ECMWF and
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610 ($\tau_{2\times}$ is the doubling time). 32

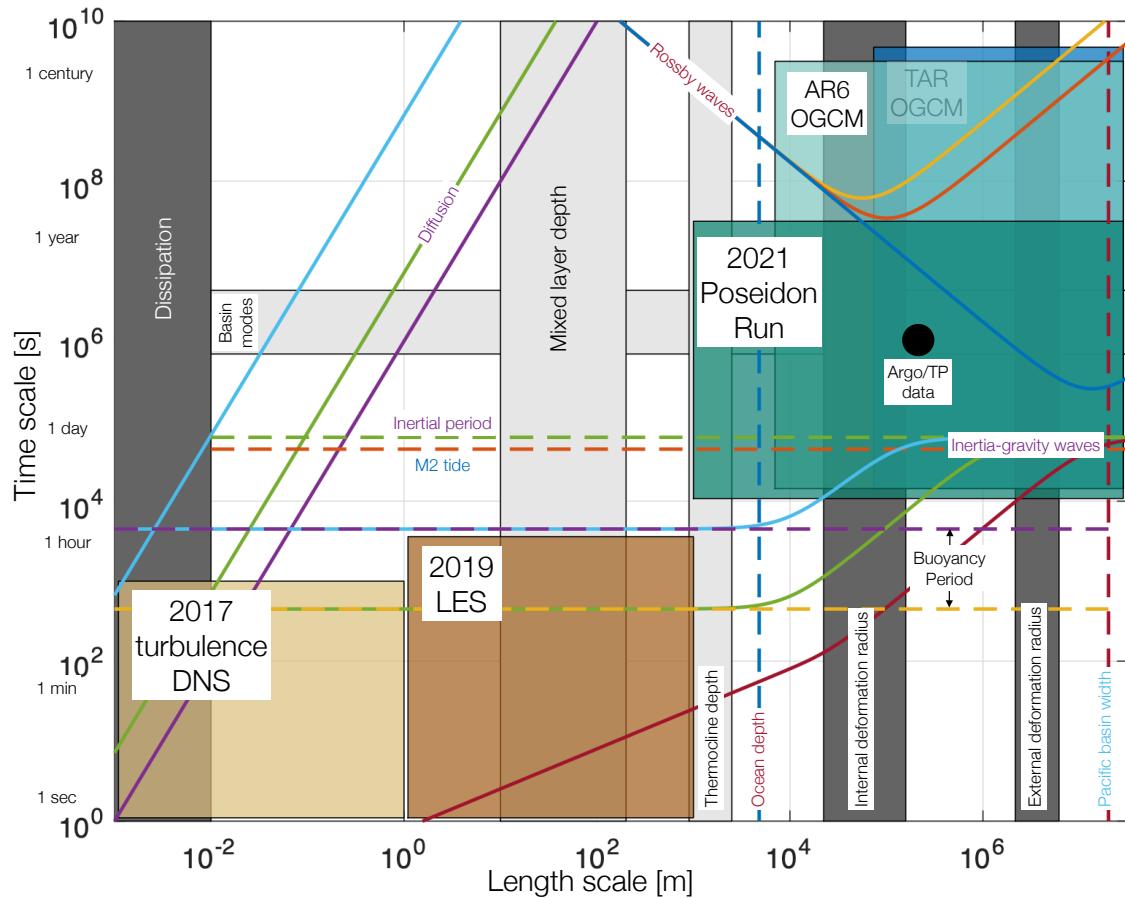
611 **Fig. 3.** Characteristic space and time scales of the ocean general circulation. Various geophysical
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614 (direct numerical simulation of turbulence, large eddy simulation, the Poseidon Project run,
615 AR6 HighResMIP, and TAR OGCMs). The black dot shows the sampling characteristics
616 of the Argo profiling floats, and the TOPEX/Poseidon-Jason altimeters. The diagram is
617 indicative, not definitive, because it suppresses the anisotropies and inhomogeneities present
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625 FIG. 2. Unequal exponential growth. (a) History of deep (deeper than 1000 m) ocean temperature mea-
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635 FIG. 3. Characteristic space and time scales of the ocean general circulation. Various geophysical and
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