

# Erosion Rates on Newly Uplift Marine Terraces Following the 2016 Kaikōura Magnitude 7.8 (Mw) Earthquake

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November 22, 2022

## Abstract

Since 1973 micro-erosion meters (MEM) have been used at Kaikōura Peninsula to determine lowering rates on inter-tidal shore platforms. Rates measured over two, two year periods (1973-1975 and 1994-1996) and at decadal scales (20-30 years) demonstrate that platform surface lowering is on average 1.1 mm/yr. The 14 November 2016 Kaikōura magnitude 7.8 (Mw) earthquake caused an instantaneous uplift of 0.8-1.0 m of the peninsula. The uplift offers the rare opportunity to examine how such an event alters processes and rates of erosion on these shore platforms, since these are now partially marine terraces as the inner margins of some platforms are now above high tidal levels (but perhaps not storm surge). Since the earthquake, 42 MEM sites have been measured seven times at 3 monthly intervals. Most recently in September 2018. MEM sites show widely varying responses to the uplift. Erosion rates are at some MEM sites three times the previous annual rate while other sites show significant amounts of rock swelling (3-4 mm in 6 months), or aggradation as weathered rock fragments are no longer removed by wave action. The coseismic uplift has fundamentally changed the process regime operating on these platforms. Zones of maximum wetting and drying have migrated seaward causing previously slow eroding (< 1 mm/yr) MEM sites to accelerate to twice the pre-earthquake rates. Erosion rates demonstrate rapid adjustment of the platform surface to this disturbance and illustrate how uplifted marine terraces can be rapidly eroded despite being above sea level. The preservation of the new marine terrace is probably dependent on further uplift within the next 300-400 years, otherwise erosion by lowering and backwear will likely remove the new surface. This scenario has significant implications for marine terrace preservation and the recording of coseismic events in the landscape.

# Erosion Rates on Newly Uplifted Marine Terraces Following the 2016 Kaikōura $M_w$ 7.8 Earthquake

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## Introduction

Since 1973, micro-erosion meters (MEM) have been used at Kaikōura Peninsula, New Zealand (Fig 1) to determine lowering rates on inter-tidal shore platforms. Rates measured over two, two year periods (1973-1975 and 1994-1996) and at decadal scales (20-30 years) demonstrate that platform surface lowering is on average 1.1 mm/yr. The 14 November 2016  $M_w$  7.8 Kaikōura earthquake caused an instantaneous uplift of 0.8-1.0 m of the peninsula (Fig. 2). The uplift offers the rare opportunity to examine how such an event alters processes and rates of erosion on these shore platforms (Fig. 3), since these platforms are now partially marine terraces as the inner margins of platforms are now above high tidal levels (but perhaps not storm surge, Fig. 3). Here we report 21 months of erosion monitoring since the earthquake, with a view to establishing the longevity of the newly uplift surface and the altered erosion rates on the shore platforms.

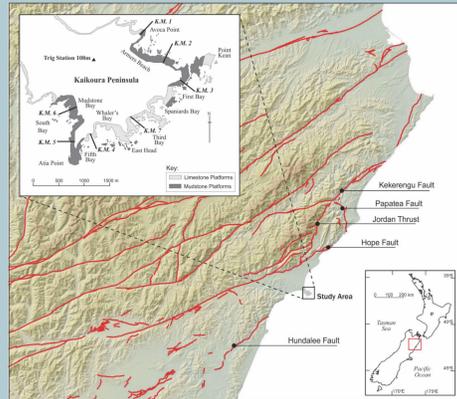


Figure 1. Location of major faults in the upper South Island, New Zealand and Kaikōura Peninsula showing location of MEM profiles and Trig station. Named faults are some of those that ruptured during the November 14 2016 earthquake (Langridge et al. 2016; Stephenson et al. 2017; Litchfield et al. 2018).



Figure 2. Seaward edge of a raised shore platform. Holdfasts of bull kelp (*Durvillaea antarctica*), previously located in the subtidal zone (18/12/2016).



Figure 3. View of a raised and inter-tidal shore platform. North side of Kaikōura Peninsula. Profile KM2 (see Fig. 5), crosses centre of platform (18/12/2016).

## Methods

Surface lowering on shore platform have been measured using the MEM (High and Hanna 1970). The MEM It comprises an engineering dial gauge attached to an equilateral triangle base, with three legs. The instrument sits atop three bolt permanently installed in the rock surface and precisely relocates using an Kelvin Clamp. The base is rotated 120° three times and the three readings are averaged, this average is then converted to a mean annual erosion rate (or swelling) for the site. The bolt sites were installed in 1973 on 6 profiles around the Peninsula and a seventh profile was added in 1993 (Fig 1.). Five profiles are located on Oligocene mudstone and two on Paleocene limestone. These profiles were resurveyed in December 2016 to determine the amount of uplift based on static GNSS observations (Stephenson et al. 2017). Since the earthquake, 45 MEM sites have been measured seven times at 3 monthly intervals. Most recently in September 2018.



Figure 4. The Micro-erosion Meter, used at Kaikōura since 1973 (Kirk, 1977). The MEM sits atop three bolts installed in the rock (KM2B). Note these sit above the surface, when installed these were recessed. Evidence of the long term lowering of the surface.

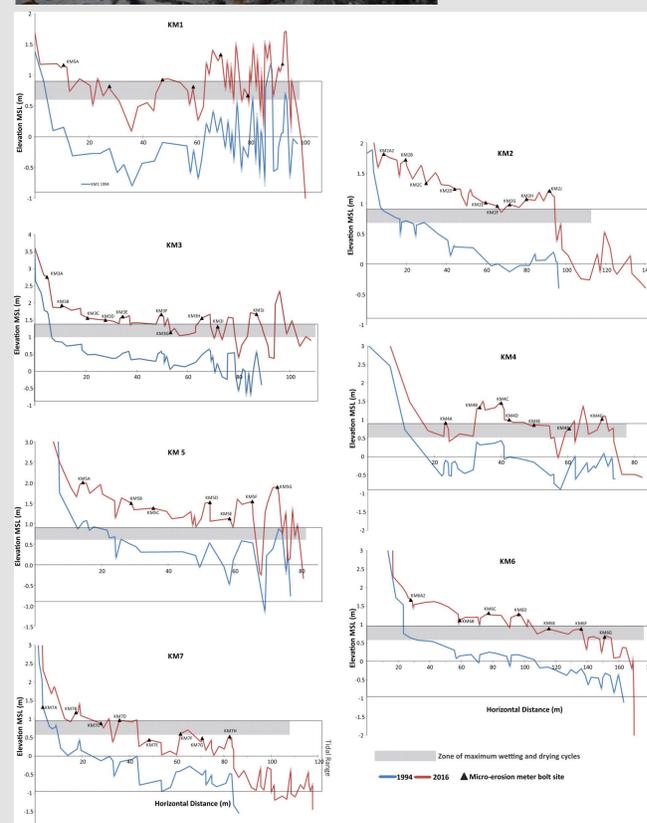


Figure 5. Surveyed platform profiles at Kaikōura and locations of MEM bolt sites. The blue profile was surveyed in 1994 and the red profile in December 2016, 4 weeks after the earthquake, following  $\approx 1$  m of uplift. Grey zone – elevation of maximum wetting and drying (Stephenson et al. 2017).

## Results

Mean erosion rates (mm/yr) at each measurement period by profile are shown in Table 1. These are compared to pre-earthquake rates from 2 (1993-1996) and 10 years (1993-2014). Mean rates mask highly variable inter-survey rates at individual bolt sites. Figure 6 illustrates this variability for three of the seven profiles, where both surface lowering and swelling are shown. Illustration of that variability are shown in the photo panels of Fig. 6. Rapid break down is evident (Fig 6A.), as is swelling and burial as weathered debris accumulates on the surface (Fig 6B), then larger scale (cm rather than mm) break down is evident in Fig. 6C. Erosion rates from before the earthquake were compared to those after using nonparametric tests of median distribution. Table 2 shows that these significantly different, with different means and distributions.

## Implications

The coseismic uplift has altered the process regime operating on these platforms. Zones of maximum wetting and drying have migrated seaward causing previously slow eroding (< 1 mm/yr) MEM sites to accelerate to 2.5 times the pre-earthquake rates. Erosion rates demonstrate rapid adjustment of the platform surface to this disturbance and illustrate how uplifted marine terraces can be rapidly eroded despite being above sea level. Based on the uplift of 0.8 to 1.0 m and assuming a constant lowering rate of 2.5 mm/yr, the new marine terrace could be removed within the next 320-400 years (not accounting for backwear). Preservation of the new terrace is dependent on further uplift beyond the reach of the sea, within that timeframe. Furthermore the amount of terrace preserved will depend on the timing of that uplift within that time frame. This scenario has significant implications for marine terrace preservation and the recording of coseismic events in the landscape. Since the erosion rates we have measured show that newly uplift surface can be rapidly removed from the landscape and there is potential for the record of coseismic uplift to be lost.

Table 1 Mean erosion rates (mm/yr) for each profile.

| Profile | 1993 to 1996 | 1993 to 2004 | 30/03/2017 | 24/06/2017 | 9/10/2017 | 18/12/2017 | 13/03/2018 | 28/06/2018 | 9/09/2018 |
|---------|--------------|--------------|------------|------------|-----------|------------|------------|------------|-----------|
| KM1     | 0.614        | 0.671        | 2.199      | 0.100      | 0.925     | 1.663      | 1.684      | 1.520      | 2.031     |
| KM2     | 1.740        | 1.825        | 2.912      | 4.228      | 3.426     | 2.912      | 3.271      | 1.482      | 3.036     |
| KM3     | 0.754        | 0.931        | 9.509      | 3.797      | 4.781     | 4.215      | 3.577      | 3.907      | 2.158     |
| KM4     | 0.910        | 0.824        | 0.483      | 2.121      | 1.813     | 1.539      | 0.717      | 0.689      | 0.583     |
| KM5     | 0.614        | 0.869        | 1.955      | 1.007      | 1.648     | 2.332      | 2.539      | 2.370      | 2.341     |
| KM6     | 2.226        | 0.691        | 11.589     | 4.001      | 4.389     | 3.521      | 3.041      | 3.774      | 5.375     |
| KM7     | 0.839        | 0.794        | 0.681      | -0.023     | -0.478    | -0.057     | 3.470      | 2.027      | 2.057     |
| Mean    | 1.100        | 0.944        | 4.190      | 2.176      | 2.358     | 2.304      | 2.614      | 2.253      | 2.512     |

## References

High, C.J., Hanna F.K. 1970. A Method for the Direct Measurement of Erosion on Rock Surfaces. *British Geomorphological Research Group Technical Bulletin*, 5:1-25.  
 Kirk, R.M. 1977. Rates and Forms of Erosion on Intertidal Platforms at Kaikōura Peninsula, South Island, New Zealand. *New Zealand Journal of Geology and Geophysics*, 20:571-613.  
 Langridge, R.M., Ries, W.F., Litchfield, N.J., Villamor, P., Van Dissen, R.J., Barrell, D.J.A., Rattenbury, M.S., Heron, D.W., Haubrock, S., Townsend, D.B. and Lee, J.M., 2016. The New Zealand active faults database. *New Zealand Journal of Geology and Geophysics*, 59:86-96.  
 Litchfield, N.J., Villamor, P., Dissen, R.J.V., Nicol, A., Barnes, P.M., A. Barrell, D.J., Pettinga, J.R., Langridge, R.M., Little, T.A., Mountjoy, J.J. and Ries, W.F., 2018. Surface Rupture of Multiple Crustal Faults in the 2016  $M_w$  7.8 Kaikōura, New Zealand. *Earthquake Bulletin of the Seismological Society of America*. 108(3B).  
 Stephenson, W.J., Dickson, M.E. and Denys, P.H., 2017. New insights on the relative contributions of coastal processes and tectonics to shore platform development following the Kaikōura earthquake. *Earth Surface Processes and Landforms*, 42:2214-2220.

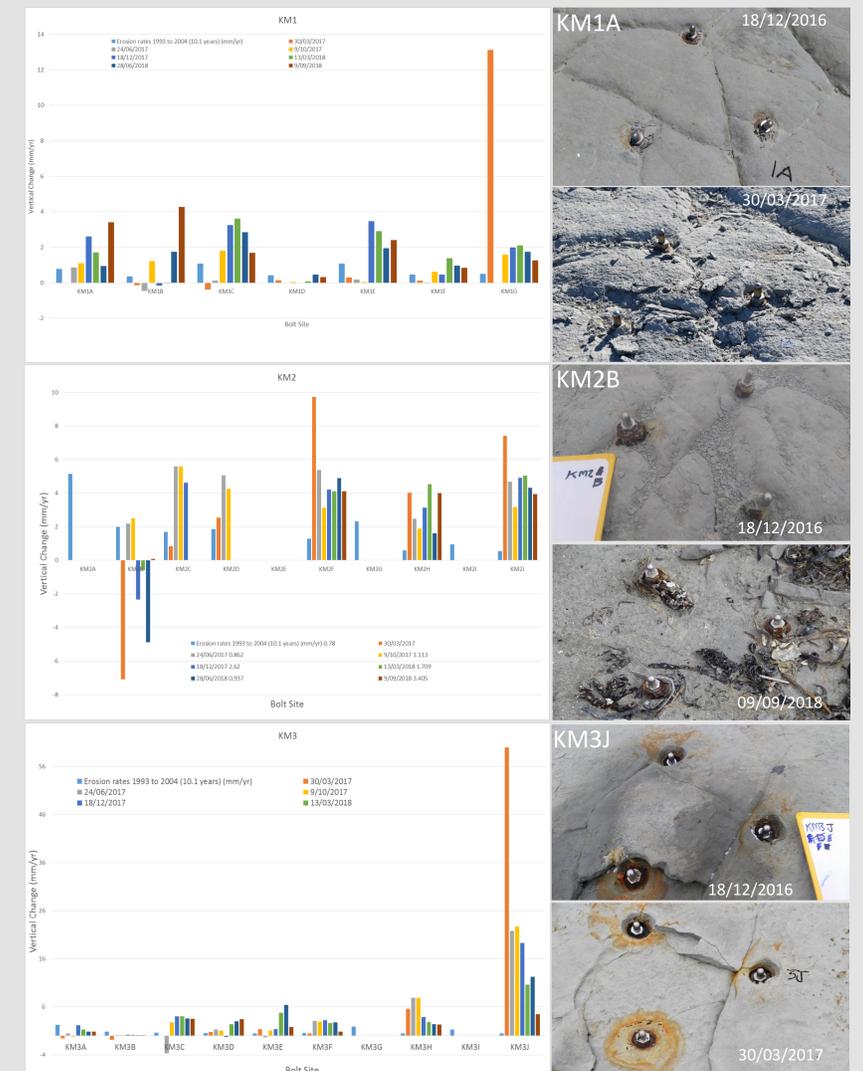


Figure 6. Erosion and swelling rates from three mudstone profiles at Kaikōura; KM1, 2 and 3. Cross shore variability in rates at each measurement period is evident. Side panels illustrate variety of surface change; rapid breakdown KM1A; erosion, swelling then burial of site KM2B; and larger scale block removal KM3J.

Table 2 Statistical tests of differences between pre and post earthquake erosion rates.

| Null  | Test  | Sig   | Decision        |
|---|---|-------|-----------------|
| The medians of Erosion Rate are the same across groups          | Independent Samples Median Test             | 0.004 | Reject the Null |
| The distributions of Erosion Rate is the same across categories | Independent Samples Mann-Whitney Test       | 0.002 | Reject the Null |
| The distribution of Erosion Rate is the same across categories  | Independent Samples Kolmogorov-Smirnov Test | 0.000 | Reject the Null |