

16 **Abstract**

17 In situ exploration of Uranus has been limited to a single flyby encounter by the Voyager 2
18 spacecraft in January 1986. Nonetheless, new investigation has led to significant questions about
19 the origin of energetic ions observed in the region between its moons Miranda and Ariel. Radial
20 and pitch angle diffusion modeling suggests that typical magnetospheric sources cannot explain
21 the observed characteristics of these energetic ions. We suggest that these are likely being
22 introduced by a source from one of these moons and give rise to waves that could result in the
23 observed particle distribution characteristics. This may reveal that internal plasma sources in the
24 system may be important for Uranus' magnetospheric dynamics and may contribute to its
25 unexpectedly strong radiation belts.

26 **Plain Language Summary**

27 Uranus is an oddity in the solar system for a variety of reasons, but mostly as a result of its
28 perpendicular rotation relative to the rest of the planets in the solar system. During its
29 approximately three-day flyby of Uranus in 1986, Voyager 2 captured the only in-situ
30 observations of the planet and its system. New analysis of these three-decade-old observations
31 have revealed a mysterious source of energetic ions in the planet's magnetosphere. These ions
32 were originally explained by dynamics of the system, but new understanding suggests that this is
33 probably unlikely. New simple modeling of the expected behavior of such energetic particles
34 show that sustaining such a population requires a very strong source and specific energization
35 mechanism. Both would potentially be consistent with the ions originating from either Miranda
36 or Ariel. This potentially hints that the Uranian magnetosphere may harbor an ocean world like
37 those known or believed to exist at the other Giant Planets.

38 **1 Introduction**

39 The single glimpse of the Uranus system provided by Voyager 2 revealed surprises and
40 provided observations that have led to new questions (e.g., Paty et al., 2020; Kollmann et al.
41 2020). Perhaps most significantly, the Voyager 2 flyby revealed that Uranus has an offset and
42 highly-tilted intrinsic magnetic field, with a tilt of $\sim 60^\circ$ from the rotational axis (Fig. 1) (Paty et
43 al., 2020). Voyager 2 also observed the energetic particle populations at Uranus to be very
44 similar to Earth (Mauk & Fox, 2010), despite the fact that Voyager 2 found much lower densities
45 of plasma in the system than Earth and the Gas Giants (McNutt et al., 1987). Thus, it is unclear
46 from which source plasma the radiation belts could be populated. However, transformational
47 advances in understanding of magnetospheric processes—afforded by observations from
48 missions at several planets in the ensuing decades since the initial analyses of the Voyager 2
49 observations—provide a new lens through which to reanalyze observations from the Uranian
50 magnetosphere.

51 The mystery surrounding Voyager 2's discovery of unexpectedly strong radiation belts at
52 Uranus motivated a revisiting of the LECP dataset. This fresh survey revealed a previously
53 underappreciated signature in the energetic particle observations from the Low Energy Charged
54 Particle (LECP) instrument (Krimigis et al., 1977). Specifically, LECP observed a significant
55 (several orders of magnitude) discrepancy between the intensity of energetic particles (both ions
56 and electrons) observed during the outbound leg of the Uranus flyby encounter in the region
57 between Miranda and Ariel compared to the inbound leg (Fig. S1). The Plasma Experiment

58 (PLS) (Bridge et al., 1977) also reported signatures suggestive of anomalous spacecraft charging
59 in the same region (McNutt et al., 1987).

60 Such an “asymmetry” in particle fluxes between the inbound and outbound legs of the
61 trajectory may have been expected if the spacecraft encountered these radial distances at very
62 different magnetic latitudes, but that was not the case – the actual difference in magnetic latitude
63 was $<20^\circ$. While this “asymmetry” was noted in initial analysis of the LECP observations (Mauk
64 et al., 1987), its extreme nature was explained as an artifact of the flyby trajectory because at the
65 time no other solution (e.g., very intense substorm-like injection dynamics) seemed plausible.
66 Fig. S2 shows average PADs observed at three different >300 keV ion energy channels created
67 by using the offset tilted dipole magnetic field model from Ness et al. (1986) to extrapolate the
68 local pitch angle pointing of LECP inbound (24 Jan 1986 17:15-17:30 UT) and outbound (24 Jan
69 1986 19:45-20:00 UT) to the same magnetic latitude (i.e., $\sim 20^\circ$ from the outbound leg). Based on
70 this mapping in pitch angle space (using conservation of the first invariant) it was actually
71 determined that because of the differences in magnetic latitude between the inbound and
72 outbound legs of the flyby, the near- 90° pitch angle particles seen on the outbound leg could not
73 be seen by Voyager 2 at the higher latitudes on the inbound leg (i.e., they were mirroring at
74 lower latitudes than Voyager 2). It should be emphasized that ion measurements at energies
75 $\lesssim 300$ keV are significantly contaminated by penetrating energetic particles (Mauk et al., 1987).

76 Transient intensity enhancements in both ions (Roussos et al., 2008) and electrons
77 (Roussos et al., 2018) due to events in the solar wind have been observed at other giant planets,
78 but between different orbits. There was also no evidence of drift dispersion and/or extended
79 signatures of such an injection at higher L, both of which might be expected if an injection had
80 occurred sometime during the periapsis pass, as the lack of injection signature on the inbound leg
81 of the trajectory might suggest. Previous work (Cheng et al., 1987) has looked at the LECP
82 energetic particle observations at Uranus and calculated phase space density (PSD) profiles.
83 However, the values of the second adiabatic invariant (K), which they confined their analysis to,
84 restricted the portions of the flyby encounter that they included in their analysis—i.e., $L < 10$ for
85 the inbound leg and $L > 9$ for the outbound leg. The authors concluded that there had to be a
86 source of energetic particles in the inner magnetosphere: specifically, they—like others (Mauk et
87 al., 1987)—suggested substorm-like injections. Further study (Paranicas et al., 1996) focused on
88 the PSD profiles measured by LECP at Uranus, but combined the inbound and outbound legs of
89 the trajectory together, thus averaging out the asymmetry of interest here.

90 **2 Exploring and Confirming the Source**

91 Of particular interest are the pitch angle distributions (PADs)—i.e., the angle of the
92 charged particle velocity vector relative to the background planetary magnetic field vector—
93 measured by LECP, which display extremely steep gradients (Fig. S2). It should be noted that
94 LECP was severely limited in its pitch angle coverage in the region between Miranda and Ariel
95 (see Figure 4 of Mauk et al. (1987)). The nature of these pitch angle (α) distributions is curious.
96 The $\sin^n(\alpha)$ fits (dashed lines) in Fig. 2a show that the n values of the fits get extremely large
97 with n – a parameter representing the slope in pitch angle space – increased toward higher
98 energies. Such a steep gradient is not due to an enhanced bounce loss cone size, since
99 observations are at a relatively large L-shell ($L \sim 7$). Maintaining such a steep pitch angle gradient,
100 which would be even steeper in terms of equatorial pitch angle, is difficult since any waves—
101 ranging in frequency from ultra-low frequency (ULF) to electromagnetic ion cyclotron

102 (EMIC)—would act to scatter the particles and isotropize the distribution. In the presence of
103 such waves, this would require a significant and relatively constant source of energetic particles,
104 specifically for those at near-90° pitch angles, at rates that can balance or even overcome any
105 loss/scattering processes from waves. Incidentally, this region between the orbits of Miranda and
106 Ariel was also exactly where Voyager 2 observed intense whistler-mode wave emissions (Kurth
107 & Gurnett, 1991). Maintaining such a steep PAD would require a significant and relatively
108 constant source of energetic particles, specifically for those at near-90° pitch angles, at rates that
109 can balance or even overcome any loss/scattering processes from waves.

110 To illustrate this point, Figs. 2b-d show results from one-dimensional pitch angle
111 diffusion model simulating the expected evolution of the 1.45-MeV ion (assuming protons) pitch
112 angle distribution observed by LECP at L=6. The model assumes that the inbound and outbound
113 LECP observations place a 6-hour constraint on the evolution of the PAD - i.e., the distribution
114 does not change appreciably over 6 hours. This assumption is supported by additional LECP
115 observations of ions and electrons across the instrument's energy range and the lack of any
116 signatures suggesting a sudden injection. As the initial simulation in Fig. 2b shows, the steep
117 pitch angle gradients observed by Voyager 2 (dashed line) are quickly reduced by pitch angle
118 diffusion over six hours—i.e., approximately the time between the inbound and outbound LECP
119 measurements in the region between Miranda and Ariel. Fig. 2c adds complexity and realism by
120 considering the same pitch angle diffusion scenario but adding rapid (~10-min) losses to the
121 moons at all local pitch angles less than 80° (i.e., those protons that would intersect the moons
122 during a single drift period considering bounce motion along field lines; relative to the
123 Voyager 2 outbound trajectory). As can be seen, this additional loss alongside the ongoing pitch
124 angle diffusion quickly reduces the proton intensity at all α over the 6-hr simulation; while this
125 simulation results in more trapped distributions (i.e., strong peaks centered around $\alpha = 90^\circ$), the
126 proton intensity is reduced by approximately two orders of magnitude compared to what was
127 observed by LECP. While it cannot be ruled out that the inner magnetosphere was
128 serendipitously enhanced during the singular Voyager 2 flyby - i.e., the observed intensities are
129 the final condition of a very slow loss environment instead of the initial condition of a very fast
130 loss environment - it seems incredibly unlikely. Finally, Fig. 2d takes a final step to match
131 diffusion simulations to the observed distributions by adding in a localized, near-90° (i.e., near-
132 equatorial) proton intensity source, which results in a relatively time-stable (i.e., over >6-hr)
133 solution for the 1.45-MeV proton pitch angle distribution that is very similar in magnitude and
134 shape to that observed by LECP. The model uses a pitch angle diffusion coefficient ($D_{\alpha\alpha}$)
135 distribution as a function of α (Fig. S3) and an arbitrarily long (36 hr used here) e -folding loss
136 time constant (τ)—approximating slower localized losses due to radial transport and neutral
137 interactions—based on 1-MeV protons in electromagnetic ion cyclotron (EMIC) waves at L = 6
138 at Earth (e.g., Glauert & Horne, 2005). While it should be noted that this assumption is several
139 orders of magnitude larger than the electron $D_{\alpha\alpha}$ calculated from Voyager 2 observations by
140 Selesnick & Stone (1991), it is still believed to be a reasonable assumption given that previous
141 studies have shown that a distinct temperature anisotropy can lead to enhanced wave growth and
142 that growth mechanisms for EMIC waves are the same for protons as right-handed whistler-
143 mode growth is for electrons (e.g., Usanova et al., 2008, 2012, 2013, and references therein); as
144 previously mentioned, strong whistler-mode wave activity was observed in this region by
145 Voyager 2 (Kurth & Gurnett, 1991); as such, EMIC waves may have been present in this region

146 as well, but would have been outside the frequency range covered by the Voyager 2
147 instrumentation.

148 To assess whether such an energetic particle source is in fact present in the Uranian
149 system, measurements of the phase space density (PSD) profiles of the ions versus L-shell as a
150 function of the first adiabatic invariant (μ) were investigated (Fig. 3) (Green & Kivelson, 2004)
151 using the same LECP response function used by Mauk et al. (1987). These PSD distributions
152 were calculated for local 90° pitch angles, which correspond to approximately $K = 0$ given the
153 low magnetic latitude. The Voyager 2 trajectory information was processed and converted into
154 several relevant coordinate systems—including the “U1” west-longitude (Ness et al., 1986) and
155 dipole (Mauk et al., 1987) coordinate systems—following the same approach applied in another
156 recent study (DiBraccio & Gershman, 2019). The peak in PSD between 18:00 and 19:00 UT
157 (i.e., planetward of Miranda) cannot be ruled out as a signature resulting only from changes in
158 magnetic latitude and/or moon losses. However, the maximum between Miranda and Ariel at
159 $L \sim 7$ clearly suggests a source of energetic ions in this region. This peak between Miranda and
160 Ariel is much higher than the PSDs at near the same magnetic latitude (within approximately
161 $\pm 5^\circ$) as the other side of those moons and the latitudinal effects expected from μ - and K -
162 conservation are mostly negligible for the $< 10^\circ$ latitudinal change at lower latitudes observed
163 here. The clear maximum between Miranda and Ariel at $L \sim 7$ clearly suggests a source of
164 energetic ions in this region. Similar PSD peaks for protons at Saturn are a common example to
165 illustrate the effect of a local source (e.g., Kollmann et al., 2011, 2017). Other peaks in energetic
166 ion PSD were found just outward of Io at Jupiter (Anglin et al. 1997, Woch et al. 2004), which
167 may serve as an ideal analogy of a geologically-active moon releasing unstable plasma
168 distributions into its environment.

169 **3 Discussion of Potential Sources**

170 It is challenging to definitively determine the source of these energetic particles given the
171 limited—in range, duration, and system coverage—in-situ measurements of the Uranian
172 magnetosphere. In particular, composition measurements of ion species—both mass and charge
173 state—in the suprathermal (10s to 100s keV) energy range are lacking, which could help identify
174 particle sources and acceleration processes. Likewise, field measurements at the relevant
175 frequencies required to assess ion-mode wave processes are missing. However, even the limited
176 measurements obtained at Uranus can be combined with understanding informed by more
177 comprehensive observations in other planetary magnetospheres to assess the feasibility of several
178 sources based on whether they could explain different aspects of the LECP observations—
179 specifically 1) the > 300 keV energies, 2) the substantial particle intensities, and 3) the strongly
180 peaked near 90° PADs. Potential energetic particle sources that will be considered here for
181 Uranus are magnetospheric particle injections, cosmic ray albedo neutron decay (CRAND), and
182 sourcing from a moon (i.e., a potentially active moon or sputtering).

183 “Injections” are sudden enhancements of energetic particles caused by magnetospheric
184 dynamics (e.g., Arnoldy & Chan, 1969; Thomsen, 2013). However, several pieces of
185 observational evidence (or lack thereof) argue against this potential source. First, as previously
186 mentioned, the LECP electron observations display neither any injection-like signature, nor do
187 the injections result in a similarly steep pitch angle gradient. The energy dependence of both the
188 inbound-outbound flux asymmetry and PAD steepness provide further evidence against an

189 injection source. Additionally, the LECP ion observations show no evidence of drift echoes as
190 might be expected (e.g., Blake et al., 1992) nor any similar injection signatures extended over
191 other L-shells. At Earth, dispersionless injections are most commonly isotropic in pitch angle
192 and extended over a broad range of L-shells (Turner et al., 2017). However, ion distributions at
193 Saturn also exhibit extremely steep PADs, which at low L-shells result from an inflated loss cone
194 and at high L-shells were suggested to be maintained by the CRAND source and the absence of
195 intense EMIC waves (Kollmann et al., 2022). Another potential argument against an injection
196 source is that moon macrosignatures—i.e., longitude averaged depletions of energetic ions at the
197 (minimum) moon L-shells—are observed during both the inbound and outbound legs of the
198 trajectory. Previous work explain that such macrosignatures develop over timescales of moon-
199 particle reencounters, which amount to many drift periods, or days. For example, Paonessa &
200 Cheng (1987) give that the timescale of moon losses at Miranda and Ariel are ~280 hours, or 11-
201 12 days. After such a time scale, any dispersion effect or drift echoes would have disappeared.
202 Though no evidence in the Voyager 2 data nor literature suggests it, there is a possibility that this
203 enhancement may have been a transient radiation belt triggered by a very fortuitously-timed
204 interplanetary coronal mass ejection, as observed at Saturn (e.g., Roussos et al., 2008).

205 Fig. 4a demonstrates the shape of the PSD profile versus L-shell that would be expected
206 from only inward transport of plasma from the outer magnetosphere ($L > 11$)—i.e., without any
207 local source in the vicinity of Miranda or Ariel. Note that the modeled PSD values are arbitrary
208 as the model is using simplified, but realistic, physical constraints to simulate the radial shape of
209 the PSD profiles. As can be seen this overall positive slope is very different from the observed
210 PSD shown in Fig. 3, suggesting that some additional source must be playing a role. Fig. 4b
211 shows the expected PSD profile adding a CRAND-like source falling off with distance from the
212 planet (dark blue). CRAND is a well-established and significant source of energetic protons in
213 the radiation belts of Earth and Saturn (e.g., Blake et al., 1992), primarily coming from ring
214 material and planetary atmospheres. This simulation assumes the planet to be the dominant
215 CRAND source in the system and falls off as L^{-2} (e.g., Kollmann et al., 2013); while a
216 contribution from moon CRAND cannot be ruled out, it is considered unlikely. Finally, Fig. 4c
217 shows that the addition of a local source in the region between Miranda and Ariel to the
218 simplified CRAND source results in a PSD profile much closer to that observed by Voyager 2,
219 specifically with the maximum in the region between Miranda and Ariel significantly higher than
220 that inside of Miranda.

221 The final potential source to be considered is a supply of ions from a moon. Since the
222 1986 Voyager 2 flyby of Uranus, several moons throughout the solar system have been revealed
223 to be geologically active, often cryo-volcanic, ocean worlds (e.g., Nimmo & Pappalardo, 2016).
224 While debate remains in the Jovian system as to whether Europa may contribute neutrals/ions to
225 the magnetosphere solely from sputtering without the need for active plumes (e.g., Blöcker et al.,
226 2016; Huybrighs et al., 2020; Smyth & Marconi, 2006; Szalay et al., 2022), prior work in the
227 Saturn system found it difficult to make sense of the observed plasma distributions before
228 Enceladus was discovered to be an ocean world (e.g., Burger et al. 2007; Jurac et al., 2002;
229 Jurac & Richardson, 2005; Shi et al., 1995). While sputtering from the surface cannot be ruled
230 out, the conclusions made here are consistent with either or both of Miranda and/or Ariel being

231 an ocean world, as has been recently suggested (e.g., Hendrix et al., 2019; Ćuk et al., 2020;
232 Schenk & Moore, 2020; Cartwright et al., 2021; Beddingfield et al., 2022a,b).

233 It must be emphasized that the narrow pitch angle source required to obtain the results in
234 Fig. 2d matches that expected from newly-created pickup ions (Coates & Jones, 2009).
235 Specifically, for the Uranian system, this effect is enhanced by the fact that the electric field
236 magnitude, responsible for giving the initial acceleration to the pickup ions, is strongest when the
237 source is located at the magnetic equator (Cochrane et al., 2021). These effects combined
238 contribute to the highly localized and narrow pitch angle nature of the expected source. As
239 previously mentioned, the introduction of a preferentially trapped population of thermal and/or
240 suprathermal plasma from either sputtering from the surface or an active moon and this pickup
241 ionization process could be expected to generate a significant temperature anisotropy that would
242 drive wave growth that is expected to result in high intensities of both the observed whistler-
243 mode waves (e.g., Kurth & Gurnett, 1991) as well as EMIC waves (unobservable by Voyager 2)
244 that could potentially accelerate the newly-sourced thermal or suprathermal population up to the
245 >1 MeV energies observed by Voyager 2 via strong drift- and/or gyro-resonance, which would
246 isotropize the lower-energy source population while contributing to the very trapped energetic
247 particle population. For example, with gyro-resonant acceleration in quasi-linear diffusion
248 theory, accelerated particles are preferentially accelerated at higher energies and diffuse towards
249 90° equatorial PA (e.g., Horne et al., 2005, and references therein) by waves generated by strong
250 anisotropies in the lower-energy, suprathermal particles of the same species.

251 In either case, the location of the maximum in the PSD profiles at L~6 (Fig. 4) makes it
252 unclear which of these two moons may be the most likely potential source, since such PSD
253 values usually fall off closer to the planet (evidence for Ariel) but moon material also tends to
254 move outwards (evidence for Miranda). An expected source from either moon could be
255 estimated using the source required to sustain the pitch angle distributions shown in the simple
256 diffusion model shown in Fig. 2d (see Supplementary Material). Other Voyager 2 observations
257 provide potential further evidence to support the case for a moon source. Unfortunately, at
258 Uranus, the LECP instrument could not measure ion composition below ~500 keV/nuc (Mauk et
259 al., 1987)—e.g., 2 MeV for helium, 8 MeV for oxygen; thus, it is unsurprising that LECP did not
260 see any non-proton populations that may have been sourced by an active Uranian moon as it was
261 unlikely that these species were easily energized to such extreme energies; however, evidence
262 exists that such heavy ions can be accelerated to suprathermal energies by EMIC waves (Wang
263 et al., 2019). Additional measurements from PLS suggest that there are heavier ions in the region
264 between Miranda and Ariel (McNutt et al., 1987). Similarly, the aforementioned PLS signatures
265 of anomalous spacecraft charging could possibly be a result of flying through a dense torus or
266 cloud of ionized particles, including salts (Behannon et al., 1977).

267 **4 Conclusions**

268 New analysis of the LECP observations from the Voyager 2 flyby of the Uranus system
269 have revealed a surprisingly intense source of energetic ions in the region between the moons
270 Miranda and Ariel. Consideration of the unique pitch angle distributions observed in this
271 population has resulted in the conclusion that the observed energetic ions are most likely
272 originating from cold neutral particles coming from one of the nearby moons – i.e., Miranda or
273 Ariel. The leading hypothesis is that one or both of the moons source low-energy neutrals into

274 the magnetosphere via sputtering or an active subsurface ocean. This population is then ionized
275 and energized by wave-particle interactions driven by anisotropies driven in the origination of
276 the plasma itself or by the influence of secondary processes like pickup ionization. Ultimately,
277 the accelerating wave-particle interaction processes sculpt the observed preferentially-trapped
278 energetic ion population. However, further investigation of this mysterious energetic ion PAD
279 will require additional observations from the Uranian system, preferably from an orbiter mission
280 that is equipped with instruments to measure the thermal plasma properties and composition,
281 suprathermal (tens to hundreds of keV) ion composition, with both mass and charge-state, and
282 wave activity extending into the ion-cyclotron modes (i.e., EMIC waves).

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289 **Open Research**

290 The data utilized in this study was obtained from the NASA Planetary Data System (PDS)
291 Planetary Plasma Interactions Node. The respective files can be found at:
292 <https://doi.org/10.17189/1520024> (LECP energy distribution), <https://doi.org/10.17189/1520025>
293 (LECP pitch angle information), <https://doi.org/10.17189/1520034> (MAG), and
294 <https://doi.org/10.17189/1520040> (spacecraft position). The processed data, related code, and
295 simulations used in the analysis are available at <https://doi.org/10.5281/zenodo.7308665>.

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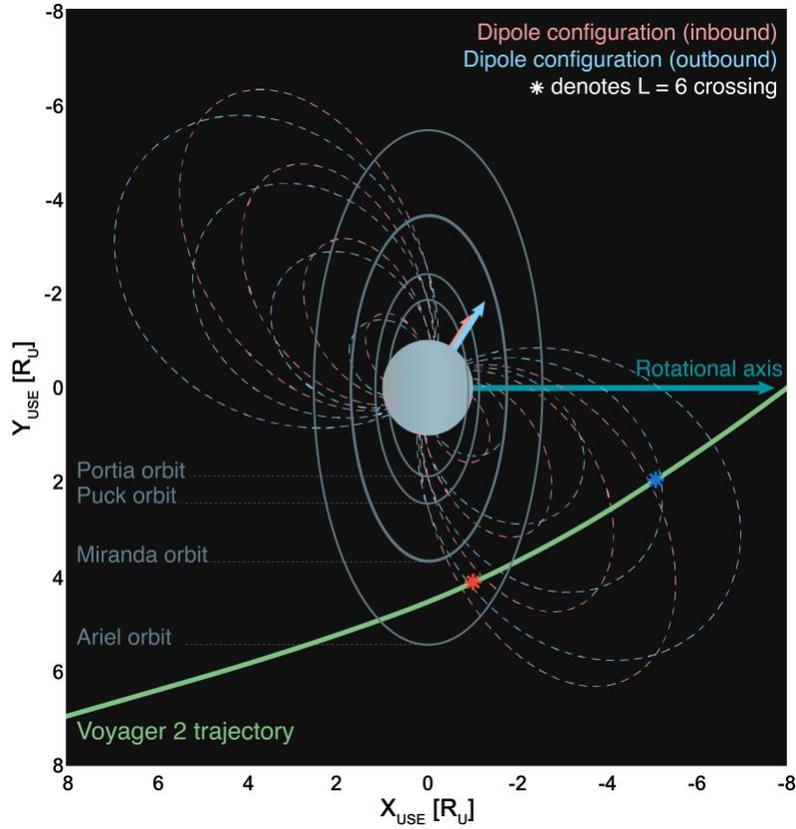
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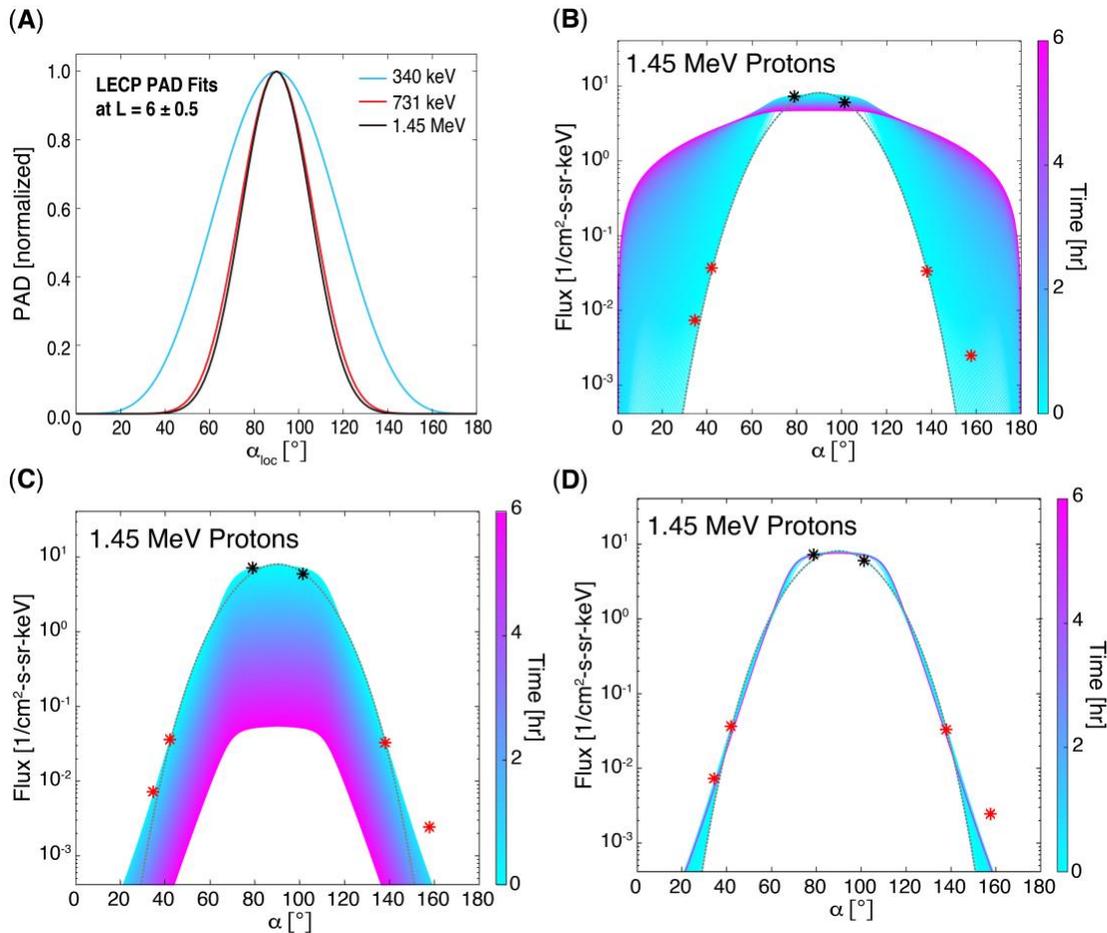
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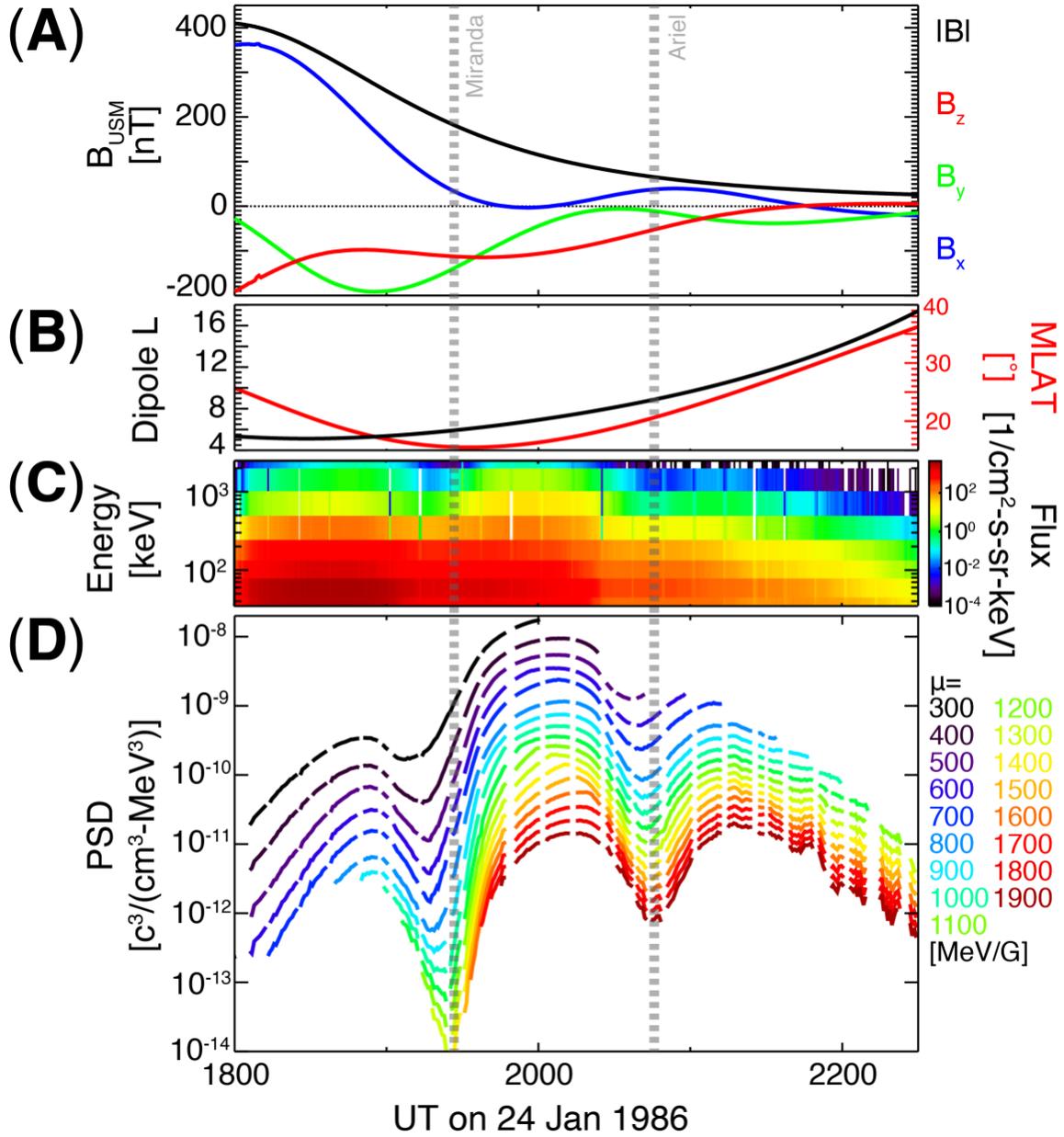
461 **Fig. 1.** Simulation of the 1986 Voyager 2 flyby encounter of Uranus. The complexity and
 462 dynamics of the Uranian magnetosphere are demonstrated by the variation in magnetic
 463 configuration between the inbound orbit (red) and outbound orbit (blue). Dipolar magnetic field
 464 lines at L-shells of 2, 4, 6, and 8 are shown for each configuration.
 465



466 **Fig. 2.** Simple one-dimensional pitch angle diffusion modeling suggests that an energetic particle
 467 source is required to sustain the steep pitch angle gradients observed by LECP. (A) Comparison
 468 of $\sin^n(\alpha)$ fit for the average PADs at three different >300 keV ion energy channels created by
 469 mapping ions at $L = 6 \pm 0.5$ from both the inbound (red; 24 Jan 1986 17:15-17:30 UT) and
 470 outbound (black; 24 Jan 1986 19:45-20:00 UT) legs of the flyby trajectory to the local pitch
 471 angle and combining them with the local pitch angle for the outbound trajectory using
 472 conservation of the first adiabatic invariant, μ . (B) The expected evolution of the 1.5-MeV ion
 473 (assuming protons) pitch angle distribution observed by LECP at $L=6$ assuming an Earth-like
 474 pitch angle diffusion coefficient ($D_{\alpha\alpha}$). (C) The same scenario but adding losses to the moons at
 475 all local pitch angles $<80^\circ$ on a 10-min loss timescale. (D) Adding a Gaussian pitch angle source
 476 centered at $\alpha = 90^\circ$ with $\sigma = 1.5^\circ$ and an amplitude of $S_0 = 1.26 \times 10^{11} \text{ cm}^{-2}\text{-s}^{-1}\text{-sr}^{-1}\text{-MeV}^{-1}/\text{s}$. This

477 analysis shows that pitch angle diffusion would eliminate the observed steep pitch angle
478 gradients unless a very narrow source is present.

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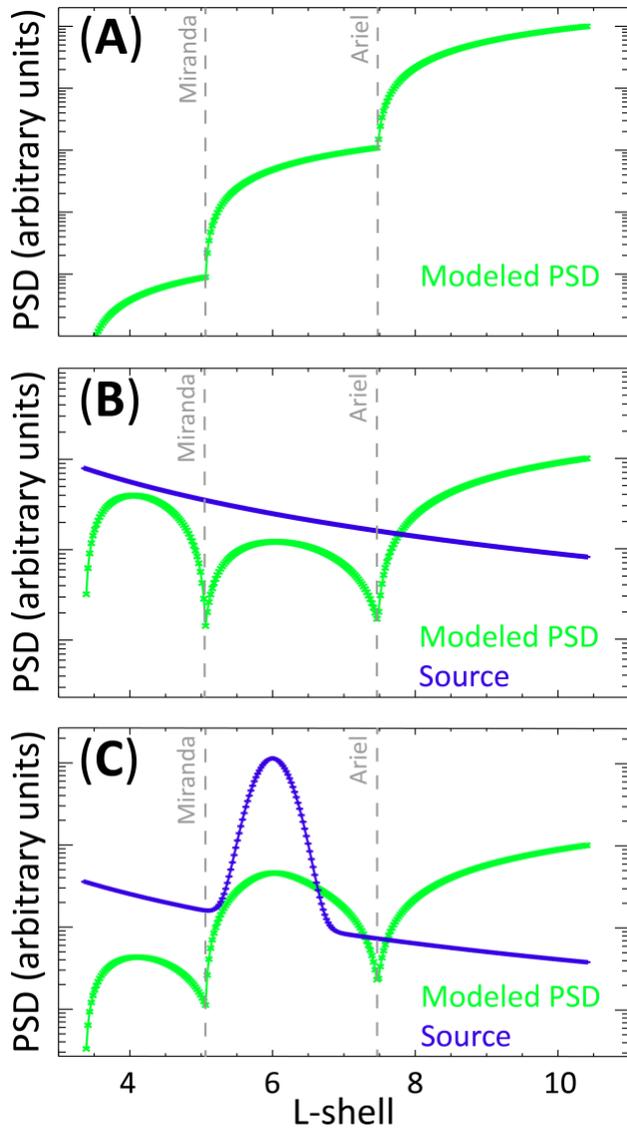


480

481 **Fig. 3.** New phase space density analysis clearly indicates a source of energetic ions in the region
 482 between Miranda and Ariel. (A) The observed magnetic field from the Voyager 2 magnetometer
 483 [34]. (B) The dipole-L calculated using the offset tilted dipole model from (Ness et al., 1986).
 484 (C) The ion flux observed by the LECF instrument (Krimigis et al., 1977). (D) The calculated
 485 PSD profiles determined using the LECF and MAG observations versus time for several values

486 of constant the first adiabatic invariant (μ). The minima surrounding the moons Miranda and
487 Ariel are due to energetic particle losses from interactions with the moons (e.g., Paonessa &
488 Cheng, 1987). However, the very clear maximum between Miranda and Ariel at $L \sim 7$ clearly
489 suggests a source of energetic ions in this region.

490



491

492 **Fig. 4.** Modeling of the expected radial distribution of ion phase space density shows that the
 493 observed global maximum between Miranda and Ariel requires a significant local source. (A)
 494 The expected PSD profile versus L-shell expected from only inward transport of plasma from the
 495 outer magnetosphere ($L > 11$), combined with absorption at the moon orbits. (B) The profile
 496 assuming a CRAND-like source falling off with distance from the planet (dark blue). (C) The
 497 profile adding a local source in the region between Miranda and Ariel to the simplified CRAND
 498 source.