

Influence of the Interdecadal Pacific Oscillation on Super Cyclone Activities over the Bay of Bengal During the Primary Cyclone Season

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Abstract

An obvious interdecadal change is measured in the super cyclones (SCs, categories 4 and 5) that occur from Oct-Nov over the Bay of Bengal (BoB). This change may be modulated by the interdecadal Pacific oscillation (IPO). A La Niña-like difference between the 1977–1998 (IP1) and 1999–2014 (IP2) periods forces a local Hadley circulation in the eastern tropical Indian Ocean by strengthening the Walker circulation, causing a plummeting upper-level temperatures and ultimately creating favorable thermodynamic conditions to enhance the cyclone intensity. Meanwhile, an equatorial downwelling Kelvin wave caused by heating and westerly wind differences enters the BoB rim along the coast and aptly intensifies the cyclone such that the downwelling Kelvin wave and Rossby wave generated by its reflection deepen the thermocline in the BoB. The favorable atmospheric and oceanic conditions in IP2 jointly and preferentially form far more SC activities from Oct-Nov over the BoB compared to IP1.

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Influence of the Interdecadal Pacific Oscillation on Super Cyclone Activities over the Bay of Bengal During the Primary Cyclone Season

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Key Points:

A La Niña-like pattern strengthens the Walker circulation causing a local Hadley circulation and Kelvin wave in the tropical Indian Ocean.

The local Hadley circulation affects upper-level temperatures over the BoB and further modulates the different SCs between IP2 and IP1.

The downwelling Kelvin wave and its reflected Rossby wave deepen the thermocline in the BoB to favor SC formation in the IP2.

Abstract

An obvious interdecadal change is measured in the super cyclones (SCs, categories 4 and 5) that occur from Oct-Nov over the Bay of Bengal (BoB). This change may be modulated by the interdecadal Pacific oscillation (IPO). A La Niña-like difference between the 1977–1998 (IP1) and 1999–2014 (IP2) periods forces a local Hadley circulation in the eastern tropical Indian Ocean by strengthening the Walker circulation, causing a plummeting upper-level temperatures and ultimately creating favorable thermodynamic conditions to enhance the cyclone intensity. Meanwhile, an equatorial downwelling Kelvin wave caused by heating and westerly wind differences enters the BoB rim along the coast and aptly intensifies the cyclone such that the downwelling Kelvin wave and Rossby wave generated by its reflection deepen the thermocline in the BoB. The favorable atmospheric and oceanic conditions in IP2 jointly and preferentially form far more SC activities from Oct-Nov over the BoB compared to IP1.

Plain Language Summary

The interdecadal Pacific oscillation (IPO) modulates pronounced interdecadal variations in super cyclones (SCs, categories 4 and 5) over the Bay of Bengal (BoB) from Oct-Nov, introducing serious casualties and heavy economic losses to nearby coastal regions. Compared with previous studies on interdecadal variations in tropical cyclones (TCs) over the BoB, this study investigates the

possible processes by which the IPO affects the TC intensity in the BoB from the perspective of air-sea interactions. A La Niña-like difference observed between two IPO phases (1977–1998 (IP1) and 1999–2014 (IP2)) enhance the local Hadley circulation in the eastern tropical Indian Ocean by strengthening the Walker circulation, modulating the upper-level temperature to increase atmospheric instability, and ultimately causing favorable conditions for SC formation in IP2. Meanwhile, the intensified Walker circulation cell over the Indian Ocean generates a downwelling Kelvin wave and deepens the thermocline while propagating along the BoB rim. Furthermore, the Rossby wave caused by the reflecting Kelvin wave enters the BoB interior, deepening the thermocline. This thickened thermocline is conducive to TC intensification. All these factors work together to ultimately lead to the significant interdecadal variation observed in SCs between IP1 and IP2.

Introduction

Tropical cyclones (TCs) are one of the most common severe weather systems and among the most destructive natural disasters (Emanuel 2003; Pielke et al. 2008; Blake et al. 2011). Although fewer TCs form over the BoB than over other tropical basins, such as the Northwest Pacific (NWP) and the Northeast Pacific (NEP) basins (BoB TCs account for only approximately 4% of the TCs that occur globally), the TCs, particularly the super cyclones (SCs, categories 4 and 5), which hit the BoB rim incur numerous casualties, accounting for approximately 80% of TC-induced fatalities worldwide (Needham et al. 2015). For example, TC Nargis (Webster 2008; Kikuchi et al. 2009; Lin et al. 2009; McPhaden et al. 2009; Yanase et al. 2010) struck the southern coast of Myanmar and caused a death toll greater than 130000 people, becoming one of the deadliest natural disasters ever recorded in Myanmar.

The TCs that occur in the BoB have a very unique bimodal feature throughout the annual cycle: TCs mainly occur during the pre-monsoon season (Pre-MS, Apr-May) and post-monsoon season (Post-MS, Oct-Nov) (Yanase et al. 2012; Li et al. 2013). The Post-MS and Pre-MS periods are the primary season (PS) and secondary season (SS) in which TCs occur, as the TC frequency is approximately 4 times higher during the PS than during the SS. Li et al. (2013) attributed this asymmetric TC frequency to the mid-level relative humidity (RH) difference that occurs between the PS and SS. Akter and Tsuboki (2014), from the perspective of dynamics, analyzed the vertical wind shear (VWS) and local trough regulated TC genesis difference between the PS and SS.

Apart from the annual cycle described above, BoB TCs also undergo obvious interannual variations. The El Niño/Southern Oscillation (ENSO) (McPhaden 2002) and Indian Ocean Dipole (IOD) (Saji 1999; Webster 1999) are the dominant interannual modes of coupled ocean-atmosphere systems in tropical regions, and these modes exert great influence on the frequency, intensity and tracks of TCs over the Indian Ocean by inducing large-scale atmospheric circulation (Wang and Chan 2002; Ho et al. 2006; Camargo et al. 2007; William and Young 2007; Eric and Chan 2012; Li et al. 2015a). However, previous studies have

given less attention to the effects of these modes in the BoB. Girishkumar et al. (2012) reported that the enhanced low-level cyclonic vorticity and enlarged upper-ocean thermal energy that occur under the La Niña regime are conducive to TC activity, while under the El Niño regime, the situation is opposite to that observed under La Niña. There is a phase-locking relationship between the mature IOD phase and the primary peak BoB TC phase. In response to the negative IOD (NIOD) mode, a cyclonic anomaly occurs over the BoB. This anomaly not only strengthens the low-level vorticity but also increases the middle-level moisture by enhancing vertical advection. All of these conditions are favorable to TC genesis. The influences of the positive IOD (PIOD) mode and NIOD are approximately mirror images (Singh et al. 2008; Yuan and Cao 2013; Li et al. 2015b).

ENSO and IOD are modulated by background field conditions with longer time scales. Based on the relatively short-term historic observation datasets available, some studies analyzing data collected from 1950 found large-scale atmosphere-ocean field regime shifts in the tropical Pacific in 1977 and 1999 (Wang 1995; Mantua et al. 1997; McPhaden et al. 2011). Such regime shifts were found to be associated with the interdecadal Pacific oscillation (IPO), which could be regarded as a mechanism by which the background conditions required for ENSO change (Hu et al. 2018). In other words, different IPO phases could regulate the structure and evolution of ENSO and further control TC genesis in some sense.

Motivated by the above idea, we conducted a preliminary statistical analysis. The 1977–2014 period was divided into two interdecadal periods according to the IPO phase: the two periods comprised 1977–1998 (IP1) and 1999–2014 (IP2) (Yao et al. 2017; Hu et al. 2018). The most pronounced difference between the periods is that the annual average number of SCs that occurred in the BoB during the PS was approximately 8 times greater in IP2 than in IP1 (Fig 1). An investigation of what and how environmental factors caused this very significant difference in TC intensity between IP2 and IP1 would be very interesting.

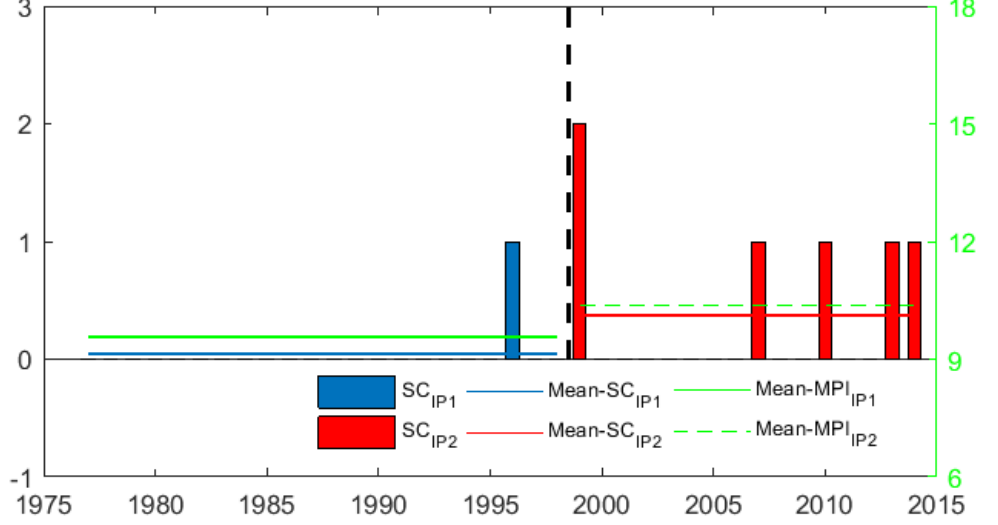


Fig 1. Statistical analysis of SCs over the BoB during the PS in IP1 (1977–1998) and IP2 (1999–2014). The blue bar is the SC frequency in IP1, and the blue line is the annual average number of SCs. Similar to the symbols used for IP1, the red symbols denote IP2. The solid and dashed green lines denote the average maximum potential intensity (MPI) values in the PS of IP1 and IP2, respectively.

The objective of this study is to reveal physical processes that led to the remarkable difference in SC activities observed in the BoB during the PS between the two interdecadal periods described above. The results of this study will advance the understanding of interdecadal SC variations in the BoB and provide a meaningful reference for improving SC predictions. The remainder of this paper is organized as follows. In Sect. 2, the data and analytical methods are introduced. The potential physical processes responsible for the significant interdecadal variation observed in SC activities in the BoB during the PS are discussed in Sect. 3. Finally, a summary and conclusions are provided in the last section.

Data and Methods

The data used in this study include (1) a TC best-track dataset obtained from the International Best Track Archive for Climate Stewardship (IBTrACS) (Knapp et al. 2010), (2) monthly sea surface temperature (SST) data obtained from the Extended Reconstructed SST product of the National Oceanic and Atmospheric Administration (NOAA), (3) monthly wind, air temperature (T), specific humidity (SH), RH, vertical motion (), precipitation rate and geopotential height (GH) data obtained from a National Centers for Environmental Prediction (NECP)-National Center for Atmosphere Research (NCAR) reanaly-

sis product (Kalnay et al. 1996) and (4) monthly sea surface height (SSH) data retrieved from the Global Ocean Data Assimilation System (GODAS) product of the National Centers for Environmental Prediction (NCEP). All the data are analyzed over the period from 1977–2014. The SST and SSH data have horizontal resolutions of 2° latitude by 2° longitude and 0.33° latitude by 1° longitude, respectively. All other datasets have a resolution of $2.5^\circ \times 2.5^\circ$.

The MPI can denote the maximum possible TC intensity (Bister and Emanuel 2002; Emanuel and Nolan 2004; Camargo and Sobel 2005). The TC intensity approaches the MPI during the PS over the BoB since the bottom-level (1000 hPa level) SH is close to the saturated SH at this time. Under these conditions, the MPI is available to characterize the TC intensity during the PS over the BoB. Thus, we can use the MPI to diagnose the difference in TC intensities between IP2 and IP1. Here, the MPI formula can be written as follows:

$$V_{\text{pot}}^2 = C_p \bullet \frac{C_k}{C_D} \bullet T_{\text{vd}} \bullet T_{\text{rd}} \bullet T_{\text{ed}} \quad (1)$$

$$T_{\text{vd}} = T_s - T_o$$

$$T_{\text{vr}} = \frac{T_s}{T_o}$$

$$T_{\text{ed}} = \ln \theta_e^* - \ln \theta_e$$

where V_{pot} is the MPI, C_p is the heat capacity at a constant pressure, C_k is the exchange coefficient for enthalpy, C_D is the drag coefficient, T_s is the surface temperature, T_o is the mean outflow temperature, θ_e^* is the saturation equivalent potential temperature at the ocean surface, and θ_e is the boundary-layer equivalent potential temperature.

We adopted a total differentiation method similar to that developed by Li et al. (2013) to examine the relative contribution of each term. Then, the contribution of each term to the difference in V_{pot}^2 between the PS of IP2 and IP1 can be written as follows:

$$V_{\text{pot}}^2 = \alpha_1 \cdot T_{\text{vd}} + \alpha_2 \cdot T_{\text{vr}} + \alpha_3 \cdot T_{\text{ed}} \quad (2)$$

$$V_{\text{pot}}^2 = V_{\text{potIP2}}^2 - V_{\text{potIP1}}^2$$

$$T_{\text{vd}} = T_{\text{vdIP2}} - T_{\text{vdIP1}}$$

$$T_{\text{vr}} = T_{\text{vrIP2}} - T_{\text{vrIP1}}$$

$$T_{\text{ed}} = T_{\text{edIP2}} - T_{\text{edIP1}}$$

Herein, an approximation has been made in deriving Eq. (2) by assuming that α_1 , α_2 and α_3 are constant coefficients, such that the following terms are satisfied:

$$\begin{cases} \alpha_1 = T_{\text{vrIP1}} \bullet T_{\text{edIP1}} \\ \alpha_2 = T_{\text{vdIP1}} \bullet T_{\text{edIP1}} \\ \alpha_3 = T_{\text{vdIP1}} \bullet T_{\text{vrIP1}} \end{cases}$$

The box difference index (BDI) (Fu et al. 2012; Peng et al. 2012; Li et al. 2019) and T-test are used herein to quantitatively contrast the differences in the environmental factors associated with TC development between IP2 and IP1 to ascertain the substantially different factors and their significance levels. The BDI is defined as follows:

$$\text{BDI} = \frac{(\text{Mean}_{IP2} - \text{Mean}_{IP1})}{(\sigma_{IP2} + \sigma_{IP1})}$$

where ‘Mean’ denotes the mean of all samples for one particular variable and σ denotes the standard deviation. The subscripts ‘IP1’ and ‘IP2’ represent the 1977–1998 and 1999–2014 periods, respectively.

A composite analysis was conducted based on the monthly data to reveal the predominant physical processes that modulate the observed differences in SC activities over the BoB during the PS between IP2 and IP1.

Analysis and Results

As was previously determined, 1977 and 1999 were key shift points for the IPO. Compared with the 1977–1998 period, the 1999–2014 period was characterized by the negative-phase IPO, and an obvious La Niña-like pattern could be observed in the Pacific Ocean at this time. Six SCs occurred over the BoB during the PS of IP2, and the annual average number of SCs that occurred during the PS of IP2 was in stark contrast with that recorded in IP1 (Fig 1), which is approximately 8 times more in IP2 than in IP1. According to the tracks of these SCs, the region of 7.5–20°N and 85–92.5°E is regarded as the reference area in this study to carry a diagnostic analysis on the differences in SC activities between the two periods (Fig 2).

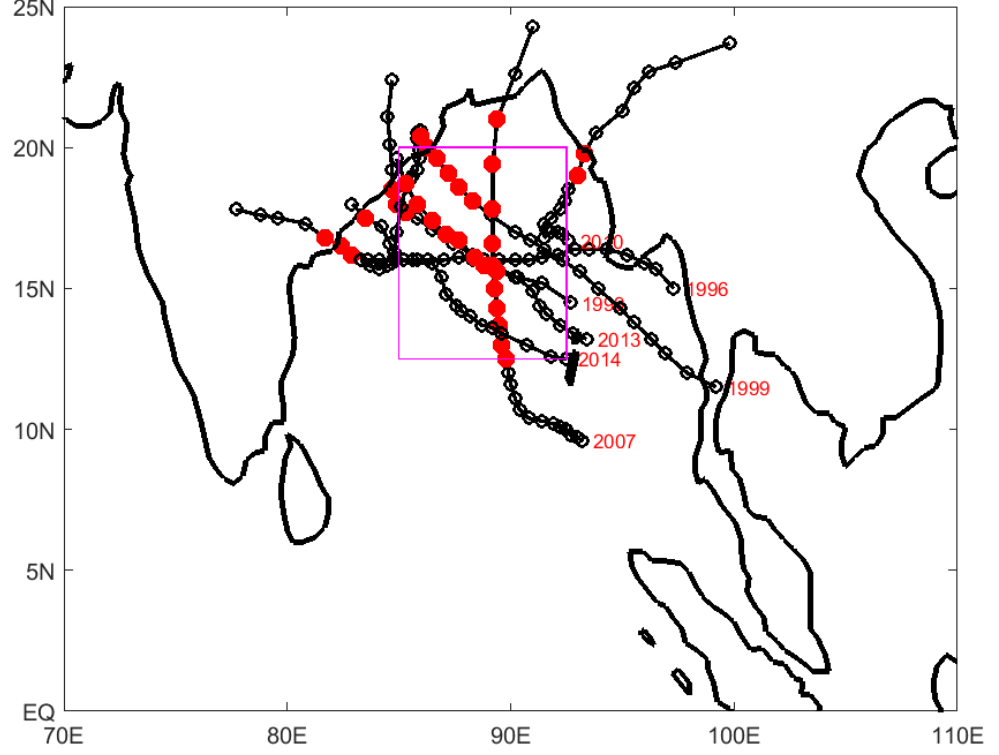


Fig. 2. The center tracks of SCs during PS of IP1 and IP2 over the BoB. The red dot means TC speed reaches category 4 and 5.

Similar to the difference observed in the average annual number of SCs between IP1 and IP2, the reference-box-averaged V_{pot}^2 values, which could effectively mirror the MPI values (Bister and Emanuel 2002), also significantly differed between IP2 and IP1 at the 95% confidence level. The components of V_{pot}^2 , such as T_{vd} , T_{vr} and T_{ed} , also reveal some differences at significance levels greater than 99%, more than 99% and less than 90%, respectively (Table 1). Hence, V_{pot}^2 can be regarded as a diagnostic equation with which the contributions of environmental factors associated with MPI to the SC differences between IP2 and IP1 can be quantitatively analyzed.

Table 1. Differences in the studied terms between IP2 and IP1, their corresponding significance levels, and the relative contributions of the changes in various terms to ΔV_{pot}^2 .

	V_{pot}^2	T_{vd}	T_{vr}	T_{ed}
	V_{pot}^2	T_{vd}	T_{vr}	T_{ed}
IP2-IP1	>0	>0	>0	>0
Significance of difference	>95%	>99%	>99%	<90%
The proportion of the contribution to V_{pot}^2	$T_{\text{vd}} \& T_{\text{vr}}$ 89%	T_{ed} 11%		

The diagnostic results show that the greatest contribution to the observed difference in SC activities between the PS periods in IP2 and IP1 was provided by the T_{vd} and T_{vr} terms. Because T_{vd} and T_{vr} are determined only by the surface and outflow temperatures, their effects could be regarded together. Table 1 illustrates that most of the contribution came from T_{vd} and T_{vr} , which together accounted for approximately 89% of the contribution (Table 1). T_{ed} provided only approximately 11% of the contribution to ΔV_{pot}^2 (Table 1). We could also conclude that surface and outflow temperatures may be the most important environmental factors modulating the remarkable observed difference.

Table 2. BDI value of each environmental factor considered in the reference at multiple atmospheric levels during the PS

BDI	1000	925	850	700	600	500	400	300	200	100
(IP2- IP1)										
Vor			*	*	*					*
T	*	*	*	*	*	*		*	*	*
RH					*	*	*	*		
SH						*	*	*		

Note: the * symbol indicates that the significance level is >95%.

To further verify the diagnostic results described above and quantitatively contrast the changes observed in other environmental factors, such as the T, vorticity, VWS, RH and SH, affecting TC genesis and development (Gray 1968, 1979; Emanuel and Nolan 2004; Murakami and Wang 2010), we introduced the BDI to quantitatively adjust the difference of the environmental factors with different unit. The diagnosis results are listed in Table 2. T was significantly different between IP2 and IP1 at almost every level, especially at the 100-hPa level (Table 2). Although the mid-level RH and SH were lower in IP2 than in IP1, they mainly affected the genesis of TCs rather than the intensity. Given that the boundary-layer SH was slightly more abundant in IP2 than in IP1, this

factor naturally had little influence on the observed difference in TC intensity (Table 2).

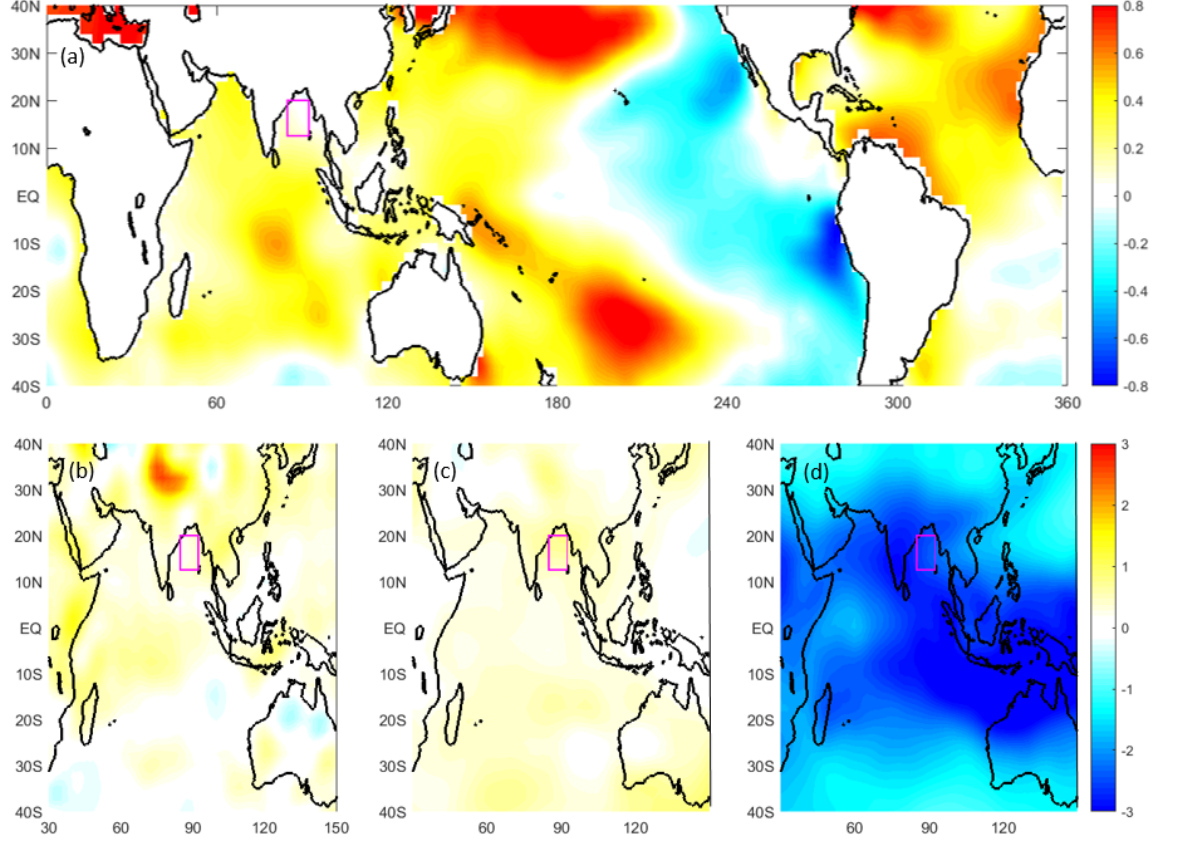


Fig 3. Composite differences in SST and T during the PS between IP2 and IP1. Panel (a) shows the SST difference (unit: °C). Panels (b), (c) and (d) show the T differences at the 1000-hPa, 500-hPa and 100-hPa levels, respectively (unit: °C). The red box is the reference area in the BoB.

Based on the above analysis of the diagnostic results, we think T, especially upper-level T, may be the most critical factor that determined the different SC formation mechanisms between IP2 and IP1. Hence, we further analyzed the SST data and the 3-dimensional T structure. Except for an obvious La Niña-like pattern in the Eastern Pacific, warming could be observed almost basin-wide in the tropical Indian Ocean between the PS of IP2 and IP1 (Fig 3(a)). The bottom- and middle-level (1000- and 500-hPa-level) T warming nearly fit the SST differences observed in the tropical Indian Ocean (Fig 3(b) and (c)), and the corresponding BDIs were 0.50 and 0.73, respectively (Table 2). The upper-level (100-hPa-level) T was evidently colder in IP2 than in IP1, particularly in the Indo-Pacific warm pool region (Fig 3(d)), and the corresponding BDI

was -1.28. Therefore, among the studied factors, the difference in upper-level T might play the most important influential role in the atmospheric stability differences observed between IP2 and IP1.

In the reference area, the composites of T with climatology removed for IP1 and IP2 showed values of 0.14°C and 0.47°C, respectively, at the 1000-hPa level and of 1.09°C and -1.42°C, respectively, at the 100-hPa level (Fig 4). The T differences between IP2 and IP1 were 0.33°C and -2.51°C at the 1000-hPa and 100-hPa levels, respectively. The vertical temperature difference (VTD) composite values obtained for IP2 and IP1 were approximately -0.95°C and 1.89°C, respectively, and the VTD difference between IP2 and IP1 was 2.84°C. This obvious disparity in the VTD could induce a remarkable difference in the MPI between IP2 and IP1 with a significance level above 95% and might further modulate the differential SC formation between IP2 and IP1. It is worth noting that this remarkable VTD difference should be mostly derived from the upper-level T difference between the PS periods of IP2 and IP1 (Fig 4). The vertical temperature ratio (VTR) has a similar result because the VTR is also determined by the lower and upper levels. These analysis results further confirm the above conclusion that the upper-level T difference is thought to be the most important environmental factor affecting the differential SC formation between the PS periods in IP2 and IP1.

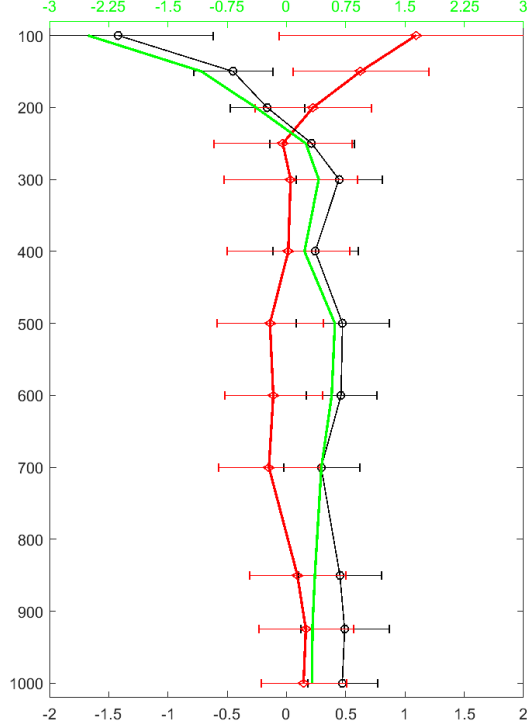


Fig 4. Composite profiles of T with climatology removed, showing the standard errors and differences between IP2 and IP1 (unit: $^{\circ}\text{C}$). The red line is the box-averaged temperature measured in IP1, the red stem is the standard error, the black line and stem are the corresponding values measured in IP2, and the green line is the temperature difference between IP2 and IP1.

Determining the reason for the very significantly negative difference observed in T at the 100-hPa level between the PS periods of IP2 and IP1 naturally becomes the key objective we will investigate next. It is known that a La Niña-like pattern could strengthen the Walker circulation in the Pacific (Collins et al. 2010; Dinezio et al. 2010). We found that under the La Niña-like SST difference that occurred during the PS periods between IP2 and IP1, the Walker circulation was enhanced in the tropical Pacific Ocean. In response to this enhanced Walker circulation in the Pacific, the other Walker circulation cell over the Indian Ocean also became stronger in IP2 than it was in IP1. The strengthening Walker circulation cells induced bottom-level westerly and easterly wind differences over the tropical Indian Ocean and Pacific Ocean and increased precipitation over the Java Sea and its adjacent regions in the maritime continent (Fig 5(a)). Unlike

the increase in precipitation observed over the eastern tropical Indian Ocean, especially in the region west of the island of Sumatra, the precipitation lessened in the BoB in IP2 compared to that measured in IP1 (Fig 5(a)). Simultaneously, the bottom-level wind spanned the equator from the BoB and converged in the tropical southeast Indian Ocean with a positive precipitation difference (Fig 5(a)). All of these results imply that the local Hadley circulation might play an important role in modulating T change over the BoB.

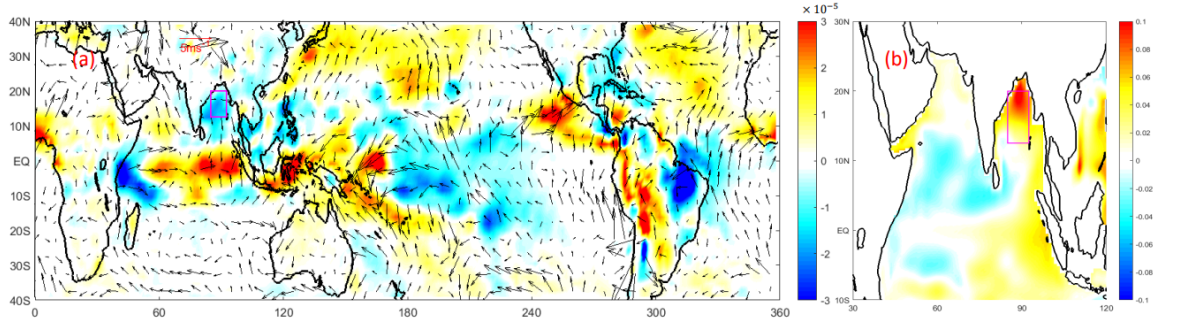


Fig 5. Composite differences in the precipitation rate (shading, unit: $\text{kg m}^{-2} \text{s}^{-1}$), 1000-hPa-level wind (vector, unit: m s^{-1}) (a) and SSH (shading, unit: m) (b) during the PS between IP2 and IP1. The red box is the reference area in the BoB.

A detailed analysis of the along-latitude vertical-motion profile is therefore carried out next. The $85^{\circ}\text{--}92.5^{\circ}\text{E}$ -averaged composite results show an anomalous ascending motion that is consistent with the enhanced convection observed in the western region of Sumatra. This ascending branch is well connected to a descending branch over the BoB through a local anomalous Hadley circulation (Fig 6 (a)). The high pressure difference at the 200 and 300 hPa level coincides with the descending difference this correspondence may result from convergence in the upper branch of the high troposphere. In contrast with the 200-hPa level, low pressure is observed at the 100-hPa level above the descending low pressure centers arise corresponding to positive and negative potential height changes, respectively, according to the hydrostatic relationship. Fig 6(b) shows a negative potential height change at 100 hPa but a positive potential height change at 200 hPa (Fig 6(b)). Naturally, the 100-hPa level becomes anomalously cold due to the reduced thickness between the 100-hPa and 200-hPa levels.

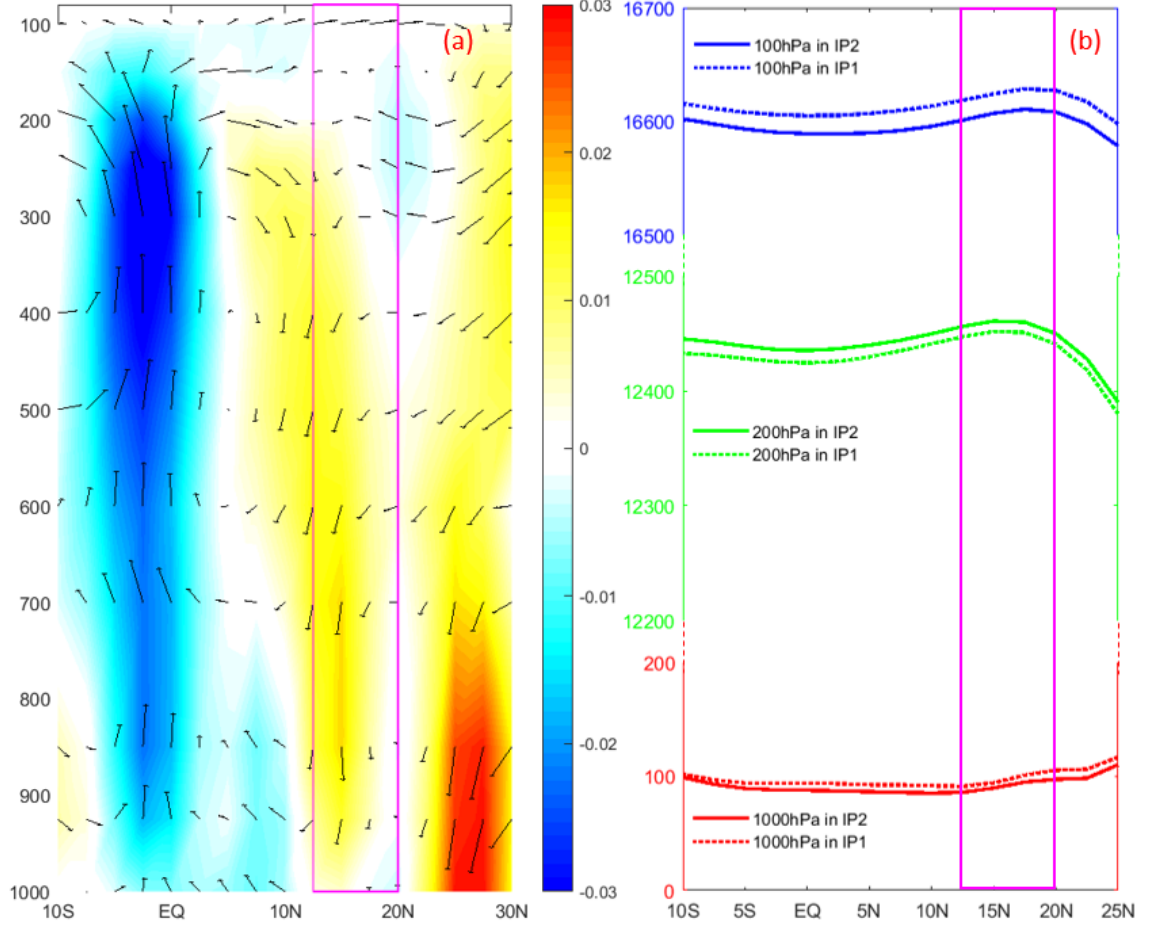


Fig 6. Meridional-vertical section of the 85°-92.5°E-averaged composite difference in vertical motion (shading) and the overturning circulation (vector) in Oct-Nov between IP2 and IP1 (a). The vectors represent the zonal velocity (unit: m s^{-1}) and vertical velocity (unit: Pa s^{-1}). Panel (b) shows the potential heights of 1000, 200 and 100 hPa levels in IP1 and IP2 (unit: m).

Through the above analyses, we comprehend the mechanism by which the observed high-level negative temperature difference formed between IP2 and IP1 and then sought to understand the physical process that modulated the significant increase in the number of SC occurrences during the PS period in IP2 from the atmospheric dynamics and thermodynamics perspectives. In addition to the possible effects of atmospheric processes, the ocean, as the underlying surface of TC activities, also established favorable conditions for enhanced TC intensities over the BoB during the PS in IP2 compared to those in IP1. Heating in the central and eastern tropical Indian Ocean and the westerly wind difference below this heating force a downwelling Kelvin wave to propagate eastward

along the equator; then, this Kelvin wave moves along the rim of the BoB as a coastal Kelvin wave (Fig. 5(b)). Furthermore, the Kelvin wave in the eastern BoB radiates westward as a downwelling Rossby wave and enters the interior of the BoB. This poleward-propagating coastally downwelling Kelvin wave and the westward-propagating downwelling Rossby wave effectively deepen the thermocline in the BoB during the PS periods in IP2 (Fig. 5(b)). This thickened thermocline results in an enlarged upper-ocean heat content (Yu 2003; Girishkumar et al. 2012) and a more stable stratification of the upper Ocean; these conditions are favorable to TC intensification. The processes described in this study jointly construct favorable atmospheric and oceanic conditions that ultimately result in the significant difference observed in SC activities over the BoB during the PS periods between IP2 and IP1.

Conclusions

The IPO experienced a positive and negative phase during the 1977–1998 (IP1) and 1999–2014 (IP2) periods, respectively. Compared to IP1, an obvious La Niña-like pattern occupied the tropical Pacific during IP2, and this pattern even impacted the global climate in some sense through teleconnections. As a possible response to this La Niña-like pattern, approximately 8 times more annual average SC incidences occurred over the BoB during the PS periods in IP2 than in IP1.

Our analysis indicates that the La Niña-like pattern strengthens the Walker circulation and then enhances convection in the central and eastern tropical Indian Ocean, especially in the area west of Sumatra. The enhanced convection in the southeast Indian Ocean induces a local anomalous Hadley circulation, and the descending branch of this circulation is located over the BoB, causing a high-pressure center at the 300-hPa level. Simultaneously, a low-pressure center occurs above the descending branch at approximately the 100-hPa level. According to the hydrostatic relationship, the 100-hPa low-pressure center and 200-hPa high-pressure center could effectively reduce the thickness between the 100- and 200-hPa levels and cause a remarkable negative T difference at the upper level between IP2 and IP1. Furthermore, a relatively small positive T difference can be observed at the bottom level, and this difference may be associated with the positive SST difference observed in the BoB between IP2 and IP1. All of these results support the positive differences observed in T_{vd} and T_{vr} between IP2 and IP1. Certainly, these positive differences can be mainly attributed to the upper-level T difference between the two periods. Finally, these differences could affect the MPI difference and even modulate the differential annual average number of recorded SCs between the PS periods of IP2 and IP1.

In addition, the positive heating and westerly wind differences observed between IP2 and IP1 caused the downwelling Kelvin wave to propagate eastward along the equator. This Kelvin wave traverses the rim of the BoB as a coastal Kelvin wave and deepens the thermocline in the BoB. Then, the Kelvin wave is reflected at the eastern coast of the BoB into the downwelling Rossby wave that enters the interior of the BoB and increases the thermocline depth in this region. This

deep thermocline is also favorable for TC intensification.

The combination of atmospheric and oceanic conditions analyzed in this study ultimately modulates the significant difference in SC activities observed over the BoB between the PS periods of IP2 and IP1. This study has important significance as a reference for understanding and predicting SCs over the BoB.

Data Availability Statement

The data that support the findings of this study are publicly available and can be derived from the following sources: TC best-track dataset (<https://www.ncdc.noaa.gov/ibtracs/index.php?name=ib-v4-access>), and NCEP/NCAR and NOAA reanalysis data (<https://psl.noaa.gov/data/gridded/reanalysis/>).

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