

Impact of including the longwave scattering effect of clouds on the Arctic energy budget and climate in winter

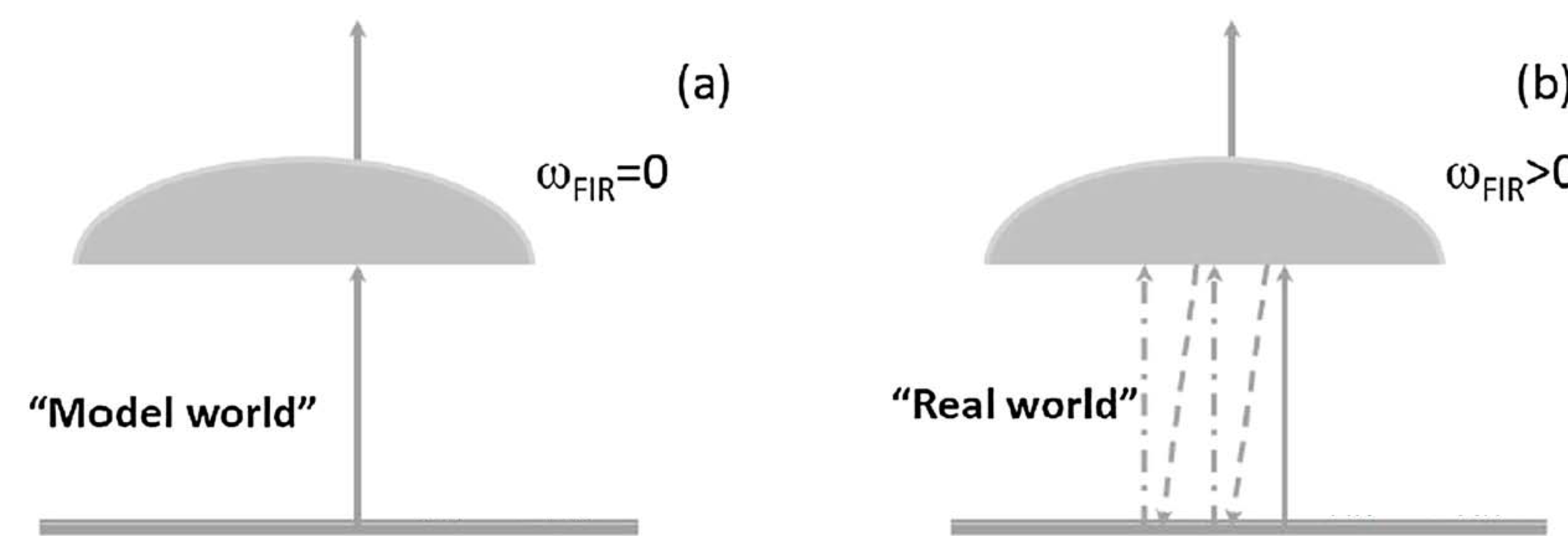
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Introduction and Objective

- Scattering of longwave radiation by clouds is regarded unimportant and neglected in global climate models.
- However, recent studies demonstrated the **nonnegligible role of cloud longwave scattering**, especially by ice clouds, in modulating the energy budget of the Earth System.
- It's not known **how this could affect the simulated Arctic climate**, considering the prevailing ice phase clouds and the dominance of longwave radiation during winter darkness.
- We implemented the longwave scattering of ice clouds into the DoE newest earth system model, E3SM and assessed the impacts on the simulated Arctic winter (DJF) energy budget and climate.**



Chen et al. (2014)

Method

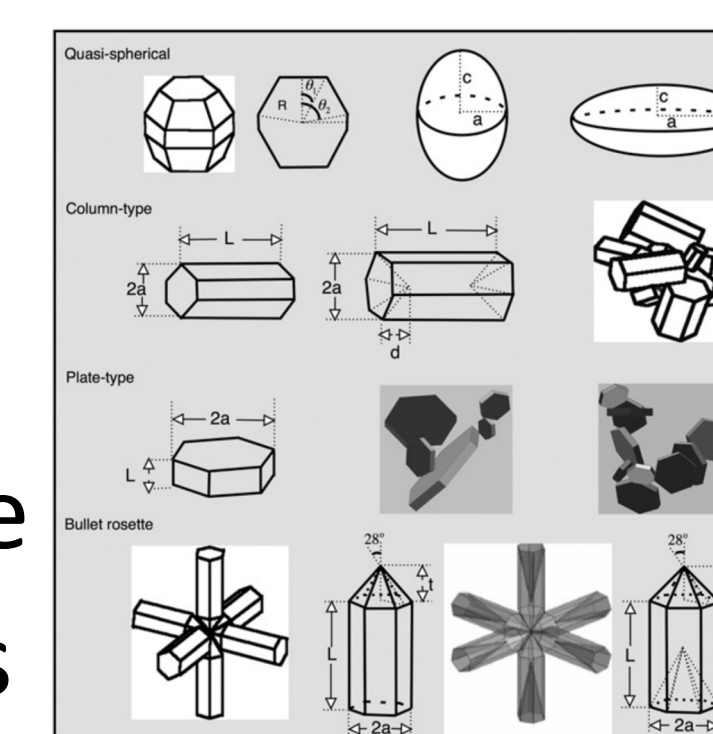
Model

- DoE Energy Exascale Earth System Model (E3SM v1; Golaz et al., 2019).



Modifications:

- Longwave Scattering Database for Ice Clouds:** derived from the MODIS collection 6 (MC6) ice cloud habit model (Yang et al., 2013) by varying ice cloud habits with treatments of severe roughness by aggregation.
- Longwave Scattering Radiative Transfer:** The hybrid two- and four-stream radiative transfer solver (Kuo et al., 2020) was implemented to incorporate scattering calculation.



Yang et al. (2013)

Experiments:

- 15 years of fully-coupled simulations with historical forcings (2000-2014) with (**Scat**) and without (**noScat**) longwave scattering. Output for Arctic (north of 60°N) winter are analyzed here.

Impact on Arctic energy budget

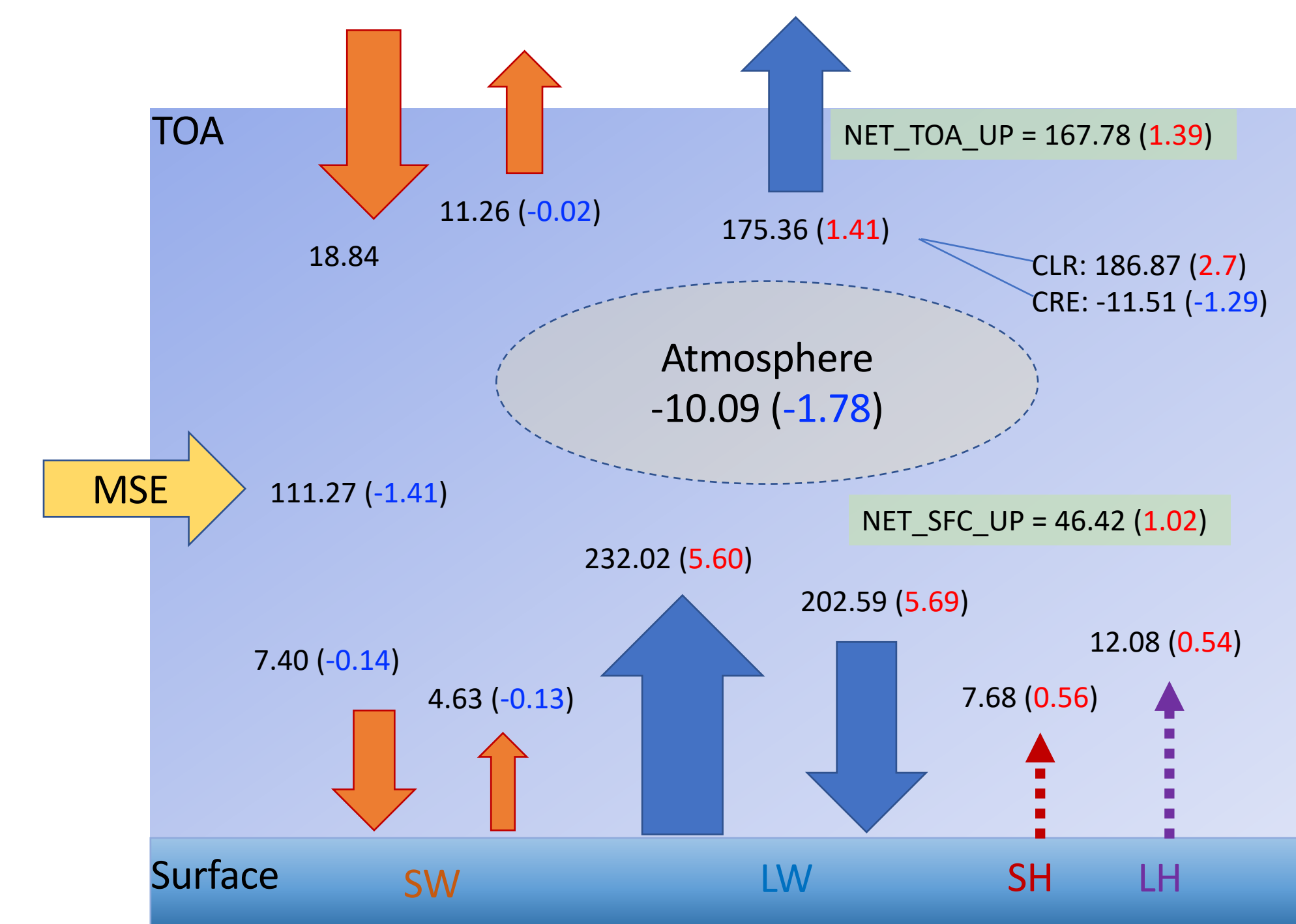


Fig. 1 Arctic domain-average energy budgets as simulated by the noScat version of E3SM (black numbers), as well as the differences between the Scat and noScat simulations (numbers in parentheses). MSE, moist static energy; CLR, clear sky; CRE, cloud radiative effect.

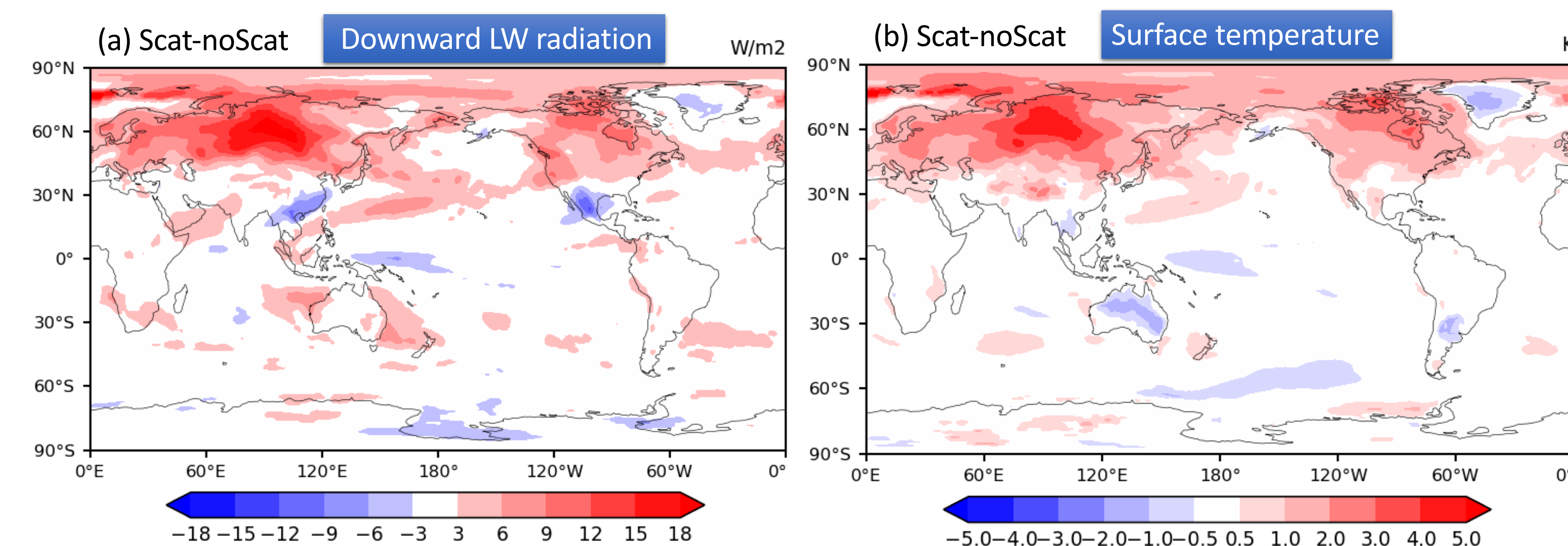


Fig. 2 Differences in the simulated (a) downward longwave radiation flux at surface and (b) the surface temperature.

Lower tropospheric stability and clouds

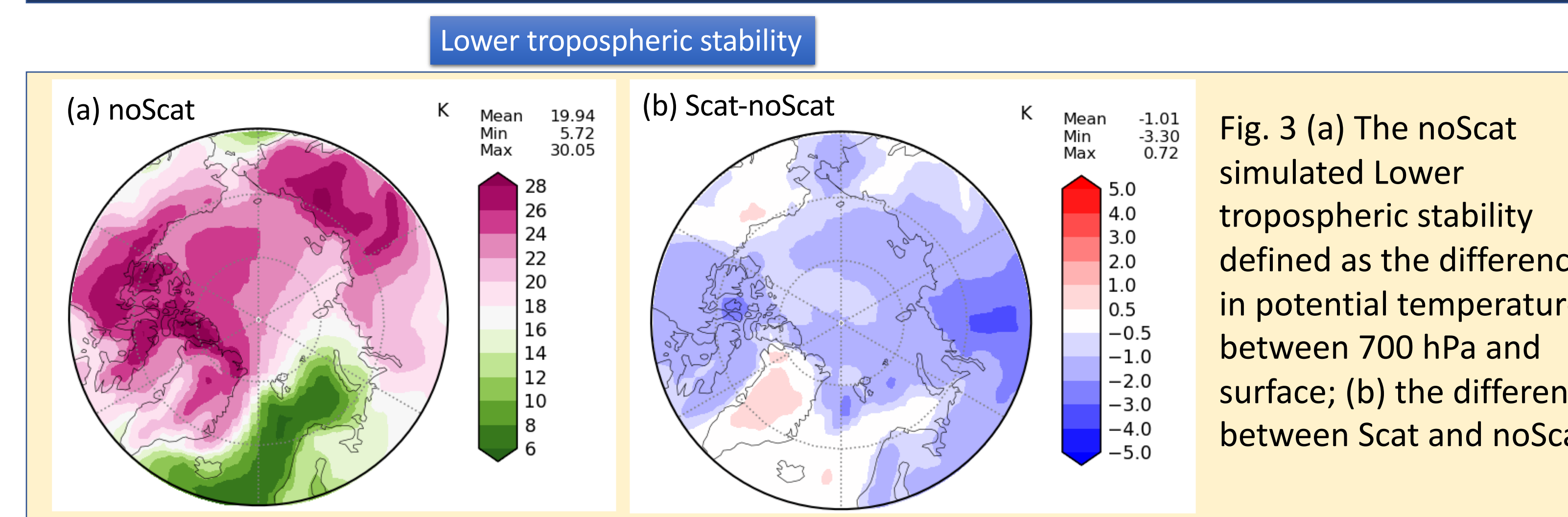


Fig. 3 (a) The noScat simulated Lower tropospheric stability defined as the difference in potential temperature between 700 hPa and surface; (b) the difference between Scat and noScat.

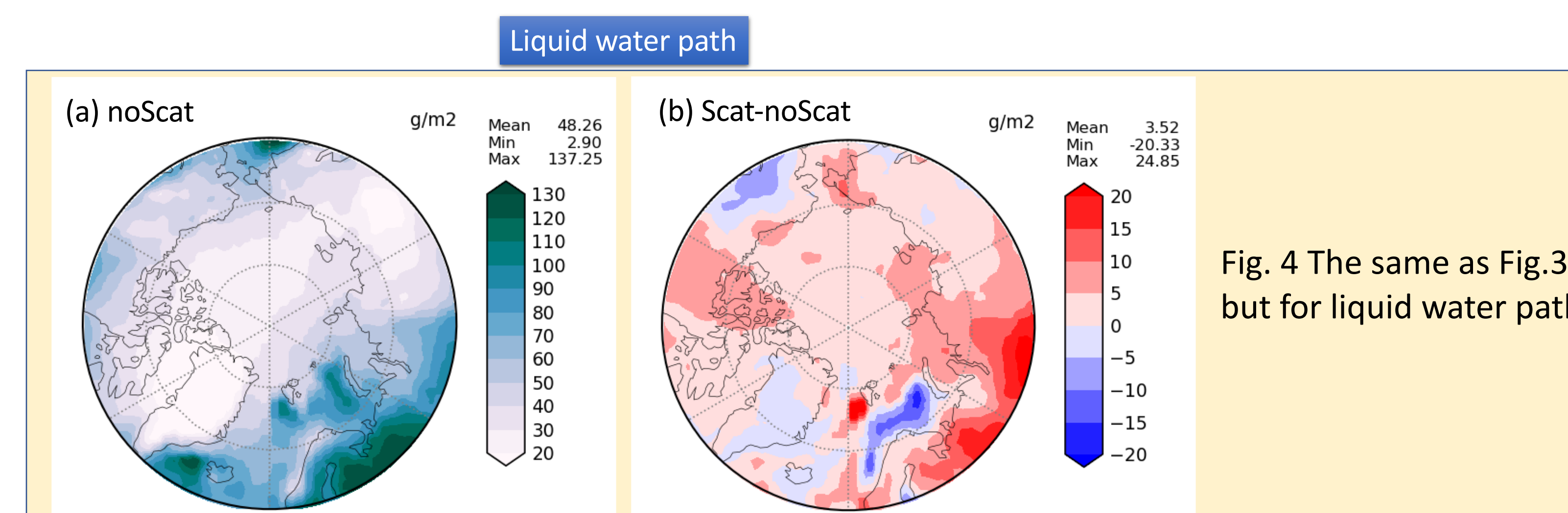


Fig. 4 The same as Fig.3 but for liquid water path.

- Increased surface temperature significantly **reduced lower tropospheric stability**, especially over high-latitude continent and Arctic ocean (Fig. 3).
- These regions features high stability and resilience of boundary layer mixed-phase clouds, thus the reduction in stability lead to more turbulence and uplift of water vapor, **favoring cloud growth (Fig.4b) as well as the potential phase shift.**

Circulation and water vapor transport

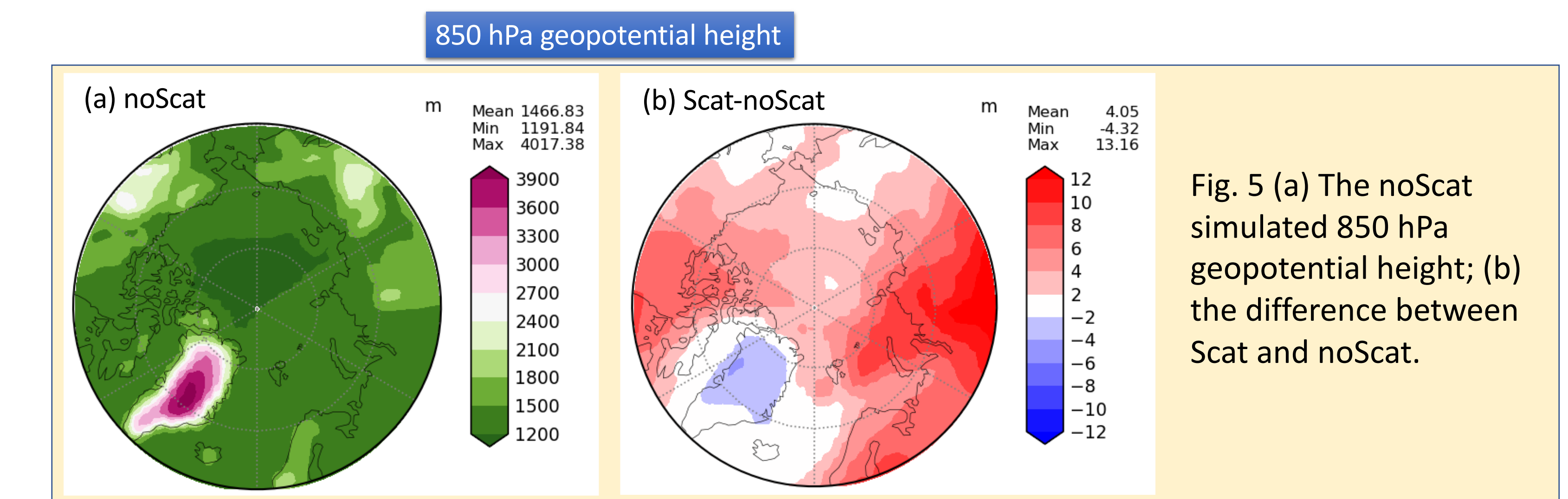


Fig. 5 (a) The noScat simulated 850 hPa geopotential height; (b) the difference between Scat and noScat.

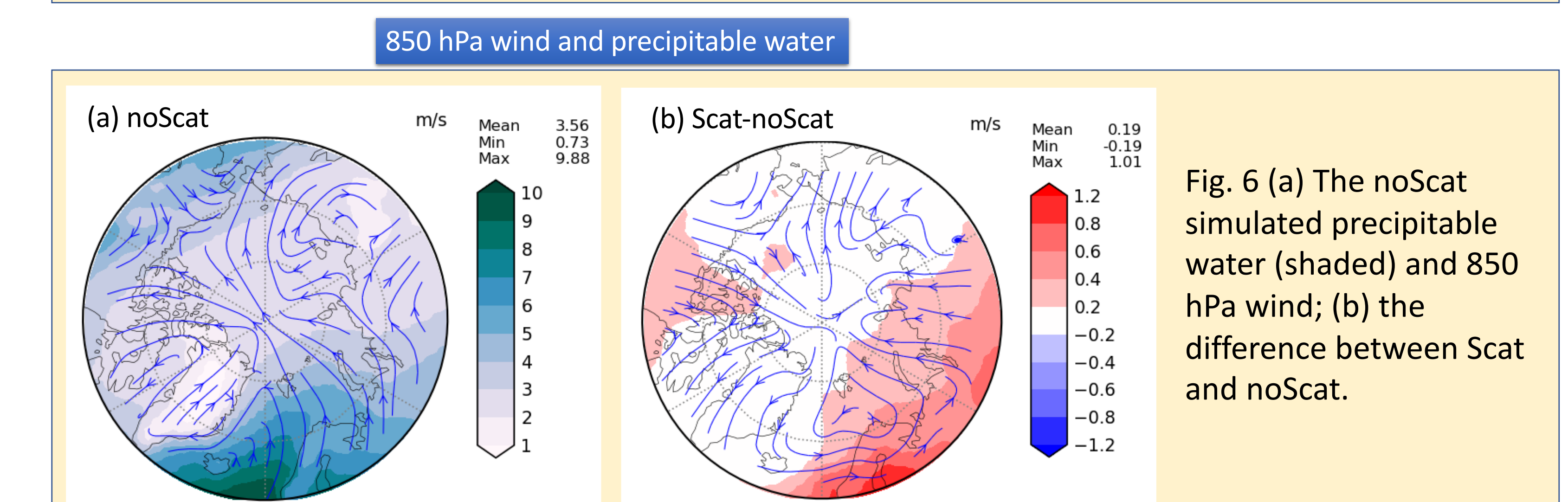


Fig. 6 (a) The noScat simulated precipitable water (shaded) and 850 hPa wind; (b) the difference between Scat and noScat.

- The **850 hPa geopotential height increased** over most of the Arctic region owing to the expansion of lower troposphere resulted from warmer surface.
- The larger increase of geopotential height over continental areas than over the Arctic oceans led to **poleward trend of wind** (Fig. 6b), conveying the increased water vapor at lower latitudes into the Arctic Circle.
- Enhanced water vapor transport** further increased FLDS to warm the Arctic surface, manifesting a positive feedback.

Conclusions

- Longwave scattering of clouds increases downward LW flux, thus **increasing wintertime polar surface temperature.**
- Reduce lower troposphere stability** by surface heating, facilitating turbulence and cloud development.
- Enhance water vapor transport** to Arctic from continental region.
- Local thermodynamics and large scale circulations interplay with each other** to affect Arctic energy budget.

References:

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